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Frontispiece.


Sketch Map to show approximate positions of South Magnetic Pole.
A. "Discovery" Expedition.
B. Lieut. Shackleton's Expedition.
C. "Southern Cross" Expedition.

## NATIONAL ANTARCTIC EXPEDITION 1901-1904

## MAGNETIC OBSERVATIONS

PREPARED UNDER THE SUPERINTENDENCE OF THE ROYAL SOCIETY
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## PREFACE.

The volume of "Physical Observations" of the National Antarctic Expedition, 1901-1904, under Captain R. F. Scott, R.N., which was published in the summer of 1908 , included a report on one portion of the Magnetic work. This Report consisted mainly of a reduction of the absolute and relative observations by Commander L. W. P. Chetwynd, R.N. It contained, also, an account by Mr. Breracioh of the Observatory site in McMurdo Sound, of the geological features of the district, of the instruments employed, and of other matters, likewise Tables of the Hourly Values of Declination, Horizontal Force and Vertical Force on term-days at different observatories during 1902-1903, and a report by Mr. R. C. Mossman and Dr. Chree on the Magnetic Observations of the "Scotia," under Dr. W. S. Bruce.

The present volume, with its detailed Tables of Hourly Values and its exhaustive discussion of the Observations, completes the presentation of the Magnetic work of the "Discovery." The Royal Society, in arranging for this investigation, was fortunate, with the sanction and co-operation of Dr. Glazebrook, F.R.S., Director of the National Physical Laboratory, in obtaining the invaluable services of Dr. Chardes Chree, F.R.S., of the Kew Observatory. His pre-eminent qualifications for the onerous task imposed upon him, and the unwearied zeal with which he has prosecuted it to the end, give to this portion of the Reports of the Expedition a special scientific interest and importance.

Dr. Chree has mentioned in his Historical Note (p.5) the various Institutions and individuals who have contributed their services towards the preparation of the materials of this volume. To all of them the thanks of the Royal Society are due. Special allusion, however, should here be made to the assistance generously given by Mr. Bernacchi, who was in charge of the Magnetographs of the Expedition. Not only has he taken a large share in the tabulation of the magnetic curves, but he has been always ready to help during the preparation of this Report. A reference is also called for to the great zeal with which, as Dr. Chree has heartily acknowledged, the successive Directors of the Christchurch Observatory, New Zealand, Dr. Coleridge Farr and Mr. H. F. Skey, entered into the co-operative scheme of observations suggested by the Royal Society, and to the interest and importance of the information which they were eo good as to transmit. Magnetic Stations in the Southern Hemisphere are so few in number that it was particularly useful to obtain so ample a record of observations from Cbristchurch, which is the nearest observatory to the Antarctic Winter Quarters of the "Discovery."

From the map, which forms the frontispiece, it will be seen how closely the positions agree which bave been assigned by three successive recent expeditions to the South Magnetic Pole. The position found by the "Southern Cross" was about Lat. $72^{\circ} 40^{\prime}$ S., Long. $152^{\circ} 30^{\prime} \mathrm{E}$. That obtained from the observations of the "Discovery" was Lat. $72^{\circ} 51^{\prime}$ S., Long. $156^{\circ} 25^{\prime}$ E. ("Physical Observations" of National Antarctic Expedition, 1908, p. 156). Lieutenant Shackleton has been so good as to furnish the exact position found by him, which is Lat. $72^{\circ} 25^{\prime} \mathrm{S}$., Long. $155^{\circ} 16^{\prime} \mathrm{E}$.

Arch. Geikif.
Royal Society, Burlington House, 27th September, 1909.

# INTRODUCTORY NOTE. 

## BY

## L. C. BERNACCHI, F.R.G.S.

The self-recording Magnetographs supplied to the National Antarctic Expedition were of the delicate transportable type devised by the late Professor von Eschenhagen and made by the firm of O. Tofpfer, of Potsdam, Germany. They were received at the National Physical Laboratory, Surrey, England, in August, 1901, after the departure of the "Discovery."

The set consisted of three instruments, viz, a Declinometer, a Horizontal Force Magnetometer, and a Vertical-Force Magnetometer, with a self-recording photographic apparatus. No instructions were supplied by the makers, and as they were of a type absolutely strange in this country, some little difficulty was at first experienced in erecting and adjusting the instruments.

Dr. Harker and Mr. F. E. Smith, of the National Physical Laboratory, succeeded in adjusting them satisfactorily, and had them recording for a few days before my departure from England by mail boat to join the "Discovery" in New Zealand.

The Horizontal-Force instrument proved to be considerably more sensitive than the Kew form of differential magnetometers-a difference which was due to the fineness of the quartz suspension fibre employed. Unfortunately, it was not discovered until too late that the boxes of spare quartz-fibres supplied contained even finer threads than that already in the instrument.

On arriving at Melbourne, at the end of October, 1901, the instruments were taken to the Government Observatory and erected on a wood bench in a cellar. From November lst to November 10th they worked in a satisfactory manner. They were subsequently conveyed to the "Discovery" at Lyttelton, New Zealand, carefully secured in my cabin, and were not again removed until the ship reached her Winter Quarters in McMurdo Sound in February of 1902.

Winter Quarters.-The Winter Quarters were situated in latitude $77^{\circ} 50^{\prime} 50^{\prime \prime} \mathrm{S}$., longitude $166^{\circ} 44^{\prime} 45^{\prime \prime}$ E. of Greenwich, and to the south of a narrow peninsula extending in a south-west direction from the base of an island formed by Mounts Erebus and Terror. The "Discovery" remained frozen-up in her Winter Quarters from February, 1902, until February, 1904.

Observatory Site.-The spot selected for the Observatory, although the best available, was hardly an ideal one for magnetic observations. From a magnetic point of view, an observatory of this kind should be placed in a position undisturbed by the presence of magnetic rocks; but it would be difficult to find such an undisturbed locality in Victoria Land, unless it were on the surface and near the seaward edge of one of the extensive ice-floes, far from the actual coast line, such as the Great Ice Barrier.

Geological Formation.-A description of the Geological Formation in the neighbourhood of the station will be found in "Physical Observations," pp. 129-130.

Observations for Local Attraction.-A comparison of the results of absolute observations made in the Magnetic Hut with some taken on the ice in McMurdo Sound at a considerable distance from land will be found in "Physical Observations," p. 134.

The site selected for the houses was a low and fairly level piece of rocky ground close to the extremity of the peninsula, and at a distance of about 300 yards from the ship (see Frontispiece). The peninsula is about 10 miles long by a mile broad, and has an average height of 600 to 700 feet, although the extremity where the Observatories were placed was only 30 feet above the mean sea level.

Observation Houses.-The Observation Houses were constructed of large asbestos slates, screwed on to the outside and inside of a wood framework. The larger of the two used for the Variation House was 11 feet 6 inches by 11 feet 6 inches, and 6 feet 8 inches high.

Although, perhaps, small $\log$ houses would have been more suitable, they certainly would not have
beon so light, compact, and easily portable. The asbestos houses were fairly satisfactory, but had some disalviantages.

Instellution.-The instruments were erected on a strong firm wood bench (see p. 3), 8 feet 5 inches in length and 1 foot 6 inches in brearth, and suppocted at one end by a drain pipe, 1.6 feet in diameter, sunk into the frozen ground, and at the other end by a thick pillar of wood sunk in the same manner. The thickness of the wood slab forming the bench was 3 inches, and it was $2 \cdot 6$ feet above the floor of the house. The bench was carefully erected in the magnetic Meridian, the magnetometer and Declination magnet being used for finding the Meridian, and the instruments fixed upon it in the following order:-

Magnetic North (S.S.E. true) extremity the recording cylinder and clock apparatus, nearest the cylinder the declinometer, then the Horizontal-Force instrument, and at the magnetic South extremity the Vertical-Force instrument.

The suspension fibre of the declinometer was the original one employed at Kew; it was used throughout the two years and returned to England with the instrument.

On examining the Horizontal-Force fibre originally supplied with the instrument it was found to be broken; this unfortunately happened to be the stoutest fibre supplied, and it became necessary to replace it by a succession of fibres which gave the magnet an unnecessary and rather troublesome degree of sensitiveness.

The magnets for the declinometer and Horizontal-Force instrument consist of well-hardened laminar pieces of watch-spring steel, 25 mm . in length and weighing about $1^{\circ} 5$ gramme. A light aluminium frame supports the mirror and the magnet, and this is hung by means of a double hook on a smatl crosspiece attached to the bottom of the quartz-fibre suspension.

The Vertical-Force instrument is a modification of the Lloyd balance, and as it was only completed shortly before leaving England, very little was known of its behaviour even at Potsdam. This instrument was a source of constant trouble. The needle was balanced for a dip of about $70^{\circ} \mathrm{N}$, The magnetic dip at Winter Quarters being nearly $85^{\circ} \mathrm{S}$., the pull on the S . end of the needle could not be overcome by the small weights and auxiliary magnets supplied for the purpose, and therefore additional weights had to be added to the N . end, which increased the temperature coefficient of the balance.

The chief feature of the recording apparatus is that all three elements, base lines, and a temperature curve are on the same photogram for the day. On disturbed days-and they were frequent-and in the Summer when the movements of the magnets are large the result was a considerable confusion of the curves.

From May, 1902, until January, 1904, the declinometer was never interfered with, to the knowledge of the observer, nor its zero mirror altered.

During the first year the Horizontal-Force instrument was two or three times found to be out of adjustment and altered, but remained untouched after April 1, 1903, while the Vertical-Force instrument was altered from time to time during both years. The method of determining sensitiveness was by deflecting the suspended magnets with one of the unifilar collimating magnets at certain known distances, and then carefully finding the moment of the deflecting magnet by a set of absolute observations.

Routine.-The general routine was usually as follows:-The Observatory was entered at hetween 11 a.m. and noon each day, the light-shutter of the magnetograph closed, and the time of doing so noted by means of a chronometer watoh. The thermometer inserted in the Vertical-Force instrument was then read.

After changing the paper on the recording cylinder, filling and trimming the lamps, the thermometer was again read, the light-shutter dropped, and the time of doing so noted as before by means of the chronometer. The whole operation occupied about 30 minutes, and times of stopping, starting, temperatures, and orror of watch on mean time were entered in a note book.

Temperature of House-During the first year the walls only of the Variation House were banked with snow, and a large brass heating lamp was kept burning within, so as to maintain as uniform a temperature as possible. This lamp was frequently a source of danger and inconvenience of an aggravating nature, and required constant watching. Unless there was a draught underneath the lamp it emitted dense smoke which on calm days filled the room and extinguished the small oil lamp which provided the beam of


B 2
light for the sensitive paper on the recording drum. This accounts for the loss of a few days' records during the first year. Nor was the lamp successful in keeping a uniform temperature and giving out a fair proportion of heat for the amount of oil burned. During the second year the lamp was dispensed with, and the house was entirely buried under snow, and although at times the temperature within was low, vǐ: $-30^{\circ}$ F., its diurnal variation was seldom large throughout the 24 hours.

General.-During the first year the curves are finer and sharper than during the second, on account of a more sensitive bromide paper being employed, and, consequently, a smaller light-slit. The magnetograms were usually developed once a week by means of ortol-soda developer, which has the advantage of being exceptionally clean to use and gives rich dark tones to the curves.

Towards the end of the second year the supply of recording paper became very short, and from the end of September, 1903, had to be distributed equally over the subsequent months, amounting to about 4 days in each month. This is the only serious break in the two years' record. In all there are records for about 600 days.
My thanks are due to the New Zealand Government for the courtesy in placing the Christchurch Observatory at our disposal, to Dr. Coleridge Farr and Mr. H. F. Skey, of that Observatory, for their valuable assistance, to Mr. P. Baracchi, Government Astronomer of Victoria, and to Dr. Charles Chree, F.R.S., who has been so closely associated with the "Discovery" magnetic work from the beginning.

# HISTORICAL NOTE. 

## BY

DR. CHARLES CHREE, F.R.S.

The tabulation of the Magnetic Curves registered at the Winter Quarters of the National Antarctic Expedition between March, 1902, and January, 1904, was commenced in the Observatory Department of the National Physical Laboratory in December, 1904. The material consisted of photographic records of Magnetic Declination, Horizontal Force and Vertical Force for some 600 days. For the first six months the work was carried out by Mr. L. C. Bernacchi, who had been in charge of the magnetographs of the Antarctic Expedition. He went carefully through the curves, and decided the exact times of starting and stopping registration on each day by reference to the daily notes he had made in the observation books and the register of watch- and chronometer-rates. He also tabulated, with the assistance of Mr. B. Jounson, a large portion of the records of Magnetic Declination. On Mr. Bernaccur's relinquishing the work it was entrusted to Mr. H. A. Maudinng, who carried it on, with Mr. Johnson's assistance, until he left for another post in the summer of 1906 . He was succeeded by Mr. A. E. Gendie, who continued the work until his appointment, early in 1908, to Eskdalemuir Observatory. The completion of the work was then entrusted to Mr. B. Francis, Librarian in the Observatory Department, who, with the assistance of Mr. F. Levin, brought it to a conclusion.

In addition to the tabulating work, the assistants mentioned carried out a great deal of arithmetical work, in counection more especially with the tables of Diurnal Inequalities and the calculation of Fourier Coefficients. Whilst the repeated changes in the personnel tended to delay the work, and interfered, possibly, to some extent with its continuity, there was the compensating advantage that the checking of the measurements almost invariably fell to a new and unprejudiced observer. The fact that under these circumstances few serious mistakes were discovered encourages the belief that a high standard of accuracy was maintained. During 1908 a good deal of photographic work in connection with the reproduction of the Antarctic curves was done by Mr. W. J. Boxall, a senior assistant in the Observatory Department, and a good many other curves were copied with the Schmidt tracer at Bushy House, by Mr. W. H. Brookes, under the supervision of Mr. F. J. Selby.

My work has been much facilitated by the care exercised by all these gentlemen. To Mr. Bervacchi I am particularly indebted for the trouble which he has taken throughout the whole course of the work in assisting in the removal of sources of uncertainty on which he alone could throw any direct light.

Thanks are also due to Dr. Coleridge Farr and Mr. H. F. Skey, successive Directors of Christchurch Observatory, New Zealand, to Mr. C. T. F. Claxton, Director of the Royal Alfred Observatory, Mauritius, to Mr. N. A. F. Moos, Director of the Colaba Observatory, Bombay, and to Mr. E. Kitto, Superintendent of Falmouth Observatory. In response to an appeal issued by the Royal Society's Antarctic Magnetic Committee, these gentlemen put at my disposal, in the most generous way, copies or originals of the records of a number of magnetic disturbances, synchronous with disturbances recorded at the Antarctic Winter Quarters. The extent to which the work has benefited by the co-operation of these gentlemen will be appreciated only after a study of Chapters VIII and IX. Valuable as was the contribution of copies of disturbed curves from Christchurch Observatory-the nearest observatory in existence to Winter Quarters-it represents but a small part of what has been done by Dr. Farr and Mr. Skey. In his anxiety to utilise to the full the opportunities presented by the Antarctic Expedition, Dr. Farr extended the scheme of co-operation arranged with the German Antarctic Expedition and co-operating Observatories, to the extent of running the Christchurch magnetographs at high speed during the whole 24 hours of the two monthly "term" days, so as to secure a very open time scale. Mr. Skey had all these quick-run curves tabulated at 1 -minute intervals and part at 20 -second intervals, and the results of all these measurements were sent to this country, along with photographic copies of the
original curves. The measurements were received by the Hydrographic Department of the Admiraltywho had undertaken the "term" day observations along with the absolute observations-only a short time before the printing of the volume of "Physical Observations" was completed. Under these circunstances the Hydrographer and Captain Chetwynd agreed that the material had better be handed over to me, to be discussed in the present volume. The discussion will be found in Appendix $\mathbf{A}$.

In view of the absence of reference in the text to the recently published 'The Norwegian Aurora Polaris Expedition, 1902-1903,' vol. 1, of Professor Kr. Birkeland, I should explain that Professor Birkeland's volume did not appear until the present work had been practically completed. The disturbed curves had been selected and copies had been made and sent to the engraver, and all the mathematical calculations of Chapters IX and X had been carried out and the discussion written. The preparation of the present volume had already taken much longer time than was originally anticipated, and to have deferred the printing, pending an examination of Professor Birkeiand's large volume, did not appear advisable.

The text has thus been left unaltered, and any conclusions or theoretical views which it may contain are absolutely independent of any similar or conflicting results which Professor Birkeland has reached.

On studying Professor Birkeland's volume, however, I found that almost all the disturbances which he had selected for discussion were represented in the Antaretic. The opportunity for comparing Arctic and Antarctic results appeared so unique that it was decided that a special Appendix, B, should be written dealing with the subject. It is hoped that the interest attaching to the results will be deemed sufficient justification for the three or four months' delay which the preparation of the Appendix has entailed in the appearance of the present volume.

My part of the work has had to be carried out with due regard to the claims of official duties, which at times leave very little unoccupied leisure.

A great many difficulties had to be dealt with, some of them calling for very delicate discrimination. Under these circumstances the ordinary tendency of hamanity to err is pretty certain to have asserted itself, but, at all events, no pains have been spared to aim at that measure of accuracy which it is given to erring mortals to secure.

Observatory Department of the
National Physical Laboratory,
Richmond, Surrey,
July, 1909.

## EXPLANATION OF TABLES OF HOURLY VALUES.

Pages 8 to 70 contain tables giving the hourly values of Declination, Horizontal Force and Vertical Force at Winter Quarters (lat. $77^{\circ} 50^{\prime} 50^{\prime \prime} \mathrm{S}$., long. $166^{\circ} 44^{\prime} 45^{\prime \prime} \mathrm{E}$. ). The way in which these valnes were arrived at is explained later (see Chap. III., \& 20 ).

As a rule, each page contains data for a single month for one of the elements, but November and December, 1903, and Jauuary, 1904, contained so few days of registration that a single page suflicerl for these three months. The time shown in these tables is local mean time, which was 11 h . 7 m . fast on Greenwich. As is usual in magnetic tables, there appars at the top of the Declination tables at value in degrees common to all the hourly values, and at the top of the Horizontal- and Vertical-Force tables the commencing figures or figure of the values expressed in C.G.S. units. For instance, the March, 1902, Declination table is headed $\mathrm{D}=151^{\circ}+$, and the value entered under $1 \mathrm{a} . \mathrm{m}$. of the 2nd is $118^{\prime} 1$. 'This means a Declination of $151^{\circ}+118^{\circ} 1$, or $152^{\circ} 58^{\circ} 1$, at the hour stated. Owing to the large size of the daily range of Declination, it was impossible to avoid entries exceeding 60'.
The last four columns give the absolutely highest and lowest values of the element shown on the day's trace, and the time or times of their occurrence. Thus on March 2, 1902, the absolutely largest or maximum reading on the Declination trace was $151^{\circ}+179^{\prime} 6$, i.e. $153^{\circ} 59^{\prime} 6$, and its hour of occurrence was 6 h .7 m . in the morning, local time.

In the case of the Vertical Force, information as to maxima and minima is confined to days when the photographic record was complete except for the interval occupied in changing the papers.

In general the cause of absence of hourly readings is indicated in the tables, the letters $a, a^{\prime}, b, c, d$, heing employed for brevity with the following meanings:-
$a$ no trace of the element, usually from no paper being on the drum;
$a^{\prime}$ (in case of Vertical Force) no temperature trace;
$b$ (special case of $a$ ) a gap between successive sheets which conld not be filled in satisfactorily ;
$c$ trace existent, but either too faint or too confused to tabulate;
$d$ instrument recording, but not working satisfactorily.
When the magnetograph in question was working, but the trace was beyond the limits of registration, the fact is indicated by inserting the value answering to the extreme limit of registration, followed by a plus or a minus sign according to circumstances. For instance, in December, 1902, the Declination is given in the table as $\left(150^{\circ}+\right) 326^{\prime} 0+$ at $6.5 \mathrm{a} . \mathrm{m}$. on the 2 nd , and as $\left(150^{\circ}+\right) 30^{\circ} 5$ - at $8.10 \mathrm{a} . \mathrm{m}$. on the 10 th . This means that on the former accasion the Declination was beyond the limit $\left(150^{\circ}+\right) 326^{\circ} 0$ of possible registration in the direction of high Declination, while on the latter occasion it was beyond the limit $\left(150^{\circ}+\right) 30^{\circ} .5$ of possible registration in the direction of low Declination. This explanation applies both to maxima and minima and to hourly values.

In the case of Declination and Vertical Force the trace was registered right up to either edge of the sheet. In the case, however, of Horizontal Force the light from the lamp began to be eclipsed whilst the trace was still considerably short of that edge of the sheet which answered to low values of the force, and the trace became fainter and fainter, eventually vanishing at a distance from the edge which varied sensibly with the brightness of the lamp and the development of the photographic paper.

The most usual time to change papers was within an hour or two of noon, so that a day's trace appeared usually on two different sheets (sometimes on more, when there were quick runs).

In the case of the Vertical Force there was the complication that the edge of the sheet had a constantly varying value unless temperature were absolutely steady. When Declination or Horizontal Force was highly disturbed, the trace might be repeatedly off and on in the course of an hour. The most appropriate entry in these cases was only decided after investigation of the special features. That absolute consistency prevailed in the treatment is hardly likely, but uniformity was aimed at, so far as possible (see Chap. III).
DECLINATION, MARCH, 1903.

DECLINATION, APRIL, 1902.

| Day. | D $=151^{+}+$ |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Midt | Forenoon. |  |  |  |  |  |  |  |  |  |  | Noon. | Aiternoon. |  |  |  |  |  |  |  |  |  | Midt. | $\begin{gathered} \text { Maximum } \\ \text { reading and } \\ \text { time. } \end{gathered}$ |  | $\begin{aligned} & \text { Minimum } \\ & \text { reading and } \\ & \text { time. } \end{aligned}$ |  |
|  | 0. | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11. | 12. | 1. | 2. | 3. | 4. | 5. | 6. | 8. | 9. | 10. | 11. | 12. |  |  |  |  |
| 1 | : | ' | $\dot{a}$ | , |  | a | a | ' | , | ' | i | ' ${ }^{\text {a }}$ | ' | a | $151 \cdot 1$ | $122 \cdot 4$ | 117.3 |  |  |  |  | $1{ }^{1} 108$ |  |  | ${ }_{165}{ }^{\prime} 0$ |  |  |  |
| 2 |  | $135 \cdot 3$ |  | ${ }^{7} 135 \cdot 1$ | $135 \cdot 4$ | $140{ }^{\circ} 6$ | $135 \cdot 3$ | 31 | $182 \cdot 5$ | $5168 \cdot 8$ | $170 \cdot 0$ | 1517 | 0 |  | 117.0 | $115 \cdot 8$ |  | 1178 | $120 \cdot 0118{ }^{-0}$ | 5 | 129-3 | $130 \cdot 5$ | 128-5 | $125 \cdot 0$ | 184.5 | 8.42 \&.m. | 108.0 | 4.4 |
| 3 | 128 | $134 \%$ | 133.1 | $1132 \cdot 0$ | 1340 | 138.5 | $143 \cdot 6$ | 148.5 | . 148.5 | $5142 \cdot 5$ | 136.5 | 128.8 | 141.0 | $117 \cdot 5$ | 10 | 6 | 0 | ${ }^{89} 5$ | $104 \cdot 8$ | $\cdot 0$ | $105 \cdot 6$ | 122.0 | 1170 | $120 \cdot 5$ | 15 | 9.28 " | ${ }_{1} 1$ | 3.6 |
| 4 | 120.5 | 1370 | $139 \cdot 1$ | $1138{ }^{\circ}$ | 134.6 | $145 \cdot 1$ | 1798 | , | $1766^{\circ}$ | $161{ }^{\circ}$ | $133 \cdot 1$ | 144.5 | 1372 | 128.7 | 125-3 | 131.0 | 124 | $113 \cdot 1$ | $117 \cdot 8113 \cdot 3$ | $\cdot 6$ | 125.0 | - 6 | 1 | 1337 | $206 \cdot 4$ | 8.5 , | $101 \cdot 6$ | 5.44 |
| 5 | 133.7 | 1296 | $139 \cdot 1$ | 12490 | $142 \cdot 5$ | 1486 | 1538 | 1454 | $180 \cdot 5$ | $5165 \cdot 8$ | 140-3 | 1358 | 122 3 | $125^{\circ}$ | $118 \cdot 5$ | $115 \cdot 5$ | 1 | 1 | $122.0138 \cdot 5$ | 5 | 1 | 139.0 | $135 \cdot 0$ | 132.5 | $219 \cdot 0$ | 8.28 | $100 \cdot 5$ | 5.34 |
| ${ }^{6}$ | $132 \cdot 5$ | $130 \cdot 5$ | 6 | $6^{1} 142 \cdot 1$ | $140 \cdot 0$ | $148 \cdot 6$ | $150 \cdot 0$ | $150 \cdot 5$ | $149 \cdot 6$ | 148.5 | 151.5 | 1275 | 123.0 | $121 \cdot 1$ | 117 ¢ ${ }^{\text {b }}$ | 122.0 | 12 | 0 | 128.0118 .5 | $\cdot 0$ | 12 | $124 \cdot 1$ | 12 | 13 | 175 | 9.41 , | 114.0 | 4.18 |
| 7 | $130 \cdot 1$ | 5 | $134 \cdot 6$ | 6 148.5 | $147 \cdot 5$ | $145 \cdot 1$ | $147 \cdot 5$ | $140 \%$ | $145 \cdot 5$ | 5149.6 | 148.5 | $133 \cdot 1$ | 1280 | $127 \cdot 5$ | 124-1 | $120 \cdot 0$ | 118 | $121 \cdot 1$ | 123-9126 0 | $132 \cdot 5$ | 13 | 1 | 132 | 135.5 | $10 \% 0$ | 9.50 | $100 \cdot 5$ | 4.34 |
| 8 | $135 \cdot 5$ | 133.5 | 133.5 | 5. $145 \cdot 1$ | 143.0 | $150 \cdot 5$ | $151 \times 1$ | $142 \cdot 5$ | 42.5 | 150.5 | 149-6 | $159 \cdot 5$ | 134.0 | -9 | $125 \cdot 9$ | 122.0 | $120 \cdot 5$ | $128 \cdot 6$ | 128.5 $123 \cdot 5$ | $2 \cdot 0$ | 10 | 1190 | 123 | 128.0 | 5 | 10.45 | 90.0 | 9.8 |
| 9 | $128 \cdot 0$ | 129.0 | 138.5 | 51470 | $142 \cdot 5$ | $143 \cdot 6$ | $151 \cdot 1$ | 148.5 | 156.5 | $158 \cdot 6$ | 15 | 16 | 11816 | 148.8 | 141.0 | , $125 \cdot 6$ | 12 | 5 | $115 \cdot 1{ }^{-1} 88{ }^{\circ} 0$ | 6 | 125.0 | 1320 | $133 \cdot 5$ | 138.5 | 201.0 | 11.6 | 63.5 | \%. 32 |
| 10 | . 5 | 145.5 |  | $5142 \cdot 5$ | $142 \cdot 1$ | 139.5 | $148 \cdot 3$ | 183.1 | 16 | 1770 | 18 | 177.5 | 1 | 137.7 | 133.5 | ${ }^{133} 1$ | 13 | $133 \cdot 1$ | $138.5135 \cdot 5$ | 138.0 | 135 | 36.51 | 139 | $131 \%$ | 196 | 10.25 | 124.1 | 11.57 |
| 11 | $131 \cdot 6$ | 158 | 1275 | -5 | 156.0 | 153.5 | 13 | 138.0 | $132 \cdot$ | - $132 \cdot 8$ |  | c | c | c | c | c |  | 10 |  | $3 \cdot 0$ |  | - 32.0 | 74.6 | 139.1 | 175.5 | 1.15 , | $0 \cdot 0-$ | i-8 ., |
| 12 | $138 \cdot 1$ | $176 \cdot 3$ | $155 \cdot 3$ | 3. 141.5 | 182.6 | ${ }^{206} \cdot 6$ | 1650 | 18 | $154 \cdot 5$ | 174.0 | $173{ }^{6}$ | 1580 | 174.5 | of | , | 5 | $139 \cdot 5$ | 10 | 109.1140 0 | 5 | , | $1180 \cdot 5$ | 156 | $160^{\circ} 1$ | 2258 | 8.52 | 90. | 5.43 |
| 13 | $186 \cdot 1$ | 183 | 1 | $1184 \cdot 0$ | 161.0 | 184.0 | 163.1 | 183.5 | $163 \cdot 1$ | 162.0 | $188 \cdot 5$ | $181 \cdot 1$ | $180 \cdot 1$ | $165{ }^{\circ}$ | 157.5 | $154 \cdot 1$ | $153 \cdot 3$ | , $145 \cdot 5$ | 150.1470 | $146 \cdot 6$ |  | 141.5 | 1610 | 153.0 | 189.0 | 11.35 | 1260 | 10.20 |
| 14 | 153.0 | 148 | $163 \cdot 1$ | $1159 \cdot 5$ | $184 \cdot 8$ | $168 \cdot 5$ | 1760 | $172 \cdot 1$ | 162.0 | $202 \cdot 1$ | $176{ }^{\circ}$ | 176.0 | 176.0 | 1610 | 156.0 | $148 \cdot 1$ |  | $134 \cdot 0$ | $144.0137^{\circ}$ | \% | $156^{\circ} 0$ | $156 \cdot 0$ | $155^{\circ}$ | 162.0 | 225.0 | 9.15 | 1160 | 6.37 |
| 15 | 162.0 | ${ }^{1}$ | 162.0 | 162.0 | 18 | $187 \cdot 6$ | $172 \cdot 1$ | $172 \cdot 1$ | 173 | $175 \cdot 3$ | $171 \cdot 5$ | 173 | $168 \cdot 2$ | $161 \cdot 0$ | 152.0 | $141{ }^{\text {d }}$ |  |  | 138 | + |  | $136 \%$ | $140 \cdot 3$ | 1418 |  | 8.47 | 129.5 | 6. 23 |
| 16 | 141.8 | 144 |  |  |  |  |  | $130 \cdot 1$ | $148 \cdot 8$ | 155 | 157.6 | 156 | $151 \%$ | 1511 | b | $143 \cdot 5$ |  | - | $140 \cdot 5141.5$ |  |  |  |  |  | -8 | 10.55 | 137 | 6.12 |
| 17 | a | ${ }^{\text {a }}$ |  |  |  |  |  |  |  |  |  |  |  | - | 181.6 | 148.6 |  | $132 \cdot 8$ | $134 \cdot 0137 \cdot 5$ | $135 \cdot 1$ | $135 \cdot 5$ | 146 .0\| | $145 \%$ | 145.0 | 2015 | 1.48 p.m. | $12 \cdot$ | 4.46 |
| 18 | 1450 | 148.0 | $147 \cdot 1$ | $1150 \cdot 1$ | $150 \cdot 1$ | 1510 | $151 \cdot 6$ | $154 \cdot 8$ | $147 \cdot 1$ | $180 \cdot 8$ | 149.0 | $150 \cdot 5$ | 159.5 | $167^{\circ}$ | 148.6 | $143{ }^{\circ}$ | $134 \cdot 0$ | ${ }^{135} 5$ | $140 \cdot 5$ | 143.0 | 142 | $135 \cdot 1$ | 140 | $139 \cdot 6$ | $177 \cdot 8$ | 0.45 . | 1235 | 5.40 |
| 19 | $139 \cdot 6$ | $148^{\circ}$ | $146^{\circ} 0$ | $0150-1$ | $151 \cdot 6$ | ${ }^{151} 10$ | $180 \cdot 8$ | 15 | $159 \cdot 5$ | $5182 \cdot 5$ | 185.5 | $159 \cdot 1$ | $149{ }^{\circ}$ | 1430 | $142 \cdot 0$ | 140 | 140.0 | $137 \cdot 3$ | $1399^{\circ 818300}$ | $132 \cdot 1$ | $125 \cdot 5$ | 132 | 1415 |  |  | 9.59 a | 113. | 9.9 |
| 20 | - | $147 \cdot 5$ | 14.5 | $5150 \cdot 1$ | 153.1 | 1520 | 16 | 154. | $158 \cdot 5$ | $180 \cdot 8$ | 1 | 180.6 | 158.5 | 1 | 148.6 | $141 \cdot 1$ |  | 1415 | $125 \times 127$ •0 | 128.5 | 118.0 |  | 94.6 |  | $186^{\prime} 8$ | 8.36 | 68.5 | 10.35 |
| 21 | 109.0 | 178.0 | $138 \cdot 1$ | 11570 | $171 \cdot 1$ | 187.0 | 1830 | 188 | 1810 | $184 \cdot 6$ | 179.5 | 184.0 | 1715 | 148.0 | 143.0 | 132 |  | 8.0 | 108.5 150.5 | 140 | 143.0 |  | $150 \cdot 1$ |  | 221.0 | 5.52 | *.0- | S.u |
| 22 | 1875 | 1480 |  | 5. $144 \cdot 1$ | 183.0 | $198 \cdot 1$ | $167 \cdot 5$ | $172 \cdot 8$ | $183 \cdot 1$ | $188 \cdot 1$ | $188^{\circ}$ | $154 \cdot 6$ | 1428 |  | $139 \cdot 6$ | $189{ }^{\circ} 8$ | 13 | 138.5 | 128 | 123•1 | 1340 |  | 44-5 |  | 225 5 | 7.48 | ${ }^{10} \cdot$ | 8,0 |
| 23 | 152.5 | 1450 |  |  | 15 | 184.0 | $168 \cdot 1$ | $155 \cdot$ | $8 \cdot 8$ | 185.1 | ${ }^{183}{ }^{6}$ | 16 | 15 | 153.1 | $140 \cdot 5$ | 146.0 | 147 | $144 \cdot 5$ | ${ }^{14}$ | 143.0 | 144 | 142 | 151 *if | 145 | 191.0 | 10.28 | 120 | 1.; 8.m. |
| 24 | 1 | $147 \cdot 1$ |  | 15 | 148.5 | $150 \cdot 1$ | 152.5 | 133.1 | 183.6 | 15 | 171.5 | 1675 | $155 \cdot 1$ | $\cdot$ | 1430 | $145 \times 6$ | $142 \cdot 6$ | $11^{1}$ | 144 | $142 \cdot 0$ | 1 | 144.5 | 151 v | $150 \cdot 3$ | 1970 | 8.32 | 19 | \%.15 |
| 25 | $150 \cdot 5$ | 150-5 |  |  | 158.5 |  |  | 1550 |  | $159 \cdot 5$ | $158{ }^{\circ}$ | 154 | 1471 | 149-5 | 145. | 143 | $141 \cdot 1$ | 00 | 14 | $141{ }^{1 / 1}$ | 141 . | 142 | $1+4 \cdot 1$ | 14505 | $165 \%$ | 3 | $1: 3$ | 2.as p.m. |
| 26 | $146 \cdot 3$ | 1485 |  |  | 148.5 | \% 0 | , 5 | $150 \cdot 1$ | $150 \cdot 1$ | 154.0 | 1546 | $155 \cdot 0$ | $152 \cdot 3$ | $140^{\circ} 2$ | ${ }^{139 \cdot 1}$ | 138.1 | $133 \cdot 1$ | $134 \cdot 9$ | $913 \cdot 3$ | 131 | 135 | 136 | 138 | 188 | 81 | ${ }^{11.35}$ | 123:5 | $s t$ |
| 27 | $138 \cdot 9$ | $138 \cdot 1$ |  | 240-5 | $13900$ | $136.4$ | b | $138 \cdot 7$ | .2 | $142 \cdot 7$ | $144 \cdot 5$ | 143 | 24-5 | 1453 | 1445 | 1456 | 140 | 48.2 | $145 \cdot 8145 \cdot 8$ | 1474 | 148.6 | $148 \cdot 6$ | 148 | 150 | $150 \cdot 3$ | 11.49 Prm | 1: | nit am. |
| 28 | $150 \cdot 7$ | $150 \cdot 5$ | $148^{\circ} 6$ | $\text { 6] } 150 \cdot 5^{\prime}$ | 151.0 | $151 \cdot 5$ |  | 153.] | $154 \cdot 3$ | 1581 | 158.0 |  | 148.0 | 14 | $43 \cdot 5$ | $142 \cdot 1$ | 14.21 | 33. | $143 \cdot 8150 \cdot 7$ | 151 | $153 \times 2$ | $152 \cdot$ | $153 \cdot \mathrm{~s}$ | 2 | 161 | 9.53 m | $111 \cdot 1$ | 5.9 |
| 29 | 15 | 152 |  |  |  | a |  | , | a | a | $\stackrel{ }{8}$ | a | 155.2 | 150.4 | 153.2 | 150.4 | 14901 | $150 \cdot 8$ | 159.0152.2 | 153 | 132 | 152 :5 | $123 \cdot 1$ | 5. | 161 | $1.39 \mathrm{pm} . \mathrm{m}$. | $146 \cdot 1$ | 34 |
| 30 | 155.2 | 153.5 | 153.7 | 7154.3 | $155 \cdot 3$ | 155:5 | a | ) a | . | a | a | $\cdots$ | a | a | , | a | a | - | 134 $3128 \% 3$ | 148 | 8 | 143.8 | 149 | 142 | $161^{\circ}$ | 4.s.anm. | 1250 | -1. |

DECLINATION, JULY, 1902.

DECLINATION, AUGUST, 1902.

| Day. | D $=150{ }^{\circ}+$ |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Midt. | Forenoon. |  |  |  |  |  |  |  |  |  |  | $\frac{\text { Noon. }}{12 .}$ | Afternoon. |  |  |  |  |  |  |  |  |  |  | $\qquad$ | $\begin{aligned} & \text { Maximum } \\ & \text { reading and } \\ & \text { time. } \end{aligned}$ |  |  | $\begin{aligned} & \text { Minimum } \\ & \text { reading and } \\ & \text { time. } \end{aligned}$ |  |
|  | 0. | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11. |  | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11. |  |  |  |  |  |  |
| 1 | 148.8 | 148.1 | 147.2 | $147 \cdot 1$ | $147 \cdot 8$ | 148.7 | $149 \cdot 8$ | 151.3 | 150.5 | 5150.3 | 158.3 | $153 \cdot 5$ |  | $148{ }^{\circ}$ | 149 "8 | $150 \cdot 7$ | 149.5 | $148 \cdot 6$ |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  | $158 \cdot$ | 153 | 149.0 |  |  | $150 \%$ |  | $148 \cdot 6$ | 147 -1 | 7.1 |  | $140 \cdot 8$ | 148.0 | 152.0 | $155 \cdot 9$ | 158 | 10.46 | p.m. | $123 \cdot 8$ | 7.39 p.m. |
| 2 | 155 '8, | 151.1 | $154 \cdot 0$ | $153 \cdot 5$ | $152 \cdot 3$ | $152 * 6$ | 152.3 | $153 \times 8$ | 154 * | $154 \cdot 1$ | 153.7 | 153 * | 148.6 | $148{ }^{\circ} 0$ | $145 \cdot 6$ | $145 \cdot 7$ | 1460 | 1468 | $144 \cdot 1$ | 148.3 | 45.6 | $147 \cdot 5$ | $149{ }^{\circ}$ | $146 \cdot 2$ | 15 | $157 \cdot 8$ | 10.3 |  | $137 \cdot 6$ | 11.25 |
| 3 | 15 | 154.8 | 148.9 | 15 | $151 \cdot 7$ | 158.8 | $\bigcirc$ | $157 \cdot 0$ | - | 156.8 | 182 | $157 \cdot 6$ | 156.5 | 147.8 | 3 | $149 \cdot 3$ | 125.0 | $126^{1}$ | 140.8 | 142.3 | 148.1 | 150 '8 | $3 \cdot 1$ | $153 \cdot 1$ | 152.0 | 167:3 | 10. | a.m. | 105.5 | 4.35 |
| 4 | 1520 | $151 \cdot 3$ | 151 \% | 1 | $\cdot 5$ | . 6 | . 8 | 153.1 | 160 '6 | 183.6 | $154{ }^{\circ} 0$ | $160 \% 6$ | 1475 | $\cdot 8$ | $\cdot 6$ | -8 | 137 '5 | 131.0 | 3.0 | 1276 | 108.5 | $6 \cdot 6$ | '9 | $152 * 9$ | $150 \cdot 1$ | 171 | 9.43 |  | 98.5 | 7.42 |
| 5 | $150 \cdot 1$ | 152.0 | 151.0 | 1540 | $152 \cdot 0$ | 154.6 | $\cdot 5$ | $158 \cdot 5$ | $159 \cdot 1$ | $158 \cdot 4$ | 157.6 | $158 \cdot 2$ | $148 \cdot 3$ | $149 \cdot 8$ | .0 | $146 \cdot 9$ | $144^{\circ} 4$ | $142 \cdot 9$ | $142 \cdot 6$ | 143.0 | $146 \cdot 9$ | $5 \cdot 7$ | 146.5 | 149.5 | 140 | $167 \cdot 2$ | 7.26 | , | 134.0 | 7.23 |
| 6 | - | 153.1 | $152 \cdot 3$ | 1471 | 151.0 | $155 \cdot 5$ | -8 | 153.5 | $154 \cdot 1$ | $153 \cdot 5$ | 153.1 | $152^{\circ} 0$ | $150 \cdot 5$ | $150 \cdot 5$ | 13 | $150 \cdot 1$ | $149{ }^{\circ} 3$ | $148{ }^{\circ}$ | $150 \cdot 8$ | 151.4 | $150 \cdot 1$ | $151 \cdot 6$ | 152.0 | $152 \cdot 5$ | 156.5 | $163 \cdot 3$ | 7.37 | " | 1415 | 2.55 am . |
| 7 | 156 | $153 \cdot 8$ | $153 \cdot 5$ | $154 \cdot 3$ | 154 '6 | 156.8 | $158 \cdot 1$ | $155 \cdot 5$ | 1570 | 158.8 | 159.5 | 158.9 | $156 \cdot 3$ | 156.5 | $153 \cdot 5$ | 151 '6 | 148.6 | $150 \cdot 2$ |  | - | a | a |  | a |  | 162.8 | 8.25 | " | 146.8 | 4.10 p.m. |
| 8 |  |  | a | a | a |  | a | a | a | a | a | a | $162 \cdot 8$ | $182 \cdot 1$ | $155^{\prime 9}$ | 153.1 | $155 \cdot 2$ | 153.4 | 5 | 157 \% | $156 \cdot 1$ | 157.0 | $6 \cdot 5$ | 157 \% | $161 \cdot \frac{1}{2}$ | 1700 | 11.55 | , | $150 \cdot 5$ | 2.58 |
| 9 | $161 \cdot 2$ | 162.4 | 162 -8 | 161.9 | $162 \cdot 5$ | 162 -8 | $\cdot 1$ | 187.0 | $169 \cdot 3$ | $166 \cdot 3$ | 168.3 | 162 '5 | 162 | $159 \%$ | $155 \cdot 8$ | $152 \cdot 9$ | $150 \cdot 1$ |  | $153 \cdot 5$ | $152 \cdot 8$ | 149 | $140 \cdot 0$ | $138 \cdot 3$ | 138.5 | 148 | 174.5 | 8.16 | , | 123.5 | 10.10 |
| 10 | -0 | $158 \cdot 3$ | 159-8 | 15 | $1.56 \cdot 1$ | 157.7 | 5 | $178{ }^{\circ}$ | $157{ }^{\circ}$ | $158 \cdot 8$ | $161{ }^{\circ} 0$ | 168.5 | 1610 | $157 \cdot 6$ | 7 | 158.8 | 157 -1 | $155 \cdot 5$ | $154 \%$ | 153.7 | 156.2 | $156 * 4$ | 156 | $157 \cdot 1$ | 15 | $187 \cdot 3$ | 6.50 | " | 144.5 | 8.35 mm |
| 11 | 156 | 156.5 | $158 \cdot 5$ | 159 '8 | 159.2 | $169 \cdot 0$ | $169 \cdot 6$ | 170.0 | $196{ }^{\circ} 6$ | $198 \cdot 1$ | $188{ }^{\circ} 0$ | 208 | $186 \cdot 1$ | 164.8 | $9 \cdot 6$ | $138 \cdot 1$ | 126.5 | 13 |  | $140 \cdot 5$ | $133 \cdot 6$ | $4 \cdot 5$ | 156 | 158.0 | 153.8 | $16 \cdot 3$ | 11.25 |  | 104 | 3.56 p.m. |
| 13 | $153.8{ }^{\prime}$ | 162.5 | $160 \% 6$ | 175.0 | $166 \cdot 7$ | 163.4 | 184.5 | 1 |  | $168 \cdot 1$ | 169.0 | 1 | $165 \cdot 5$ | $170 \cdot 8$ | $2 \cdot 1$ | $160^{\circ} 0$ | $159 \cdot 5$ | 158.0 | $\cdot 0$ | 9.7 | 159.1 | 161.0 | $159 \cdot 5$ | $160 \cdot 6$ |  | 195-2 | 3.18 |  | 149 | 0.2 a.m. |
| 13 | 159.5 | 6 | 1640 | $166^{\circ} 0$ | $188 \cdot 3$ | -5 | 188.7 | $185 \cdot 5$ | $166 \cdot 3$ | 16 | $168 \cdot 6$ | 165 | (165-9) | 186.0 | $1 \cdot 3$ | $161 \cdot 3$ | 159 -8 | ${ }^{1} 1$ | 154.7 | $155 \cdot 5$ | $159 \cdot 1$ | $155 \cdot 6$ | $59 \cdot 1$ | $162 \cdot 5$ | $163 \cdot$ | 171.5 | 7.6 | , | 147.5 | $5.42 \mathrm{pm.m}$. |
| 14 | $164 \%$ | $167 \cdot 9$ | $166 \cdot 9$ | 188.2 | $168 \cdot 2$ | 169.7 | $170 \cdot 5$ | 16 | 169.0 | 16 | 169.0 | 16 | $164 * 0$ | $158 \%$ | $8 \cdot 5$ | 159 •5 | $160 \cdot 3$ | 159 | $157 \%$ | 156.4 | $147 \cdot 1$ | $160 \cdot 6$ | 159.5 | 162.5 | $163 \cdot 3$ | 173.8 | 5.35 |  | $135 \cdot 3$ | 8.3 |
| 15 | 163.3 | 166 | $186 \cdot 2$ | $165 \cdot 8$ | $186 \cdot 0$ | $167 \cdot 3$ | 167.0 | 167.0 | 1 | 171.5 | 172.0 | 10 | 164.9 | 163.9 | -8 | $159 \cdot 1$ | $160 \cdot 8$ | 158.0 | 155.0 | 156.2 | 158.2 | $162 \cdot 2$ | $159 \cdot 7$ | $161 \cdot 9$ | 163 | 1760 | 10.28 |  | $149 \cdot 6$ | 8.58 |
| 13 | 16 | $168 \cdot 3$ | $164 \cdot 6$ | $165 \cdot 2$ | $166 \%$ | -6 | . 2 | $170 \cdot 0$ |  | 170-5 | 171.5 | $170 \cdot 2$ | 169.0 | $164 \cdot 0$ | $166^{\circ} 0$ | $165{ }^{\prime}$ | $164 \cdot 5$ | $161 \cdot 3$ | $160 \cdot 6$ | 157 '6 | 154 | 143.3 | 153.2 | 159.1 | 10 | 188 | 7.5 | p.m | 133.4 | 7.11 |
| 17 | 163.0 | 157 • | $185 \cdot 4$ | 169.7 | $170 \cdot 0$ | $171{ }^{4}$ | $171 \cdot 1$ | 171.1 |  | 171.5 | 171.5 | 174.5 | $167 \cdot 9$ | 164 '5 | $163 \cdot 4$ | $165 \cdot 1$ | $163 \cdot 4$ | $162 \cdot 5$ | $165 \cdot 2$ | 16\%.7 | 185.5 | . 5 | $104 * 0$ | $184{ }^{\circ} 0$ | 16 | 178.0 | 2.6 | a. | 154.3 | 0.6 a.mb. |
| 18 | $1{ }^{\circ} 0$ | $165 \cdot 5$ | $183 \cdot 4$ | $171 \times 8$ | 173.0 | '6 | . 5 | $183 \cdot 5$ |  | 178.7 | 189.6 | 167 '5 | $171 \cdot 1$ | $159 \cdot 5$ | - | 1378 | 142 | 141.5 | $143 \cdot 0$ | 121.0 | $120 \cdot 8$ | 34.5 | $143 \cdot 2$ | 158 | $165 \cdot 1$ | $189 \%$ | 7.48 | , | 108 | $7.5 \mathrm{p} . \mathrm{m}$ |
| 19 | $165 \cdot 1$ | $173 \cdot 3$ | $174 \cdot 3$ | 175 ¢ | $173 \cdot 5$ | 171.5 | . | $173 \cdot 3$ | 173.0 | $174 \cdot 1$ | $173 \cdot 6$ | - 1 | $171 \cdot 5$ | 163.3 | 159.4 | $159 \cdot 2$ | $158 \cdot 6$ | 159.4 | $160 \cdot 9$ | $161 \cdot 3$ | 160.3 | $160 \cdot 9$ | $161 \%$ | 161.5 | 10 | 186:5 | 2.25 |  | 157 | 3,20 |
| 20 | 162.2 | $160 \%$ | $160 \cdot 9$ | $161^{\circ} 0$ | $162 \cdot 1$ | $163 \cdot 6$ | 31 | $163 \cdot 7$ | $186 \cdot 9$ | $166^{\circ} 0$ | $169 \cdot 7$ | 162 \% | $162 \cdot 1$ | 159.2 | $155 \cdot 3$ | $155 \cdot 8$ | $157 \%$ | $157 \%$ | $158{ }^{\circ}$ | $155 \cdot 3$ | $150 \cdot 8$ | - | 159 | $160 \cdot 1$ | 161 | 1768 | 9.38 | " | 1530 | 1.28 |
| 21 | 3 | 141.9 | 163.0 | $181 \cdot 3$ | $162 \cdot 2$ | $2 \cdot 5$ | $1 \cdot 8$ | 184.0 | 171.2 | 188.8 | 171.5 | 172.4 | 159.5 | 169 |  | 153.5 | $149 \cdot 5$ | $147 \cdot 5$ | 147.5 | $140 \cdot 5$ | $120 \cdot 1$ | 4 | $5 \cdot 6$ | $125 \cdot 5$ | 155 | 183-1 | 8.30 |  | 113'3 | \%.35 |
| 22 | 155.0 | 5 | $179 \cdot 0$ | 17 | $188 \cdot 1$ | 205 -6 | $249 \cdot 5$ | 253.0 | 248.0 | $270 \cdot 5+$ | $258{ }^{\circ} 5$ | $218{ }^{\circ} 0$ | 165.5 | 143.5 |  | . 0 | $115 \cdot 8$ | 103.0 | $4 \cdot 0$ | $140 \cdot 5$ | $93 \cdot 1$ | -0 | 163.6 | *\% | $167 \cdot 5$ | $270 \cdot 5+$ |  |  | 59.9 | 3.15 |
| 23 | 10 | $171 \cdot 1$ | 164.5 | 1 | $160^{\circ} 0$ | "3 | . 3 | 167.0 |  | 178. | $170 \cdot 5$ | $173 \cdot 0$ | 170 | $182 \cdot 0$ | $173 \cdot 5$ | 181.0 | $150^{1} 1$ | $152 \cdot 5$ | $149 \cdot 0$ | 119.5 | $155{ }^{\circ} 0$ | - 0 | $161{ }^{\circ} 0$ | 160 | 1 | 1463 | 2.5 | p.in | 9 | 2.5 |
| 24 | $15 \% \%$ | $162 \cdot 1$ | $159 \cdot 5$ | 10 | 17 | 173.0 | . 0 | $168 \cdot 1$ |  | 197 | 193.0 | 188. | $172 \cdot 6$ | 167.0 | 149.0 | 123.5 | $137 \cdot 5$ | $129 \cdot 1$ | $126 \cdot 1$ | $135 \cdot 1$ | $150 \cdot 1$ | 158.5 | 158.0 | 162 | 155 | 113*8 | 5ns | a.m | 115 | cild |
| 25 | $155 \cdot 5$ | 161.0 | $155 \cdot 5$ | $154 *$ | 158.5 | $168 \cdot 1$ | -6 | $3 \cdot 5$ | $180 \cdot 5$ | 185.0 | 18 | $172 \cdot 6$ | 162.5 | $156 \cdot 1$ | $162 \cdot 5$ | $133 \cdot 1$ | 1487 | 152 | $137^{\circ} 0$ | 151.0 | 1340 | $150 \cdot 1$ | 158.1 | 185 | 1 | $24^{*}$ | 8.54) | . | 124.4 | ${ }^{1,1}$ |
| 28 | $157 \cdot 6$ | - | 163.6 | $170 \cdot 8$ |  | $169 \cdot 6$ | 171.5 | 190.0 | - | $104 \cdot 3$ | 5 | 162-5 | (163.3) | 184.0 | $147 \cdot 5$ | $142 \cdot 6$ | $151 \cdot 3$ | 152.5 | $154 \cdot 8$ | $180 \%$ | 155.5 | $155 \cdot 5$ | $159 \cdot 5$ | $156 \cdot 5$ | $159 \cdot 1$ | $217 \cdot 3$ | 8.10 | .. | $1: 3 \cdot 3$ | 1.50) |
| 27 | 15 | 159.5 | 16 | $160 \cdot 3$ | $165 \cdot{ }_{1}$ | $100{ }^{\circ} 0$ | . 3.6 | 0.6 | $186 \cdot 5$ | $161 \cdot 0$ | 162-1 | $180 \cdot 6$ | $150 \cdot 5$ | 153.5 | ${ }^{-0}$ | $148 \cdot 8$ | $145 \cdot 8$ | 154.5 | 153.1 | $140 \cdot 0$ | $141 \cdot 1$ | $143 \cdot 0$ | 80 | 180 | 160 | 143:5 | \%.55 |  | 131.0 | 7.24 |
| 28 | $160 \cdot 7$ | 161.3 | 15 | 157 \% |  |  | 2.0 |  |  | $165 \cdot 1$ | 104.5 | $163 \cdot 3$ | $162 \cdot 8$ | 153.5 | ' $157 \cdot 6$ | $5^{1}$ | 155-5 | 155 | $157 \cdot 0$ | $150 \cdot 5$ | $156 \cdot 1$ | 180.0 | $180 \%$ | 162 | $159 \cdot 1$ | 1940 | \%.50 |  | 1:3\% | 8.15 |
| 29 | 159.1 | 15 | 184.5 | $171 \cdot 1$ | 165.5 | 164.5 | 188.0 | $170 \cdot 5$ | 170 | $160^{\circ} 0$ | $168 \cdot 7$ | $164 \cdot 8$ | $162 \cdot 5$ | 155.0 | 154.0 | 153.5 | $154 \cdot 6$ | 155.5 | $153 \cdot 8$ | 154.6 | $157 \cdot 6$ | $156 \cdot 1$ | 157 | 1 mo | $13:$ | 233.3 | 9.2) |  | 140 | (1.2: |
| 30 | 157.3 | $184 \cdot 5$ | $181 \cdot 0$ | $184 \cdot 5$ | $165 \cdot 1$ | $165 \cdot 1$ | $165 \cdot 1$ | $130 \cdot 0$ | 174.1 | $169 \cdot 8$ | 1843 | $184 \cdot 5$ | 164.0 | $152 \cdot 5$ | 151.0 | 149.0 | $150 \cdot 1$ | 153.0 | $151 \cdot 6$ | $133 \cdot 5$ | $159 \cdot 1$ | $158 \cdot 3$ | $159 \cdot 5$ | 15N | 1.8 | $1: 93$ | S. 43 | . | 12tis | 8.40 |
| 31 | 158.0 | 1 161) 0 | 158.5 | $100^{\circ} 0$ | $180 \cdot{ }^{\prime \prime}$ | $162 \cdot 1$ | $162 \cdot 5$ | 184.0 | 163.0 | $182 \cdot 1$ | 154.0 | $160^{\circ} 3$ | $155{ }^{\circ} 0$ | 153.1 | 152.0 | 152:3 | $150 \cdot 1$ | 148 | $151 \%$ | $14 \mathrm{R}^{\circ} \mathrm{A}$ | 1:4\% | 98 | 129-1 | $134 \cdot 5$ | 149 3 | 1:1\% | I0.4 |  | 4 | 9.0 |

DECLINATION, SEPTEMBER, 1902.

DECLINATION, OCTOBER, 1902.

DECLINATION, NOVEMBER, 1902.

| Day, | Midt. Forenoon. |  |  |  |  |  |  |  |  |  |  |  | $\frac{\text { Noon }}{12 .}$ | Afternoon. |  |  |  |  |  |  |  |  |  |  |  | $\begin{aligned} & \text { Maximum } \\ & \text { reading and } \\ & \text { time. } \end{aligned}$ |  | $\begin{gathered} \text { Minimum } \\ \text { readizug and } \end{gathered}$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 0. | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11. |  | 1. | 2. |  |  | 5. |  |  | 8. | 9. | 10. | - | 12. |  |  |  |  |
| 1 | 154 | $139 \cdot 1$ | 18 | ' 221.6 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 2 | $192 \cdot 3$ | 191.81 | $179 \cdot 1$ | $179 \cdot 1$ |  |  |  |  |  | 2 |  | 35 6 |  | $158 \cdot 6$ |  |  | $145 \cdot 4$ |  | 17 |  |  | 178 |  |  | 191.0 | 263.3 |  |  |  |
| 3 | $191 .{ }^{\prime}$ | 191.0 | $195 \cdot 51$ | $192 \cdot 6$ | 2 | ${ }^{14} \cdot 7$ | 1.5 | $239 \cdot 1$ | 216.5 | $172 \cdot 4$ | 179.0 | 177.5 | $169 \cdot 5$ | 159.2 | $166^{\circ} 1$ | $148 \cdot 4$ | 1317 | $143 \cdot 9$ | 8 | $0 \cdot 7$ | 162 | $165 \cdot 2$ | $175 \cdot 8$ | $179 \times 6$ | $185 \cdot 3$ | 263.3 | 7.24. |  | 3,37 .. |
| 4 | $185 \cdot 3$ | $175 \cdot 2$ | 174.0 | 5 | $182 \cdot 91$ | $177^{\circ}$ | 174.5 | 202.5 | 193.5 | $199 \cdot 1$ | 187.1 | 171.9 | $155 \cdot 6$ | 15 | 14 | 1499 | 154.1 | 1517 | 158.3 | 162.5 | 188.0 | $171 \cdot 6$ |  | 8 | $65 \cdot 5$ | 234.8 | 8.15 , | 141.6 | 2.14 |
| 5 | -5 | $160 \cdot 8$ | 163.1 | 172-9 |  | -6 | 5:0 | 184.5 | 171.8 | 168.0 | 168.8 | 3.5 | $155 \%$ | 152-9 | $153 \cdot 8$ | 153.8 | 154.1 | $152 \cdot 9$ | 153.8 | 148.2 | 1 | $164 \cdot 0$ | $172 \cdot 5$ | $173 \cdot 6$ | $174 \cdot 8$ | 203.0 | 6.15 . | $144 \cdot 3$ | 7.1 ., |
| 6 | $174 \cdot 8$ | $180 \cdot 8$ | $177 \cdot 5$ | 5 | 1 | $167^{\circ} 0$ | $189 \cdot 3$ | $184 \cdot 5$ | 191.0 | $190 \cdot 1$ | 191 '6 | $188 \cdot 9$ | $187^{\circ}$ |  | $185 \cdot 6$ | 1571 | $150 \cdot 5$ | 3 | $149 \cdot 3$ | $53 \cdot 5$ | 156 | 0 | $174 \cdot 9$ | 0 | $187 \cdot 1$ | $222 \cdot 5$ | 9.3 " | 141.2 | 5.43 |
| 7 | 1 | 186.5 | 18801 | $170 \cdot 6$ | 1 | . 5 |  |  | +288.5+ | $268 \cdot 5+$ | $2811^{\circ}$ | b | - | 2 | $170 \cdot 1$ | $127 \cdot 1$ | $131 \cdot 4$ | 5 5 | 166.5 | 180.4 | $170 \cdot 8$ | 176.4 | 178-5 | 5 | $175 \cdot 1$ | $268.5+$ | 6-10 , | 116 | 8.23 |
| 8 | 1 | 13 | 5 | 01 | 174.81 | $188 \cdot 0$ | 181.8 | 2096 | 218.4 | $224 \cdot 6$ | 174.3 | $35^{\circ}$ | 8 | 17 | $156 \cdot 3$ | $46 \cdot 0$ | $150 \cdot 9$ | 142.8 | 138.6 | $\cdot 6$ | 154 |  | 168.3 | -9 | $8 \cdot 5$ | 2490 | 9.18 | 123.0 | 3.34 |
| 9 | $186 \cdot 5$ | $188^{\circ}$ | $196 \cdot 11$ | $5 \cdot 9$ | $195 \cdot 51$ | - 8 | $194 \cdot 7$ | $225{ }^{\circ}$ | $232 \cdot 8$ | 203.0 | 4.9 | 177.0 | - | $161 \cdot 4$ | 7 | $145 \cdot 7$ | $142 \cdot 2$ | ${ }^{1} 1$ | 163.5 | 188.8 | 162.8 | 170.0 | 163.1 | $173 \cdot$ | $2 \cdot 1$ | - | 7.17 | $138 \cdot 9$ | 4.25 |
| 10 | 17 | $172 \cdot 1$ | $179 \cdot 4$ | $195 \cdot 51$ | 19502 | $200 \cdot 6$ | 192.5 | $202 \cdot 3$ | 89.9 | 258.0 | $\cdot 1$ | 18.0 | 213.5 |  | 192. |  | $\cdot 5$ | 164-6 | 143.7 | $145{ }^{1}$ | $146^{\circ}$ |  | 168.8 | $1 \cdot 6$ | $\cdot 1$ | 264.0 | 8.42 | 1370 | 6.5 |
| 11 | 188 | $9 \cdot 1$ | 160.21 | 1 | 185.91 | 1893 | $202 \cdot 2$ | 179.0 | $175 \cdot 5$ | 28.5 | $2 \cdot 0$ |  |  |  | $188 \cdot 8$ | $158 \cdot 3$ | 1610 | $155 \cdot 3$ | $154{ }^{\text {] }}$ | $162 \cdot 5$ | 164.7 | $167 \cdot 4$ | 178.1 | 5 | 178.5 | $230 \cdot 4$ | 9.26 | 151.8 | 2.330 b.m. |
| 12 | 178.5 | \% | 188-2 | 188.2 | 182.91 | 8 | 192.5 | 231.0 | $229 \cdot 5$ |  |  | a | 179.0 |  | 158.7 | $150 \cdot 5$ | 2 | 1458 | $9 \cdot 8$ | $150 \cdot 3$ | , | 1 | 149.9 | - | 159.5 | $252 \cdot$ | ¢. 50 | 139 | 4.30 p.m. |
| 13 | 15 | 155'9 | 1 | 149.7 | 167.91 | $150 \cdot 3$ | 141.0 |  | 154.4 | $180 \cdot 1$ | $171{ }^{\circ} 0$ | a |  |  |  |  |  |  | $131 \cdot 6$ | $160 \cdot 7$ | $158 \%$ | $165{ }^{\circ}$ |  | 154.1 | 161.0 | $198 \cdot 5$ | . 23 |  | 7.3 a.m. |
| 14 | 16 | - 5 | 4.7 | 71 | 190 | 20.8 | 230 |  | 211.1 | 224 | $211 \cdot 4$ | $3 \cdot 3$ | 0 | $200 \cdot 8$ | 180.8 | $173 \cdot 3$ | 1610 | $155 \cdot 7$ | 1385 | $122 \cdot 6$ |  | $138 \cdot 0$ | 170.6 | $\cdot 0$ | $185 \cdot 7$ | 261.0 | 4.44 , | 108 | 6. 45 p.m |
| 15 |  |  |  | 23369 | 202.5 | 86 | 83.0 | 85.6 | 108.8 | 140.6 | $104 \cdot 6$ | 18 | $139 \cdot 1$ | 118.7 | $12 \cdot 4$ | 1113 | 108.6 | $114 \cdot 0$ | $113 \cdot 9$ | $117 \cdot 3$ | $132 \cdot 9$ | $146^{\circ} \theta^{\prime}$ | $162 \cdot 6$ | $5 \cdot 1$ | 6. | 2550 | 2.30 " | 28.8 | 7.8 a.m. |
| 16 |  | 159.0 |  | 1 |  | T | $142 \cdot 5$ | 91.3 | 1870 | $137 \cdot 9$ | 7.5 | 78 | $155 \cdot 3$ |  |  |  | $119{ }^{\circ}$ | $105 \cdot 0$ | $122 \cdot 1$ | 134.0 | $138{ }^{\circ}$ | $136 \cdot 2$ |  |  |  | $232 \cdot 5$ | 9.45 | $42^{\circ} 0$ | 6.45 |
| 17 | - | 1 | 11 | 1 |  | $165 \cdot 5$ |  | $2 \cdot 8$ | 218.7 | $144 \cdot 6$ | 4.1 | $200 \cdot 3$ | $190 \cdot 5$ |  | 141.6 | $190 \cdot 1$ | $170 \cdot 1$ | $157 \%$ | $151 \cdot 5$ | $156 \cdot 9$ | 159.5 | $167 \cdot 4$ | $173 \cdot 9$ |  | 178.2 | $256 \cdot 5$ | 8.25 . | $109 \cdot 7$ | 9.25 |
| 18 | 178.2 | 168.4 | 81 | $161 \cdot 1$ | $168 \cdot 6$ | 139.5 | 87.1 | $3 \cdot 8$ | $160 \cdot 1$ | 77 | 198•3 | 171.8 | 156.0 |  |  |  | $149 \%$ | 151.7 | $154 \cdot 8$ | 2 | $170 \cdot 9$ | $175 \cdot 5$ |  |  | 180.5 | 0 | 10.22 , | 117.0 | 5.5 |
| 19 | $180 \cdot 5$ | 183.0 |  | $208 \cdot 1$ |  | 2015 | 204.0 | 194.3 | 97-1 | 2078 | 200 | 142.7 | 149.1 |  |  |  |  |  | $113 \cdot 1$ | - 9 | 9 | 16 | $31 \cdot 6$ | 54. | 152.0 | $256 \cdot 5$ | 9.21 , | $85 \cdot 5$ | 5.22 p.m. |
| 20 | $\bigcirc$ | 1814 | $158 \cdot 21$ | 158.0 |  | $149^{\circ}$ | $148^{\circ} 0$ | 128.6 | $184 \cdot 1$ | 174.0 | $186 \cdot 5$ | $38 \cdot 2$ | 1 | 3 | $135 \cdot 6$ |  | 12 | $131 \cdot 8$ | 1271 | $140 \cdot 0$ | $154 \cdot 1$ | $182 \cdot 8$ | $176 \cdot 9$ | $169 \cdot 2$ | $170 \%$ | 195.0 | 7.35 | 104 | 7.15 am. |
| 21 | $170 \cdot 9$ | 0 |  | $185 \cdot 6$ |  |  | 137.0 | 131.7 | $159 \cdot 3$ | 2430 | $29 \cdot 2$ | ${ }^{228} \cdot 7$ | $228 \cdot 1$ |  | $210 \cdot 9$ |  | $187 \cdot 4$ | $154 \cdot 1$ | $148 \cdot 4$ | $152 \cdot 1$ | $145 \cdot 1$ | $1 \cdot 2$ | 1518 | $158 \cdot 3$ |  | 31.3 | 8.52 , | 85.9 | 6.50 \% |
| 22 | 175.5 | 0 | $100 \cdot 51$ | 181.7 |  | 12438 | 208.0 | 1845 | 292.8 | $223 \cdot 1$ | $180 \cdot 1$ | 32. | 7 | $130 \cdot 8$ | 184.5 |  |  |  | $138 \cdot 3$ |  | '5 | 119.1 | $132 \cdot 5$ | 131.0 |  | $300 \%+$ | 8.0 " | 49.5 | 6.55 " |
| 23 | 142.8 | 1 |  | 5 |  | $232 \cdot 1$ | $225 \cdot 5$ |  | $254 \cdot 3$ |  | $265{ }^{5}$ | 231 | $209 \cdot 6$ |  | $173^{\circ}$ |  |  |  |  |  | -1 | $103 \cdot 8$ | 8.7 | 128.0 |  | $298 \cdot 5+$ | 7.50 | 75.0 | 10.35 p.m. |
| 24 | $195 \cdot 8$ | $220 \cdot 5$ |  | 158.1 |  |  | $170 \cdot 6$ | 201.0 | 2550 |  |  | 288.5+ | $252 \cdot 0$ | $189 \cdot 0$ | $240^{\circ}$ |  |  |  |  |  | $88^{\circ}$ | 70 | $69 \cdot 3$ |  |  | 298 '5 | 9-11 | $4 \cdot 5$ | 8.40 |
| 25 | 127.8 |  |  |  |  | 2220 |  |  |  |  |  | 228.8 |  |  | 15 |  |  |  | 1070 |  |  |  | 85.1 |  |  | $298 \cdot 5$ | 8.40 " | $46 \cdot 5$ | 8.26 |
| 28 |  | 1301 |  |  |  | , |  |  |  |  | $298.5+$ | 270.0 |  |  | 2 |  |  | $123 \cdot 3$ | 1 |  | $124 \cdot 5$ | $143 \cdot 8$ | $158 \cdot 8$ |  |  | $298 \cdot 5$ | 8-8 | 102 | 4.10 |
| 27 |  |  |  |  |  |  | $0 \cdot 6$ | 1710 | $9 \cdot 8$ | $183 \cdot 1$ | $204 \cdot 0$ | 154.5 |  |  | $173 \cdot 8$ |  | $151 \cdot 2$ |  | 128.0 | 4 |  | $154 \cdot 1$ | $162 \cdot 3$ | $136 \cdot 1$ |  | $243 \cdot 8$ | 9.45 | $85 \cdot 5$ | am. |
| 28 |  | 5 | $154 \cdot 11$ | $159^{\circ}$ |  | $185^{\circ} 0$ | 174.8 | 152 | 178.7 | $70^{3}$ | 238.0 | 175 | 144.5 | $175 \cdot 4$ | 190.1 |  | $144 \cdot 8$ | 14 | 12 | $115 \cdot 1$ |  | 1 | $135 \cdot 8$ | $142 \cdot 1$ | $154 \cdot 2$ | $253 \cdot 5$ | 10. | 88.0 | 6.43 p.m. |
| 29 | 15 | $174 \cdot 6$ | 188.0 | 162.0 |  | 178-9 | 180.2 | $200 \cdot 0$ | $209 \cdot 6$ | 1 | $7 \cdot 8$ | 10 | $186^{6}$ |  | 184. |  | $144 \cdot 8$ | $150 \cdot 5$ | 159.3 |  | $141 \cdot 8$ | $140 \cdot 3$ | $14 \cdot 2$ | 148 | $183 \cdot 7$ | 2370 | 8.17 ., | 127.5 | 7.53 |
| 30 | 18 | 178.1 | 1 | $188 \cdot 1$ | 184.3 | $1150 \cdot 5$ | 169.1 |  | $202 \cdot 5$ | 183.0 | 159.4 | 131.3 |  |  |  |  | $105 \cdot 3$ | 108.9 | $113 \cdot 8$ | $124 \cdot 5$ | 18 | 20.51 | 112.7 | $128 \cdot 6$ | 4 | 2180 | 8.1 | 91.5 | 4.23 |

DECLINATION, DECEMBER, 1902.


DECLINATION，FEBRUARY， 1903.

| Day． | Midt． <br> 0. | D $=151^{\circ}+$ |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | Midt． | Maximum reading and time． |  |  | $\begin{aligned} & \text { Minimum } \\ & \text { reading and } \\ & \text { tinge. } \end{aligned}$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | Forenoon． |  |  |  |  |  |  |  |  |  | Noon． | Afternoon． |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11.12. | 1. | 2. | 3. | 4. | 5. | 8. | 7. | 8. | 9. | 10. | 11. | 12. |  |  |  |  |  |
|  | ， |  | ， | ， |  | 1 ＇ |  | ＇ | ＇＇ | ＇ |  |  |  |  | ＇ |  |  | ， |  | ， |  |  |  |  |  |  |  |  |  |
| 1 | $110 \cdot 9$ | $102 \cdot 9$ | $123 \cdot 8$ | 124＇2 | $2128{ }^{\circ} 0$ | $130 \cdot 4$ | 4 112.4 | 4119.8 | 3 108＇2 | 104.0 | $121-5$ | 96.090 | $87 \cdot 2$ | 88.7 | $87 \cdot 0$ | 74.4 | 94.5 | $121 \cdot 2$ | $110 \cdot 3$ | 117.0 | $121 \cdot 1$ | 118.1 | 1278 | $117 \cdot 2$ | 1470 | 5.22 a | a．m． | 58.5 | 11.7 a．m． |
| 2 | $117 \cdot 2$ | 103.8 | $107 \%$ | $126 \cdot 8$ | $8110 \cdot 1$ | $145 \cdot 1$ | $147 \cdot 3$ | $3137 \cdot 0$ | $143 \cdot 1$ | 150.0 | $191 \cdot 7$ | 161．7136．2 | 121.7 | $120 \cdot 3$ | 113.7 | 104－9 | 98.7 | 91.1 | $100 \cdot 7$ | $116^{\circ} 4$ | 118.5 | 106.3 | $104 \%$ | 132.5 | 209.0 | 9.44 | － | 82.5 | 6.11 p．m． |
| 3 | 132.5 | 124－2 | $132 \cdot 2$ | 140．7 | $784 \cdot 3$ | $107{ }^{\circ}$ | 9， $127 \cdot 5$ | 168.0 | 13 | $178 \cdot 5$ | $202 \cdot 4$ | a |  | a | a | a | a | 98.5 | 95.6 | 108.0 | $100 \cdot 1$ | $95 \cdot 1$ | 1199 | $130 \cdot 1$ | 221.4 | 9.41 | ， | 42.0 | $4.22 \mathrm{ar.m}$ ． |
| 4 | $130 \cdot 1$ | $136 \cdot 1$ | $136^{\circ} 4$ | $120 \cdot 5$ | 5122.4 | $119 \cdot 9$ | 9 127 －5 | 128.1 | 138.5 | 146.7 | $135 \cdot 3$ | $145 \cdot 8.134 \cdot 6$ | $125 \cdot 91$ | 132.0 | 124.8 | $111 \cdot 3$ | a | a | ${ }^{8}$ | a | $101 \cdot 6$ | 99.0 | 103 5 | 126．5 | $161 \cdot 1$ | 10.43 | ， | $85 \cdot 2$ | 4.51 p．m． |
| 5 | $126 \cdot 5$ | $132 \cdot 6$ | $133 \cdot 4$ | $144 \%$ | 3． $148 \cdot 1$ | $156 \cdot 2$ | 2 191 \％ | 6） $182 \cdot 4$ | 1976 | 203.9 | 187.2 | $174.0150 \cdot 0$ | 143.9 | 1295 | $82 \cdot 1$ | 87.2 | 678 | 58.5 | $54 * 3$ | 87.6 | 75.3 | 69.5 | 85.4 | 102 ＇5 | $330 \cdot 0+$ | 8.33 | ， | $36^{\circ} 0-$ | 6.40 |
| 8 | 102.5 | $99 \cdot 2$ | 9. | 132 5 | 5） 111.5 | 109.5 | 5． 74.0 | （0） 69.8 | 9 | $117 \cdot 6$ | $105 \cdot 3$ | $115 \cdot 497 \cdot 1$ | 104.4 | ． 0 | a | 82.5 | 78.8 | $73 \cdot 4$ | $74 \cdot 1$ | 83.0 | $90 \cdot 9$ | 98.6 | 98.1 | 104 | 145．0． | 3.56 | ＂ | 37.5 | 4.18 |
| 7 | $104 \cdot 7$ | 118.5 | $105 \cdot 8$ | 123．5． | 5． $136 \cdot 1$ | $127 \cdot 2$ | 232.2 | $135-2$ | $129 \cdot 6$ | $127 \cdot 8$ | $137 \cdot 9$ | $146^{\prime} 4129^{\circ} 3$ | 125.0 | $117 \cdot 8$ | 1118 | 98.7 | 105.0 | 95.7 | $93 \cdot 7$ | $89 \cdot 3$ | $89 \cdot 1$ | $80 \cdot 0$ | 98.6 | 0 | $160 \cdot 5$ | 10.42 | ＂ | 70－2 | 9.44 |
| 8 | 100.5 | 111.0 | $110 \cdot 3$ | 119.0 | 116.0 | 110.0 | 126－5 | 111.9 | 5 | 141.3 | $123 \cdot 3$ | $123.036 .0-$ | $72 \cdot 2$ | －5． | $75 \cdot 2$ | 72.0 | 9 | $36.0-$ | $48 \cdot 0$ | 36.0 | $36 \cdot 0-$ | 43.0 | 85.2 | 115.2 | $160 \cdot 3$ | 10.50 | ， | $36^{\circ} 0-$ | 8－9 |
| 9 | $115 \cdot 2$ | 107．9 | $139 \cdot 1$ | 133＊8 | 8145 | 139．5 | 5． 180 \％ | 5283.7 | $210 \cdot 0$ | $134 \cdot 9$ | 94.1 | $128 \cdot 6 \cdot 122 \cdot 9$ | a | a | a | a | ． 0 | $92 \cdot 6$ | 84.5 | 91.1 | 86.7 | $01 \cdot 4$ | 118.4 | 117.9 | 2. | 7． 11 | ， | 50.4 | 9.45 m m．m． |
| 10 | $117 \cdot 9$ | 118.4 | $121 \cdot 1$ | $124 \cdot 1$ | 1121.4 | 133.7 | 140.0 | $180{ }^{\circ} 0$ | 165. | $188 \cdot 3$ | 168.0 | 165．0162．2 | 147．31 | $145 \cdot 1$ | $126 \cdot 9$ | $97 \cdot 4$ | 75.0 | 84.0 | $80 \cdot 1$ | $92 \cdot 0$ | $92 \cdot 3$ | 96.5 | $94 \cdot 1$ | $102-9$ | $212 \cdot 3$ | 9.23 | － | 62：3 | 5.3 p．m． |
| 11 | 102.9 | 112.7 | 1151 | $123 \cdot 5$ | $5.109 \cdot 1$ | 121.5 | 51445 | 5164.0 | 201.0 | 192.0 | 187 \％ | $183.8170^{\circ} 3$ | 163.7 | 1479 | $129 \cdot 5$ | $100 \cdot 4$ | 71.7 | 59.7 | $66^{6} 6$ | $77 \cdot 9$ | 93.5 | $100 \%$ | 9 | 113 | $5 \cdot 0$ | 8.1 | ＂ | 42.0 | 6.51 |
| 12 | 113.9 | 114.0 | $128 \cdot 8$ | 117.8 | $8102 \cdot 9$ | 134.3 | ） | ． 8 | －96\％3 | 148.2 | 1200 | $117 \cdot 9120 \cdot 5$ | $117 \cdot 2$ | $93^{\circ} 0$ | $85^{\circ} 6$ | $84 \cdot 9$ | $92 \cdot 1$ | $91 \cdot 4$ | 96.0 | $97 \%$ | 104.3 | 98.0 | $106 \cdot 1$ | 59 | $198 \cdot 9$ | 8.25 | ， | $3{ }^{3} 5$ | 8．5 m．mm， |
| 13 | 99.3 | $104 \cdot 1$ | $107 \%$ | 110.0 | 0103.5 | 81.5 | $153 \cdot 5$ | 5 255.8 | ． 0 | a | a | 125.6 | $7 \times 2$ | 154.2 | 1349 | 153.0 | $105 \cdot 0$ | $82 \cdot 1$ | 101.7 | 120.0 | $115 \cdot 8$ | 1076 | 114 | 126 | ／8 | 7.30 | ＂ | 51 \％ | 6.1 |
| 14 | －126．5 | 121.7 | $128 \cdot 5$ | 128．9 | ${ }^{9} 123 \cdot 9$ | 114.8 | $8{ }^{91} 1$ | $163 \cdot 1$ | 163．5 | b | $186 \%$ | $198 \cdot 0175 \cdot 2$ | $155 \cdot 6$ | 1347 | 110.0 | 87 | 4 | $92 \cdot 4$ | 104.0 | 98.6 | 83.2 | $93 \cdot 1$ | $146 \cdot 2$ | 1197 | 3.0 | 10.32 | － | $34 \cdot 5$ | 5.47 |
| 15 | 119.7 | $\cdot 6$ | $119 \%$ | $134 \cdot 6$ | 6128.0 | $131 \cdot 9$ | $145 \cdot 8$ | 145．1 | 193.5 | 201.5 | 184.4 | 178.5179 .6 | $161 \cdot 9$ | 134.3 | 122 \％ | 143.5 | 1143 | $115 \cdot 2$ | 1108.5 | 107.1 | 93.9 | $103 \cdot 8$ | 119 | 105 | 2185 | 8.46 | － | $\because$ | 4．2．3 p，m． |
| 16 | $105 \cdot 9$ | $109 \cdot 1$ | $139 \cdot 2$ | 148.4 | ${ }_{4} 177 \cdot 8$ | 182.8 | $145 \cdot 2$ | 167.0 | $180^{\circ} 0$ | $191 \%$ | $181 \%$ | $181 \cdot 5162 \cdot 8$ | －8 | $149 \cdot 1$ | 124.7 | 114.3 |  | $13 \cdot 1$ | $107 \cdot 4$ | 94 | 107 | 119.0 | 128.0 | 12 | 48.5 | 9.20 | $\cdots$ | 52.3 | ＊． 21 |
| 17 | 126.8 | $129{ }^{\circ}$ | $125 \cdot 9$ | $130 \cdot 5$ | 5.1470 | $165 \cdot 5$ | $163 \cdot 8$ | $183 \cdot 3$ | $186{ }^{\circ}$ | 192.0 | 194.0 | $175 \cdot 8181 \cdot 2$ | 165.5 | $159 \cdot 0$ | $153 \cdot 0$ | $122 \cdot 4$ | 104.1 | 118.0 | $106 \cdot 8$ | $92 \cdot$ | 99.0 | $102 \cdot 1$ | 10.0 | 9 | $24 \%$ 为 | 10.15 | ， | 87 | $\times 1$ |
| 18 | 97.7 | $107 \cdot 4$ | $130 \cdot 1$ | 1415 | 5 5 151－2 | ＇4 | 132 －6 | i $146^{\circ} 0$ | 158.4 | 153.5 | 164.9 | $160 \cdot 4150 \cdot 9$ | $115 \cdot 1$ | 91.5 | 83.4 | 770 | 71 | 64．5 | 45 | 75.8 | $80 \cdot 3$ | 108.2 |  | 119 | 185\％ | （4．20） | ， | 为0－ | ． 2.2 |
| 19 | 119.0 | 118.8 | 123.0 | 1316 | $6138{ }^{\circ} 0$ | 146.0 | $149 \cdot 1$ | $1155 \%$ | 168.5 | $163 \cdot 1$ | $161 \%$ | $14.5 \cdot 6145 \cdot 1$ | 141.91 | $126 \cdot 3$ | $110 \cdot 9$ | 11001 | $111 \cdot 9$ | $110 \cdot 4$ | 1095 | 897 | 101. | $110 \cdot 3$ | 1193 | 111 | 1594 | 8． 16 | － | ¢ | 8.17 |
| 20 | $111 \%$ | $114 \cdot 8$ | 123.5 | $122 \cdot 0$ | $0128 \cdot 3$ | 1388 | $3142 \cdot 8$ | $8149 \cdot 0$ | 143.1 | 186.6 | 169.4 | $15{ }^{\circ} 0154 \cdot 5$ | $132 \cdot 3$ | $107 \cdot 6$ | \＃i | 356 | $60^{\circ} \mathrm{O}$ | 53：3 | $72 \cdot 2$ | $\pi$ | ． | 41. |  | 1230 | $4 \cdot 1$ | $\cdots$ | ． | $11 \cdot 4$ | 二．3＊ |
| 21 | 1260 | 115.3 | 109.4 | $110 \cdot 6$ | 6） 109 万 | 101.0 | 0 i4．4 | $4112 \cdot 5$ | 5 $144 \cdot 6$ | $155 \cdot 5$ | 1760 | $172.2148 \cdot 1$ | 116.9 | $158 \cdot 0$ | 89. | 11 | a | a | a | 117 | $1 \cdot 1$ | ． |  | an 3 | ！as | 11.23 |  | is |  |
| 22 | $0 \cdot 3$ | $125 \%$ | $129 \cdot 2$ | $12 \% 1$ | 1138.0 | 1.0 | 1143.7 | 7148.2 | $\cdot 5$ | 124.4 | $141 \%$ | $141 \cdot 8 \cdot 130 \cdot 5$ | 174.6 | $167 \cdot 6$ | 1.20 | 12. | 57.2 | 34．5－ | $34 \cdot 5-$ | 3 | 53： | $59 \cdot 3$ |  | 隹碞 | 1．8． | 1.31 | m．m． | 3： | \％pm． |
| 23 | 108.0 | 1197 | $117 \times$ | 3． $115 \cdot 1$ | $1100 \cdot 5$ | 58.8 | －8 $95 \cdot 4$ | ＋ $112 \%$ | 107 ＊ | 114.2 | $91 \times$ | ${ }_{12} 2 \cdot 0 \cdot 88 \cdot 1$ | $114 \%$ | 84.8 | x $4 \cdot 6$ | 89.4 | 8 ¢ ${ }^{5}$ | 88.5 | 86.6 | 88.8 | 89.1 | 20 | $102 \%$ | J14： | ： | 1．．74 | A．11 | ：480 | 11．2．s．m |
| 24 | 104.4 | 10\％ 3 | $101-9$ | 97．1 | $1105 \cdot 0$ | 112.4 | $4105 \cdot{ }^{-1}$ | ${ }^{-1} 102 \cdot 5$ | 5 $111 \%$ | 100．7 | 1：31\％ | $123 \cdot 1.140 \cdot 1$ | $140 \cdot 1$ | 1196 | 128.0 | $121 \cdot 3$ | 10s．0 | $48 \cdot 1$ | $100 \cdot 1$ | $102 \cdot 5$ | 113.1 | 11 tit | 11.15 | 115： | 171： | 11．40 | $\cdots$ | $\therefore$ | －．．3 |
| 25 | 115.8 | 1187 | 118.8 | 81187 | 7118.1 | $139 \cdot 1$ | $1138 \cdot 3$ | 3） $245 \cdot 2$ | $162 \cdot 9$ | － | $168 \cdot{ }^{\text {ch }}$ | $1 \cdot 7157 \%$ | 7•3 | $133 \cdot 1$ | 158.3 | $121 \cdot 1$ | x9\％ | 61\％ | tin＇ 8 | $64 \%$ | 7 | 43 | $10 \cdot 6$ | 120 | $\cdots$ | 4．2\％ |  | ：413 |  |
| 26 | 1290 | 125.9 | $140 \cdot 0$ | 148．3 | $2140 \cdot 4$ | 131.4 | 4143.2 | $215 \% 1$ | 1：n\％$\%$ | $1 \%$ | $1+8$. | $\therefore 1066^{\circ}$ | 97.5 | $\pm 2 \cdot 1$ | 91.5 | \％ |  | $4 \cdot 1$ | 96.5 | 14.9 | 16\％${ }^{\text {\％}}$ | $14 \mathrm{n} \cdot 9$ | 1119 | $112 \cdot 1$ | －429 | \％．16 | $\cdots$ | 11：$\%$ | い乐 |
| 27 | $112 \cdot 1$ | 114.5 | 1170 | 1160 | $0115 \cdot 1$ | 108.3 | 3 105．8． | $8.110 \%$ | 12x＇0 | 122 | $12 \times 4$ | 122＊3127．1 | －3 | ：3 | 11×\％ | 78\％ | 34 | N13 | 933 | 103．2 | 14\％ | $110 \cdot \%$ | 114 | 11：1 | 1：\％1 | －69 |  | 404 | 1．2＊ |
| 28 | $13 \cdot 1$ | 119．9 | 121＊8 | ＋ $124 \cdot 1$ | $128 \cdot 1$ | $123: 9$ | 9120＇3 | 3 $137 \cdot 1$ | $142 \times 2$ | 118 如 | $149 \cdot 13$ | 14.1138 .8 | $119 \% 1$ | $110 \cdot 1$ | Lu9 ${ }^{3}$ | lutis | 101\％ | 46－2 | 4，9－5 | 13.2 | 140 |  |  | 121 | 1，－ | 9．3： | － |  | $\cdots$ |

$$
\text { p } 2
$$

DECLINATION, MARCH, 1903.

DECLINATION, APRIL, 1903.

| Day. | $\mathrm{D}=150^{\circ}+$ |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Midt. |  |  |  |  | Forenoon. |  |  |  |  |  |  | Noon. | Afternoon. |  |  |  |  |  |  |  |  | Midt. ${ }^{\text {' }}$ | $\begin{aligned} & \text { Maximum } \\ & \text { reading and } \\ & \text { time. } \end{aligned}$ |  | $\begin{aligned} & \text { Minimum } \\ & \text { rearling and } \\ & \text { time. } \end{aligned}$ |  |
|  | 0. | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 8. | 10. | 11. | 12. | 1. | 2.3. | 4. | 5. | 6. | 7. | 8. $\quad 9$. | 10. | 11. | 12. |  |  |  |  |
|  |  |  |  |  |  | ' |  | ' ${ }^{\prime}$ | , 15 |  | 2 | ' |  |  |  |  |  | ${ }_{1}$ | , | ' ${ }^{\prime}$ | ' 12 | ${ }^{\prime}$ |  |  |  |  |  |
| 1 | 123.7 | 127.0 | $130 \cdot 0$ | $150 \cdot 1$ | 144.8 | 139.0 | $139 \cdot 3$ | $136 \cdot 9$ | $150 \cdot 5$ | 157.0 | $142 \cdot 9$ | $133 \cdot 3$ | 111.2 | 118.3 | $134 \cdot 5116.0$ |  | 4107.5 | $105 \cdot 1$ | 103.0 | $115 \cdot 3118 \cdot 7$ | $117 \times 2$ | $120 \cdot 2$ | 1160 | 185.5 | $8.59 \mathrm{a} . \mathrm{m}$. | 85.4 | 3.5 p p.m. |
| 2 | 116.0 | $120 \cdot 4$ | $134 \cdot{ }^{5}$ | 128.3 | $131 \%$ | $130 \cdot 0$ | 135.3 | $135 \cdot 4$ | $138 \cdot 1$ | 146.0 | 175\% ${ }^{\circ}$ | $168 \cdot 6$ | 141.1 | 135.2 | $135 \cdot 1112 \cdot 4$ |  | $114 \cdot 1$ | 96.7 | $91 \cdot 1$ | $82 \cdot 01108$ | 113.5 | $120 \cdot 4$ | $125 \cdot 3$ | $180^{\circ} 2$ | 10.19 | 75:\% | 7. ¢ $^{6}$ |
| 3 | $125 \cdot 3$ | 131.0 | $130 \cdot 9$ | $130 \cdot 4$ | $132 \cdot 1$ | 14.1 | 168.8 | 152-2 | $158{ }^{\circ}$ | 151.0 | $135 \cdot 5$ | 170.5 | 134.0 | $129 \cdot 2$ | $115 \cdot 3109 \%$ |  | $1 \cdot 8$ | 111.5 | 117.4 | $1210119 \%$ | 1171 | 123.2 | $119 \cdot \%$ | $182^{\circ} 0$ | 11.22 | $105 \cdot 1$ | 3.2 |
| 4 | 119.8 | $126 \cdot 1$ | 125.9 | $124 \cdot 6$ | 128.6 | 129-4 | 137.5 | 146.2 | $139 \cdot 6$ | $130 \cdot 6$ | $130 \cdot 6$ | 131.8 | 1324 | 120 | 3:1117 7 | 99 | 3 | 102.5 | 103.7 | $106 \cdot 6110 \cdot 6$ | 1944 | 123.7 | $117 \cdot 1$ | $155 \%$ | 8.15 | $80 \cdot 8$ | 4.28 |
| 5 | $117 \cdot 1$ | $128 \cdot 9$ | $131 \cdot 2$ | $135 \cdot 4$ | $134 \cdot 8$ | $137 \cdot 5$ | $152 \cdot 5$ | 128.5 | $152 \cdot 0$ | 153.8 | $165 \cdot 1$ | - 8 | 1376 | $123 \cdot 5$ | $109 \cdot 3 \mid 94.4$ | 93 | $4 \cdot 6$ | 72.2 | $93 \cdot 1$ | $120 \cdot 8108.8$ | 129.5 | 129 | 131\% | 1970 | 4.32 |  | 5.42 |
| 6 | 131 5 | 128.91 | 168.4 | 139-7 | 159.5 | 145.6 | 130.6 | $121{ }^{\circ} 9$ | 128.5 | 135.4 | $146 \%$ | $2 \cdot 7$ | $\checkmark$ | 225.5 | 168.1, 36.5- |  | 4188 | $142 \cdot 9$ | 134.5 | $72.1365-$ | $\cdot 5$ | 8\%- | $90 \cdot 5$ | 263.5 | 0.15 p.m. | $3{ }^{3} \cdot 5$ | 9-11 |
| 7 | 90-5 | 180.6 | 161.0 | $167 \cdot 6$ | 178.0 | 175.9 | $153 \cdot 1$ | $143 \cdot{ }^{\circ}$ | 141-1 | $150 \cdot 1$ | $159 \times 2$ | $177 \cdot 1$ | 179.9 | $167 \%$ | 158.0142 .0 | 138. | $1132 \times 2$ | 1205 | 109.6 | 103.01172 | 13119 | $130 \cdot 3$ | 12n-1 | 205.3 | 11.11 | 23-8 | \%.17 |
| 8 | 126.1 | $157 \cdot 31$ | $125 \cdot 5$ | $135 \cdot 4$ | 146.0 | $158 \cdot 3$ | 146 | 157.0 | $155 \%$ | 154 '6 | 139-3 | $153 \cdot 1$ | 152 | $144 \cdot 8$ | 144.5137 3 | 136 | 131.5 | 128.6 | 125.5 | $124 \% 128 \cdot 5$ | 133.9 | $132 \cdot 2$ | $139 \cdot 0$ | 213 \% | 4.2 a.m. | $143: 3$ | 4.25 a.r |
| 9 | 139.0 | 128.81 | $133 \cdot 1$ | 153.4 | $135 \cdot 1$ | $141 \cdot 1$ | $145 \%$ | 151.6 | 137.8 | 138.5 | $170 \cdot 3$ | $194 \cdot 3$ | $179 \cdot 3$ | 180* | 134.5128 .0 | 113.5 | $41^{10}$ | $41^{\circ} \mathrm{O}$ | 1490 | $77 \% 10036$ | 125.6 | $137 \cdot 5$ | $142 \cdot 3$ | $2.4 \%$ | 11.3 | $41^{\circ} \mathrm{C}-$ | inti p.m. |
| 10 | $142 \cdot 3$ | 143.51 | 140.0 | 144 -3 | $165 \cdot 8$ | $175 \cdot 9$ | 176 | 191.3 | $166 \cdot 1$ | $205 \cdot 9$ | $207 \cdot 2$ | 159.2 | $145 \cdot 6$ | $142 \cdot 0$ | 138.5124 .0 |  | $128 \cdot 3$ | 112.7 | $115 \cdot 1$ | $84.5110 \cdot 8$ | $124 \cdot 1$ | $142 \%$ | $11 \cdot 1$ | 2534 | 6.40 | 53.30 | 7.45 |
| 11 | $117 \cdot 1$ | $131 \cdot 81$ | $161 \cdot 3$ | $143 \cdot 5$ | $138 \cdot 5$ | $149 \cdot 0$ | $163 \cdot 1$ | $193 \cdot 1$ | $192 \cdot 5{ }^{\prime}$ | $172 \cdot 3$ | 156.2 | 143.0 | $157 \%$ | $135 \%$ | 125.3138 .1 | 129 | $121 \cdot 3$ | 115.4 | 98.8 | $105 \cdot 4108 \cdot 1$ | 112.4 | 143.5 | 18434 | $203: 3$ | 6.16 | 74.0 | -. 10 |
| 12 | $138 \cdot 4$ | 140.01 | $138 \cdot 1$ | 134.0 | $136{ }^{\circ} 8$ | $142 \cdot 9$ | 142.0 | $140 \cdot 0$ | $142 \cdot 9$ | 172.6 | $157 \cdot 6$ | 163.7 | 168.5 | 181 \% | $170-8144 \cdot 1$ | 124.0 | $122 \%$ | 133.1 | 1313 | $183 \% 134 \cdot 9$ | 130) 1 | 1376 | 1388 | 2140 | 1.14 p.m. | $110 \cdot 3$ | 4.13 |
| 13 | $137 \cdot 3$ | $135 \cdot 81$ | 133.0 | $145 \cdot 3$ | 1474 | 150 ${ }^{\text {a }}$ | 162.5 | $180 \cdot 6$ | $172 \cdot 6$ | 185.0 | $201 \cdot 8$ | 176 | 154-3 | $164 \%$ | $140 \cdot 6133 \cdot 1$ | $1: 35 \cdot 8$ | 1 | $107 \cdot 8$ | $102 \cdot 1$ | $121 \cdot 9110 \cdot 2$ | 133-\% | 140.0 | 139.7 | 216.5 | 9.31 arm | $8 \%$ | 6.35 |
| 14 | 139.7 | 133.91 | $135 \cdot 5$ | $135 \cdot 5$ | $135 \cdot 5$ | 141.7 | 148.3 | 158.5 | $1 \mathrm{H}^{\circ} \mathrm{C} 4$ | $154 \%$ | 154 \% | 154.0 | $152 \cdot 8$ | $141 \cdot 7$ | 138.2132 .5 | 130.0 | 124.9 | $124 \cdot 1$ | $1.11^{19}$ | $134 \cdot 6136 \cdot 9$ | 13389 | $137 \%$ | 1370 | 122.3 | 8.14 | 84.5 | 6,3.35 |
| 15 | 137.0 | $145 \cdot 1$ | 139.9 | $2 \cdot$ | $138 \cdot 1$ | 138.5 | 143.5 | $145 \cdot 7$ | 151.9 | 144*8 | $131 \cdot 8$ | $139 \cdot 9$ | 147.8 | $156 \cdot 7$ | 151.0142 .9 | 130 \%6 | [128-9 | $124 \%$ | 114.2 | $81 \cdot 878 \cdot 1$ | (wis 5 | $105 \cdot 2$ | 122 | 174\% | 110.10 | $44^{-3}$ | 9.30 |
| 16 | 122.0 | $138{ }^{7} 7^{\prime} 1$ | 186.0 | 174.2 | $183 \cdot 1$ | $166 \cdot 3$ | 171'5 | $164 \cdot 6$ | 176.0 | $144 \cdot 1$ | 1379 | 127 '3 | $140 \cdot 5$ | $142 \cdot 7$ | $138 \cdot 8.136 \cdot 1$ | $142 \cdot 0$ | $133^{-9}$ | $134 \cdot 8$ | 132.8 | $131 \cdot 8135 \cdot 5$ | $134 \cdot 3$ | $134 \cdot 1$ | $136 \%$ | 2168 | 8.11 | 9 sin | (1.3) a.m. |
| 17 | $134 \cdot 8$ | 138.51 | 0.0 | 141.2 | $144 \cdot 1$ | 146 | $147 \cdot 2$ | $148 \%$ | 147.7 | $154 \cdot 1$ | $160 \%$ | 1574 | 151.0 | 141 '4 | $132 \cdot 1124 \cdot 3$ | 128 | $131 \cdot 9$ | $130 \cdot 6$ | 1315 | $132 \cdot 5134 \cdot 9$ | 133.7 | 139.0 | 134* | $171 \%$ | 10.5m | 129 | $2 . \mathrm{in}$ p.m. |
| 18 | 134.0 | 133.01 | 2.0 | $172 \cdot 6$ | $167 \%$ | 193.0 | 179 | $153 \cdot 8$ | $150 \cdot 4$ | 148.6 | 150)4 | $146 \cdot 3$ | $134 \%$ | 1264 | $126 \cdot 1122 \cdot 0$ | 12. | 1:8.8 | 127 \% | 1246 | $122 \cdot 3101 * 3$ | $91 \cdot 0$ | $93 \cdot 1$ | $125 \%$ | 2150 | 5.35 | 75.5 | 110.53 |
| 19 | $125 \cdot 9$ | $142 \cdot 11$ | 122.0 | $164 \cdot 9$ | 188.0 | 194 | 201 \% | 178.6 | 1765 | $158 \times 3$ | 181 \% | 183.7 | 180-5 | 155.0 | $133 \cdot 6143 \cdot 8$ | 131 | $113 \cdot 3$ | '115.9 | $130 \cdot 7$ | $133 \cdot 4139 \cdot 4$ | 1:990 | $139 \cdot 6$ | 180.10 | 214.4 | 5.34 | $\because 6$ | 1.3* a,m. |
| 20 | 140.0 | $139 \% 1$ | 141.8 | 142.9 | $142 \cdot 1$ | 145\% | $145 \%$ | 144.4 | 145.0 | $150 \%$ | 158.0 | $161: 3$ | $151 \%$ | $137 \%$ | $135 \cdot 1143 \cdot 8$ | 1336 | 127 3 | $122 \cdot 9$ | $119 \cdot 3$ | 130.0140-5 | $142 \cdot 7$ | $140 \cdot 5$ | 140* | 1:10\% | 11.1\% | 1100 | 5,43, $3^{3} .411$. |
| 21 | $140 \cdot 8$ | $138 \cdot 1$ | 142.0 | $143 \cdot 5$ | $145 \%$ | $144 \cdot 4$ | $144 \%$ | $147 \cdot 2$ | $145 \cdot 1^{\prime}$ | 155.5 | $161: 3$ | $160 \cdot 1$ | $154 \cdot 3$ | $142 \cdot 4$ | $132 \cdot 5126 \cdot 1$ | 123 | $131 \cdot 2$ | 124.6 | 88.9 | $100 \cdot 012 i 48$ | 146-2 | 14.8 | 143'3 | 1700 | 14.3is | N-N | -. 1 |
| 22 | $144 \cdot 5$ | $144 \cdot 2^{\prime} 1$ | $140 \%$ | $148{ }^{\circ} 8$ | 148.9 | $145 \%$ | 148.7 | 148.5 | 1519 | $154 \% 3$ | 1468 | 148.7 | $145 \cdot 4$ | 1414 | $136 \cdot 1133 \cdot 0$ | $135 \cdot 5$ | 1336 | $128 \cdot 8$ | 139.4 | $135 \times 135 \cdot 2$ | $142 \cdot 9$ | 141** | 122.6 | 154\% | 11.35 | 123:5 | -.ivi |
| 23 | $142 \cdot 6$ | $144 \cdot 1$ | 143.3 | $140 \cdot 8$ | 139.6 | $141 \cdot 8$ | $145 \cdot 4$ | 154.3 | 1594 | $167 \cdot 2$ | 1670 | $151 \%$ | 148: | $151 \%$ | $133 \cdot 3131 \cdot 3$ | $133 \cdot 1$ | 1329 | 1230 | 1378 | 130-31:99-7 | 1:3\% | 1:4\%\% | 14** | 1.015 | 8.49 | 120 | T.41 |
| 24 | 148.0 | 139 'r' | 148.3 | $145 \%$ | $152 \cdot 2$ | $158-2$ | 154.1 | $155 \cdot 5$ | $1580^{\circ}$ | $155 \%$ | $150 \cdot 1$ | 1490 | $146 \%$ | 49 | a a | a | a | a | a | A a | a | a | a | 1619 | $\therefore \times$ | 1325 | 1.int $3 . \mathrm{m}$ |
| 25 | a | a | a | a | , ${ }^{\text {a }}$ | a | a | a | a | a | a | a |  | a | $131 \% 133 \%$ | 127.4 | 134\% | 124.4 | $131 \%$ | $1311 \times 1299 \cdot 8$ | 122 | 123 | 14.4 | 154.3 | 11.24 1.m. | $116: 3$ | 10.sen n m. |
| 26 | $148 \cdot 6$ | $134 \cdot 1$ | 148.0 | 147.8 | $143 \cdot 3$ | 148.3 | 148.3 | 148.4 | 14.4 | 151.4 | $151 \%$ | $157 \%$ | $179{ }^{\circ} 6^{\prime \prime}$ | $136 \cdot 1$ | 124.313337 | $136 \cdot 1$ | 11373 | [13\% 11 | $140 \%$ | 128:5140\%8 | 1454 | 148.4 | 14:3 | 1ris.t | -. 23 ar am. | $124 \%$ | 1.1 |
| 27 | 149 \% | $145^{-5}{ }^{\text {c }}$ | $145 \cdot 9$ | 1441 | 148.0 | $150 \cdot 2$ | 154 $1^{\prime}$ | -8 | $3 \cdot$ | 9 | 208 | $7 \cdot 1$ | 289.1 | 157 '9 | a a |  |  | a | B | a a | a | a | a | 243: | C.4 |  | $2.43 \mathrm{ar} . \mathrm{m}$ |
| 28 | a |  | a | a | a | a | a |  | a | a | a | a | a | a | a 1390 | 1175 | $12 \times 9$ | 133.0 | 138.4 | 145'\% $145 \%$ | 1473 | 147: | 14:-7 | 149: | -13pm. | 1102 | $6.9 \mathrm{nm}$. |
| 29 | 147 | $151 \cdot 3$ | 149.3 | $1570^{\circ}$ | $154 \cdot 3$ | 1574 | 1571 | 154.7 | 154.3 | $183 \%$ | 1624 | $149 \cdot 3$ | $14: 5$ | 142.0 | $142 \% 14 \%$ | 1399 | $1233 \cdot y$ | 12:303 | 132: | 189-81453 | 14.8 | 1469 | 1490 | 1-3:2 | 4.140.m. | $1: 111$ | at |
| 30 | 1490 | $143 \cdot 51$ | $147 \cdot 4$ | $155 \%$ | 184.9 | $161 \cdot 5$ | $134 \cdot 9$ | 1366 | $170 \cdot 8$ | $179 \cdot 2$ | 164* | $147 \cdot 8$ | 142 见 | $144^{\prime \prime}$ | 1460142.1 | 122-8 | 1:398 | 145:3 | 145:3 | 14.115:3 | 14.7 | $145 \%$ | 14.9 | 192\% | * 40 | 12\% 4 | 8.8 |

DECLINATION, MAY, 1903.

DECLINATION, JUNE, 1903.

DECLINATION, JULY, 1903.

DECLINATION, AUGUST, 1903.

DECLINATION, SEPTEMBER, 1903.

DECLINATION, OCTOBER, 1903.



DECEMBER, 1903.

HORIZONTAL FORCE, MARCH, 1902.

HORIZONTAL FORCE, APRIL, 1902.

HORIZONTAL FORCE, MAY, 1902.







 HORIZONTAL FORCE, JUNE, 1902.
Afternoon.






HORIZONTAL FORCE, JULY, 1902.


HORIZONTAL FORCE, SEPTEMBER, 1902.

| Day. | $\mathrm{H}=06 . .$. c.a.s.s. |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Miat. | Forenoon. . |  |  |  |  |  |  | Nooni: |  |  | Afternoon. |  |  |  |  |  |  | $\mid$ | $\underset{\substack{\text { Maximum } \\ \text { reasimgand } \\ \text { time. }}}{\text { and }}$ |  | $\begin{gathered} \text { Minimum } \\ \text { readina and } \\ \text { time. } \end{gathered}$ |  |
|  | 0. |  | 3.4. | 5. |  |  | 8. 9. | 10. 11. | 12. |  | 3. | 4. |  | 7. |  |  | 10. | ${ }^{11}$ |  |  |  |  |  |
| 1 | ${ }^{64}$ | 700 | ${ }^{690}$ | ${ }^{881}$ | 690 | 70. | 716 | 723 : 723 | 723 | 719 | 727 | 738820 | ${ }^{214}$ | 711 | 704 | 705 | 708 | 306 | 705 | 746 | 0.3 p.m. | 670 | ${ }^{5} .5$ am |
| 2 | 705 | 207.705 | 701 | 702 | 704 | 700 | $707 \quad 710$ | ${ }^{74}$ | 22 | $730 \quad 737$ | ${ }_{4} 4$ | $750 \quad 719$ | ${ }^{73} 8$ | 717 | 700 | ${ }^{89}$ | 701 | 693 | 690 | $765+$ | 3.40 | 675 | 11.15 p.m. |
| 3 | 690 | 679 689 | ${ }_{672}$ 657- | 676 | ${ }_{688}$ | 689 | 703 700 | 729 740 | b | 745 | 758 | 756747 | ${ }^{3} 4$ | 718 | ${ }^{13}$ | 200 | 698 | 700 | 890 | $765+$ | 2.20 |  | - 4.0 a.n |
| 4 | 690 | 694707 | 704800 | 698 | 704 | 709 | 711815 | 718825 | 726 | 738 : 723 | 224 | 7238 | 249 | 739 | ${ }^{2} 21$ | 6988 | 203 | 710 | :06 | ${ }^{88}$ | 5.50 | 683 | 0.25 " |
| 5 | 709 | 708808 | 707 [998 | ${ }_{691}$ | 698 | 700 | 722 | ${ }_{721} 733$ | 728 | $733 \quad 753$ | ${ }^{738}$ | 723817 | 724 | 217 | 710 | \% 4 | \%u3 | 698 | 698 | $\mathrm{itb}_{6}$ | 2.4 | 683 | 4.50 . |
| - | 698 | 700804 | 705803 | 702 | 703 | 702 | 703870 | ${ }_{712}{ }^{1} 13$ | 218 | ${ }_{718}{ }^{17}$ | 720 | 72885 | ${ }^{2} 24$ | 704 | 700 | $0^{103}$ | +00 | 693 | 670 | 739 | 4.17 | 635 | 11.58 p.nn. |
| 7 | 670 | 670 \|697 | 701699 | 690 | 701 | 700 | 708808 | $718{ }^{1725}$ | ${ }^{223}$ | 731781 | ${ }^{754}$ | 73883 | 115 | 709 | 705 | 704 | 702 | 699 | 694 | \%65 | 2.44 | $633-$ | - 1.2 a.m. |
| 8 | 694 | ${ }_{694} 702$ | \%00 | 703 | ${ }^{2} 2$ | 705 | 705 705 | 710800 | 719 | 7097 | 709 | 714 | 706 | 702 | 705 | 702 | 702 | \%22 | \%02 | T22 | 0.24 | ${ }_{688}$ | 0.30 |
| 9 | 702 | 702702 | 702 | 701 | 702 | 707 | 7098 | $\tau_{2} 2$ | \%28 | $728 \quad 724$ | 722 | $724 \quad 726$ | ${ }^{224}$ | 708 | \% 2 | 201 | ${ }^{2} 1$ | 203 | 700 | ${ }^{3} 2$ | 2.53 , | ${ }^{695}$ | 5.32 .. |
| 10 | 706 | 706805 | 700705 | 703 | 698 | 698 | 705809 | 714816 | i25 | 724.721 | 710 | 708809 | ${ }^{13}$ | ${ }^{710}$ | $0^{0}$ | 708 | ${ }^{487}$ | 688 | 689 | ${ }^{730}$ | 1.1 " | ${ }_{8}^{86}$ | 10.101 pm. |
| ${ }^{11}$ | 889 | ${ }_{688} 687$ | 679685 | ${ }^{884}$ | 692 | 702 | 200807 | ${ }^{714}$, 718 | ${ }^{72 \%}$ | 73478 | ${ }^{737}$ | $7^{738} \quad 725$ | ${ }_{19}$ | 712 | ${ }^{69}$ | ${ }^{696}$ | 694 | 691 | 645 | ${ }^{746}$ | 2.24 | 869 | 2.38 a.m. |
| 12 | 695 | ${ }^{697}$; 695 | ${ }_{892} 890$ | ${ }_{691}$ | 690 | ${ }_{697}$ | 706819 | ${ }^{737}$ | 40 | 740870 | T28 | 728 , 728 | ${ }^{119}$ | ${ }^{711}$ | ${ }^{11}$ | 711 | 213 | 213 | TH0 | 751 | 0.8 | 64 | 4.30 |
| ${ }^{13}$ | 70 | ${ }^{684} 685$ | 672873 | 882 | 671 | ${ }^{89} 4$ | 711820 | 727 b | $\checkmark$ | ${ }^{\text {z32 }}$ | 754 | 743822 | 725 | 719 | 71 | 207 | 708 | T08 | 710 | ${ }^{66}$ | 2.23 | ${ }^{655}$ | 1.40 |
| ${ }^{14}$ | ${ }^{710}$ | ${ }^{009} 712$ | 712712 | 712 | 712 | 215 | 7198 | 731 | 331 | 728.724 | 725 | $723 \quad 17$ | 717 | 71.5 | 715 | 715 | 715 | ${ }^{16}$ | 717 | ${ }^{336}$ | 0.45 | iow | 1.0v .. |
| ${ }^{15}$ | 717 | 716715 | 714713 | ${ }_{74}$ | 814 | ${ }^{714}$ | $71{ }^{1} 722$ | 725 " ${ }^{\text {\% } 28}$ | 227 | 727 \% 728 | 118 | 718 | ${ }^{116}$ | ${ }^{710}$ | 212 | ${ }^{712}$ | 72 | ${ }^{73}$ | 2 | 788 | ${ }^{1.57}$. | 9 | 6.47 1 1, 12. |
| 16 | 712 | 711.710 | 711810 | 71 | ${ }^{713}$ | ${ }^{713}$ | 715.711 | 7311733 | ${ }^{441}$ | 744737 | 732 | $\begin{array}{ll}726 & 827\end{array}$ | ${ }^{24}$ | 723 | \%18 | 713 | 312 | ${ }^{13}$ | 22 | ${ }^{3} 4$ | 0.16 | ${ }^{62} 8$ | 9.13 |
| 17 | 727 | $712{ }^{214}$ | ${ }_{713} 708$ | 71 | 707 | ${ }^{113}$ | 716 | 118 |  | 781 | 761 | $\begin{array}{ll}758 & 78\end{array}$ | ${ }^{729}$ | ${ }^{727}$ | ${ }^{24}$ | 75 | 723 | ${ }^{238}$ | 727 | ${ }^{\text {isio }}$ | 3.30 | ${ }^{697}$ | a.m. |
| ${ }^{18}$ | ${ }^{227}$ | $7_{22} 717$ | 702707 | 709 | ${ }^{21}$ | ${ }^{737}$ | 733872 | 740, 347 | ${ }^{7} 35$ | 73984 | 739 | 7338 | 722 | 729 | 238 | 729 | $\because$ | ${ }^{699}$ | \% | $7^{7} 8$ | 10.32 | ${ }^{679}$ | 12.0.m.m. |
| 19 | ${ }^{998}$ | 730829 | $7277^{733}$ | ${ }^{238}$ | ${ }^{22}$ | 216 | 734 | ${ }^{250} 854$ | a | a a |  | a a | . | a | a | a | a | a |  | $i^{6} 5^{5}+$ | 11.20 | ${ }^{63}$ | 0.11 am . |
| 20 | a | a a | ${ }^{\text {a }}$ a | a | a | a | a a |  |  | 785 | $\mathrm{i}_{65}+$ | $765+{ }^{165+}$ | 762 | ${ }^{3} 5$ | 671 | 709 | 4\%80 | ${ }^{699}$ | 702 | ${ }^{6} 58$ | $\underline{2-4} \mathrm{p} . \mathrm{m}$. | 659 | Nosp.un |
| ${ }^{21}$ | 802 | ${ }^{2} 25$ | 204818 | 714 | 725 | 730 | 728 | $740 \quad 388$ | 780 | $755 \quad 765+$ | 765+\| | $785+757$ | 31 | ${ }^{736}$ | 734 | 725 | ; ${ }^{\text {a }}$ | 22 | \%1 | ;ins | $2-4$ | sis | ${ }^{11.45}$ |
| ${ }^{22}$ | 701 | 7158 | ${ }_{72}{ }^{2} / 23$ | 733 | 741 | 742 | 743 745 | ${ }^{7} 483$ | 750 | 750 746 | 447 | $743 \quad 341$ | ${ }^{7} 3$ | 335 | 7:3\% | 738 | 17 | 733 | \%,33 | 283 | ${ }^{1112,3}$ | 98 | $\cdots$ |
| 23 | 738 | 732832 | 730 | ${ }^{715}$ | ${ }^{10}$ | 22 | 72887 | ${ }^{755} 8386+$ | 768 + | $7^{768}+788+$ | \% $768+$ |  | 347 | ${ }^{738}$ | 818 | 210 | : 20 | 721 | ${ }^{183}$ |  | cinimen | six | $\checkmark$ |
| 24 | ${ }^{99}$ | 702727 | 748305 |  | ${ }^{71}$ | ${ }^{16}$ | 713 724 | 241839 | 757 | 73585 | 761 | 758788 | 7ts | 780 | 7335 | 73 | 732 | 733 | :31 | $768+$ |  | dis | (27 an. |
| 25 | 731 | 7288 | 7228.330 | ${ }^{731}$ | 734 | 738 | $\left.{ }^{736}\right\|^{736}$ | 7388 | \% | 3478 | \%43 | 338438 | ${ }^{4} 4$ | 739 | 720 | :2s | $i^{2 \times 4}$ | ies | :20 | :*3 | 1.10 | : | 1 |
| ${ }^{26}$ | i22 | 719716 | 707813 | 716 | 733 | 34 | 7451745 | ${ }^{753}$ | 757 | $767+761$ | :30 | 二83 ${ }^{\text {ans }}$ | 44 | :38 | $8: 30$ |  | \% 2 | ${ }^{7 \times 9}$ | '3 | :nition | 1.30 | is: | 230..n |
| ${ }^{27}$ | 731 | 729730 | 731728 | 723 | 720 | ${ }^{731}$ | 732 \%41 | ${ }^{743} 1705$ | 759 | 736 | ${ }^{\text {i }}$ | 836, 833 | $7^{3} 3$ | ;33 | 208 | 17 | \%1 | 819 | \%*3 | i8i+ | 0.50 | \% |  |
| ${ }^{28}$ | 74 | 707719 | 718718 | 13 | ${ }^{13}$ | 720 | $722{ }^{24}$ | 729 | 731 | 3848 | 74s | :in | :30 | : | 7 | - | * | a | , | :\%m | 2.2 | \% | (19, a.m. |
| $\because$ | " | a a | a | ${ }^{\text {a }}$ | . | - | a ${ }^{\text {a }}$ | a $7: 8$ | \% | \%as in | ${ }_{3}+1$ | and iss | :3: | :3 | \% | \% | \% | \% 2 | :2 | $\cdots$ | …s . | \% | \% 6 |
| 3, | 74 | 7238 | 722 \| ${ }^{\text {\%21 }}$ | \%19 |  | ${ }^{3} 1$ | 72783 | ${ }^{74} 138$ | 73 |  | 738 | ;313 ${ }^{123}$ | 71 | ; 3 | 816 | 214 | :13 | :11 | :\% | :* | 1.3. 4.10 | ; ${ }^{\text {a }}$ | [...... |


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HORIZONTAL FORCE, NOVEMBER, 1902.

| Day. |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Mide | Forenoon. |  |  |  |  |  |  |  |  |  |  |  | Noon. |  |  | Atteroon. |  |  |  |  |  |  |  | Midt |  |  |  |  |  |
|  | 0. | 1. |  | 2. | 4 |  | B |  |  |  | 10. | 10. 11. | 11. 12 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  | 500 |  |  |  |  |  | 357 | - 587 |  |  | ${ }^{138} 884$ | sest + ost | Sist |  |  |  | 58 |  | 183 | \%45 | 48 | 48 | ${ }_{68} 613$ | ${ }_{613}$ | 21821 |  |  |  |  |  |
|  | 621 |  | ${ }_{62}$ | 6 | 128 | ${ }^{\text {ar }}$ | ${ }^{2}$ | 6 | ${ }^{810}$ | -629 | ${ }_{628} 1837$ | ${ }_{637}{ }^{\text {eto }}$ | ${ }_{40}{ }^{\text {¢12 }}$ |  | 67 | 84 | . | 487 | \%83 + | 3+ ${ }_{6}$ | 368 | ${ }_{83} 818$ | ${ }_{616} 619$ | ${ }_{619} 612$ | ${ }^{12} 8$ |  |  |  |  |  |
| 3 | ${ }_{611}$ | ${ }_{615}$ | ${ }_{11}$ | ${ }_{\text {\%38 }}$ | 577 | 576 | 588 | 588 | ${ }^{3}$ evs | ${ }^{5} 621$ | 8218 | 83\% ${ }_{810}$ | ${ }_{40}$ ) | ${ }^{203+} \times 650$ |  |  | 4688 |  | н 4 | ${ }^{49}$ |  | ${ }_{122} 182$ | ${ }^{22}{ }^{183}$ | ${ }^{23} 824$ | ${ }^{24} 182$ |  | bist |  |  |  |
|  | 622 |  | ${ }^{23}$ | ${ }_{17} 17$ | 821 | ${ }^{83}$ | ${ }^{27}$ | 613 | ${ }^{819}$ |  | ${ }_{88} 8^{84}$ | on ext | ${ }_{01} \mathbf{H}_{68}$ |  |  |  | \%1 |  | *1 |  |  | ${ }_{627} 821$ | ${ }_{82}{ }^{182}$ | ${ }^{62} 1627$ | ${ }^{22}$ |  |  |  |  |  |
|  | 625 |  | ${ }^{19}$ | ${ }_{31}$ | 618 | 615 | ${ }^{3} 4$ | ${ }^{23}$ |  |  | ${ }_{821} 835$ | ${ }^{235} \mid 612$ | 42188 | $\left.{ }^{477}\right\|^{646}$ |  |  | ${ }_{45}$ |  | 42 |  |  | ${ }^{227}$ 428 |  | ${ }^{64} \cdot 623$ | ${ }^{23}$ | 438 | ${ }^{633}$ |  | 14.8 |  |
| - | 62 |  | *21 | ${ }_{118}$ | 817 | ${ }_{6} 18$ | ${ }_{6} 14$ | 614 | ${ }_{81}$ | \%38 | ${ }_{838} 832$ | ${ }_{83}{ }^{\text {ce38 }}$ | ${ }^{388} 86$ | ${ }^{48} / 848$ | 853 | 880 | 183 |  | ${ }_{6} 8$ | $)_{\text {*2 }}$ | ${ }^{63}$ | 833828 | ${ }^{238} 862$ | ${ }^{62}$ | ${ }^{627}$ : 82 |  | ${ }_{683}+$ |  |  |  |
| \% | 624 | ${ }^{23}$ | ${ }_{62}$ |  |  | ${ }^{\text {spo }}$ |  | 58 | ${ }_{\text {s93 }}$ |  | ${ }_{817} 1$ le8 | ${ }_{\text {c38 }}{ }^{\text {a }} 18$ | 491) | *31 ${ }_{63}$ | 648 | 1 ens+ | 350 |  | 825 |  |  | 821815 | ${ }^{115}$ 618 | ${ }^{618}$ | ${ }^{32}$; 62 |  | ${ }_{\text {cis }}+$ |  |  |  |
|  | ${ }_{625}$ | ${ }^{23}$ | ${ }^{19}$ | ${ }^{62}$ | ${ }_{80} 8$ | 827 | 820 | ${ }^{39}$ | ${ }_{3} 3$ | ${ }_{60}$ | 840838 | ${ }_{82}{ }^{2}$ ¢48 | ${ }^{48} 1638$ | ${ }_{683}{ }^{683}+$ | 683- | ${ }^{683}+$ |  |  | ${ }_{\text {e83+ }}$ |  |  | ${ }_{84} 88$ | ${ }^{87}{ }^{\text {820 }}$ | ${ }^{20}$ | ${ }^{13} \cdot 6$ |  | *303- |  |  |  |
|  |  |  | ${ }^{2} 3$ | ${ }_{63}$ | ** |  |  |  |  |  |  | $8_{80} 8$ | ${ }^{4}+$ |  | 682 |  |  |  | 48 |  |  | ${ }^{22} 888$ | 128 838 | ${ }_{822} 162$ | 82 |  | ${ }_{605}{ }^{2} 1$ |  |  |  |
| 10 | 620 | 3 | 619 | ${ }_{618}$ | 829 | ${ }^{38}$ | ${ }_{4} 8$ | ${ }^{288}$ | ${ }^{3}$ | ${ }_{839}$ | ${ }_{838} 838$ | ${ }_{835} 183$ | ${ }^{337}{ }^{183}$ | ${ }^{39}$ ) 483 | 6sast | *63+ |  |  |  |  |  | \%9 1 ¢8 | ${ }_{64} 681$ | 61 | ${ }^{20} .42$ |  | wh | wi | \% 2.5 |  |
|  |  |  | ${ }_{68}$ |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | *3\% **3 | *33 |  |  | 801+ 2.0 |  | - 7 |  |
| ${ }^{12}$ | ${ }^{1839}$ | ${ }^{388}$ | \%38 | ${ }_{4} 8$ | 84 | 4 | 158 | \%\% | ${ }_{47}$ | \% | - ${ }^{\text {a }}$ | a ${ }^{\text {a }}$ | 450 | 450 |  |  |  |  | 4, |  |  | $4_{2}{ }^{2} 12$ | A2 2 H6 |  |  |  |  |  |  |  |
|  | Q34 |  |  |  |  |  |  |  |  |  | ${ }_{\text {m90 }}{ }^{\text {a32 }}$ | ${ }^{632}$ |  |  |  |  |  |  |  |  |  | 819 | ${ }_{\text {w6\% }}{ }^{102}$ | ${ }_{612}$ |  |  |  |  |  |  |
| ${ }^{14}$ | 3645- | 583 | ${ }_{585}$ |  |  | 535 | 5 | 00 | 618 |  | 811818 | 816 |  |  | 158 |  | 673 |  | tho |  |  | ${ }_{813} 807$ | ${ }_{\text {dit }}$ | 588 |  |  | 890183 |  |  |  |
|  |  |  |  |  | ev8 |  |  |  |  |  |  | 48880 | ${ }^{50} 845$ |  |  |  |  |  |  |  |  |  | $\mathrm{SH}_{6}$ wes | ${ }^{308} \mid$ eni | S07 581 |  | 68 1.23 | St | -2: |  |
| ${ }^{16}$ | ${ }_{5} 50$ | en | 598 | ${ }^{560}$ | 611 | 3 | 43 |  | 575 |  | ${ }_{332} 537$ | 597 623 | ${ }_{22} 840$ |  | 114 |  | 635 |  |  |  |  |  | ${ }_{3}^{25}$ Wen | \%4 816 | ${ }^{116}$ |  | 6st g.lva |  |  |  |
|  |  |  |  |  | ${ }_{812}$ |  |  |  |  |  | S88 ${ }^{\text {we }}$ | 208 83 | ${ }^{23} 880$ |  |  |  |  |  |  |  |  |  | en 80 | ${ }_{615} 613$ |  |  | 701 2.27p |  | 。 |  |
|  |  | ${ }^{613}$ | ${ }_{11}$ | ${ }^{* 1}$ | 591 | 82 | ${ }^{19}$ |  |  |  | ${ }_{827} 838$ | ${ }^{839} 883$ | ${ }_{43} 845$ |  |  |  |  |  | ${ }_{812}$ |  |  |  | *2\% *22 | \%2\% 000 | no |  | \% ${ }^{\text {\% }}$ 2.42 |  |  |  |
|  |  |  |  | ${ }^{\text {se8 }}$ | 555 |  |  |  |  |  | ${ }_{887} 817$ | ${ }_{817} 824$ |  |  |  |  |  |  |  |  |  | ${ }_{804} 801$ | ${ }_{801}$ Bis | ${ }_{8 \rightarrow 0}$ |  |  | \%91 :318 |  | 3- |  |
| 20 | 591 | 588 | 592 | ${ }_{588}$ | 612 | ${ }^{23}$ | 20, | ${ }_{58} 8$ | ${ }^{508}$ |  | 5818 | 810 | 814 | ${ }^{238} 182$ |  |  |  |  | th |  |  | 812 mi | \%if wis | di3 \%8 | ${ }^{488} 80$ |  | 67\% 3.3 sa |  | 3- |  |
|  |  |  | 801 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | alo aub |  |  |  | \%090 |  |  |  |
| 22 | $598$ | 1 ssb | 591 | ${ }_{570}$ | ${ }_{\text {sab }}{ }^{\text {a }}$ | 575 | ${ }_{5 \text { sp }}$ | ${ }^{115}$ | $583-$ | 848 | x4 ${ }^{613}$ | ${ }_{813}^{618}$ | ${ }^{18} 8^{183}$ | ${ }^{34} 8$ |  |  | 452 |  | ${ }^{6} 83$ |  |  | 182480 | \%is ssi | s97 816 | 818.812 |  | -02 2.18 |  |  |  |
|  | ${ }^{12}$ | 54 | ${ }^{570}$ | 37 |  |  |  |  |  |  | *u8 598 |  |  |  |  |  |  |  |  |  |  | w6 68 | 6;2 ${ }^{182}$ | ax2 81 | 11 |  | ${ }^{\text {abe }}$ |  |  |  |
|  | 580 | 1 Se2 | sts | 375 | 558 | S88 |  | 588 | ${ }^{576}$ |  | \%0 115 | ${ }^{15}$ | 723 |  |  |  |  |  | ${ }^{4} 3$ | ${ }_{8}$ |  | \%or | \%00 ${ }^{\text {eng }}$ | en8 sex | $4{ }_{4} 3$ | 3- | :98+ |  | \%- |  |
| ${ }^{25}$ | - | S83- | - ${ }_{\text {Sn3 }}$ | ${ }^{3}-{ }^{\text {sas }}$ - | W- sss - |  |  |  |  |  |  |  |  |  |  | ${ }^{\text {in }}$ |  |  | 4 $2023+$ |  |  | \%se be | ${ }^{\text {suv }}$-3s | ${ }^{04}$ 61x |  |  |  |  | - |  |
| * | sen |  |  | ss3- | s33- | 3- sas- | S ${ }^{\text {sase }}$ | $3-583-$ |  |  |  | ${ }_{438}^{4}$ ! | ${ }^{108}$ |  |  | 887 | 063 |  |  |  |  | ass $\operatorname{sis}$ | ${ }_{\text {dx }}{ }^{618}$ | ${ }^{618}$ | 4, mex | \% | \% $2+11.3$ |  | - |  |
|  | - $0^{8}$ | d13 | si | s88- | stex- | Smb |  |  | 6s |  | ${ }_{838}{ }^{3} 85$ | ${ }_{33}{ }^{1335}$ | ${ }^{35}$ \%45 | ${ }_{4}^{4} 8632$ |  |  |  |  | ${ }_{8} 81$ |  | ats | [35 ${ }^{\text {dil }}$ | flv ${ }^{\text {an }}$ | en ${ }^{38}$ | ${ }_{38}$ +22 | $2{ }^{\text {a }}$ | a3 : $: \times$ |  | - |  |
| $\pm$ | ${ }_{82}$ | ${ }_{82}$ | 18 | 818 | Sus | ${ }^{83}$ | *1 | ${ }_{6 \times 3}$ | . |  | ${ }_{\text {w3 }}{ }^{\text {, w4 }}$ | ${ }_{94}{ }^{128}$ | ${ }^{226}$ [42 | 42 ${ }^{4} 43$ | \%4 | 187 | \%i" |  | 473 |  |  |  | *x ${ }^{38}$ | 0 | 03 | A) | axe $\quad 1.50 \mathrm{pm}$ |  |  |  |
| 20 | gip | 4 | ${ }^{619}$ | 11 | ${ }_{*}{ }^{\circ}$ |  | ${ }^{15}$ | 3 \%es | ${ }^{614}$ | 40 | \$40 448 | \%4 ${ }^{\text {wi }}$ | wi 10.105 | \$57 :871 | \$00 | ess |  |  | dis |  | "ss | .ss $\times$ | "s3 48 | at | 0 | $\cdots$ | $\cdots$ |  |  |  |
|  |  | 820 | ${ }_{82}$ | 182 | 1820 | 182 |  |  |  |  | \%4 ${ }^{\text {a }}$ \% | us 98 | 45 'wi | \$1 ** |  | Wes |  |  | 3 in |  |  |  | eno | eno us | 38 | :2a | \%as+ |  |  |  |

HORIZONTAL FORCE, DECEMBER, 1902.

HORIZONTAL FORCE, JANUARY, 1903.

HORIZONTAL FORCE, FEBRUARY, 1903.

HORIZONTAL FORCE, MARCH, 1903.

HORIZONTAL FORCE, APRIL, 1903.

HORIZONTAL FORCE, MAY, 1903.

HORIZONTAL FORCE, JUNE, 1903.


HORIZONTAL FORCE, AUGUST, 1903.

HORIZONTAL FORCE, SEPTEMBER, 1903.

HORIZONTAL FORCE, OCTOBER, 1903.




| Day. | Midt. | Forenoon. |  |  |  |  |  |  |  |  |  |  | Noon. | Afteruoon. |  |  |  |  |  |  |  |  |  |  | Midt. | $\begin{gathered} \text { Maximum } \\ \text { reasing ind } \\ \text { time. } \end{gathered}$ | $\begin{aligned} & \text { Minimutu } \\ & \text { reating ani } \\ & \text { time. } \end{aligned}$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 0. | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11. | 12. | 1. | 2. | 3. | 4. | 5. | 6. | \%. | 8. | 9. | 10. | 11. | 12. |  |  |  |
| 12 | a | a | $a$ | a | a | a | : | * | a | a | a |  | a | 560 | 576 | 581 | 568 | 585 | 568 | 530- | 520- | $520-$ | $520-$ | 520- | $520-$ | 595 t.in 1 pm. |  | -13 p.m. |
| ${ }^{13}$ | ${ }^{520}$ | $520-$ | 520- | $520-$ | 520- | $520-$ | $520-$ | 535 | 53 | 538 | 346 | 569 | b | 571 | ${ }^{626}$ | ${ }^{632}$ | ${ }^{614}$ | 581 | 579 | 571 | 551 | 533 | 520- | 530 | 530 | $651 \quad 3.4$ |  |  |
| 14 | 520- | 520- | 520- | $520-$ | 520- | $520-$ | 520- | 520- | 520- | 520 | 644 | ${ }^{621}$ | 580 | 593 | 595 | ${ }^{636}$ | ${ }^{680}+$ | 678 | 553 | $520-$ | :20- | $520-$ | 548 | 3;9 | : 40 |  |  |  |
| 15 | 540 | ${ }^{520}$ | ${ }^{\text {s20- }}$ | $520-$ | 520- | $520-$ | $520-$ | 567 | ${ }^{5}+5$ | $520-$ | 531 | 534 | ${ }^{520}$ | 541 | ${ }^{571}$ | 596 | ${ }^{15}$ | 588 | 539 | $5{ }^{5}$ | $320-$ | 320- | 520-1 | :20- | 520- | ${ }^{636}$ |  |  |
| 16 | ${ }^{520}$ | - | 52 | $520-$ | $520-$ | $520-$ | 520- | - | 595 | $520-$ | 550 | 59 | ${ }_{616}$ | 』 |  |  |  | a | a |  |  | a | $\cdots$ | a |  | 6;5 0.15 | \%0- | u- м.... |

JANUARY, 1904.
$\mathrm{H}=\mathrm{v}$. . .c.c.s.s.
Atternuon.




HORIZONTAL FORCE, OCTOBER, 1903.


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VERTICAL FORCE, MARCH, 1902.

VERTICAL FORCE, APRIL, 1902.


VERTICAL FORCE，JUNE， 1902.

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VERTICAL FORCE, JULY, 1902.


VERTICAL FORCE, SEPTEMBER, 1902.

VERTICAL FORCE, OCTOBER, 1902.


VERTICAL FORCE, DECEMBER, 1902.

VERTICAL FORCE, JANUARY, 1903.

VERTICAL FORCE，FEBRUARY， 1903.

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VERTICAL FORCE, MARCH, 1903.

VERTICAL FORCE, APRIL, 1903.

VERTICAL FORCE, MAY, 1903.

VERTICAL FORCE, JUNE, 1903.

VERTICAL FORCE, JULY, 1903.

VERTICAL FORCE, AUGUST, 1903.


VERTICAL FORCE, OCTOBER, 1903.

VERTICAL FORCE, NOVEMBER AND DECEMBER, 1903, AND JANUARY, 1904.

DECEMBER, 1903.


\footnotetext{
JANUARY, 1904.

days made use of in calculating diurnal inequalities.

(Days marked * used for "Quieter Days.")

days made use of in calculating diurnal inequalities


## DISCUSSION OF THE OBSERVATIONS.

BY
DR. C. CHREE, F.R.S.

## CHAPTER I.

## Instruments and Records.

§1. The magnetographs used in the Antarctic were of the Eschenhagen pattern, constructed by O. Toepfer und Sohn, of Potsdam. As full descriptions are readily accessible,* it is unnecessary to go into details here. The Declination, Horizontal Force and Vertical Force are recorded on a single drum. Answering to each element there is a separate base line, and there is further a record of temperature from a metallic thermometer inside the box of the Vertical-Force magnet. There are thus seven traces heing simultaneously produced on each sheet. The photographic paper has a width of nearly 20 cms , and so long as there is little magnetic disturbance and only small variations of temperature, and the sensitiveness is similar to that customary in Europe, the difficulty of keeping the traces separate and all on the sheet is not very serious. In the Antarctic, however, the conditions were much less favourable than is usual. The magnetic elements possessed unusually large variations, both regular and irregular, whilst the changes of the external atmospheric conditions were such as to cause large fluctuations of temperature even inside the hut where the magnetographs were in action. The magnetographs arrived in this country from Germany at so late a date that no time remained for the observer or the staff at Kew to become really familiar with them, and it was unfortunately not discovered that the quartz-fibre suspensions for the Horizontal-Force magnet were all so fine as to give unduly high sensitiveness.
The sensitiveness of the Declination magnetograph is determined by the greater or less distance of the magnet from the photographic paper. In the Antarctic instrument it was approximately $l^{\prime} \cdot 5$ per mm. of ordinate. Though rather low for ordinary latitudes, this proved much too high a sensitiveness for the conditions experienced.

The Vertical-Force magnet, as actually used, was by no means too sensitive; in fact, during most of the time its sensitiveness might with advantage have been at least doubled. It possessed, however, a high temperature coefficient, and this, in consequence of the large temperature changes encountered, tended to make the Vertical-Force trace shift across the paper. The large variations of temperature also gave a wide range to the trace from the metallic thermometer, which possessed a high sensitiveness.

The observer had to attempt to keep on the sheet three magnetic traces and a temperature trace, all four tending to shift their positions on the sheet, and two of them (the Declination and Horizontal-Force traces) subject to almost incessant fluctuations, which were not at all infrequently of the order of half the width of the sheet. Under these circumstances it is not surprising that traces not infrequently got off the sheet, and that those from different elements tended at times to become confused. Troubles of this kind were most serious near Midsummer, when the magnetic movements were largest.
§2. The source of light was an oil lamp which had to be filled daily. Any movement of the lamp moved the positions of all seven spots of light on the sheet simultaneously. The Vertical-Force trace, the temperature trace and the base line common to these two traces were altered together if the Vertical-Force instrument was moved. These traces could also be shifted independently by means of screws. Any movement of the Declination or the Horizontal-Force instrument altered the corresponding base line, and also, to some extent, the trace of the corresponding magnetic element. Thus the number of ways in which traces might have their position changed on the sheet was very large, while the instruments were so light that a slight touch might cause movement. As a rule, mechanical disturbances can be distinguished from true magnetic changes. The latter are seldom, if ever, absolutely sudden, and when magnetic changes occur which appear sudden on the ordinary slow-run magnetograph trace they are seldom confined to one element; they also, of course, have no effect on the base-line traces. Thus changes due to mechanical causes in the relative positions of the spots of light, though causing extra trouble to those

* 'Terrestrial Magnetiem,' vol. 5, 1900, p. 59 and Plate IV.
reducing the curves, are seldom a source of real uncertainty, provided they occur when the instrument is running. It is, unfortunately, otherwise with changes that take place when the instrument is not in action. At an ordinary station, it is true, supposing only a few minutes trace lost whilst changing the sheets, an artificial shift of a trace can usually be detected at a glance, and at least a very approximate estimate be olitained of its amount. But in tho Antarctic the conditions were not favourable for detecting artificial changes. The lamp had to be filled and trimmed, and there was not infrequently an interval of over 20 minutes between successive days' records. At times a trace was off the sheet when papers were changed. It will thus be readily understood that with elements perpetually altering-an alteration of $1^{\prime}$ per minute in the Declination was quite an ordinary one-a very appreciable artificial shift in the trace of an element might occur during the changing time without the traces themselves suggesting it. In the case of the temperature trace and the Vertical Force, the observer on a good many occasions made a shift intentionally when changing the papers. He would, for instance, find the temperature trace off the sheet and bring it back. The extent of these changes was seldom much in doubt, because readings from a mercury thermometer affordod a check on the temperature record, while the Vertical Force was seldom much disturbed.
§3. When a change was detected or suspected, the conditions were examined into by myself, and when the existence of a change was accepted, allowance was made for it after full consideration of the circumstances.

In March, April, and September, 1902, and September, 1903, there were a good many interruptions in the record. Subsequent to September, 1903, only a few records were taken, as the photographic paper was nearly exhausted. At times, in very cold weather, the oil did not burn properly, and the trace became gradually fainter and finally invisible. On one or two occasions, during a blizzard, snow got into the magnetograph room and interrupted the record, and on one occasion a magnetograph suffered through the fall of some ice which had formed on the ceiling. The only satisfactory way of treating discontinuities in the base-line values that may occur on such occasions is by reference to the absolute observations. The primary object, of course, of absolute observations at a magnetic station furnished with a magnetograph is to determine base-line values. Under ordinary conditions, if one suspects a discontinuity, one simply compares the values given by the absolute observations for the base line before and after the date of the supposed discontinuity. In the Antarctic, unfortunately, with Declination changing half a degree or more during the course of an absolute observation, the results of a single observation are of no very high precision. To obtain information as reliable as that existing, for instance, at Kew-where absolute observations are taken once a week-it would have been necessary to observe at least once a day. As will be seen, however, on reference to Commander Chetwynd's discussion of the absolute observations, "Physical Observations," p. 133, the number of absolute observations available was really small.
§4. After these remarks it is perhaps unnecessary to say that the uncertainties entering into the absolute values of the elements as given in the various tables are much greater than would be the case at an ordinary European station. They are probably greatest in the case of the Vertical Force, V. This is derived from the Horizontal Force, H, and the Inclination, I, by the formula

$$
\mathrm{V}=\mathrm{H} \tan \mathrm{I} .
$$

The absolute value of V corresponding to an observed Inclination I is deduced by multiplying tanI by the corresponding value of H , as obtained from measurement of the Horizontal Force curve at the time of the dip observation. If $v$ be the ordinate of the V curve in centimetres at this same time, and $s$ the scale value (i.e. the equivalent in C.G.S. measure of 1 cm . of curve ordinate), then the base-line value $V_{0}$ is given by

$$
\mathrm{V}_{0}=\mathrm{H} \tan \mathrm{I}-v \varepsilon_{0}
$$

H was a quantity not far from 065 , while I averaged about $84 \frac{1}{2}^{\circ}$. Thus an error of $1^{\prime}$ in $I$ means an error of some $200 \gamma(1 \gamma \equiv 0.00001$ C.G.S. $)$ in the value of $\mathrm{H} \operatorname{tanI}$. Accuracy to 0.5 in the dip derived from observations with two needles is considered good under the most favourable conditions, at places where the dip is from $60^{\circ}$ to $70^{\circ}$. What it is reasonable to expect from a dip circle in the Antarctic it is impossible to say on our present knowledge. There can, however, be little doubt that even if we neglect the uncertainties in the values of H -and the uncertainty in an absolute observation was probably nearer $1 / 500$ than $1 / 1000$ of H -the uncertainty in the base-line values of the V curves was too large to admit of discontinuities of the order of $100 \gamma$ being detected by reference solely to the base-line values.
§5. The Declination magnetograph remained at a fixed distance from the recording drum and its suspension was unaltered, so the scale value was accepted as constant. The light from the lamp, cmerging from a narrow slit, traverses a slightly convex lens forming a window to the magnet box, is reflected from a plane mirror carried by the magnet, passes a second time through the lens window, and eventually fablls on the photographic paper, after passing through a hemi-cylindrical lens a little in front of the fraper. The thickness of the hemi-cylindrical lens, approximately 0.5 cm ., and the focal length of the lens window, about 120 cm ., exercise a slight effect on the scale value; but for practical purposes it depents cssentially on the distance of the plane mirror from the photographic paper. The scale value calculated from the optical conditions in the Antarctic was $1 \mathrm{~mm} .=1^{\circ} \cdot 43$. This requires, however, a small correction for the torsion of the suspension, a very fine quartz fibre. This suspension returned safely to Europe, and independent observations by Mr. Brrnacchi and the Kew staff -in which the torsion head was twisted through $\pm 180^{\circ}$-made the torsion correction very approximately $4 \frac{1}{4}$ per cent. The scale value for the Antarctic curves was thus taken as $1 \mathrm{~mm} .=1^{\prime} 50$.
§6. The scale values of the Horizontal- and Vertical-Force magnetographs were determined from time to time in the Antarctic by means of one or other of the collimator magnets 25 A and 25 D supplied with the unifilar magnetometer. In the case of the Horizontal-Force magnetograph, the magnet was placed horizontally in front of the recording drum in the magnetic Meridian. A paper being on the drum, the position of the spot of light was recorded photographically with the magnet as originally laid down and when turned end for end. The whole operation was repeated at least once on each occasion. The double deflection, i.e., the distance between the two deffected positions of the spot of light, answers to a change of very approximately $4 \mathrm{~m} / r^{3}$ in the Horizontal Force, $r$ being the distance between the centres of the deflecting and the deflected magnets (usually at least 150 cms .), and $m$ the moment of the magnet at the time. The change in $m$ was small and slow, so that a sufficiently approximate value was derivable from the ordinary absolute observations. The procedure in the case of the Vertical Force was practically the same except that the collimator magnet was held vertically over the Vertical-Force magnet. The scale values thus found and the values actually employed are recorded in Table I.

Table I.--Scale Values.
Value of 1 cm . of ordinate in terms of $1 \gamma(1 \gamma \equiv \cdot 00001$ C.G.S. $)$.

§7. On December 31, 1902, an important change was made in the Vertical-Force magnetograph. That instrument is fitted with three auxiliary magnets. Two are thin bars, almost wires, one on either side of the Vertical-Force magnet, with their centres near its level. The third is a short, much thicker har, in a brass piece which screws on to the base of the Vertical-Force box, at some distance under the Vertical-Force magnet. These auxiliary magnets are intended apparently to assist in reducing the temperature coefficient. The use of auxiliary magnets is one of which the wisdom is open to much doubt. The results are of a somewhat complicated character, and any unrecognised change of moment or position in an auxiliary magnet may cause error and confusion. The auxiliary magnets in the present case may be described as of "soft" iron, but still they possess very appreciable "permanent" moments. In December, 1902, the temperature changes in the Antarctic were large, and Mr. Bernacchi, noticing a large temperature effect on the Vertical-Force trace, made a serious attempt to reduce it. After trying several less heroic remedies, he finally removed the short magnet from its place below the magnet box and placed it above, where it remained until the instruments were finally dismounted.

When the tabulations were commenced at Kew, it was feared that nothing could be made of the VerticalForce records. There were occasions when the instrument was obviously out of action, while on many occasions the Vertical-Force trace appeared so extraordinarily quiet compared to the others as to raise suspicions. Eventually I decided to have two months' trace measured, selecting December, 1902, and February, 1903, as months when the Vertical-Force changes seemed specially large and the information as to the scale values most complete. Even a superficial inspection of the sheets at times of large temperature change showed that the influence of temperature on the Vertical-Force trace must be considerable. But having had the 2-hour readings of Antarctic temperature through my hands, I knew that the regular diurnal variation of atmospheric temperature had a range of at most $3^{\circ}$ or $4^{\circ} \mathbf{F}^{\circ}$., and I thence inferred-wrongly, as it proved-that inside the Magnetic Hut the regular diurnal inequality of temperature would be so small even at Midsummer that its effect on the Vertical-Force diurnal inequality might be neglected without risk of serious error. I realised, of course, that this neglect might prejudice seriously individual curve readings, but for my immediate object that did not matter. Accordingly, diurnal inequalities were got out for December, 1902, and February, 1903, from the Vertical-Force curves uncorrected for temperature. The ranges appeared surprisingly large, and the inequality in the one month was inverted as compared to the other. This pointed to something being wrong. On inspecting the curves it was apparent that rise of temperature was associated with movement of the Vertical-Force trace down the sheet in 1902, but movement up the sheet in 1903. The explanation that may appear most natural-viz., that there had occurred an actual change in sign in the temperature coefficientpostulated conduct so contrary to my previous experience of magnetographs, that after reflection I inclined to the view that the Vertical-Force magnet must somehow or other have got turned end for end between December, 1902, and February, 1903. Minute inspection of the curves fixed the date of this supposed occurrence as December 31, 1902. Investigation of the written records then showed that on this afternoon Mr. Bernacchi had made the alteration mentioned above in the position of the short magnet, and as the alteration entailed a rebalancing of the magnet whose ends are closely alike, he agreed with me in regarding the occurrence as at least a possible one. Experiment showed that the magnet worked equally well whether its $N$. end were east or west. The natural way of settling the question was to refer to the information on the sheets of scale-value determinations as to the position of the deflecting magnet, whether N. pole up or S. pole up. Information on this point was given, however, only on some of the sheets, and the different sheets for 1903 contradicted one another. When making the deflection experiment, the observer had to hold the magnet close up to the ceiling in the dark, and had to rely on his memory as to how he held it. Naturally it never occurred to him that circumstances might introduce an uncertainty as to the sign even of the Vertical-Force change.

The difficulty, of course, would not have arisen if temperature coefficients had been determined in the Antarctic before and after changes of the instrument, but no determinations had been found practicable.
§8. Though I did not at first think it possible that differences so large as those presented by the diurnal inequalities found for December, 1902, and February, 1903, could arise from temperature alone, inspection of the curves led me to suspect that the regular diurnal inequality of temperature in the Magnetic Hut
during these months must have been much larger than I had supposed, and I decided that some means must be devised for eliminating its effects. The absolute sign of the Vertical-Force changes during 1903 was, as already explained, in doubt, but there was no ambiguity either in 1902 or 1903 as to the direction in which rise of temperature deflected the Vertical-Force trace. It was thus apparent that if one could determine the change of Vertical-Force ordinate answering to a rise of $1^{\circ}$ in temperature, one could-from the trace of the metallic thermometer-eliminate the effects of temperature, whatever their sign might prove to be.

Thanks to the large irregular changes of the Antarctic temperature, the task proved simpler than anticipated. The method adopted was one which I have found to work successfully in several cases. The principle is that if one can get a number of instances in which there is a large change of temperature in the course of 24 hours, a coefficient calculated by assigning the 24 -hour apparent change of Vertical Force to temperature alone will not be much in error. There are, of course, individual occasions when the true values of Vertical Force at the same hour on two successive days differ considerably; but still, if one is dealing with 20 or 30 days on which no specially large magnetic disturbance has occurred, the effects of natural magnetic changes will in most instances be very nearly eliminated. In most months there were fairly copious temperature data from the trace given by the thermometer inside the Vertical-Force box, standardised by reference to the readings of the mercury thermometer. Numerical values having been obtained for the temperature coefficient, hourly measurements were made of the temperature trace, and corrections were thus obtained to the diurnal inequalities already calculated for December, 1902, and February, 1903. Considerably to my surprise, the result was not merely a large reduction of the range in each case, but a complete inversion of the inequality for February.
§9. A difference of sign between the true inequalities of Vertical Force for December and February appearing highly improbable, I came to the conclusion that the hypothesis that the magnet bad been changed end for end on December 31, 1902, must be wrong, and that there must in reality have been a change of sign in the temperature coefficient. To obtain further light on the subject, two direct determinations were made at Kew of the temperature coefficient, the soft-iron bar being on one occasion in the position it occupied during 1902, on the other occasion in the position it occupied during 1903. The coefficients obtained differed notably in size, but they agreed in sign. At first sight this was rather staggering, but surmising the true explanation of the phenomenon, I had a further determination made under conditions as similar as possible to those at Winter Quarters. With the aid of a number of bar magnets it proved possible to produce at the position of the Vertical-Force magnet a fairly uniform vertical field of about -72 , the natural field at Kew being about $+\cdot 44$, and now, much to my relief, the expected difference in sign appeared. With the soft-iron bar below the Vertical-Force magnet, as in 1902, the trace went down the sheet as temperature was raised; with the soft iron on the top of the magnet box the reverse happened. The experiments were made with both rising and falling temperatures, the readings being taken by Mr. T. W. Baker and Mr. G. W. Walker quite independently of me, and the results appeared quite decisive in favour of the view that the Vertical-Force magnet was not altered in position in December, 1902-movement up the sheet meaning increase of force throughout-but that the temperature coefficient was negative in 1902, and positive in 1903. This view was thus finally accepted. If any additional evidence in its favour is thought necessary, it will be found in the general consistency of the results for the diurnal inequalities in Table XVIII. In several individual months the corrections from temperature to the mean diurnal inequality exert but a trifling effect on its nature, and it is out of the question that the diurnal inequalities in corresponding months of 1902 and 1903 should be the antitheses of one another.

## CHAPTER II.

## Base-Line Values, Annual Inequality, and Secular Change.

§10. Table II gives particulars of the values given by the absolute observations for the base line of the Declination curves and the mean values for individual months as calculated and used. In the calculations regard was sometimes paid only to the observations of the particular month, as in May, 1902 and 1903; but in most cases the observations in adjacent months were also considered. Thus the calculated value for June, 1902, is the arithmetic mean of the observed results on May 26 and June 30. In no case was anything taken into account except the results of the absolute observations.

The base line was assumed to have a constant value for each month. In passing from one month to the next the change in the assumed value of the base line introduces a discontinuity, the amount of which is shown in the last column of Table II. If the base-line value accepted for the second of two consecutive months is higher than that accepted for the preceding month, then the first midnight of the second month appears with a correspondingly higher Declination than the last midnight of the preceding month. The

Table II.-Values of the Declination Base Line.


* Applied also during March and April, 1902.
two hours are of course the same, and the apparent difference between the Declinations assigned to them is wholly fictitious.

The existence of such apparent discontinuities is of course undesirable. What they really imply in the present case it is impossible to say. They are contributed to, no doubt, by observational errors ; but even in the Antarctic it seems unlikely that errors of more than $5^{\prime}$ or $6^{\prime}$ would arise in absolute observations
taken by a practised observer like Mr. Bernaccin. On the other hand, it is difficult to imagine what natural cause could lead to alterations of $50^{\prime}$ or even of $30^{\prime}$ per month in the basc-line value of a Declination magnetograph, the alteration continuing for several months in the same direction. The base line, it need hardly be mentioned, is due to light reflected from a mirror, which is supposed to be fixed, on to photographic paper on a drum which is also fixed. In the Kew magnetograph it is doubtful whether the total alteration in the base line between 1890 and 1900 attained to as much as $1^{\prime}$.
§11. Table III gives the values deduced for the base line of the Horizontal-Force curves from individual observations, the values thence calculated for individual months, and the values actually used. The last column gives the excess of the base-line value for each month over that for the previous month. As already explained in the case of the Declination, the differences between the base-line values accepted for successive months appear as discontinuities-modified in April, 1903, by a change of scale value-between the values assigned respectively to the first midnight of a month and the last midnight of the previous month. As with the Declination, it is difficult to account for the large apparent variations in the base-line value. Owing to the limited number of absolute observations, observational exrors doubtless come in ; but they can hardly account for any large fraction of the larger discontinuities. The collimator magnets used for the absolute observations were old, and there was but little change in their magnetic moments, especially in that of 25 A , the magnet chiefly employed. If the absolute observations had been faultya circumstance improbable in view of Mr. Bernacchi's experience-this would have shown itself through irregularity in the values given by the individual observations for the moment of 25 A . The accordance, however, in the values for the moment is satisfactory and suggests that the probable error in individual absolute observations of Horizontal Force was at most from $10 y$ to $20 \%$. Moreover, errors in the absolute determinations would naturally vary irregularly in sign, while the apparent monthly changes in the values of the base line in Table III appear on the whole systematic.

Table III.-Values of the Horizontal-Force Base Line.

| Date. | Obscrved values. | Monthly mean values. |  | Jiscontinuity. (Unit 1 $\gamma$.) |
| :---: | :---: | :---: | :---: | :---: |
|  |  | Calculated. | Accepted. |  |
| 1902 |  |  |  |  |
| April $17 . . . . . .$. | -06371 | -06371 | -06370* |  |
| May 12. . . . . . | -06228 |  |  |  |
| June $36.0 .:$. | .06314 | .06271 | .06270 .06345 | -100 +75 |
| July $23 . . . . .$. | $\cdot 06414$ | $\cdot 06394$ | -06395 | + 50 |
| August . . . . . . |  | -06511 | -06510† | $+115$ |
| September 5. . . . . . | -06608 | -06608 | -06610 | $+100$ |
| October 21. . . . . . . | -06626 | -06626 | -06625 | $+15$ |
| November 12. . . . | $\cdot 06515$ |  |  |  |
| " 12. . . . . . | -06504 | - 06510 | . 06510 | $-115$ |
| December $27 . . . . .$. | -06247 | -06298 | -06300 | -210 |
| 1903 January 7 . . . . . . | -06207 |  |  |  |
| 30 . . . . . | off sheet | -06207 | -06205 | - 95 |
| February . . . . . . | - | -06210 | -06210 | + 5 |
| March 10. . . . . | . 06214 | -06214 | -06215 | + 5 |
| April . . . . . . | - | -06368 | . 06370 | +155 |
| May $19 . . . . .$. | -06523 | -06523 | -06525 | $+155$ |
| June $28 . . . . . .$. | -06619 | -06600 | -06600 | + 75 |
| July . . | - | -06656 | -06655 | + 55 |
| August 31 . . . . . | $\cdot 06768$ | -06731 | -06730 | + 75 |
| September . . . . . . | - | -06778 | -06780 | + 50 |
| October . . . . . . | - | -06800 | -06800 | + 20 |
| November 2. . . . . . . | -06810 | -06697 | $\cdot 06695$ | -105 |
| December . . . . . . | - | -06470 | -06470 | -225 |
| 1904 January 17 . . . . . . . | -06244 | *06244 | .06245 | -225 |

[^0]§12. Table IV shows two sets of mean values of Declination for individual months derived from the hourly readings, accepting the base-line values given in Table II. The first set of data are derived from all the days of complete registration; the second set are derived from a smaller number of days, selected as the least disturbed of the month. The differences between these two sets of values given in the last column would be unaffected by any alteration in the monthly values accepted for the base lines. The smallness of these differences seems to justify the conclusion that they are not seriously affected by the incidence of magnetic disturbances or by uncertainties in the measurements of individual days' curves. Whilst the differences are small, the all-days' mean appears in excess in so large a majority of cases as strongly to suggest that the phenomenon is not a purely accidental one. Differences between mean values derived from all and from quiet days have been observed elsewhere, though not of so large a size.

Table IV.-Mean Monthly Values of Declination (from the Curves), employing the Base-Line Values accepted in Table II.

§13. Table V gives the mean monthly values of the Horizontal Force as derived from the hourly values in days of complete registration, accepting the base-line values given in Table III.

Table V.-Mean Monthly Values of Horizontal Force (from Curves), employing the Accepted Base-Line Values in Table III.

| Month. | 1902. | 1903. |
| :---: | :---: | :---: |
| Junuary . | - | '06344 |
| February | - | -06334 |
| March . | - | -06335 |
| April . . . . . . . . | -06558 | -06534 |
| May . . . . . . . . . | $\cdot 06503$ | .06673 |
| June . . . . . . . . | '06584 | .06735 |
| July . | -06622 | -06780 |
| August . . . . | -06721 | .06838 |
| September . . . . . . . | '06718 | *06873 |
| October . . . . . . | . 06749 | - |
| November | -06627 | - |
| December . . . . . . | .06437 | - |

§14. A comparison of the mean values of an element for corresponding months of consecutivo years enables an estimate to be formed of the rate of secular change.
Table VI shows the results thus found for 1 and $H$, accepting the monthly mean values given in Tables IV and V for the five months May to September. The data subsequent to September, 1903, were too few to give representative results, while during March and April, 1902, matters were still somewhat in a preliminary stage.

Table V1.-Secular Change.


The results for the secular change in H show a rather unexpected consistency. The only suspicious feature in the figures, as figures, is the extraordinarily large size of the apparent annual change, representing as it does nearly $2 \frac{2}{2}$ per cent. of the absolute value of the Horizontal Force. If the mean is a true measure of the secular change, the natural inference is that the south magnetic Pole is receding from Winter Quarters-i.e., is moving northwards-at a rapid rate.

The Declination figures are no less remarkable, but appear much less consistent. Starting with an apparent decrease of $103^{\prime} \cdot 1$ in the year ending with May, 1903, we finish with an apparent increase of $8^{\prime} \cdot 7$ in the year ending with September, 1903.

Such a phenomenon seems hardly credible, and one cannot but suspect some instrumental source of error. In the case of the Declination, as already stated, there is no apparent reason why the base line should change, and it is conceivable that some seasonal change may have influenced the absolute observations, especially in view of the apparent inconsistencies in the azimuth readings obtained for the distant mark ("Physical Observations," p. 139). I have thus thought it worth while to ascertain what results would be obtained for the secular change of Declination if the base-line value were assumed to be invariable. On this hypothesis the results given in Table VI are replaced by those in the following Table VII.

Table VII.-Secular Change of Declination.

| Mean value from month of 1903 - Mean value from same month of 1902. |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| May. | June. | July. | August. | September. | Mean. |
| $+4^{\prime} \cdot 8$ | $+8^{\prime} \cdot 8$ | $+1^{\prime} \cdot 7$ | $+11^{\prime} \cdot 2$ | $+16^{\prime} \cdot 6$ | $+8^{\prime} \cdot 6$ |

The results given in Table VII are the antithesis of those in Table VI, and are more consistent amongst themselves. Whilst Table VI suggests that the south magnetic Pole is moving towards the west, Table VII suggests that it is moving towards the east.*
$\S 15$. In even the best European stations two years is too short a period to give results of a really representative character for the annual inequality of the magnetic elements, i.e. the variation that remains in the

[^1]mean monthly values after an allowanco has been made for the effects of secular change, assumed to proceed at a uniform rate throughout the year. This is due, at least in the case of the Declination, to the fact that the regular annual inequality in temperate Europe if not absolutely nil is exceedingly small. Table VIII gives the results obtained for the annual inequality at Winter Quarters. In obtaining the Horizontal-Force results, the mean monthly values were taken from Table V, and the secular change accepted was the mean given in Table VI. The Declination results under both (i) and (ii) are based on all the days of registration. The results under (i) accept the mean monthly values given in Table IV, and the mean value given in Table VI for the secular change, i.e., they answer to a variable base line as given by the absolute observation. The results under (ii) assume the base-line value to be invariable, and the true value of the secular change to be the mean given in Table VII.

Table VIII.-Annual Inequality.

| Month. | Declination. |  | Horizontal Force. <br> (Unit 1 $\gamma$.) |
| :---: | :---: | :---: | :---: |
|  | (i.) | (ii.) |  |
|  | ' | ' |  |
| January . | + 2 . 5 | $-28.5$ | -237 |
| February . . . . | +21.1 | -49.5 | -260 |
| March . . . . | 9.2 +94 | $-52.0$ | $-271$ |
| ${ }_{\text {May }}^{\text {April . . . . . . . }}$ | -14.6 -0.4 | -37.4 -16.6 | $+\quad 3$ $+\quad 32$ |
| Jame . . . . . . | - 8.9 | -10.8 +0.8 | + 91 |
| July . . . . . . . | $-15 \cdot 3$ | $+17.8$ | +120 |
| August . . . . . . | $-5.8$ | $+33 \cdot 7$ | +186 |
| September . | + 7.4 | +42.3 | +189 |
| October . . | + 2.9 | +44.8 | + 205 |
| November . | + 0.8 | $+37.0$ | + 71 |
| December . . . | + $0 \cdot 6$ | + 7.2 | -132 |

The range given for the annual inequality by any set of figures in Table VIII is simply enormous compared to anything that exists in ordinary latitudes. The two sets of figures for the Declination are far from similar. Those under (i) look at first sight the less improbable, as giving the smaller range. It should, however, be remembered that at Winter Quarters 1' of arc in Declination answered to only about $1 \cdot 9 \gamma$ in force, so that the range under (ii) when converted into force is only about $184 \gamma$, or 40 per cent. of the apparent range in Horizontal Force. Thus if a large range in an annual inequality is regarded as too improbable a result to be accepted, the argument is not so strong against the Declination results under (ii) as against the Horizontal-Force results.
§16. If we take a Midsummer mean from the months November, 1902, to February, 1903, and a Midwinter mean from the months May to August of both years combined, both the epochs concerned centre at January 1, 1903, so that the results are free from the uncertainty as to the real value of the secular change.

The results thus obtained are as follows :-

| Season, | Declination on hypothesis of |  | Horizontal Force. |
| :---: | :---: | :---: | :---: |
|  | Variable base line. | Fixed base line. |  |
|  | - , | , |  |
| Midsummer . . . . . . | $15250 \cdot 3$ | - | -06436 |
| Midwinter . . . . . . . | $15236 \cdot 6$ | - | -06682 |
| Excess of Midsummer . . | +13.8 | $-17 \cdot 4$ | - 000246 |

In view of what has been already said and of the difference between the Declination results under (i) and (ii) in Table VIII, it is, perhaps, unwise to say more than that it is desirable that the attention of the observers of the next Antarctic Expedition should be called to the importance of making a careful study of the annual inequality. If an inequality with a range of the order suggested by Table VIII should be established, it would be a most important result, strongly suggestive of an annual oscillation in the position of the $S$. Magnetic pole.

The question may be asked why a second set of results answering to those in Table VII and to those under (ii) in Table VIII has not been given for the Horizontal Force. The reason is simply that the Horizontal-Force base line inevitably changes with time as the moment of the suspended magnet alters, and there is the further reason that a discontinuity arose more than once through breakage of the suspension, or similar disturbing cause. A diminution in the moment at the rate of 1 per cent. per annum would have had at Winter Quarters the same apparent effect as a secular change of $65 \gamma$ per annum, and it is by no means improbable that the change of moment amounted to several per cent. as the magnet was exposed to large and numerous changes of temperature.
§17. Table IX shows the Vertical-Force base-line values observed (i.e. derived by combining observed values of dip with the corresponding values given by the Horizontal-Force curves) and those actually employed. As already explained, the probable error in the value of V (Vertical Force) derived from a single observation of dip is very large. This made it advisable to derive base-line values from the observations of a number of months combined, when this appeared feasible, allowing for apparent curve discontinuities. Thus the base-line values up to September 12,1902 , were determined by combining the absolute observations of dip made in April, May, June, July and September. The base-line values from October 1 to November 12 depend on the absolute observation of October alone. The base-line values for the latter part of November, 1902, and up to the end of January, 1903, depend on the observation of December. The base-line value for February, 1903, is from the absolute observation of that month. From March to October, 1903, the base-line values depend on the observations of March, May, June, August and September combined, allowing for apparent curve discontinuities. The base-line values for November, 1903, and January, 1904, depend on the absolute observations of these respective months, and the base-line value for December, 1903, was interpolated.

It is obvious from what has been already stated that the Vertical-Force base lines for individual months are affected by uncertainties which would render any deductions as to secular change or annual inequality of very problematical value.

Table IX.-Vertical-Force Base-Line Values.

| Month. | 1902. |  |  | 1903. |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Date of observation. | Observed value. | Accepted value. | Date of observation. | Obsorved value. | Accepted value. |
| January . | - | - | - | - | - | -7348 |
| February . . . . | - | - | - | 10 | 73407 | -7341 |
| March . . . . . | - | - | -7349 | 10 | $\cdot 72508$ | '7301 |
| April . . . . . . | 17 | - 72720 | -7349 | - | - | 27414 |
| May . . . . . | 12 | '73506 | 7331 | 19 | -73933 | -7386 |
| June . . . . | 30 | -73229 | '7302 | 28 | -73574 | -7368 |
| July . . . . . . | 23 | - 72430 | -7290 | - | - | 7357 |
| August . . . . | - | - | $\cdot 7293$ | 31 | 73740 | $\cdot 7336$ |
| September . . . . | 5 | $\cdot 73548$ | 7273 | 30 | -73711 | $\cdot 7355$ |
| October . . . | 21 | $\cdot 73111$ | .7311 | - | - | -7352 |
| *November . | 28 | '73278 | -7311 | 2 | -3348 | $\cdot 7335$ |
| December. . . . | 27 | $\cdot 73142$ | .7314 | Jan. $\overline{17}, 1904$ | .72770 | $\begin{aligned} & 7306 \\ & \cdot 7277 \end{aligned}$ |

* Of the two ralues giren for November, 1902, the first was applied up to and including the 12th, the second for the rest of the month.
§18. Table X contains data relating to the temperature correction to the Vertical Force. They were obtained by intercomparing the temperatures recorded by a mercury thermometer adjacent to the VerticalForce instrument, at the times when the sheets were put on and taken off, with the readings of the temperature trace and Vertical-Force curve at the beginning and end of each sheet. Sometimes, of course, the natural values of V at the begiming and end of a day's trace differ considerably owing to magnetic disturbance. Again, at times when temperature was changing rapidly it is improbable that the temperature of the mercury thermometer, the metallic thermometer and the Vertical-Force magnet were really identical. Uncertainties also arose from discontinuities in the temperature trace and Vertical-Force curve.

Comparing the changes during the 24 hours in the temperature trace with the corresponding changes in the readings of the mercury thermometer, one obtained the number of millimetres in the temperature trace answering to $1^{\circ} \mathrm{C}$. The mean results thus obtained from the curves of individual months form the first column of Table X. Taking everything into account, I concluded that from March, 1902, to the end of February, 1903, there was no evidence of any real change in the scale of the metallic thermometer. Accordingly the mean value 16 mm . per $1^{\circ} \mathrm{C}$. was accepted for these twelve months.

Subsequent to February, 1903, for some unknown reason, the temperature scale was much less open and at the same time more variable. This continued until nearly the end of the year, when the scale again reverted to nearly its original value. During the last three months a common value of 14.8 mm . per $1^{\circ} \mathrm{C}$. was accepted, but from March to September, 1903, use was made of the individual monthly values actually found.

Table X.-Vertical Force. Temperature Correction.


Comparing the daily changes in the temperature trace with the corresponding changes in the VerticalForce ordinate, one got for each month-assuming the effects of magnetic disturbances to neutralise one another-the apparent change of Vertical-Force ordinate in millimetres answering to a change of 1 mm . in temperature trace; while comparing the daily changes in the readings of the mercury thermometer with the corresponding changes in the Vertical-Force ordinate, one got the apparent change of Vertical-Force
ordinate in millimetres for a change of $1^{\circ} \mathrm{C}$. From a mathematical standpoint the two seta presults just mentioned are not independent, the one being deducible from the other hy means of the relationships given in the first column of Table X. In reality, however, the two sets of results were to some extent independent, because there were in all months days-sometimes a gool many days-in which information was lacking as to the mereury temperature, the temperature trace or the Vertical-Force curve, and the loss was sometimes of one element, sometimes of another. The third column of Table $X$ is in most months based on both sets of results (i.e on Vertical-Force and mercury-temperature changes as well as on Vertical-Force and temperature-trace changes), allowing most weight to the direct Vertical-Force and temperature-trace comparisons. The second and fourth columns in Table X are calculated from the third column, employing the accepted scale values of the Vertical-Force curve and the temperature trace. Any uncertainty in the scale value of the Vertical-Force curve thus affects these two columns, but it does not-and this is important-have any influence on the accuracy of the temperature correction actually applied to the readings of the Vertical-Force curve. The ordinates of the Vertical-Force curve and temperature trace were read in millimetres, and for each millimetre change of ordinate of the temperature trace a corresponding correction in millimetres was applied to the Vertical-Force ordinate. There is not the slightest doubt that the application of the temperature corrections immensely improved the accuracy of the Vertical-Force results-in fact, in some of the Midsummer months the results if uncorrected for temperature would have been practically useless-but considering the many changes of scale value and other sources of uncertainty it would be too much to hope for complete success.
§19. The uncertainties remaining after the application of the temperature correction increase, of course, with the probable error in the calculated value of the temperature coefficient, which there is no satisfactory means of estimating, but they are equally dependent on the range of temperature of the Vertical-Force magnet. So far as the regular diurnal inequality of V is concerned, it is only the regular diurnal change of temperature that counts. It is the range of this regular diurnal change that appears in the first column of Table XI. It is derived from the temperature trace on those days which were actually used in deducing the diurnal inequalities of V in Table XVIII.

Table XI.

| Month. | Range, C. | Hours of |  | Equivalent of temperature range. (Unit $1 \gamma$.) | Range of corrected V-inequality. (Unit 1\%.) |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | Highest temperature. | Lowest temperature. |  |  |
| 1902. |  |  |  |  |  |
| March . . | $2 \cdot 38$ | 4 p.m. | midnight | 97 | 37 |
| April . . . . . | $1 \cdot 10$ | 2 " |  | 48 | 46 |
| May . . . . . . . | $1 \cdot 40$ | ${ }^{7}$ ", | noon | 56 | 19 |
| June . . . . . . . | $0 \cdot 53$ | $7 \mathrm{~m} . \mathrm{m}$. | $10 \mathrm{p} . \mathrm{m}$. | 19 | 22 |
| July . . . . . . | 0.47 | 6 \% | midnight | 14 | 15 |
| August . . . . | $0 \cdot 46$ | midnight | noon | 18 | 22 |
| September . | 1.37 | 7 p.m. | 11 a.m. | 46 | 26 |
| October . . . . . . | $1 \cdot 31$ | 6 , | 6 " | 29 | 27 |
| November . . . . | - | - | - | - | 46 |
| December . . . . | 4-16 | 3 p.m. | $4 \mathrm{a} . \mathrm{m}$. | 145 | 51. |
| 1903. . . . | $6 \cdot 33$ | 3 p.m. | miduight | 135 | 80 |
| February . . . . . . | - | - | - | - | 52 |
| March . . . . . . . | $4 \cdot 31$ | $2 \text { p.m. }$ | 1 a.m. | 75 | 52 |
| April . . . . . . . | $0 \cdot 40$ | $2 \text { and } 3 \text { p.m. }$ | $9 \mathrm{p} . \mathrm{m}$. | 9 | 35 |
| May . . . . . . | $0 \cdot 37$ | $2 \mathrm{p} . \mathrm{m}$. | 7 , | 7 | 25 |
| June . . . . | $0 \cdot 35$ | noon | $1 \mathrm{a} . \mathrm{m}$. | 5 | 19 |
| July . . . . | $0 \cdot 24$ | $11 \mathrm{a} . \mathrm{m}$. and $3 \mathrm{p.m}$. | 7 and $8 \mathrm{p} . \mathrm{m}$. | 5 | 33 |
| August . . . . . . | $0 \cdot 45$ | $1 \text { p.m. }$ | $2 \mathrm{a} . \mathrm{m}$. and $6 \mathrm{p} . \mathrm{m}$. | 10 | 39 |
| September . . . . | $0 \cdot 39$ | 1,5 and 6 p.m. | 7, 8, $11 \mathrm{a} . \mathrm{m}$. and noon | 6 | 40 |
| October . . . . . . | - | 1, | 7, 8, 11 a.m. | 16 | 44 |
| November . . . . | $0 \cdot 36$ | - | - | 8 | 66 |
| December . . . . | $0 \cdot 25$ | - | - | 6 | 130 |

The second and third columns give the hours at which the highest and lowest mean temperatures occurred. The fourth column shows the equivalent of the temperature range in the first column, in terms of the correction to $\mathbf{V}$; for comparison with this, the fifth column gives the range of the corrected V diurnal inequality. To illustrate the table, take the case of December, 1902. The VerticalForce magnet had its highest temperature at $3 \mathrm{p} . \mathrm{m}$., its lowest at 4 a.m., the difference between the two being $4^{\circ} \cdot 16 \mathrm{C}$. The correction to the apparent value of V necessary to eliminate the effects of temperature was greater at $3 \mathrm{p} . \mathrm{m}$. than at $4 \mathrm{a} . \mathrm{m}$. to the extent of $145 \gamma$, or about 2.8 times the total range obtained for V after applying the correction.

In the Midwinter months the temperature range was small, and the hours of highest and lowest temperature varied much from month to month. Thus, the V-inequality results for the Midwinter months, and especially that for the whole season, should be but little influenced by temperature uncertainties. Subsequent to March, 1903, the regular diurnal variations of temperature in the Magnetic Hut were small. It is thus satisfactory that the inequalities finally deduced from the 1902 and 1903 curves agree in type.

## CHAPTER III.

Diurnal Inequalities.

§20. The questions of secular change and annual inequality have been treated out of their natural order because they appeared so closely connected with instrumental questions that it was desirable to discuss them in that connection. Before passing to the diurnal inequalities, it will be convenient to deal with some points relating to the measurement of the curves and the construction of the tables of observational data (pp. 8-70). After examining the curves, I came to the conclusion that if the value assigned to a particular element at a definite hour were derived as usual from a single measurement of the curve ordinate at the exact hour, the irregularities arising from "accidental" disturlances would probally conceal the regular diurnal inequality altogether, or at least in great measure. To meet this difficulty, it is customary at some places to smooth curves by drawing a pencil trace, giving the general trend, as distinguished from "accidental " irregularities. An alternative plan, which has been sometimes suggested, is to determine mean ordinates for successive hourly intervals of time by planimeter measurements. Owing to the excessively disturbed character of the Antarctic Declination and Horizontal-Force curves, neither of these alternatives appeared feasible. Eventually I decided not to smooth the curves in any way, but to take as the ordinate at any hour the arithmetic mean of three ordinates, one exactly at the hour, the others at 20 minutes before and after. This necessitated the measurement of the curves at 20 -minute intervals throughont the day. A suitable glass scale was constructed by Mr. Foster, one of the senior assistants in the Observatory Department. The ordinates on the scale were divided at millimetre intervals and, when reading, 0.1 mm . was usually aimed at, though this degree of accuracy is not claimed in the results. Readings were in all, or nearly all, cases repeated, usually by a different observer, and in the event of serious discrepancy a third measurement was made. Owing to the time occupied in changing papers, it was a frequent occurrence for one of the 20 -minute readings to be lost, and, at some seasons, intervals exceeding 20 minutes were not unusual. The way of dealing with these gaps varied according to circumstances. Supposing, for instance, the reading at an exact hour missing, the value assigned to that hour was usually the mean from the readings at 20 minutes before and after; but if a gap commenced or ended within a few minutes of an exact hour, the readings taken 20 minutes before and after the hour might be combined with an interpolated value, intended to represent the missing reading. The precise procedure to be adopted was determined by myself with the curves before me. In the case more particularly of the Horizontal Force, the limits of registration were exceeded rather frequently, especially in Summer. If this happened at, say, 20 minutes before the hour, the curve coming on the sheet at, say, 10 minutes to the hour and continuing on for some time, the value given in the table of hourly values was the mean from three readings which corresponded to the exact hour and to 10 minutes before and after. If, however, a limit of registration were exceeded for an appreciable time at the exact hour, the entry in the tables normally indicated an excess of the limit, precisely as it would have done if the limit had been exceeded not merely at the hour but at 20 minutes before and after as well.
§21. In the case of the Declination, measurements were always taken to 0.1 mm .—answering to $0^{\circ} \cdot 15$ -but decimals are discarded in the hourly values in the tables as not really warranted by the accuracy attainable. In the case, however, of the diurnal inequalities (pp. 90-99), the data being means from a number of readings are given to the nearest $0^{\prime} \cdot 1$.

In the Horizontal-Force tables hourly values are given to $1 \gamma$, and diurnal-inequality data (pp. 94 and 95 ) to $0.1 \gamma$. This degree of accuracy was fairly warranted, so far as mere uncertainties of reading are concerned, owing to the very open scale. It is, however, I freely admit, open to criticism on the ground that no temperature correction has been applied. What the exact degree of uncertainty on this ground may be it is difficult to say. None of the suspensions used in the Antarctic survived, and no inference seemed capable of being drawn by application of the method employed in the case of the Vertical Force. All that inspection showed was that the temperature coefficient was small, its sign even not being disclosed. Even if it had been large, the continually disturbed state of the Horizontal Force would have rendered any high accuracy in its determination impossible. We know that the rigidity of the quartz suspension
would rise with increase of temperature, while the magnetic moment of the magnet would naturally diminish; thus, presumably, the correction required was positive when temperature was above its mean.

At Kew, quartz suspensions used in the Watson type of magnetograph have had temperature coefficients of from $3 \gamma$ to $6 \gamma$ per $1^{\circ}$ C. But in the Antarctic the Horizontal Force was only about a third of that at Kew, so that correspondingly less torsion was put into the suspension, and the presumption, accordingly, is that the temperature coefficient was considerably lower than those found at Kew. The absence of a temperature correction may appreciably influence the diurnal inequality in two or three of the Midsummer months. At other seasons its influence is hardly likely to be of any importance, except in the case of individual hourly readings on days of large temperature change.

In the Vertical-Force tables the hourly values are given only to the nearest $10 \gamma$, but the absolute maximum and minimum are given to the nearest $1 \gamma$. The reasons for this are as follows:-Owing to dislocations in both the Vertical-Force and the temperature traces there were at times somewhat large uncertainties as to the relative values of the base line at different parts of the same month. It was thus felt that the retention of five significant figures in hourly values represented far more than the accuracy attainable. There was also the consideration that the scale was very contracted in most months, and the further consideration that readings to the nearest $10 \gamma$ sufficed to give diurnal-inequality data, going to $1 \gamma$. When the Vertical-Force curves were first tabulated it seemed unlikely that anything beyond diurnal inequalities would be attempted. In fact, it was not until the absolute ranges in Declination and Horizontal Force had been analysed that it was decided to attempt to obtain corresponding data for the Vertical Force. It was obvious, however, that in most cases accuracy to nearer than $10 \gamma$ was obtainable in the ranges-because the uncertainties attending absolute values of the base line largely disappear when considering differences of readings taken on the same day-and accordingly it was decided to retain five significant figures in the values of the daily maximum and minimum. It was often exceedingly difficult to say which of several very nearly equal ordinates (allowance being made for temperature) represented the true maximum or minimum. The times of occurrence were not measured very exactly.
§22. Before dealing with the data obtained for the diurnal inequalities reference must be made to one source of uncertainty to which attention has already been drawn by Commander Chetwynd ("Physical Observations," p. 134). Winter Quarters was a station at which there was appreciable local magnetic disturbance. Observations made on the ice in McMurdo Sound, about $1 \frac{3}{4}$ miles from the Magnetic Huts (see Frontispiece), gave for the Horizontal Force a value of • 0433 C.G.S., or only about two-thirds of that found at Winter Quarters. The differences in the other elements were much less. The Declination at the Ice Station was about $5^{\circ}$ less and the Inclination $13^{\circ}{ }^{\circ}$ greater than at Winter Quarters. As the Ice Station was in fairly deep water, the presumption is that these differences represent local disturbance at Winter Quarters. To the question what the effect of such local disturbances is on diurnal inequalities one cannot give a positive answer. If we suppose diurnal inequalities to be due to overhead electric currents, as Dr. Schuster's mathematical calculations indicate, and as is most generally supposed, and if no sensible diurnal variation is caused by solar radiation or similar cause in the local disturbing field, then the presence of the local disturbance will influence only the Declination diurnal inequality, the amplitude of which will vary inversely as the value of the Horizontal Force. If this view is correct, then the only effect of the local disturbance at Winter Quarters would be to reduce the amplitude of all Declination changes in the ratio approximately of $2: 3$. It must be admitted that there is little, if any, positive evidence of the correctness of the view. There is some evidence, on the contrary, that in highly disturbed regions the effect on the diurnal inequality is much more complicated. But the local disturbance at Winter Quarters was not really large. An increase of 50 per cent. in the Horizontal Force is not, of course, small from one point of view, but an increase of 02 C.G.S. does not mean anything very much out of the way at a place where the total force exceeded $\cdot 7$ C.G.S. If the source of the disturbance were basaltic rock close to the surface-as seems most probable from the description of the stationthere would not appear to be much reason to fear anything more than a reduction in the Declination range. Under such circumstances, no doubt, direct heating by the sun would have some effect at Midsummer, but it could hardly be large. Even at the surface the range of the regular diurnal inequality of temperature, when largest, was only about $4^{\circ} \mathrm{F}$.
§23. Tables XII to XXII (pp. 90 to 99 ) give the results obtained for the diurnal inequalities. In the case of the Declination two sets of results are given, the one derived from all days available, the other from the quieter only of these days. Table XII gives the all-days' Declination results for individual months. Table XIII gives the corresponding results for the twelve months of the year, taking a mean from 1902 and 1903 in cases when data from both years were available ; it also gives diurnal inequalities for the year as a whole and for three seasons, Midwinter (May to July), Equinox (March, April, September, and Octoher), and Midsummer (November to January). Tables XIV and XV give corresponding results from the quieter days. For November, 1902, an additional inequality is added in Table XIV which exclurles only the five most highly disturbed days, and so is intermediate between the inequalities based on "all" and on "quieter" days. Midwinter was limited to three months, so as to include only days throughout the whole of which the sun was below the horizon; in Midsummer, on the other hand, the sun never set. The Horizontal Force and Vertical-Force data were treated similarly to the Declination, except that no quicter days' inequalities were formed. Tables XVI and XVII relate to the Horizontal Force, Tables XVIII and XIX to the Vertical Force. In the case of the Inclination, inequalities have to be calculated from the corresponding Horizontal-Force and Vertical-Force inequalities. It appeared sufficient to give a single table, Table XX, containing results for the twelve months of the year, and the three seasons, employing data from both 1902 and 1903 when available. Tables XXI and XXII give the diurnal inequalities of the components of the Horizontal Force, respectively in and perpendicular to the astronomical Meridian; results are given only for the three seasons and the year. These were calculated from the corresponding inequalities $\Delta \mathrm{D}$ and $\Delta \mathrm{H}$ in Declination and Horizontal Force by means of the formulæ

$$
\Delta \mathrm{S}=\cos \phi \Delta \mathrm{H}+\sin \phi \Delta \mathrm{D}, \quad \Delta \mathrm{~W}=\cos \phi \Delta \mathrm{D}-\sin \phi \Delta \mathrm{H}
$$

where $\phi$ is the supplement of the easterly Declination, and $\Delta$ denotes the departure at any given hour from the mean value of the day. The mean values accepted for $\phi$ and for H , and the force equivalent to $1^{\prime}$ in Declination, were as follows:-

|  | Midwinter. | Equinox. | Midsummer. | Year. |
| :---: | :---: | :---: | :---: | :---: |
| $\phi$ | $27^{\circ} 23^{\prime}$ | $27^{\circ} 12^{\prime}$ | $27^{\circ} 13^{\prime}$ | $27^{\circ} 17^{\prime}$ |
| H : . ${ }^{\circ}$. | '06650 | '06606 | .06469 | -06573 |
| Equivalent of $1^{\prime}$ in D . . | $1.93 \gamma$ | $1.92 \gamma$ | $1.88 \gamma$ | $1.91 \gamma$ |

The working equations thence actually deduced, $\Delta \mathrm{H}, \Delta \mathrm{S}$, and $\Delta \mathrm{W}$, being measured in terms of $1 \gamma$ as unit, and $\Delta \mathrm{D}$ in terms of $1^{\prime}$, were as follows :-

|  | Midwinter. | Equinox. | Midsummer. | Year. |
| :---: | :---: | :---: | :---: | :---: |
| $\Delta \mathbf{S}=$ | $0 \cdot 89(\Delta \mathrm{H}+\Delta \mathrm{D})$ | $0.89 \Delta \mathrm{H}+0.88 \Delta \mathrm{D}$ | $0.89 \Delta \mathrm{H}+0.86 \Delta \mathrm{D}$ | $0.89 \Delta \mathrm{H}+0.88 \Delta \mathrm{D}$ |
| $\Delta \mathbf{W}=. . .$. | $-0 \cdot 46 \Delta \mathrm{H}+1 \cdot 71 \Delta \mathrm{D}$ | $-0 \cdot 46 \Delta H+1 \cdot 71 \Delta \mathrm{D}$ | $-0 \cdot 46 \Delta H+1 \cdot 67 \Delta \mathrm{D}$ | $-0 \cdot 46 \Delta \mathrm{H}+1 \cdot 70 \Delta \mathrm{D}$ |

§24. In all the diurnal inequality tables the algebraically largest and least of the hourly values are in heavy type. The tables also contain the ranges as derived from the largest and least of the hourly readings, and the sum of the 24 -hourly differences from the mean for the day.
In calculating the diurnal inequalities use was made in general only of days in which the record was complete from midnight to midnight. In some months, where the number of days of complete record was small, hourly values were interpolated when a gap of only an hour or two occurred in the record during a relatively quiet day. In some cases where material was scanty, use was made of periods of 24 successive hours which commenced at hours other than local midnight. This happened, for instance, with the Vertical Force in November and December, 1903. In the latter month the record extended from 2 p.m. on the 12 th to 1 p.m. on the 16 th, and by making the day start at 2 p.m. four complete days' record were obtained.
Particulars of the days employed for the various inequalities are given in the table on pp. 71 and 72.
Table XII．－Declination．＂All Days＂Diurnal Inequality．

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TABLE XIII.-Declination. "All Days" Diurnal Inequality.

Table XIV.-Declination. "Quieter Days" Diurnal Inequality.

Table XV.-Declination. "Quieter Days" Diurnal Inequality.

|  | Forenoon. |  |  |  |  |  |  |  |  |  |  |  | Afternoon. |  |  |  |  |  |  |  |  |  |  |  | Range. |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11. | 12. | 1. | 2. | 3. | 4. | 5. | 6. | 7. |  | 9. | 10. | 11. | 12. |  |  |
|  |  |  | , |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | , | 1 , |  |  |  |
| January | -12.3 | -3.7 | +3.2 | + 59 | +18.4 | +24.0 | $+24.0$ | +33.0 | +89.6 | +37 4 | +28.8 | +12.4 | $+2.0$ | -8.4 | -11.5 | -18.1 | -19.3 | -28.4 | -31-1 | -31.5 | -20.0 | -13.8 | 8 -13.5 | -10.1 | 71.0 | $442 \cdot 1$ |
| February | - $3 \cdot 6$ | + 2.7 | + $7 \cdot 1$ | +12.8 | +14.8 | $+9 \cdot 2$ | +19.8 | +25 11 | +25.9 | +28.4 | +17.9 | +18.4 | + 8.8 | - 4.2 | -6.3 | -19.3 | -24.3 | -25.6 | $-27 \cdot 8$ | -27.8 | -20.5 | -14.6 | 6 - 7.5 | - 7.3 | 56.0 | ${ }^{376 \cdot 3}$ |
| March | - $2 \cdot 0$ | + 1.0 | + 3.3 | + 4.2 | + $9 \cdot 6$ | +15 8 | +18.7 | +21.2 | +18 8 | +14.8 | +11.5 | +1.9 | -0.6 | -6.2 | -11.2 | -14.8 | -16.9 | $-17.2$ | -15.9 | -14.5 | -8.9 | -4.4 | 4-3.1 | -2.1 | 38.4 | $236 \cdot 6$ |
| April. | - $1 \cdot 2$ | -1.4 | + $2 \cdot 3$ | +3.0 | + 4.8 | + 6.9 | + $8 \cdot 2$ | +12.3 | +16.0 | +16.7 | +10.5 | +3.6 | -0.5 | -8.9 | -8.3 | -9.7 | -11.4 | -11.1 | -10.9 | -8.4 | $7 \cdot 3$ | - 4.5 | $5-2.1$ | -1. | $28 \cdot 1$ | $170 \cdot 4$ |
| May | + 0.2 | + 0.8 | + 1 19 | + $3 \cdot 2$ | + 3.5 | $+6.2$ | +9.5 | + 9.5 | + 78 | +6.8 | + $5 \cdot 2$ | + $3 \cdot 1$ | +0.9 | - 2.7 | -3.7 | -8.2 | -7\% | -8.6 | -8.5 | - 9.0 | -6.4 | - 4.3 | 3-2.3 | -1 | 18.5 | $120 \cdot 3$ |
| June . | +1.2 | + $1 \cdot 2$ | + 1.2 | +1.7 | + 3.6 | +3 | + 40 | +3.1 | $+5.0$ | $+4.2$ | + 2.8 | + $2 \cdot 0$ | $+0.5$ | -0.2 | -1.1 | - 2.7 | - 4.3 | - 6.8 | -3.8 | - 4.3 | - 4.4 | -4.3 | 3-2.7 | -0. | $10 \cdot 6$ | $67 \cdot 7$ |
| July | $+0.2$ | + $1 \cdot 5$ | + | +3.1 | +4.1 | + $5 \cdot 5$ | + 5.2 | + 5.8 | + | + 6.5 | +5.6 | $+4.5$ | +2.0 | $+1 \cdot 3$ | -2.2 | -4.9 | $-7.1$ | -8.6 | -8.8 | - 9.4 | - 76 | - 4.1 | $1-1.3$ | $+0.4$ | $15 \cdot 9$ | 108.4 |
| August . | + 0.9 | + $2 \cdot 3$ | + | + 4.0 | $+5.0$ | +6.4 | +80 | + | + | +6.2 | $+5.7$ | +2.5 | -0.5 | -29 | - 3.5 | - 5 \% | - 6.1 | -8.1 | - 8.0 | -8.0 | - $7 \cdot 6$ | - 4.4 | 4-2.8 | + 0.4 | $16 \cdot 1$ | 114.9 |
| September. | -1.0 | + 1.3 | + | +12.1 | +11.5 | +12.8 | +13.6 | +11.9 | +14.5 | +13.7 | +11.1 | +6.8 | -1.3 | - 5.2 | -8.1 | -10.8 | -12'8 | -15.1 | -16.7 | -14.9 | $-12 \cdot 4$ | - 9.9 | 9-1.7 | -3.5 | $31 \cdot 2$ | $234 \cdot 8$ |
| October | + $8 \cdot 3$ | +11.0 | +10.5 | $+\theta \cdot 1$ | +11.4 | +10.4 | +9.1 | +10.5 | +10 ${ }^{\text {8 }}$ | + 6.0 | + 2.4 | + 2.7 | - $2 \cdot 6$ | -8.6 | -13.2 | -15.4 | $-17 \cdot 3$ | -18.5 | -14.3 | - | -6.5 | -2.3 | $3+17$ | + 7 | $29 \cdot 9$ | $221 \cdot 2$ |
| Novermber | +3.5 | +3.3 | +2.3 | +8.4 | +1.2 | +12.6 | +28.9 | +99.0 | +29.9 | +21.8 | 8.4 | + 0.9 | - 1.9 | $-14.5$ | $-20 \cdot 6$ | -29.0 | -19.8 | -18.7 | -18.0 | $-14.5$ | 8.3 | -6.3 | $2 \cdot 3$ | +1. | 61.0 | 2934 |
| December | - 5.0 | + 1.6 | + 9.4 | +14 5 | +22.6 | +28.5 | +32.1 | +38.2 | +40.6 | +27 4 | 4.3 | +2.1 | - 5.5 | -11.8 | -23.2 | $-24.3$ | $-28 \cdot 1$ | $-22 \cdot 6$ | -28.2 | $-22 \cdot 2$ | $-17 \cdot 1$ | -13.6 | 6-10.2 | -6.5 | 88.8 | 4376 |
| Year | -0.9 | + 1.8 | + 4.5 | + 6.8 | + $9 \cdot 0$ | +11.7 | +15*0 | +17 3 | +18.1 | +15 8 | + $9 \cdot 2$ | + 4.9 | $+0.1$ | -5.8 | $-9 \cdot 5$ | -12.9 | $-14 \cdot 6$ | $-15.5$ | -15.9 | $-146$ | -10'6 | \% | -4.2 | -1 | 340 | 227.9 |
| Midwinter | + 0.5 | + 1.2 | + 1.8 | + 27 | + 4.4 | + 5.1 | + 8.2 | $+6.1$ | +6.4 | + $5 \cdot 8$ | +4.5 | + $3 \cdot 2$ | $+1.1$ | -0.5 | $-2 \cdot 3$ | -4.6 | -6.4 | $7 \cdot 6$ | -7*0 | -78 | -6.1 | $4 \cdot 2$ | $)^{-2 \cdot 1}$ | -0 | 14.0 | 98 |
| Equinox. | $+10$ | + 30 | + 8.0 | $+7.1$ | + $9 \cdot 3$ | +11.5 | +12.6 | +14.0 | +14:5 | +12.8 | +8.8 | + $3 \cdot 8$ | -12 | -6.7 | -10.4 | $-12.7$ | $-14.6$ | -15.6 | $-143$ | ${ }^{-12 \cdot 3}$ | $8 \cdot 8$ | $3 \cdot 3$ |  | + 0 - | 30.0 | 20.6 |
| Midsummer | - 4.8 | $+0.4$ | + 8.0 | + $9 \cdot 8$ | +12.4 | +21.0 | $+27 \cdot 7$ | +33.4 | +86.4 | - | $+12 \cdot 4$ | + 511 | $-1.8$ | -11.6 | $-18.4$ | $-21 \cdot 8$ | -22.4 | $-22 \cdot 6$ | -25.8 | $-22 \cdot 7$ | $-15 \cdot 1$ | -11.2 | 2 | -8.1 | 62.3 | $356 \cdot 1$ |

Table XVI.-Horizontal Force. Diurnal Inequality. (Unit $1 \gamma$.)


DIURNAL INEQUALITIES.
Table XVII.-Horizontal Force. Diurnal Inequality. (Unit 1 $\mathbf{\gamma}$.)

|  |  |  |  |  |  |  | enoon. |  |  |  |  |  |  |  |  |  |  | After | rnoon. |  |  |  |  |  |  | Sum of |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11. | 12. | 1. | 2. | 3. | 4. | 5. | 6. | 7. | m. | 9. | 10. | 11. | 12. |  |  |
| January | -14.8 | $-26.7$ | $-28.1$ | -27.9 | -30.8 | -23.0 | -14.7 | $-13.7$ | 7-20.8 | -12.2 | -3.7 | +11'7 | +21.1 | +26.5 | +25.1 | +31-4 | +83.8 | +30.9 | +18.8 | +16.2 | +10.8 | + $2 \cdot 2$ | -0.6 | -12.3 | 64.6 | ${ }^{457} \cdot 8$ |
| February | -11.8 | -13.9 | -173 | -18.4 | -18.8 | -9'6 | - 3.5 | -10'6 | 3-3.3 | +2.5 | + 76 | +10•1 | +21.1 | +82 9 | $+20.4$ | +18.8 | +18.3 | +14.1 | + | -0.9 | -0.9 | - 6 | -12.2 | -12.0 | 41.7 | 280 |
| March | -12.0 | -14.4 | -12.5 | -12.5 | -13.6 | -14.8 | - 8.7 | -4.8\| | B $+1 \cdot 1$ | + 5 . 0 | +11.8 | +13.4 | +1711 | +18.2 | +15 9 | +18.5 | +14.6 | +11.2 | + $4 \cdot 8$ | +0.1 | - 4.5 | - $7 \cdot 2$ | -10.1 | -11.2 | $31 \cdot 9$ | 254.8 |
| April. | -10.8 | -13.2 | -11.0 | $-13 \cdot 5$ | -11.6 | - | - 5.4 | -0.4 | $4+1 \cdot 3$ | +6.5 | $+10 \cdot 2$ | +13.6 | +17.5 | +18.0 | +15.9 | +12.7 | +113 | + $7 \cdot 8$ | + $1 \cdot 9$ | -2.6 | -6.7 | -8.7 | -15-9 | -10.1 | $33 \cdot 9$ | $234 \cdot 2$ |
| May | 8.8 | -8 | -9.8 | $-11.4$ | -11.7 | -8.3 | - 6.1 | -2.6 | ) $0 \cdot 0$ | +3.6 | +5.3 | + $8 \cdot 6$ | +13.6 | +15*0 | +16.5 | +14.8 | +11.8 | +90 | +3.6 | -1.0 | - $5 \cdot 9$ | - $7 \cdot 8$ | - $9 \cdot 1$ | -8.7 | $27 \cdot 2$ | 201 - |
| June. | -12.3 | -8 | -11.3 | -10.7 | - 9.4 | - 7.4 | 8 | - $3 \cdot 6$ | + 0.6 | + 4.5 | + $7 \cdot 8$ | +12.0 | +15 $\cdot 4$ | +14.919 | +16 . 6 | +12.6 | +11.1 | + $7 \cdot 8$ | +4.1 | - 2.3 | -6.0 | - 8.0 | - 8.8 | -11.7 | $27 \cdot 9$ | 214.2 |
| July | 5.7 | - 7.9 | - 78 | -7\% | - | - 73 | -5•1 | -3.4 | - $0 \cdot 6$ | $+1.0$ | + 4.3 | + 77 | +10.0 | +12.8 | +18-8 | +13.1 | +11.1 | + $5 \cdot 8$ | + $2 \cdot 0$ | -0.5 | -5.5 | - 7.2 | -8.1 | - 8.2 | ${ }_{22} 2 \cdot 3$ | $164 \cdot 1$ |
| August | -9.8 | - 75 | - 7.5 | 8.0 | - 6.8 ' | - $5 \cdot 8$ | 4 | - $2 \cdot 5$ | 0.9 | + $4 \cdot 6$ | +511 | +8.4 | +10.9 | +11-3 | +12.3 | +10.9 | + 8.4 | + 74 | +2.2 | -0.3 | - 5 ¢ | 6.7 | - $8 \cdot 8$ | $-10.7$ | 23.0 | 164. |
| September. | -12.0 | - 7.3 | 9 | -10.1 | - 8.1 | -7.2 | $\cdot 6$ |  | +1.7 | $+6 \cdot 6$ | +8.7 | +11/3 | +13.9 | $+17.0$ | +18.7 | +14.8 | + 7 :3 | + 5 .8 | + 0.3 | -4.4 | -11.7 | - 78 | $9 \cdot 0$ | -11.1 | $30 \cdot 7$ | 21 |
| Octoter. | -11.1 | -11 | $-10 \cdot 9$ | - | -9.2 | - 7 •2 | -7.5 | - $4 \cdot 31$ | -0 |  | +10.3 | +13.4 | +18.2 | +16.9 | +15'3 | +15.8 | +13.5 | 2 | + | . 5 | - $7 \cdot 5$ | -10.7 | -10'T | -11.7 | $28 \cdot 6$ | 232 |
| November . | -17.7 | -18 | -19.0 | -15\% | 14.5 | -9.0 | -13.9 | -18.3 | - - $^{\text {P }} 4$ | +3.0 | + $9 \cdot 3$ | +17.4 | +24* | +29.7 | + 31.8 | +29:3 | +18.6 | +11.9 | + 3.5 | 7 | -6.7 | -10.9 | -12\% | $-16^{11}$ | $50 \cdot 8$ | 358 |
| December . | -13.1 | -174 | -22.0 | -31.0 | -36.1 | $1{ }^{1} 1$ | $-23 \cdot 2$ | ${ }^{-14.8}$ | $8 \cdot 0$ | 15 | +12.5 | +21.6 | +30.8 | +34.3 | $+29.4$ | +29.1 | +25 2 | +17.8 |  |  | -7 | -0.8 | $0 \cdot 0$ | -12.1 | 20.4 | 439 's |
| Year | -11.7 | -13.0 | -13.9 | $-14.6$ | -14.9 | $-11 \cdot 5$ | 8.7 | -6.8 | - $2 \cdot 5$ | + $2 \cdot 8$ | + 7.4 | +12.4 | +17.7 | +19 6 | +19.2 | +18.3 | +15 4 | +11.5 | $+5 \cdot$ | $0 \cdot 0$ | -4.0 | -6.7 | 9.6 | -11-2 | 34.5 | 258 |
| M 1 dwinter . | -8.9 | 8.7 | -9.8 | -9.9 | - 9.8 |  | - 6.0 | $-3.2$ | 20 | +3.0 | $+5.8$ | $+8.4$ | +13.0 | +14.2 | +153 | +13.5 | +11.3 | + $7 \cdot 5$ |  | -1.3 | $-5 \cdot 8$ |  | 8.7 | -9.2 | 3.2 | 192 |
| Equinox | -11.5 | -11.5 | -11.1 | -11.6 | $-10 \cdot 6$ | $9 \cdot 2$ | -6.8 | $-3.0$ | +1.3 | + $5 \cdot 8$ | +10.2 | +12.9 | +16.2 | +17.0 | +16'5 | +1500 | +11\% | + $8 \cdot 2$ | +2.0 | 2.6 | - -6 | -8.6 | $-11.4$ | -11.0 | 28.6 | 233.2 |
| Midsummer | -15.2 | $-20 \cdot 8$ | -23.0 | $-24 \cdot 9$ | $-87 \cdot 1$ | -91.0 | -17.3 | $-15 \cdot 8$ | -11-1 | $-2 \cdot 6$ | $+60$ |  | +25.5 | +80 | $+28.8$ | +28.9 | +25.9 | +20.2 |  |  | $2 \cdot 3$ | -3.2 | - 5 | -13'5 | $57 \cdot 3$ | 415 |


Table XIX．－Vertical Force．Diurnal Invquality．（Unit 1\％．）

|  |  |  | 等 | 9 | $\widetilde{\%}^{\text {\％}}$ | \％̈ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{aligned} & \text { 晏 } \\ & \text { 恳 } \end{aligned}$ |  |  | \％ | $\propto$ | $\%$ | 发 |
| $\begin{aligned} & \text { 号 } \\ & \stackrel{y}{2} \\ & \stackrel{y y}{4} \end{aligned}$ | ® |  | $\stackrel{\square}{+}$ | $+$ | \＃ | $\stackrel{\square}{+}$ |
|  | $\because$ |  | $\stackrel{9}{+}$ | $\begin{aligned} & \infty \\ & + \\ & \hline \end{aligned}$ | $\stackrel{\square}{\ddagger}$ | $\stackrel{\square}{+}$ |
|  | $\stackrel{\circ}{\circ}$ |  | $\stackrel{\#}{7}$ | $\begin{aligned} & \infty \\ & + \\ & \hline \end{aligned}$ | $\pm$ | $\ddagger$ |
|  | $\cdots$ |  | 7 | $\stackrel{\infty}{+}$ | $\stackrel{\%}{+}$ | $\bigcirc$ |
|  | $\infty$ |  | $\stackrel{+}{+}$ | $\stackrel{\infty}{+}$ | $\ddagger$ | $\stackrel{+}{+}$ |
|  | $\therefore$ |  | $\stackrel{+}{+}$ | $\stackrel{+}{+}$ | $\stackrel{\infty}{+}$ | $\stackrel{\circ}{+}$ |
|  | $\stackrel{\circ}{\circ}$ |  | $\stackrel{\square}{+}$ | $\pm$ | $\stackrel{+}{+}$ | $\stackrel{\infty}{+}$ |
|  | $\therefore$ |  | 7 | $\stackrel{+}{+}$ | $\cdots$ | $\cdots$ |
|  | ＊ |  | $i$ | $\stackrel{\uparrow}{1}$ | $i$ | 7 |
|  | ฑ่ | －¢ <br> 1 <br> 1 <br> 1 | $\stackrel{\sim}{i}$ | $i$ | $\stackrel{\sim}{1}$ | $\stackrel{\infty}{1}$ |
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Table XX．－Inclination．Diurnal Inequality．

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Table XXI.-Southern Component. Diurnal Inequality. (Unit 1\%.)

|  | Forencon. |  |  |  |  |  |  |  |  |  |  |  | Afternoon. |  |  |  |  |  |  |  |  |  |  |  | Range. | $\begin{aligned} & \text { Sum of } \\ & \text { 2\& dif- } \\ & \text { ferences } \\ & \text { from the } \\ & \text { mean. } \end{aligned}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10. | 11. | 12. | 1. | 2. | 3. | 4. | 5. | 6. | 7. | 8. | 9. | 10 | 11. | 13. |  |  |
| Midwinter . | -7.8 | 6.5 | -5.5 | - $2 \cdot 9$ | -1.4 | + $1 \cdot 3$ | +3.7 | + 6.7 | +10.1 | +13.0 | +14 3 | +16.8 | +15.0 | +13•4 | +10.2 | + $3 \cdot 9$ | -0.4 | -6.8 | -10.7 | -14.1 | -15.0 | $-14 \cdot 5$ | $-11 \cdot 4$ | -9.7 | $30 \cdot 3$ | 213.7 |
| Equinox | $-10 \cdot 6$ | 7.8 | -3.8 | - 1.7 | + 2.5 | $+7 \cdot 1$ | + $9 \cdot 7$ | +14.8 | +20.6 | +22•3 | +22.6 | +19 2 | +15 0 | +911 | +2.1 | - $2 \cdot 6$ | -7.9 | -12.5 | -16.9 | -184 | -19-2 | -16.3 | -14.7 | -12.4 | 41.8 | 289.8 |
| Midsummer . | $-17 \cdot 6$ | -18.1 | -16.0 | -10-4 | -11.3 | -0.9 | + $4 \cdot 1$ | \| $14 \cdot 8$ | +20.5 | +23.8 | +20.3 | +23.6 | +25 4 | +28.7 | +12.0 | + 7 .5 | + 0.8 | - 5.8 | $-14.4$ | -16.8 | -16.6 | -16.1 | -18.1 | -15*6 | 43.5 | 331.0 |
| Year. | -12.2 | - $9 \cdot 9$ | -78 | -4.7 | - 2.6 | +3.2 | + 7 :3 | +12.8 | +18.3 | +21.0 | +20.1 | +19.8 | +18.6 | +14.1 | $+7 \cdot 3$ | $+18$ | - 3.8 | - 9.2 | $-14.9$ | $-17.3$ | $-17 \cdot 6$ | $-16.1$ | $-15.2$ | $-13.0$ | $38 \cdot 6$ | 288.5 |
|  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |

Table XXII.-Western Component. Diurnal Inequality. (Unit $1 \gamma$.)

§25. In forming the diurnal inequalities, use was made of the readings for the first and last midnights of the local day. The non-cyclic increment N -i.e. the algebraic excess of the mean value for the second over that for the first midnight-was eliminated by applying to the mean value for hour, $n$, the correction $+N(12-n) / 24$, where $n$ is counted from 0 to 24 . Particulars of the mean diurnal non-cyclic change for the days actually employed in calculating the diurnal inequalities are given in Table XXIII.

Table XXIII.-Non-cyclic Changes.


In the case of the Declination, results are given both for all days and for quieter days, i.e. for the days on which depend the inequalities in Tables XII and XIV respectively. Omitting March, 1902, for which no quieter-days' results existed, we find that the mean daily non-cyclic change of Declination from the 18 months April, 1902, to September, 1903, was $+0^{\prime} \cdot 37$ for all days, as against $-0^{\prime} .56$ for quieter days. So substantial a difference suggests that the difference between the quieter days and the others is not wholly accidental. The great preponderance of individual months in which the all-days' non-cyclic change is algebraically larger than the quieter-days' change points to the same conclusion. Accepting the difference as real, we should infer that on the representative quieter day the progressive increment of the easterly Declination is algebraically less than on the representative all-day by the amount $0^{\prime} .56+0^{\prime} \cdot 37$, or $0^{\prime} 93$. We have already seen in Table IV that the mean value of the easterly Declination for the representative quieter day was less than that for the representative all-day to the extent $1^{\prime} \cdot 46$. An analogous phenomenon has been observed elsewbere in the case of the Horizontal Force. Thus according to Ellis the average non-cyclic increment of Horizontal Force at Greenwich on quiet days from 1890 to 1895 was $+4.3 \gamma$, while the absolute value of the force on the representative quiet day of the epoch 1889 to 1896 exceeded that from the average day by $3 \cdot 3 \gamma$. The apparent parallelism may be pure accident, but it merits recognition as a possibly important key to the study of the relationship between non-cyclic changes and disturbances.

Taking the all-days' Declination results, the mean non-cyclic change from the 18 months April, 1902, to September, 1903 , as already stated, is $+0^{\prime} \cdot 37$. If we include March, 1902 , the mean falls to $+0^{\prime} \cdot 20$; while if we take a mean from a year, combining common months in 1902 and 1903 , it falls to $+0^{\prime} 083$. Even at the last figure we should deduce an annual change of $+30^{\circ} \cdot 0$, which is much in excess of the secular change obtained in Table VII. Both the non-cyclic and the Table VII secular-change data can claim to be regarded as natural magnetic changes only if the base-line value of the Declination curves was invariable. But if this be granted, it would appear at first sight that they ought to give identical values for the secular change. It should, however, be remembered that the non-cyclic data in Table XXIII are based only on days on which the record was complete, so that all highly disturbed days are excluded, whereas all days, however disturbed, contribute to the secular change. There is thus no necessary inconsistency between the resựlts of Tables XXIII and VII. If, however, we are to avoid an inconsistency, we must suppose that on the average day for which no complete Declination record existed-or we may say for the
average highly disturbed day - the non-cyclic change was opposite in sign to that on the average allday of Table XXIII, and so the same in sign as that in the average quicter day. Also the non-cyclic change on the highly disturbed days must have been numerically large, at least $-0^{\prime} 2$. It must be allowed that this result looks at first sight exceedingly improbable, as one would expect a priori that the noncyclic phenomena on quiet days and on highly disturbed days would be the opposite of one another. However, probable or improbable, the fact remains that exactly the same phenomenon has been recently discovered at Kew, where instrumental uncertainties are too small to be of importance. At Kew the noncyclic change was found on the average highly disturbed day of the 11 years 1890 to 1900 to agree in sign with that on the average quiet day, while differing in sign from that on the average day of the year (and so from the secular change). The direction of the secular change of Declination has altered at most places within historical times, and presumably will alter from time to time in the future. Thus here again we appear on the threshold of a most suggestive line of inquiry.

The mean non-cyclic change given by Table XXIII for the Horizontal Force, $-0.53 \gamma$, would give in a year of average days an apparent decrease of $194 \gamma$, whereas Tahle VI gives a secular increase of $150 \%$. No necessary inconsistency, however, exists between these figures, because the difference between them may mean nothing beyond a diminution in the magnetic moment of the Horizontal-Force magnet, which would have exactly the same effect, so far as the non-cyclic change is concerned, as a real decrease of Horizontal Force. A fall of 4.3 per cent. per annum in the moment of the Horizontal-Force magnet would suffice to wholly account for the non-cyclic change observed.

In February, 1903, the diurnal inequality of Vertical Force was based on only eight "days" or periods of 24 successive hours, and only three of these commenced at local midnight. Thus no value is assigned in that month to the non-cyclic change of Vertical Force. It was, as a matter of fact, large and negative on each of the three actual days, the average amount being $-38 \gamma$. Excluding February, 1903, there were 18 months possessing a sufficient number of complete days to be worth taking into account. The results for individual months vary much, but have a decidedly negative tendency, the mean from the 18 months being $-5 \cdot 35 \gamma$. A real change of $5 \cdot 35 \gamma$ per diem in $V$ would mean a change of 01954 per annum. An apparent decrease of $5 \cdot 35 \gamma$ per diem would, however, be equally well accounted for by a fall of about 2.7 per cent. per annum in the moment of the Vertical-Force magnet. This is really a smaller change than that required to account for the non-cyclic change observed in the case of the Horizontal Force, and it seems of by no means an improbable magnitude.
§26. The inequalities for the 12 months, the year, and the three seasons are also shown graphically in figs. 1 to 13 , pp. 104 to 116.

In the case of individual months, especially in Summer, the irregularities in the hourly values arising from disturbances catch the eye whether one looks at the diagrams or the tables. They are, however, much less prominent than one would have anticipated from a survey of the magnetic curves or of the hourly tabulations, and, when one comes to the mean seasonal results, the smoothness is not a little remarkable.

Considering the Declination inequalities in Tables XII to XV more closely, it will be seen that, so far as the general type is concerned, the differences between all and quiet days, or between the different seasons of the year, are comparatively small. There is a single daily oscillation, the extreme positions being reached about $9 \mathrm{a} . \mathrm{m}$. and $6.30 \mathrm{p} . \mathrm{m}$., the whole year round.

If in the figure NS represent the geographical Meridian, $N$ being to the north, and EW the east-west line, then if $n s$ represent the direction of the Declination needle, the angle Non was about $152^{\circ} 40^{\prime}$ on the average. It was largest about $9 \mathrm{a} . \mathrm{m}$., and least about $6.30 \mathrm{p} . \mathrm{m}$.
Though the type of the diurnal inequality did not vary much with the season, the amplitude varied largely, being greatest in Midsummer (about December), and least in Midwinter (about June). Considering, however, that the sun was continually below the horizon in Midwinter, the wonder is not so much that the range at that season was markedly less than at Midsummer as that it attained the size it actually did.


Comparing Tables XII and XIV, it will be seen that with the exception of January, 1903, all months
gave a smaller range for the quieter days than for all days, the difference being usually conspicuous, especially at Midwinter. It should also be remembered that the days of the largest disturbances did not contribute to Table XII, as the traces did not remain on the sheet, while the quieter days included about half those that were tabulated, and so cannot be regarded as representing any very high standard of Antarctic quietness. Thus the differences between the amplitudes of the regular diurnal inequalities on the very quietest and the most disturbed days must presumably be considerably larger even than those presented by Tables XII and XIV.
§27. In the case of the Horizontal Force prior to April, 1902, and subsequent to September, 1903, there were practically no days during which the registration was complete, so that diurnal inequalities are limited to the 18 months April, 1902, to September, 1903. Even between these dates there were various months, especially January and February, 1903, when only about one day in three gave a complete record. Considering this fact, the inequality curves, especially the seasonal ones, are much smoother than might have been expected. As with the Declination, the type of the diurnal inequality seems to vary but little throughout the year; the tendency is towards a single maximum about 2 or 3 p.m., and a single minimum in the early morning. The annual variation in the amplitude of the diurnal inequality is well marked. But even in June, 1902-which was not far from sunspot minimum-i.e. in the depth of the Antarctic Winter, the amplitude exceeded $20 \gamma$.

As to the relative sizes of the diurnal changes in Declination and Horizontal Force, a change of $1^{\prime}$ in $\mathbf{D}$ answered roughly to a force of $1.91 \gamma$, so that taking the mean results for the year we have for the amplitudes $45^{\prime} \cdot 5$ or $87 \gamma$ in D as against $35 \gamma$ in H . It must be remembered, however, that a good many days contributed to the D inequalities which were too disturbed to contribute to the H , and if these days had been excluded from the D inequalities the amplitudes of these would probably have been reduced. Still, even if we confined ourselves to the quieter days of Table XV , we should obtain for D , from the year as a whole, an average daily range of $34^{\prime}$ or $65 \gamma$. This is nearly double the corresponding value for $H$. It is thus certain that the forces to which the diurnal inequality is due are less potent in the plane of the magnetic Meridian than in the perpendicular plane.
§28. The Vertical-Force trace itself seldom went off the sheet, but loss of temperature trace was less uncommon. Also in some months a good many days' traces which were complete were omitted because there was reason to doubt whether the instrument was working. In some cases the magnet was evidently stuck. The space in which it moved gave very little clearance, and presumably the magnet-whose azimuth was approximately east and west magnetic-tended to shift in azimuth, and so came in contact with the adjacent metal. In other cases it was impossible to feel certain from mere inspection of the trace that the magnet was stuck, while the extreme quietness of the trace suggested that it was. In such cases one was generally able to settle the question by reference to the temperature trace. If temperature changed sensibly-and it usually did-and the Vertical-Force trace showed no sympathetic movement, then one inferred that the magnet was stuck, and rejected the day's record.

Table XVIII gives inequalities for all the individual months from March, 1902, to December, 1903. The inequalities, however, for November and December, 1903, are each based on only 4 days' records, and the temperature-correction data for these were more than usually uncertain. Thus little weight has been allowed these as compared to the corresponding months in 1902 when forming the inequalities in Table XIX.

There is an apparent difference, clearly visible in fig. 11, between the types of the diurnal inequality of Vertical Force at different seasons. At Midsummer the force, instead of falling continuously from a maximum near midnight to a minimum near noon, shows a tendency to a slight increase about 6 to 8 a.m., and there is at least a suspicion of the same phenomenon in the equinoctial months. The phenomenon may be a real one, but it ought to be mentioned that disturbances in $V$ were more prevalent from 4 to 9 a.m. than at other hours, and their satisfactory elimination might require several years' data.
§29. The diurnal inequalities of Inclination in Table XX were calculated from the corresponding V and H inequalities by means of the formula

$$
\delta \mathrm{I}=\delta \mathrm{V} \times \cdot 00441-\delta \mathrm{H} \times \cdot 0477
$$

Here $\delta \mathrm{I}, \delta \mathrm{V}, \delta \mathrm{H}$ denote the departures at any, the same, hour from the mean values of $\mathrm{I}, \mathrm{V}$ and H for the day. The numerical multipliers "00441 and • 0477 answer to the mean values of $V$ and $H$ during the stay
at Winter Quarters. The contribution from H to $\delta \mathrm{I}$ is usually much the larger. Thus the curves in fig. 12 have a considerable resemblance to those in fig. 10 , with the sign of the ordinates reversed. The Midsummer curve, however, in fig. 12 shows a slackening in the rate of fall of I about 7 or 8 a.m., which is due to the influence of the V contribution.

Winter Quarters was only about 400 miles from the position which Commander Chetwynd has deduced for the south magnetic Pole. Thus, if the diurnal inequality is due to electrical currents in the upper atmosphere, we should on the whole expect no great difference to exist between the amplitudes of the diurnal inequality of I at Winter Quarters and in the immediate neighbourhood of the Pole itself. This would imply, of course, a considerable diurnal variation in the actual position of the I'ole. The probable nature of this movement may be derived from the consideration that an increased Inclination at Winter Quarters is presumably equivalent to a diminution in its distance from the magnetic Pole. Inclination was largest at Winter Quarters about 4 or $5 \mathrm{a} . \mathrm{m}$. in the morning, so that the magnetic Pole was presumably at that hour displaced towards Winter Quarters, i.e. was to the SE of its mean position for the day. Inclination at Winter Quarters was lowest from 2 to $3 \mathrm{p} . \mathrm{m}$., so that the magnetic Pole at that hour is probably to the NW of its mean position for the day.
§30. From a mathematical standpoint much is to be said for treating the diurnal variation of the components of foree in and perpendicular to the geographical Meridian as fundamental rather than those of the Declination and Horizontal Force. The arguments in favour of this course are in some respects stronger for Winter Quarters than for an ordinary station, because the magnetic Meridian there was so highly variable. Again, if the causes to which the diurnal inequality is due are related to the Earth's axis of rotation rather than to the position of the magnetic Pole, the diurnal inequalities of the components S and W, directed to geographical south and west, are likely to be much less variable round a parallel of latitude than those of D and H . If the orientation of the spot relative to an adjacent magnetic Pole is of none, or of but secondary importance, the diurnal inequalities of D and H at Winter Quarters probably differ largely from those at other stations only 50 or 60 miles to the west, whereas the diurnal inequalities in $S$ and $W$ are unlikely to differ much. On the other hand, any element of uncertainty that may attach to the absolute value of the Declination enters into the diurnal inequalities of $S$ and $W$, whereas it hardly enters into the diurnal inequality of D itself. Another drawback is that the S and W inequalities each depend on both D and H , whereas the days contributing to the D and H inequalities differ to some extent. If the $D$ and $H$ contributions to $S$ and $W$ had been derived strictly from the same days, the results would probably have been slightly different from those actually obtained. This last source of uncertainty might of course have been avoided, but, considering all the circumstances, it seemed hardly worth while to recalculate D and H inequalities from a common series of days.

The methods of calculating the inequalities in Tables XXI and XXII have been already explained in §23. Both $S$ and $W$ show only a single daily period. The extreme values of $W$ occur distinctly earlier in the day than those of S . As appears either from the ranges or the sum of the hourly differences from the mean, the amplitude of the daily changes of W is more variable throughout the year than is that of S .

The inequalities of $S$ and $W$ are combined and shown graphically in the vector diagrams of figs. 13. In these NS and EW are drawn respectively in and perpendicular to the geographical Meridian, and the ends of the lines denoted by the letters $N, S, \& c$., are each at a distance from the point of intersection which represents $10 \gamma$ on the scale to which all the diagrams are drawn. The crosses and the numbers attached answer to the hours of the day counted from 0 , local midnight. The line drawn from the intersection of NS and EW to a particular cross represents in direction and magnitude the horizontal component of the forceacting on the north pole of the magnet-to which may be ascribed the departure of the magnetic elements at the hour indicated from their mean value for the whole 24 hours. The diagrams are all described anti-clockwise, and so in the opposite direction to corresponding diagrams for English stations.

Being based on less than 2 years' data, the diagrams are naturally not very smooth, but their general form is remarkably symmetrical. The difference of type between the diagrams for the different seasons is unusually small. Another exceptional feature is the comparative smallness of the difference between day and night, the angular velocity of the vector being only very slightly less during the 12 hours centring at midnight than during the corresponding "day" hours.


Fig. 1. Declination.

DIURNAL INEQUALITIES.


Fig. 2. Declination.


Fig. 3. Declination.


Fig. 4. Horizontal Force. (Cnit ly.)
[ :


Fig. 5. Horizontal Force. (Unit 1 $\gamma$.)



Fig. 7. Vertical Forve. (Unit 1\%.)


Fig. 8. Inclination.


Fig. 9. Declination.


Fig. 10. Horizontal Force. (Unit ly.)


Fig. 11. Vertical Force: (Unit 1 $\gamma$.)


Fig. 1\%. Inclination.


Fig. 13. Vector Diagrams

## CHAPTER IV.

## Diurnal. Inequalities. Fourier Confficifats.

§31. From the diurnal inequalities, calculations were made of the Fourier coefficients answering to the "waves" whose periods are $24,12,8$ and 6 hours. The analysis of the diurnal inequality may be supposed to proceed according to either of the two equivalent series

$$
\begin{aligned}
& a_{1} \cos t+b_{1} \sin t+a_{2} \cos 2 t+b_{2} \sin 2 t+ \\
& c_{1} \sin \left(t+a_{1}\right)+c_{2} \sin \left(2 t+a_{2}\right)+.
\end{aligned}
$$

Here $t$ is time, counted from local midnight, one hour being taken as equivalent to $15^{\circ}$. The constants with suffix 1 refer to the 24 -hour term, those with suffix 2 to the 12 -hour term, and so on. The $c$ and $b$ constants are calculated directly from the inequality tables. The mean $a_{1}$, for instance, for a particular season of the year is the arithmetic mean of the values of $a_{1}$ for the months composing that season. The $c$ (amplitude) and $a$ (phase angle) constants are calculated from the corresponding $a$ and $b$ constants ly means of the formulæ

$$
\alpha=\tan ^{-1}(a / b), \quad c=a / \sin \alpha=b / \cos \alpha
$$

The $c$ (or a) derivable from a seasonal diurnal inequality is not, as a rule, the arithmetic mean of the $c$ 's (or $\alpha$ 's) of the individual months which form the season.
§32. Tables XXIV to XXXII are devoted to the $a$ and $b$ coefficients. These were in all cases really calculated to at least one figure further than is retained. It is, however, hardly necessary to remark that even as thus curtailed they cannot be regarded as physical facts freed from observational uncertainties. This reservation ought especially to be borne in mind in the case of the coefficients with suffixes 3 and 4, which relate to the 8 -hour and 6 -hour waves. The differences between the values obtained for successive months probably owe at least as much to the existence of "accidental" disturbances as to any real difference between the magnetic conditions characteristic of successive months of an average year.

In the case of the Declination, Horizontal Force and Vertical Force the values of $a$ and $b$ are recorded for the individual months of the two years, as well as for the months of a representative year in which common months of 1902 and 1903 are combined. Also two sets of values are given for Declination. Of these the first set, comprising Tables XXIV and XXV, relate to the all-days' inequality data of Tables XII and XIII ; while the second set, comprising Tables XXVI and XXVII, relate to the quieter-days' inequality data of Tables XIV and XV. For Inclination only one table is given. When the sign to be attributed to the numerical value of a constant is the same for each month and season it is indicated only at the top and bottom of the column.

Table XXIV.—Declination (All Days). Fourier Coefficients. (Unit 1'.)


Table XXV.-Declination (All Days). Fourier Coefficients. (Unit 1'.)


Table XXVI.-Declination (Quieter Days). Fourier Coefficients. (Unit 1'.)


Table XXVII.-Declination (Quieter Days). Fourier Cocfficients. (Unit 1'.)


Table XXVIII.-Horizontal Force. Fourier Coefficients. (Unit 1 1 .)


Table XXIX.-Horizontal Force. Fourier Coefficients. (Unit $1 \gamma$.)


Table XXX.-Vertical Force. Fourier Coefficients. (Unit 1\%.)


Table XXXI.-Vertical Force. Fourier Coofficients. (Unit ly.)

|  | $a_{1}$. | $b_{1}$. | $a_{2}$. | $b_{2}$. | $a_{3}$. | $b_{3}$. | $a_{4}$. | 4. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| January | + $20 \cdot 11$ | $+18 \cdot 89$ | -11.98 | -3.11 | + $7 \cdot 3 \cdot 1$ | +1.78 | $-3 \cdot 35$ | + $3 \cdot 29$ |
| February . | $17 \cdot 60$ | + 8.02 | $3 \cdot 34$ | -6. 23 | -0.11 | $+5 \cdot 65$ | -1.13 | -0.06 |
| March . | $18 \cdot 09$ | + 4.28 | $4 \cdot 38$ | -4.91 | +2.30 | 0 () () | -1.04 | +0.47 |
| April | $15 \cdot 88$ | $-6.63$ | 162 | -0.28 | $+1 \cdot 45$ | +0.63 | -1.07 | -0. 49 |
| May . | $7 \cdot 00$ | $-1.40$ | $0 \cdot 35$ | -0.72 | +0.81 | +0.39 | $+0 \cdot 12$ | +1) $\cdot 1(1)$ |
| June . | $8 \cdot 43$ | $-2 \cdot 16$ | $0 \cdot 96$ | $-0.27$ | -0.48 | +0.69 | -0.27 | +0.28 |
| July . | $9 \cdot 33$ | - 5.82 | $1 \cdot 42$ | -0.99 | -0. 0 | -0.10 | +0.23 | -0.38 |
| August. | $11 \cdot 54$ | $-6.24$ | 1.01 | +0.01 | +0.40 | $+0 \cdot 9.4$ | -1.37 | -0.13 |
| September . | $14 \cdot 10$ | - 0.73 | 1199 | $-1 \cdot 31$ | +0.27 | +0.92 | -0.13 | +0:31 |
| October. . | $13 \cdot 86$ | + 1.12 | $0 \cdot 94$ | -3.82 | $+168$ | $+1 \cdot 15$ | -1 10 | +0.22 |
| November. | $22 \cdot 07$ | - 3.81 | $3 \cdot 46$ | $+0 \cdot 67$ | +0.52 | +1.97 | -0. 29 | -0) ()2 |
| December . | $25 \cdot 85$ | + 4.53 | $5 \cdot 07$ | +0.41 | $+2 \cdot 43$ | +1.96 | -2.14 | -0 39 |
| Year. . | $15 \cdot 32$ | + 0.85 | $3 \cdot 04$ | -1 71 | +1 34 | +1.33 | -0.96 | +0.29 |
| Midwinter. | $8 \cdot 25$ | - $3 \cdot 13$ | $0 \cdot 91$ | -0 66 | -0.03 | +0:33 | +0.03 | +0.10 |
| Equinox - | $15 \cdot 48$ | $-0.45$ | $2 \cdot 22$ | -2.58 | +1.42 | + $0 \cdot 68$ | -0.84 | +0.13 |
| Midsummer . | +22.67 | + 6 .54 | $-6.84$ | -0.68 | $+3 \cdot 43$ | +1.90 | -1.92 | +0.96 |

Table XXXII.—Inclination. Fourier Coefficients. (Unit 1'.)

|  |  | $a_{1}$. | $b_{1}$ 。 | $a_{2}$. | $b_{2}$ 。 | $a_{3}$. | $b_{3}$. | $a_{4}$. | $b_{4}$. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| January. | - . . . - | +0*295 | + 1 :394 | +0.031 | -0.098 | -0.045 | +0.134 | -0.142 | -0.056 |
| February | . . . . . | $0 \cdot 709$ | $0 \cdot 648$ | -0.033 | $0 \cdot 166$ | -0.053 | +0.062 | -0.044 | -0.060 |
| March . | . . . . | $0 \cdot 680$ | $0 \cdot 506$ | -0.076 | $0 \cdot 108$ | +0.018 | -0.063 | -0.009 | +0.023 |
| April. | . . . . . | 0.702 | $0 \cdot 319$ | -0.052 | 0-104 | -0.001 | -0.020 | -0.018 | -0.050 |
| May . | - . . . | 0.516 | $0 \cdot 375$ | -0.013 | 0-129 | -0.049 | -0.030 | +0.002 | -0.028 |
| June. . | . . . . . | $0 \cdot 576$ | $0 \cdot 347$ | -0.028 | $0 \cdot 123$ | +0.005 | -0.012 | -0.009 | -0.003 |
| July . | . . . . . | 0.431 | $0 \cdot 285$ | -0.012 | $0 \cdot 151$ | -0.046 | -0.021 | +0.005 | -0.016 |
| August . | . . . . | $0 \cdot 475$ | $0 \cdot 232$ | +0.019 | $0 \cdot 102$ | $+0 \cdot 019$ | -0.023 | +0.011 | +0.003 |
| September . | . . . . | $0 \cdot 632$ | $0 \cdot 281$ | -0.063 | $0 \cdot 171$ | -0.038 | -0.002 | +0.020 | +0.012 |
| October . . | . . . . . | $0 \cdot 660$ | 0.359 | -0.036 | $0 \cdot 178$ | +0.005 | -0.027 | -0.083 | +0.026 |
| November | . . . . | 0.912 | $0 \cdot 695$ | $-0.066$ | $0 \cdot 325$ | +0.004 | $+0 \cdot 115$ | -0.015 | +0.070 |
| December | . . . . . | 0.914 | 1-149 | -0.319 | 0.170 | -0.024 | -0.028 | +0.023 | -0.076 |
| Year | - - . - | $0 \cdot 650$ | $0 \cdot 549$ | -0.054 | $0 \cdot 152$ | -0.017 | $+0.007$ | -0.017 | -0.013 |
| Midwinter. | . . . . | $0 \cdot 508$ | $0 \cdot 336$ | -0.018 | $0 \cdot 134$ | -0.030 | -0.021 | -0.001 | -0.015 |
| Equinox.' | . . . . | $0 \cdot 669$ | $0 \cdot 366$ | -0.057 | $0 \cdot 140$ | -0.004 | -0.028 | -0.010 | $+0.003$ |
| Midsummer | . . . . . | +0.807 | +1.079 | -0.118 | -0.198 | -0.022 | +0.074 | -0) 045 | -0.021 |

§33. Tables XXXIII to XL are devoted to the $c$ and $\alpha$ (amplitude and phase angle) Fourier coefficients in the diurnal inequality as derived from the $a$ and $b$ coefficients in Tables XXIV to XXXII. The tables correspond in pairs, except that there is no $c, \propto$ table for Vertical Force corresponding to Table XXX. If the phase angle a varies much throughout the year, the contributions from the individual 12 months to the mean diurnal inequality for the whole year tend to neutralise one another, and so the amplitude of the mean inequality for the year is apt to suggest that the corresponding wave is smaller throughout the year than it actually is. Some writers thus prefer the arithmetic mean of the c's and a's from the 12 monthly data for a particular wave to the corresponding results from the mean diurnal inequality of the whole year. This is the reason why Tables XXXIV, XXXVI, XXXVIII, XXXIX, and XL contain
arithmetic means in addition to the values of $c$ and $a$ derived from the mean diurnal inequality for the year. A warning may, however, be not superfluous to the effect that in individual months "accidental " irregularities are apt to influence largely the amplitude and phase angle of the Fourier waves of shorter period. When one derives a diurnal inequality by combining together the months forming a season or the year, these "accidental" effects tend to neutralise one another and disappear, but they do not do so in the case of an arithmetic mean of amplitudes derived from individual months. The arithmetic mean of the $c^{\prime}$ 's from the 12 months is necessarily larger than (or at least not less than) the $c$ assigned in the tables to the "year" (i.e. to the $c$ derived from the mean diurnal inequality for the year). The difference between the two $c$ 's is greater the more variable the phase angle throughout the year. The variation from month to month in the phase angles in the tables is partly, no doubt, natural (i.e. representative of the average of years), but it undoubtedly arises in part from "accidental" disturbances and from observational uncertainties. The "accidental" phenomena are especially apt to influence the waves of shorter period, and this is no doubt partly the reason why the values of $c$ derived from the arithmetic mean of the 12 monthly values and from the mean diurnal inequality for the year are relatively so much closer for the 24 -hour wave than for the others. In the case of the shorter-period waves an arithmetic mean could not in all cases be assigned for the phase angle. Provided this angle varies slowly and regularly throughout the year, if we get, say, $359^{\circ}$ and $1^{\circ}$ as its values in two consecutive months, we know that we must regard these either as $359^{\circ}$ and $361^{\circ}$, or as $-1^{\circ}$ and $+1^{\circ}$, in forming a mean. But when, as in Table XXXIV, the values in three successive months are $200^{\circ}, 40^{\circ}$, and $-77^{\circ}$, it is by no means clear how best to interpret the figures.

The phase angles were all calculated out to the nearest minute, though none of them can really claim that degree of accuracy. The phase angles for the 24 -hour wave are shown as calculated; for the 12 -hour wave decimals of a degree are retained; for the 8 -hour and 6 -hour waves the results are recorded only to the nearest degree. In all cases local time is used, Midnight answering to $t=0$. Declination, it will be remembered, is counted positive when the angle $n o \mathbf{N}$ of the figure on $p .101$ is above its mean value for the day; while inclination is regarded as increasing when the needle approaches the vertical.

An increase in a phase angle means that the maxima and minima of the corresponding wave occur earlier in the day. An advance of 1 hour in time requires an increase of $15^{\circ}$ in $\alpha_{1}$, of $30^{\circ}$ in $\alpha_{2}$, of $45^{\circ}$ in $\alpha_{3}$, and of $60^{\circ}$ in $\alpha_{4}$.

Table XXXIII.-Declination (All Days). Amplitudes (Unit 1') and Phase Angles.

|  | $c_{1}$. | $\alpha_{1}$. | $c_{2}$. | $a_{2}$. | $c_{3}$ | $\alpha_{3}$. | $c_{4}$. | $\alpha_{1}$. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1902 |  | - |  | - |  | - |  | - |
| April. . | $13 \cdot 01$ | 33955 | $3 \cdot 52$ | $163 \cdot 1$ | 1 -38 | 12 | $0 \cdot 38$ | 165 |
| May . | $9 \cdot 09$ | 34623 | $1 \cdot 55$ | $90^{\circ} 4$ | $1 \cdot 14$ | 220 | $0 \cdot 97$ | 159 |
| June . | $8 \cdot 07$ | 3312 | $2 \cdot 45$ | $74 \cdot 3$ | $0 \cdot 91$ | 114 | $0 \cdot 52$ | 185 |
| July . | $8 \cdot 16$ | 33447 | $3 \cdot 95$ | $88^{\circ} 2$ | $0 \cdot 61$ | 164 | 1.44 | 13 |
| August . . | 9.19 | $342^{\circ} 2$ | $2 \cdot 29$ | $124 \cdot 9$ | $1 \cdot 10$ | 46 | $0 \cdot 65$ | 131 |
| September . | $17 \cdot 06$ | 34555 | 3.00 | $160 \cdot 4$ | $1 \cdot 48$ | 149 | $0 \cdot 31$ | - 67 |
| October . | 19•34 | 35242 | $5 \cdot 73$ | $127{ }^{\circ} 0$ | C. 54 | 191 | $0 \cdot 20$ | 40 |
| November | $22 \cdot 92$ | 34926 | $5 \cdot 77$ | $137 \cdot 4$ | $2 \cdot 52$ | 34 | $3 \cdot 24$ | $-77$ |
| December | $35 \cdot 08$ | 33524 | $9 \cdot 29$ | $164 * 6$ | 2.98 | 25 | $2 \cdot 38$ | $-70$ |
| 1903 |  |  |  |  |  |  |  |  |
| January. | $25 \cdot 36$ | 33117 | $3 \cdot 05$ | $119 \cdot 8$ | 2.09 | 159 | $2 \cdot 11$ | 153 |
| February | $29 \cdot 34$ | 32932 | $8 \cdot 62$ | $124 \cdot 3$ | $0 \cdot 78$ | $-17$ | $1 \cdot 60$ | 262 |
| March . | $21 \cdot 95$ | 34141 | $5 \cdot 89$ | 158.4 | 0.44 | 60 | 0.99 | 147 |
| April . | $21 \cdot 31$ | 33857 | $2 \cdot 59$ | $111{ }^{4}$ | $0 \cdot 58$ | $-6$ | 1.33 | 180 |
| May . | $15 \cdot 42$ | 33439 | $3 \cdot 30$ | 115.4 | $1 \cdot 06$ | 103 | $0 \cdot 33$ | $-14$ |
| June . | $17 \cdot 35$ | 33411 | $2 \cdot 45$ | $100 \cdot 0$ | $1 \cdot 10$ | 183 | $0 \cdot 39$ | 138 |
| July . | 17 '39 | 33635 | $3 \cdot 78$ | $78 \cdot 7$ | $0 \cdot 65$ | 215 | $1 \cdot 10$ | 182 |
| August . | $20 \cdot 17$ | 3319 | $4{ }^{\circ} 68$ | $145{ }^{\circ} 0$ | 1.91 | 148 | $3 \cdot 18$ | 164 |
| September . | $30 \cdot 56$ | 34047 | $6 \cdot 56$ | 121.6 | 3•17 | $-76$ | $2 \cdot 63$ | 193 |

Table XXXIV.-Declination (All Days). Ampliturles (Unit 1') and Phase Angles.


Table XXXV.-Declination (Quieter Days). Amplitudes (Unit 1') and Phase Angles.

|  | $c_{1}$. | $\alpha_{1}$. | $c_{2}$. | $\alpha_{2}$. | $c_{3}$. | $\alpha_{3}$. | $c_{4}$. | $\alpha_{1}$. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1902 |  | - ' |  | - |  | - |  | 0 |
| April. . | $9 \cdot 50$ | 33931 | $4 \cdot 05$ | $168 \cdot 0$ | 1.96 | 6 | $0 \cdot 87$ | 200 |
| May . . . . . . | $5 \cdot 74$ | 3425 | $0 \cdot 90$ | $109 \cdot 3$ | $0 \cdot 07$ | 23 | $0 \cdot 51$ | 101 |
| June . . . . . . . | $3 \cdot 37$ | 33629 | 0.97 | $120 \cdot 4$ | $0 \cdot 29$ | 101 | $0 \cdot 31$ | $-37$ |
| July . * . . . . . | $5 \cdot 22$ | 34.22 | 1.85 | $95 \cdot 7$ | $0 \cdot 67$ | 120 | $0 \cdot 15$ | 128 |
| August . . . . . | $5 \cdot 57$ | 34727 | 0.84 | $157 \cdot 9$ | $0 \cdot 84$ | 29 | 0-36 | 124 |
| September . . . . | $12 \cdot 59$ | 35248 | $2 \cdot 78$ | $167 \cdot 9$ | $0 \cdot 50$ | 131 | $0 \cdot 26$ | - 81 |
| October . . . . . . | $14 \cdot 02$ | 37047 | 3•71 | $106 \cdot 2$ | $0 \cdot 09$ | 128 | $0 \cdot 47$ | 1 |
| November . . . . . . | $19 \cdot 21$ | 351.20 | $8 \cdot 67$ | $167 \cdot 1$ | $3 \cdot 61$ | 62 | $2 \cdot 11$ | - 26 |
| December * . . . . | $28 \cdot 83$ | 3478 | $7 \cdot 52$ | $194 \cdot 4$ | $2 \cdot 31$ | 87 | $2 \cdot 06$ | $-77$ |
| 1903 January . . . . . | $28 \cdot 83$ | 3305 | $7 \cdot 32$ | $160 \cdot 0$ | $2 \cdot 47$ | 62 | $3 \cdot 56$ | 202 |
| February . . . . . | $24 \cdot 24$ | 3347 | $5 \cdot 76$ | 111 '3 | $0 \cdot 50$ | - 7 | $1 \cdot 69$ | 231 |
| March . . . . . . | $15 \cdot 89$ | 34521 | $4 \cdot 73$ | $165 \cdot 0$ | $1 \cdot 61$ | 129 | $0 \cdot 15$ | - 66 |
| April . . . . | $12 \cdot 27$ | 33944 | $5 \cdot 14$ | 145 '0 | $0 \cdot 92$ | 48 | 1. 55 | 207 |
| May . . . . | $10 \cdot 43$ | 34129 | $2 \cdot 11$ | $150 \cdot 4$ | 1.56 | 116 | 0.17 | - 33 |
| June . . . . . . . | $5 \cdot 47$ | 33938 | $0 \cdot 71$ | $26 \cdot 4$ | 0.14 | 7 | $0 \cdot 37$ | 50 |
| July . . . . . . . | $9 \cdot 04$ | 33515 | $2 \cdot 36$ | $74 \cdot 8$ | 0.53 | 142 | $0 \cdot 61$ | 140 |
| August . . . . . | $9 \cdot 36$ | $344 \% 6$ | $1 \cdot 40$ | $77 \cdot 1$ | 0.79 | 109 | $0 \cdot 60$ | 104 |
| September . . . . . | $17 \cdot 86$ | 33731 | $2 \cdot 06$ | $56 \cdot 6$ | 1.96 | $-46$ | 3-31 | 173 |

Tabie XXXVI.-Declination (Quicter Days). Amplitudes (Unit 1') and Phase Angles.


Table XXXVII.-Horizontal Force. Amplitudes (Unit 1y) and Phase Angles.


Tarle XXXVIII.-Horizontal Force. Amplitudes (Unit $1 \gamma$ ) and Phase Angles.


Table XXXIX.-Vertical Force. Amplitudes (Unit 1y) and Phase Angles.

|  | $c_{1}$. | $\alpha_{1}$. | $c_{2}$. | $\alpha_{2}$. | $c_{3}$. | $x_{3}$. | $c_{4}$. | $\alpha_{1}$. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | - , |  | - |  | $\square$ |  | - |
| January . | $27 \cdot 69$ | 4647 | $12 \cdot 37$ | $255 \cdot 4$ | $7 \cdot 55$ | 76 | $4 \cdot 70$ | 314 |
| February | $19 \cdot 34$ | $65 \quad 29$ | 7.07 | $208 \cdot 2$ | $5 \cdot 66$ | $-1$ | $1 \cdot 13$ | 267 |
| March . | 18.59 | 7642 | $6 \cdot 55$ | 221.4 | $2 \cdot 30$ | 90 | 0.01 | 294 |
| April . . | $17 \cdot 19$ | 11242 | $1 \cdot 64$ | $260 \cdot 4$ | $1 \cdot 58$ | 67 | $1 \cdot 18$ | 245 |
| May . | $7 \cdot 14$ | 10117 | $0 \cdot 80$ | $206 \cdot 0$ | 0.90 | 65 | $0 \cdot 42$ | 17 |
| June . | 8.70 | 10424 | 0.99 | 254:3 | $0 \cdot 84$ | - 35 | $0 \cdot 39$ | 316 |
| July . | 11.00 | 12158 | $1 \cdot 73$ | $235{ }^{\circ}$ | $0 \cdot 57$ | 260 | $0 \cdot 44$ | 149 |
| Auguat . | $13 \cdot 12$ | 11824 | $1 \cdot 01$ | $270 \cdot 4$ | 1.03 | 23 | $1 \cdot 38$ | 26.4 |
| September . | $14 \cdot 12$ | 9258 | 2.38 | $236 \cdot 7$ | 0.96 | 16 | 0.33 | 338 |
| October. | $13 \cdot 91$ | 8524 | $3 \cdot 94$ | $193 \cdot 9$ | $2 \cdot 03$ | 56 | $1 \cdot 13$ | 281 |
| November . | $22 \cdot 39$ | 9948 | $3 \cdot 53$ | $280 \cdot 9$ | $2 \cdot 04$ | 15 | $0 \cdot 29$ | 266 |
| December | $26 \cdot 24$ | $80 \quad 3$ | $5 \cdot 09$ | $274 \cdot 7$ | $3 \cdot 12$ | 51 | 2-17 | 260 |
| Arithmetic Means . | 16.61 | 9210 | 3.92 | 241 4 | $2 \cdot 38$ | 57 | 113 | 251 |
| Year | $15 \cdot 34$ | 8652 | $3 \cdot 49$ | $240 \cdot 6$ | 189 | 45 | $1 \cdot 00$ | 287 |
| Midwinter . | $8 \cdot 83$ | 11046 | 1-12 | 234.0 | $0 \cdot 34$ | $-13$ | 0-10 | 16 |
| Equinor | $15 \cdot 48$ | 9149 | $3 \cdot 40$ | $220 \cdot 7$ | 1.58 | 65 | $0 \cdot 85$ | 279 |
| Midsummer | $23 \cdot 59$ | 7355 | $6 \cdot 87$ | 264 -3 | $3 \cdot 92$ | 61 | $2 \cdot 15$ | 296 |

Table XL.-Inclination. Amplitudes (Unit 1') and Phase Angles。

§34. Before making some general remarks on Tables XXXIII to XL, it is convenient to deal with an interesting result deducible from the Declination Tables. If we compare Tables XXXIII and XXXV, or XXXIV and XXXVI, it will be observed that the differences between the phase angles derived from all days and from the quieter days are to some extent systematic, at least in the case of the 24 -hour and 12-hour waves.

This will be more easily recognised on consulting Tables XLI and XLII, which show the algebraic excess of the phase angle from the quieter-days' inequality over that from all days. Thus, taking Table XLI, we find the quieter-days' value of $\alpha_{1}$ the larger in 13 months out of 18 , the average excess for the 18 months being no less than $4^{\circ} 29^{\prime}$. This signifies that on the average quieter day the maximum and minimum of the 24 -hour wave occurred about 18 minutes earlier than on the average day for which records existed. The quieter-days' phase angle $\alpha_{2}$ is also the larger in a substantial majority of the months ; the average excess is, however, only $2^{\circ} 42^{\prime}$, representing about $5 \cdot 4$ minutes of time. The excess would, however, have been very substantially larger but for the results from the later months of 1903 , which are based on a rather smaller number of days than usual. The results from the seasonal and mean annual diurnal inequalities in Table XLII point in the same direction. The yearly results in this table for $\alpha_{3}$ and $\alpha_{4}$ have the same sign as those for $\alpha_{1}$ and $\alpha_{2}$, but the phenomena in individual months, and even in individual seasons of the year, appear too irregular to justify our regarding the difference between all and quieter days as established for the 8 -hour and 6 -hour waves.
In the case of the 24 -hour and 12 -hour waves the difference does seem fairly established. As to its most probable size, if we take a mean from the yearly results in Table XLII and the final means in Table XLI, we find for the advance in quieter days in the time of maximum or minimum $15 \cdot 2$ minutes for the 24 -hour wave and $20^{\circ} 2$ minutes for the 12 -hour wave.

This is not the first occasion on which a difference has been noted between the phase angles on ordinary and on quiet days. Recently* it was pointed out that a substantial difference existed at Kew between the 24-hour and 12-hour phase angles derived from the Astronomer Royal's quiet days and those derived from all days of the month. There is, however, a certain remarkable difference between the Antarctic phenomena and those at Kew. In the Antarctic we have found the quieter-days' angles to be the larger, both for the 24 -hour and the 12 -hour waves; but at Kew, while the quieter-days' phase angle was the larger in the case of the 12 -hour wave, it was the smaller in the case of the 24 -hour wave.

* 'Phill. Trans.' A, vol. 208, 1908, pp. 223, \&c.

When discussing the Kew results, I hazarded the remark that the difference observed there between the phenomena in the 24 -hour and 12 -hour waves might be associated with the further difference that the 12-hour phase angle is largest in Summer, while the 24 -hour phase angle is largest in Winter. Referring to Tables XXXIV and XXXVI, it will be seen that in the Antarctic, the 24 -hour phase angle agrees with the 12 -hour phase angle there and at Kew in being larger at Midsummer than at Midwinter. This is, of course, in accordance with the suggestion which I threw out, but much weight ought not to be attached to what not unlikely may be a mere coincidence, especially as the seasonal variation of $\alpha_{1}$ in the Antarctic is not large.

The difference of phase is not the only difference between all- and quieter-days' Declination results. The amplitudes, at least in the case of the 24 -hour wave, are markedly less for the quieter days, but this we could have foreseen from simple comparison of the diurnal inequalities in Tables XIII and XV.

Table XLI.-Declination. Quieter-days' Phase Angle - All-days' Angle.

|  | $\alpha_{1}$. |  |  |  | $\alpha_{2}$. |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 1902. | 1903. | 1002-3. | 1902. | 1903. | 1902-3. |
|  | - , | - , | - | - | - , | - , |
| January . |  | $-112$ |  |  | +4012 |  |
| February . . . . . . |  | + 435 |  |  | -1255 |  |
| March . . . . . . . |  | + 340 |  |  | + 636 |  |
| April . . . . . . . . | - 024 | + 047 | + 018 | + 457 | +3339 | $+1336$ |
| May . . . . . . . | $-418$ | + 650 | + 241 | +1852 | +35 2 | +3058 |
| June . . . | +527 | + 522 | $+513$ | +46 5 | -73 36 | -425 |
| July | + 715 | - 120 | + 144 | + 731 | - 351 | + 026 |
| August . . | + 525 | +1317 | $+110$ | +3253 | $-6757$ | -32 49 |
| September . . . . . . . | +653 | $-316$ | +113 | $+733$ | $-65 \quad 2$ | $-90$ |
| October . . . . . . | +185 |  |  | -2049 |  |  |
| November . | +154 |  |  | $+2940$ |  |  |
| December | +1144 |  |  | +2946 |  |  |
| Mean from 18 months | $+4^{\circ} 29^{\prime}$ |  |  | $+2^{\circ} 42^{\prime}$ |  |  |

Table XLII.-Declination. Quieter-days' Phase Angle - All-days' Angle.

§35. The most consistent and striking phenomenon in Tables XXXIII to XL is the relatively small amplitude of the Fourier waves of shorter period as compared to the 24 -hour wave. To bring this out more clearly, Table XLIII records the ratios of the amplitudes of the 12 -, 8 - and 6 -hour waves to the corresponding amplitude of the 24 -hour wave in the diurnal inequalities for the several elements and seasons. D' relates to the quieter-days', D to the all-days' Declination results. The ratios of the arithmetic means of the 12 monthly values of $c_{2}, c_{3}$, and $c_{4}$ to the corresponding mean for $c_{1}$ are also given. The figures in the sixth row are means derived from the four elements D, H, V and I. For contrast, the seventh row gives corresponding results from the Kew Declination on ordinary days.

Comparing the results under the heading "Year," we see that relative to the 24 -hour wave the importance of the 12 -hour and 6 -hour waves is, in the Antarctic, only about a third of what it is at Kew, while the relativs importance of the 8 -hour wave in the Antarctic is only about a sixth of that at Kew.

Taking means from the four elements，we should infer that in the Antarctic the relative importance of the 12 －to the 24 －hour wave is but little dependent on the season of the year．The 8 －hour and 6 －hour waves seem，however，of greatest relative importance at Midsummer，which is the exact opposite of what occurs at Kew．

Table XLIII．－Ratios of the Amplitudes of the 12－，8－and 6－hour Fourier Waves to that of the 24－hour Wave．

|  | $c_{2} / c_{1}$ ． |  |  |  |  | $c_{3} / c_{1}$ ． |  |  |  |  | $c_{4} / c_{1}$ ． |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 㚻 | 范 | 这 | $\begin{aligned} & \text { 岂 } \\ & \underset{\sim}{\underset{\sim}{x}} \\ & \hline \end{aligned}$ | 象 | E | － | 号范 |  | 总范 |  | \％ | 家宽 | 吕 | 家最罥 |
| $\mathrm{D}^{\prime}$ 。 | $0 \cdot 264$ | $0 \cdot 227$ | $0 \cdot 187$ | 0－234 | $0 \cdot 300$ | $0 \cdot 181$ | $0 \cdot 063$ | $0 \cdot 073$ | $0 \cdot 026$ | 0－108 | $0 \cdot 077$ | 0.037 | $0 \cdot 026$ | 0.041 | 0.058 |
| D ． | ＇235 | $\cdot 218$ | $\cdot 226$ | －229 | －209 | －059 | $\cdot 017$ | ． 050 | －010 | $\cdot 054$ | ． 064 | －032 | －034 | ． 084 | ．054 |
| H ． | ＇200 | －186 | $\cdot 225$ | －195 | －169 | ．069 | .029 | －063 | $\cdot 047$ | －059 | ． 058 | －023 | －058 | ． 009 | ． 085 |
| V．． | $\cdot 236$ | $\cdot 227$ | －127 | －220 | －291 | －143 | $\cdot 123$ | ．038 | －102 | $\cdot 160$ | ． 068 | －065 | $\cdot 012$ | －055 | .091 |
| I． | $\cdot 197$ | $\cdot 188$ | $\cdot 222$ | －199 | －171 | －064 | －022 | －060 | －037 | ． 057 | ．056 | ． 025 | ．025 | ．013 | ＇087 |
| Mean | 0－217 | $0 \cdot 205$ | $0 \cdot 200$ | $0 \cdot 211$ | 0－210 | $0 \cdot 084$ | $0 \cdot 048$ | $0 \cdot 053$ | 0.049 | $0 \cdot 084$ | $0 \cdot 061$ | $0 \cdot 036$ | 0.092 | 0.028 | 0.054 |
| Kew D ． |  | $0 \cdot 62$ | $0 \cdot 49$ | $0 \cdot 65$ | $0 \cdot 65$ | － | $0 \cdot 28$ | $0 \cdot 23$ | $0 \cdot 34$ | $0 \cdot 23$ | － | $0 \cdot 10$ | $0 \cdot 14$ | $0 \cdot 13$ | 0.03 |

## CHAPTER V.

## Annual Variation. Fourier Coffficients.

§36. To throw more light on the nature of the annual variation in the amplitude of the diurnal inequalities, the ranges, the sum of the 24 hourly differences from the mean, and the amplitude $c_{1}$ of the 24 -hour wave were analysed in the Fourier series

$$
\mathrm{M}+\mathrm{P}_{1} \sin \left(t+\theta_{1}\right)+\mathrm{P}_{2} \sin \left(2 t+\theta_{2}\right)+\cdots
$$

where $t$ represents time counted from January 1st, a month in $t$ answering to $30^{\circ}$. M represents, of course, the mean value for the year, while $\mathrm{P}_{1}, \mathrm{P}_{2}$ are the amplitudes, $\theta_{1}, \theta_{2}$ the phase angles of the 12 -month and 6 -month waves. Differences in the lengths of the calendar months were neglected. The values of $P_{1}, P_{2}, \theta_{1}, \theta_{2}, P_{1} / M, P_{2} / \mathbf{M}$ and $\mathrm{P}_{2} / \mathrm{P}_{1}$ for the several elements dealt with are recorded in Table XLIV.

Means of the quantities in the last 5 columns, derived from D (all days), $\mathrm{H}, \mathrm{V}$ and I are compared with corresponding means derived from the $\mathrm{D}, \mathrm{H}, \mathrm{V}$ and I results (on quiet days during 1890 to 1900) at Kew. As the Antarctic results are based on only two years' observations, they are presumably not so close an approximation to what exactly represents average conditions as are the Kew results. Still, they present features which can hardly be regarded as the result of accident, and which seem of much interest.

We know already that the Antarctic diurnal inequality varies little in type thronghout the year, and that the 24 -hour Fourier wave is largely dominant. Thus we might have anticipated that no very large differences would exist between the values of $\theta_{1}$ or the values of $P_{1} / M$ derived from the ranges, the sum of the 24 differences and $c_{1}$ in any one element. But we could not have foreseen that the differences would be so small as appears from Table XLIV.

There is also remarkably little difference between the values of $\theta_{1}$ or between the values of $P_{1} / \mathrm{M}$ which are derived from the Horizontal Force, the Vertical Force and the Inclination. There is a somewhat conspicuous difference between the values of $\theta_{1}$ and between the values of $P_{1} / M$ derived, in the case of the Declination, from all and from quieter days. Also, somewhat curiously, whilst the all-days' value of $P_{1} / M$ accords closely with the corresponding values for $\mathrm{H}, \mathrm{V}$ and I , it is the quieter-days' value of $\theta_{1}$ that accords most closely with the corresponding angles for the other elements.

The fact that the quieter-days' Declination value of $\mathrm{P}_{1} / \mathrm{M}$ is enhanced indicates that, relatively considered, the seasonal variation in the amplitude of the diurnal inequality is greater for quieter days than for all days.

In all cases the 6 -month wave is smaller than the 12 -month wave, but on the average of the elements its relative importance appears greater in the Antarctic than at Kew.

The large difference between the values of $\theta_{1}$ at Kew and in the Antarctic arises almost entirely from the six-month difference in the season. If we add $180^{\circ}$, i.e. six months, to the Antarctic value of $\theta_{1}$ we get to within $3^{\circ}$-or roughly three days-of the Kew value. In the case of the 6 -month term the difference between the mean values of $\theta_{2}$ for the Antarctic and Kew is $188^{\circ}$, or about three months and four days. This means that this wave also is at Winter Quarters very nearly opposite in phase to what it is at Kew.

Table XLIV,-Annual Variation. Fourier Coefficients.

§37. Table XLV shows approximately the dates when the 12 -month and 6 -month waves attain a maximum. The 12 -month wave has of course only one maximum in the course of a year, separated by six months from the minimum. The 6 -month wave has two maxima, 6 months apart, with minima equidistant between them. The date recorded for this wave in Table XLV refers to that one of the two maxima which falls nearest to January 1.

Table XLV.-Annual Variation. Dates of Maximum.


## CHAPTER VI.

## Absolute Daily Ranges. Daily Maxima and Minima.

§38. By the absolute range of an element is meant the excess of the absolutely largest over the absolutely smallest value met with during the 24 hours. The term ubsolute is added to indicate that the quantity considered is not the range of the regular diurnal inequality, nor the range derived from mere hourly readings. The absolute ranges are at once derivable from the daily maxima and minima given in the tables of tabulated values, but for convenience of reference they have been collected and presented in Tables XLVI, XLVII, and XLVIII.

It is of course impossible ever to say with certainty, in the case of an absolutely isolated station like Winter Quarters, what may have been taking place during the time of changing papers. There is always the possibility that, in the course of a few minutes during which no trace was being recorded, an element may suddenly have changed and then reverted to near its primitive value. But even in the Antarctic, though large sudden changes were not of very rare occurrence, it was very unusual for them to occur singly, and one could usually feel fairly confident that neither the daily maximum nor the daily minimum had occurred whilst there was no paper on the drum. When, however, the interval between successive sheets was considerable, as occasionally happened, or when the trace was highly disturbed about the time of changing, there might be considerable doubt as to whether a maximum or minimum might not have occurred in Declination or Horizontal Force.

In Tables XLVI and XLVII, relating respectively to D and H , ranges are usually given whether the record for the day was complete or not. There are, however, omissions on a few days when the traces were confused or indistinctly visible. Figures inside [ ] brackets relate to days when the record was incomplete, but when the general appearance of the curve seemed to warrant the belief that both the maximum and the minimum for the day were actually recorded. Figures inside () brackets are for days of incomplete record, when appearances suggested that either the maximum or the minimum, if not both, was unrecorded. The combination $>+$ means that the trace went beyond the limit of registration in the direction of element increasing, while $>-$ means that the trace exceeded the limit in the direction of element diminishing. When both limits were exceeded, the combmation $> \pm$ is employed. The sign * denotes that the trace was too confused to decipher, while _ denotes that no trace, or only a few hours' trace, existed. Figures preceded by $>$ are certainly, and those in ( ) brackets probably, under-estimates of the true range.

In the case of D , and still more in that of H , the number of days when the record was incomplete was so considerable, and the cause was so frequently due to the limits of registration being exceeded, especially in Summer, that a very imperfect idea of the average amplitude would have been derived if days of incomplete record had been omitted. This consideration did not, however, apply to $V$, as loss of record of this element very seldom arose merely from the daily amplitude being large. The almost invariable cause was defective action in the magnetograph, or loss of temperature trace, and so ignorance of the temperature correction. There was thus no reason to regard the ranges derived from days of complete record as below the average size. It was thus decided to give maxima and minima and absolute ranges only for the days of complete record. These ranges appear in Table XLVIII, and means are given for the separate months. In February, October, November, and December, 1903, however, the number of days from which the means are derived are so small that the figures possess but slight significance.
Table XLVI．－Absolute Ranges．Declination．

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Table XLVII．－Absolute Ranges．Horizontal Force．（Unit ly．）

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Table XLVIII.—Absolute Ranges. Vertical Force. (Unit $1 \gamma$.)

§39. Table XLIX gives an analysis of the results obtained by grouping the ranges of Declination according to their amplitude. The first group gives the number of days in which the range did not exceed $30^{\prime}$, the second the number of days in which the range, while in excess of $30^{\prime}$, did not exceed $1^{\circ}$, and so on, the seventh and last group giving the number of days in which the range exceeded $3^{\circ}$. Days were arranged in two principal classes, the first including all days when the record was complete, or when loss of trace was due solely to one or both limits of registration being exceeded, the second including all days when there was imperfection of record through photographic failure, absence of sheet, or similar cause. Each of these two classes was sub-divided into two sub-classes, according as the trace did or did not keep within the limits of registration. The days are arranged under three seasons, Midwinter (May to July), Midsummer (November to January), Equinox (March, April, September, and October). The results for the remaining months, August, 1902 and 1903, and February, 1903, are combined.

Table XLIX.-Declination Ranges.


It will be noticed that out of a total of 111 days in Midsummer there was only one in which the range did not exceed $1^{\circ}$, while at least 44 had a range in excess of $3^{\circ}$. Taking the whole period, we find that out of 461 days for which the record was complete, except for the limits of registration being exceeded, 329 , or 71 per cent., had a range over $1^{\circ} ; 250$, or 54 per cent., had a range over $11^{\circ} ; 179$, or 39 per cent., a range over $2^{\circ}$, and 75 , or 16 per cent., a range over $3^{\circ}$. Of the 581 days, complete and incomplete, included in the table, 407 , or 70 per cent., had a range over $1^{\circ} ; 218$, or 38 per cent., a range over $2^{\circ}$; 90 , or $15 \frac{1}{2}$ per cent., a range over $3^{\circ}$; while 24 , or fully 4 per cent., had a range over $4^{\circ}$. On seven days the trace exceeded the limits of registration on both sides of the sheet, whose complete width represented from $4^{\circ} 50^{\prime}$ to $4^{\circ} 55^{\prime}$.
840. Results for H corresponding to those for D , just discussed, appear in Table L . It contains two principal classes, each with two sub-classes analogous to those in Table XLIX. The ranges are again dealt with in seven groups, the first containing days in which the range did not exceed $25 \gamma$, the second days in which the range exceeded $25 \gamma$ but did not exceed $50 \gamma$, and so on, the last group containing days when the range exceeded $150 \gamma$. Owing to the sensitiveness of the Horizontal-Force magnetograph, the limits of registration were exceeded in about one day out of two. In Midsummer the limits were exceeded
in eight days out of eleven, so that our information at this season is unfortunately very imperfect. Even during Midwinter the range exceeded $25 \gamma$ on 90 per cent. of the days.

Table L.-Horizontal Force Ranges. (Unit ly.)

|  | Days when no incompleteness or failure of record, except through the trace going beyond the limits of registration. |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Days when trace kept within limits of registration. |  |  |  |  |  |  | Days when trace went beyond one or both limits of registration. |  |  |  |  |  |  |  |
|  | Total | Range. |  |  |  |  |  | Otal | Range. |  |  |  |  |  |  |
|  | Days. |  | $\begin{array}{cc}2 .) \\ 50 . & 50 \text { to } \\ 55 .\end{array}$ | $\begin{aligned} & 75 \text { to } \\ & 100 . \end{aligned}$ | $\begin{aligned} & 10010 \\ & 125 . \end{aligned}$ | $\begin{gathered} 125 \\ 150 . \end{gathered}$ | Over 130. | Days. | 0 to 25. | 25 to 50. | 50 to 75. | $\begin{aligned} & 75 \text { to } \\ & 1000 . \end{aligned}$ | $\begin{aligned} & 100 \text { to } \\ & 125 . \end{aligned}$ | $\begin{gathered} 125 \text { to } \\ 150 . \end{gathered}$ | $\begin{aligned} & \text { Over } \\ & 150 \text {. } \end{aligned}$ |
| Midwinter . | 116 | 8 | $45 \quad 29$ | 18 | 10 | 4 | 2 | 29 | 0 | 0 | 0 | 3 | 8 | 10 | 8 |
| Equinox . | 59 | 0 | 19 21 | 13 | 4 | 1 | 1 | 94 | $\square$ | $\square$ | 17 | 25 | 29 | 7 | 16 |
| Midsummer . |  | 0 | 0 0 3 | 4 | 6 | 11 | 0 | 58 | 0 | 0 | 8 | 7 | 12 | 13 | 18 |
| February and Augusts |  | 1 | 15 13 | 5 | 0 | 0 | 0 | 43 | 0 | 0 | 4 | 16 | 13 | 2 | 5 |
| Total. | 233 | 9 | 79 : 66 | 40 | 20 | 16 | 3 | 224 | 0 | 0 | 29 | 51 | 62 | 32 | 50 |
| Midwinter . . . . | Days when record incomplete from some cause other than trace going beyond limits of registration. |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  | Days when trace kept within limits of registration. |  |  |  |  |  |  | Days when trace went heyond one or both limits of registration. |  |  |  |  |  |  |  |
|  | 25 | 10 | 8, 5 |  | 0 | 1 | 0 | 7 | 0 | 1 | 0 | 0 | $\square$ | 8 | 1 |
| Equinox . | 17 | 0 | 34 | 7 | 1 | 1 | 1 | 29 | 0 | 0 | $\pi$ | 7 | 2 | 4 | 11 |
| Midsummer . | 6 | 0 | 1 1 1 | 1 | 2 | 1 | 0 | 21 | 1 | 1 | 0 | 7 | 3 | $\square$ | 4 |
| February and Augusts | 2 | 1 | 10 | 0 | 0 | 0 | 0 | 5 | 0 | 0 | 2 | 1 | 2 | 0 | 0 |
| Total . . . | \%0 | 11 | 1310 |  | 3 | 3 | 1 | 82 |  | .$^{2}$ | 9 | 15 | 10 | 9 | 16 |

Of the 457 days for which the record was complete, except for the trace going beyoud the limits of registration, 274 , or 60 per cent., had a range over $75 \gamma$, and 183 , or 40 per cent., a range over $100 \gamma$. Of the 569 days, complete or incomplete, included in the table, 340 , or 60 per cent., had a range over $75 \gamma$, and 225 , or 40 per cent., a range over $100 \gamma$.

If we compare the D and H results we find that the percentage of days showing a range of over $100 \gamma$ in H is closely similar to the percentage showing a range of over $2^{\circ}$ in D . As $1^{\prime}$ in D range answers to about $2 \gamma$ in $H$, the natural conclusion-which one would also draw from the diurnal inequalities-is that diurnal variations of force were considerably larger in the direction perpendicular to the magnetic Meridian than in the magnetic Meridian itself. The conclusion is almost certainly correct. At the same time, it should be noticed that even when the $H$ magnetograph was least sensitive it was impossible to record a range much over $200 \gamma$, while the possible limit for D was equivalent to nearly $600 \gamma$ throughout. Thus any deduction based on Tables XLIX and L as to the relative size of the average daily ranges in D and H is practically certain to underestimate the range in H .
§41. Table LI gives some data for V analogous to those in the last two tables. Days are arranged in six groups, the first including cases where the range did not exceed $50 \gamma$, the second cases in which it

Table LI.-Vertical-Force Ranges. (Unit 1 $\gamma_{0}$ )

|  | Total of days. | Range. |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | 0 to 50. | 50 to 100. | 100 to 150. | 150 to 200. | 200 to 250. | Over 250. |
| Midwinter . . . . . . . | 123 | 47 | 53 | 13 | 7 | 1 | 2 |
| Equinor . . . . . . | 101 | 20 | 49 | 22 | 7 | 3 | 0 |
| Midsummer . . . | 47 | 3 | 10 | 18 | 6 | 4 | 6 |
| February and Auguste . . . | 46 | 10 | 24 | 9 | 1 | 2 | 0 |
| Total . . . . | 817 | 80 | 136 | 62 | 21 | 10 | 8 |

exceeded $50 \gamma$ but did not exceed $100 \gamma$, the last cases in which it exceeded $250 \gamma$. Of the 317 days included, only 5 had a range as small as $25 \gamma$, while 101 , or 32 per cent., had a range exceeding $100 \gamma$. The latter figure is rather smaller than the corresponding percentage in the case of H .
§42. In addition to the results for individual days, Table XLVIII gives means for individual months, varying from $234 \gamma$ for January, 1903, to $50 \gamma$ for July, 1902. The monthly means, however, fluctuate somewhat irregularly, and can hardly claim to closely represent average conditions. The January mean, for example, appears abnormally large. This is partly accounted for by the exceptionally large ranges on January 11 and 12. It is by no means impossible that some of the irregularities may be due to errors in the scale values, or in the temperature corrections. The former source of error is most to be feared in the months of July, August, and September, 1902, the latter in the Midsummer months of 1902-3, when the temperature range was especially large.
If we combine corresponding months from the two years, allowing equal weight to each day, we obtain the following somewhat more regular mean monthly values:-

Table LII.-Mean Absolute Ranges of V. (Unit $1 \gamma$.)

843. Table LIII gives the largest and smallest absolute ranges recorded in each individual month.

Table LIII.-Absolute Ranges.

|  | Declination. |  |  | Horizontal Force. (Unit 1r.) |  |  | Vertical Force. (Unit ly.) |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Largest recorded. | Smallest recorded. |  | Largest recorded. | Smallest recorded. |  |  |  |
|  |  | On day when record complete. | On any day. |  | On day when record complete. | On any day. | recorded. | recorded. |
| 1902 | $\bigcirc$ - | - , | - , |  |  |  |  |  |
| March . | $\geq 246$ | $45 \cdot 8$ | $15 \cdot 0$ | $>196$ | - | 66 | 162 | 44 |
| April . . . | $>333 \cdot 0$ | 22.5 | $15 \cdot 7$ | $>216$ | 36 | 36 | 190 | 50 |
| May . . . | $333 \cdot 0$ | $12 \cdot 8$ | $6 \cdot 8$ | $>224$ | 30 | 11 | 168 | 28 |
| June . . . . | 2220 | 11.1 | $11 \cdot 1$ | $>160$ | 19 | 17 | 166 | 39 |
| July . . | $>323{ }^{\circ}$ | 11.0 | $8 \cdot 6$ | 159 | 23 | 17 | 165 | 15 |
| August . . . | $>330 \cdot 7$ | $20 \cdot 3$ | 16.0 | $>153$ | 25 | 22 | 174 | 36 |
| September . . | $253 \cdot 2$ | $29 \cdot 0$ | $29 \cdot 0$ | $>112$ | 29 | 29 | 92 | 39 |
| October . . | $>4471$ | $50 \cdot 7$ | $50 \cdot 7$ | $>111$ | 40 | 40 | 155 | 31 |
| November . | $>454{ }^{\circ}$ | $58 \cdot 7$ | $58 \cdot 7$ | $>160$ | 52 | $>35$ | 303 | 35 |
| December . | $>455.5$ | 128.8 | 128.8 | $>161$ | 93 | 85 | 233 | 60 |
| $\begin{gathered} 1903 \\ \text { January } \end{gathered}$ | 4 2 7 | 122.2 | $122 \cdot 2$ | $>161$ | 84 | 53 | 415 | 94 |
| February . | $\geq 454{ }^{\circ}$ | $15 \cdot 7$ | 15.7 | $>115$ | 58 | 58 | 101 | 92 |
| March . | $>324{ }^{\circ}$ | 39 '6 | $39 \cdot 6$ | $>114$ | 32 | 32 | 232 | 25 |
| April . . | $>347^{\circ} 0$ | $47 \cdot 1$ | 29.4 | $>144$ | 30 | 30 | 142 | 47 |
| May . . . | 343.5 | $24 \cdot 0$ | $18 \cdot 7$ | $>150$ | 29 | 18 | 183 | 30 |
| June . . . | $>430 \cdot 4$ | $17 \cdot 1$ | $15 \cdot 7$ | $>154$ | 14 | 14 | 209 | 24 |
| July . . | $>400$ | $23 \cdot 7$ | $23 \cdot 3$ | $>151$ | 36 | 36 | 258 | 27 |
| August . | $>4510$ | $39 \cdot 7$ | 397 | $>155$ | 32 | 32 | 249 | 31 |
| September . | 358.8 | 112.3 | $112 \cdot 3$ | $>155$ | 61 | 51 | 161 | 33 |
| October . | >4 52. | $\geq 20 \cdot 7$ | 54.9 | $>157$ | . | $\geq 60$ | 245 | 72 |
| November . . | $\geq 4.75$ | $>325.5$ | $>2 \quad 6 \cdot 7$ | $\geq 153$ | -_ | $>23$ | 164 | 161 |
| December - . | $>454.0$ | $339 \cdot 0$ | 1210 | $>160$ | - | $>75$ | 318 | 122 |
| $\begin{gathered} 1904 \\ \text { January . . . } \end{gathered}$ | $>350 \cdot 2$ | - | 141.3 | >162 | - | 43 |  |  |

Even in the case of D there are 16 months in which the largest range of the month is underestimated owing to the limit of registration being exceeded. In the case of H every month, except July, 1902,
suffered in this way. The information as to the smallest ranges does not suffer from any such uncertainty, but subsequent to September, 1903, the number of days of observation was too small to give results of much value. The smallest recorded ranges on days when the record was complete were-

$$
\begin{aligned}
& \text { for D, } 11^{\prime} \cdot 0 \text { in July, } 1902 ; \\
& " \mathrm{H}, 14 \gamma \text { "June, } 1903 ; \\
& " \mathrm{~V}, 15 \gamma \text { "July, } 1902 .
\end{aligned}
$$

Perhaps the most natural way of comparing the different months or seasons of the year as to their relative liability to disturbance is to consider the size of the ratio
(mean absolute range)/(range of diurnal inequality).
This criterion cannot be very readily applied in the present instance, except to the Vertical Force. For it the above ratio has the following values: Midwinter 4.1, Equinox 2.5, Midsummer 2.6. The ratio is thus docidedly largest at Midwinter. The same phenomenon has been observed at Kew in the case of the Declination.

The practical equality of the ratios for Equinox and Midsummer suggests that in the Antarctic the former is not a season of specially large or frequent disturbances. If instead of taking for our criterion of disturbance the ratio of the mean absolute to the inequality range we were to take the amplitude o the mean absolute range, or the frequency of occurrence of specially large ranges, we should come to the conclusion that the equinoctial months were much less disturbed than those at Midsummer. For instance, in the case of D , taking all days, complete and incomplete, the absolute range was in excess of $3^{\circ}$ on only 11 per cent. of equinoctial days as against 40 per cent. of Midsummer days.
§44. Tables LIV to LIX show the frequency of occurrence of the daily maxima and minima in $\mathrm{D}, \mathrm{H}$ and $V$ at different hours of the day. For instance, from Table LIV we see that, taking the 24 days of April, 1902, for which data exist, the maximum in $D$ occurred eight times between 8 and 9 a.m., six times between 9 and $10 \mathrm{a} . \mathrm{m}$., three times in each of the hours $10-11 \mathrm{a} . \mathrm{m}$. and $11 \mathrm{a} . \mathrm{m}$. to noon, and once in each of the hours 5-6 a.m., 7-8 a.m., noon to 1 p.m., and 11-12 p.m. Maxima and minima were assigned for D and H even on days of incomplete record. In some cases the imperfection was of such a nature-e.g. trace beyond a limit of registration during part of one hour-as to introduce no uncertainty into the hour of occurrence of either maximum or minimum; in a greater number of cases there was uncertainty as to one only of the two quantities. The number of days in which uncertainty affected the maximum and the minimum often differed widely. Thus in October, 1902, the hour of minimum in H could be assigned on 26 days, but the hour of maximum only on 12. In the case of $\mathbf{V}$, only days of complete record were included, so that maxima and minima were equal in number. The occurrence of $\frac{1}{2}$ will be noticed in some of the entries for $D$ and $H$, especially the latter. This may mean equality in the extreme ordinates during two different hours, but it usually signifies that the trace was beyond a limit of registration during part of each of two hours. Days were omitted in which either limit of registration was exceeded during parts of more than two hours.

If irregular disturbances did not exist, the daily maxima and minima would synchronise with the maxima and minima in the diurnal inequalities. But in the highly disturbed conditions of the Antarctic one would not have been entitled to assume a priori that the absolute maxima, for instance, would have shown even a preference for the hour of the inequality maximum. Comparing, however, Tables LIV and XIII, we see that the daily maxima for D cluster thickly round the hour $9 \mathrm{a} . \mathrm{m}$., when the maximum presents itself in the mean diurnal inequality for the year. This is true of all the seasons, especially Midsummer. It is truly remarkable that at this season, when disturbances were so large, out of the 79 days for which data exist not one gave an absolute maximum during the 12 hours ending with $2 \mathrm{a} . \mathrm{m}$.

If we define the "concentration" of the frequency as the percentage which the occurrences during the three consecutive hours which, combined, give the greatest number of occurrences bear to the total number, we find in the case of the D maxima for the concentration 41 in Midwinter, 60 in Equinox, and 62 in Midsummer. The time of greatest frequency seems a trifle later in Equinox than in the other two seasons, but the difference is minute.
Table LIV.-Declination Maxima. Frequency of Occurrence

Table LV.-Declination Minima. Frequency of Occurrence.

Table LVI.-Horizontal-Force Maxima. Frequency of Occurrence.

Table LVII.-Horizontal-Force Minima. Frequency of Occurrence.

Table LVIII.-Vertical-Force Maxima. Frequency of Occurrence.

Table LIX.-Vertical-Force Minima. Frequency of Occurrence.


The law of occurrence of minima in D is not so simple. At all seasons, it is true, the greatest frequency of occurrence is from 5 to 7 p.m., the minimum in the diurnal inequality occurring at 6 or $7 \mathrm{p} . \mathrm{m}$. ; but there is, at least at Midsummer, a second maximum of frequency. This occurs, curiously enough, ahout 8 a.m., and so near the hour of the inequality maximum. The cause is probably the marked tendency observed at times for Declination disturbances to be particularly large and numerous from 5 to $10 \mathrm{a} . \mathrm{m}$. The figures for the "concentration" in the frequency, viz. Midwinter 54, Equinox 46, and Midsummer 30, show a marked seasonal contrast to those for the D maxima.

Comparing Tables LVI and XVII, we see that the H maxima mainly occur in the early afternoon, near the hour of the inequality maximum. The figures for the "concentration," viz. Midwinter 50, Equinox 54, Midsummer 51, indicate but little difference between the zeasons. The time of greatest frequency of occurrence appears to be somewhat later at Midwinter than at the other seasons; this is in accordance with what is seen in the case of the diurnal inequality.

In the case of the H minima there is a distinct indication, at least in Equinox, of a double maximum in the frequency of occurrence. The greatest concentration is found near the hour- 4 or $5 \mathrm{a} . \mathrm{m}$.-of the minimum in the diurnal inequality, but there is also a marked concentration near midnight. Taking the three consecutive hours of greatest frequency as before, the figures for the concentration are: Midwinter 28, Equinox 35, Midsummer 50.

The figures in Tables LVIII and LIX for the maxima and minima in $V$ are a little more irregular than those for D and H . The sensitiveness of the V instrument was so small that the trace, when quiet, appeared to the eye practically a straight line; also the difference between the temperature corrections to be applied at two hours was often in excess of the difference between the uncorrected values. Frequently five or six ordinates had to be measured, and the final differences obtained were not infrequently microscopic. Thus the results are less certain than for the other elements. There is clearly, however, at all seasons a tendency for the maximum to be most frequent near midnight, and the minimum near noon, i.e. near the hours of the inequality maximum and minimum. The figures for the "concentration" are-

|  | Midwinter. | Equinox. | Midsummer. |
| :---: | :---: | :---: | :---: |
| $V$ maximum | 33 | 53 | 34 |
| V minimum . . . . | 33 | 47 | 45 |

§45. A general idea of the incidence of maxima and minima will be most readily derived from Table LX, which gives the mean results from all the observations. To obtain greater smoothness, the day has been divided into two-hour periods, commencing with midnight. The figures are percentages of the total number of maxima or minima. The maxima of frequency are indicated by heavy type.

Table LX.-Percentage Frequency of Occurrence from all Observations.

| Two hours ending . . . |  | 2. | 4. | 6. | 8. | 10. | Noon. | 2. | 4. | 6. | 8. | 10. | Midt. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Declination . . $\{$ | Mexima | 1.4 | $4 \cdot 5$ | $7 \cdot 6$ | $25 \cdot 0$ | $39 \cdot 2$ | $15 \cdot 3$ | 3.8 | 0.9 | 0.2 | $0 \cdot 2$ | 0.0 | 1.8 |
|  | Mina |  | 13 |  | $4 \cdot 4$ | 31 | $2 \cdot 4$ | 31 | 8.6 | 251 | 28.4 | 13.0 | $5 \cdot 5$ |
| Horizontal Force. $\{$ | Maxima | $0 \cdot 3$ | $0 \cdot 1$ | $0 \cdot 1$ | 1.4 | $3 \cdot 0$ | $10 \cdot 6$ | 27.4 | $30 \cdot 3$ | $20 \cdot 7$ | $5 \cdot 1$ | $1 \cdot 1$ | $0 \cdot 0$ |
|  | Minima | $12 \cdot 2$ | $14 \cdot 8$ | 21.2 | $16 \cdot 4$ | $3 \cdot 9$ | $0 \cdot 3$ | 0.8 | $0 \cdot 3$ | $0 \cdot 3$ | $2 \cdot 1$ | $10 \cdot 7$ | $16 \cdot 9$ |
| Vertical Force. \{ | Maxima | 277 | 11.7 | $5 \cdot 4$ | $3 \cdot 2$ | $4 \cdot 1$ | 1 '6 | 0.9 | $1 \cdot 9$ | 5 \% | 8.8 | $12 \cdot 6$ | 17.0 |
|  | Minima | 13 | $0 \cdot 8$ | $3 \cdot 5$ | $6 \cdot 6$ | $20 \cdot 2$ | $23 \cdot 2$ | $19 \cdot 8$ | 12.0 | $5 \cdot 5$ | $0 \cdot 3$ | $2 \cdot 2$ | 4.4 |

## CHAPTER VII.

## Term-Hour and Special Records.

846. The scheme of international co-operation in magnetic work included observations to be made on certain "term days," two a month, and especially during certain term hours, one on each term day.

The results of the hourly readings taken during the term days-the 1st and 15 th of each month-have already appeared in "Physical Observations," pp. 160-179. The term days all started at Greenwich midnight. On the first of the regular term days, February 1, 1902, the term hour was $0-1$ a.m., G.M.T. On the second term day, February 15, 1902, it was 1-2 a.m., G.M.T., and so on, advancing an hour each time for the 24 term days, up to and including January 15, 1903. The scheme included two additional term days, February 1 and 15, 1903, the term hours on which were respectively $0-1$ a.m. and 1-2 a.m., G.M.T.

On the term hours it was intended that the magnetograph drums should be rotated at a higher speed than usual, so as to obtain a very open time scale.

To change the speed of rotation entailed the attendance of an observer during the term hour, and as many of the term hours-including in Europe those on the earlier term days-fell during the night, it is, perhaps, not surprising that the number of observatories which took quick runs during all the term hours was very limited. Of the results from the quick-run magnetographs those obtained at Batavia and Pola bave already appeared in the official publications of those observatories, where they can be consulted. At Christchurch Observatory quick-run records were obtained throughout the whole 24 hours of the term day and copies of the curves, together with tabulated results at 20 -second or 1 -minute intervals, were sent home by Dr. C. Coleridge Farr and Mr. H. Skey, the successive Directors of the Observatory. The labour this entailed must have been very great. At Mauritius quick runs were obtained during 10 term hours, and tracings of these and of the slow-run curves obtained during the other term hours were received from the Director, Mr. C. T. F. Claxton.

Quick-run curves were obtained at the National Physical Laboratory on all the term hours. Most of these were taken in the Observatory Department, Kew, but some were taken at Teddington, in the Physics Department.

As it so happened, the term hours were one and all exceedingly free from disturbances. When a magnetic disturbance of moderate size occurs, a quick-run curve has great advantages, as it admits of much higher precision in the measurements of time. But when magnetic conditions are very quiet, there is a great compensating disadvantage in that the gradient on the quick-run curve becomes so slight that there are no salient features left to catch the eye. There still may remain substantial advantages for tabulating purposes, but, so far as appeals to the eye are concerned, the quick-run curve becomes really inferior to the slow-run.

Even in 1902 the Kew curves, especially the Vertical-Force ones, were sensibly affected by electric-tram currents. These keep the Vertical-Force magnet in constant oscillation, and though the amplitude is not large it is sufficient to hide from the eye any natural change that is both small and slow. Even at Teddington some of the quick-run curves showed artificial disturbances.

Taking into account the total absence of any but the most insignificant of natural movements-other than the regular diurnal change-it was decided not to reproduce the curves obtained at Kew and Teddington.

An examination of the Mauritius curves showed that, as is usual at most places, the Declination and Vertical-Force curves exhibited less trace of disturbance than the Horizontal-Force curves. As all the curves were very quiet, it appeared unnecessary to go to the expense of reproducing them for all three elements, and the Horizontal Force was selected as the element most worthy of reproduction. The Horizontal-Force curves for all the term hours available appear in Plates I and II.

The copies of the Mauritius curves sent by Mr. Claxton were traced by Mr. Brookes with the curvetracer designed by Dr. Ad. Schmidt, and the results from the different term hours were arranged so ats to appear in juxtaposition.
§47. In Plate I the successive curves appear one below the other so as to have a common time scale; the vertical cross lines answer to $10,20 \ldots 60$ minutes from the commencement of the term hour. The date and the term hour-e.g. 6-7. a.m. on May 1, 1902-appear on the left-hand margin. This also shows the absolute value of the Horizontal Force at the commencement of the term hour, as given by Mr. Claxton's own measurements ( $6 f$. the table on p. 174 of "Physical Observations"). The scale value ( $1 \mathrm{~mm} .=2.93 \gamma$, or $20 \gamma=7 \mathrm{~mm}$. approx.) is also shown at the side. Motion down the sheet means increase of force, as is indicated by the arrows.

On May 1 and June 1 the trace was lacking for a few minutes at the commencement of the hour.
At Mauritius the drums, when quick run, were rotated at twelve times the usual rate. Normally the time scale is about $15 \frac{1}{2} \mathrm{~mm}$. to an hour, and so in the quick-run curves it is about 186 mm . to the hour. Plate I shows the curves natural size.

To the eye the departure of the curves from straight lines is not very conspicuous. It is, in fact, so small that it was deemed inexpedient to draw horizontal lines across the sheet to show the absolute values. If anyone wishes to measure variations in the ordinate he may employ as base line the time line at the top. The absolute value of this for the curve of any one hour can be readily arrived at by reference to the absolute value assigned to the commencement of the hour in question.

Plate II gives the slow-run Mauritius Horizontal-Force curves for the remaining term hours for which records were available. To show the trend of the curve, the record extends for about 20 minutes on either side of the term hour, except on October 1, when the original record stopped immediately after the end of the hour. On April 15, 1902, and February 15, 1903, there were small gaps in the original, representing probably the interval required to change the photographic paper on the drum. The curves in Plate II are all very quiet, but owing to the small time scale they appear less quiet to the eye than the quick-run curves in Plate I. On February 1, 1902, there was quite an appreciable amount of variation, but this unfortunately was prior to the arrival of the "Discovery" at Winter Quarters. The curves of October 15, 1902, and January 1, 1903, show some small oscillations.

The original curves had, of course, separate base lines having different absolute values. To economise space, the curves have been so placed in Plate II that a common base line serves for all in one horizontal row. There are thus only three base lines shown, the top one serving for the term hours of the four days February 1 and 15, and March 1 and 15, 1902. In addition to the base line for each group of curves there is a second straight horizontal line, answering to a value of the force exceeding that for the base line by $50 \%$. As in the case of the quick runs, the curves are as nearly as possible exact copies of the original, 1 mm . of ordinate representing $2.93 \gamma$. As indicating the truly remarkable absence of disturbance, it may be remarked that there is not one of the curves reproduced which shows a range exceeding $6 \gamma$ during the term hour.
§48. At Christchurch, as already stated, quick-run traces were available not merely for the term hours, but for the term days. It was decided, however, for reasons subsequently stated (App. A), to reproduce only two specimens of the term-hour records, and to deal with the Christchurch tabulations in a special Appendix.

The curves selected for reproduction were those taken during the term hours of June 15 and July 1, 1902. Their selection was due to the fact that on these days there were corresponding quick-run curves from the Antarctic.

The Christchurch curves for the term hours of June 15 and July 1 appear in Plate III. To economise space, a single time or base line is drawn to serve for the three elements from each hour. The Declination, Horizontal-Force, and Vertical-Force traces are distinguished by the letters D, H, and V respectively. The absolute value of each element at the commencement of the hour and the corresponding scale value are shown at the left margin. The arrows assist in indicating the direction in which the elements increase (at Christchurch, down the sheet answers in all the elements to an increase). The curves in the photographic copies sent from Christchurch were copied by Mr. Brookes with the Schmidt tracer. The
scale of ordinates is as in the originals, 1 mm . representing $1^{\prime} \cdot 125$ in Declination, $4 \cdot 53 \gamma$ in Horizontal Force, and $3.18 y$ in Vertical Force. The slow-run Christchurch time scale is practically the same as at Mauritius and Kew, but the quick-run scale was more open. This was reduced by the tracer so as to be approximately the same as in the Mauritius curves.

On July 1, trace was lacking for about a minute near the beginning of the term hour.
Of the curves themselves it is perhaps sufficient to say that it requires a very acute eye to detect their difference from straight lines.
§49. In the Antarctic, quick-run curves were obtained for the whole of the term hour on June 15, 1902 ( $9-10$ a.m., G. M.T.), for the greater part of the term hour on July 1, 1902 ( $10-11$ a.m., G.M.T.), and for a short portion of the term hour on July 15, 1902.

Plate IV shows the records obtained on June 15. They were copied from the originals by Mr. Brookes with the Schmidt curve tracer, leaving the scale of ordinates unaltered, but reducing the time scale by about 25 per cent., so as to make it approximately the same as in Plates II and III. $\mathrm{D}_{0} \mathrm{D}_{0}, \mathrm{H}_{0} \mathrm{H}_{0}, \mathrm{~V}_{0} \mathrm{~V}_{0}$ are the base lines for the Declination curve DD, the Horizontal-Force curve HH, and the temperature curve TT respectively. The Vertical-Force magnet was out of action during the hour in question. Vertical lines are drawn answering to 10 -minute intervals, the times from the commencement of the hour being marked on the Declination base line. The absolute values answering to the base lines and the scales for the Declination and Horizontal-Force curves are shown on both margins. Fine horizontal lines are drawn to aid the eye in estimating the extent of the movements, and arrows assist in indicating the direction in which the elements increase.

The alteration in the elements is sufficient to make the open time scale of real value. At the same time it should be noticed that the Horizontal-Force magnet was unusually sensitive, so that 1 mm . of ordinate answers to only $1 \cdot 4 \gamma$. In the case of the Declination the scale is less open, 1 mm . answering to $1^{\prime} \cdot 5$-in force about $2 \cdot 9 \gamma$-and the change during the hour in Declination is considerably more important than that in Horizontal Force.

Plate V shows the records obtained in the Antarctic during the term hour of July 1. They were treated in the same way as those for June 15. In this case there is a record VV from the Vertical-Force magnet.

There are two gaps in the curves. The first, occurring from about 3 to $4 \frac{1}{3}$ minutes after the commencement of the hour, was due to the light being intercepted by the clamp attaching the paper to the drum. The second, occurring from 26 to 30 minutes after the commencement of the hour, represents the interval occupied in changing papers. The Declination and Horizontal-Force scales have the same value as for June 15. The changes in these elements are somewhat larger than on the previous term day. The departure of the V trace from a straight line is not apparent to the eye. The changes in this element were really less than for the other two, but the difference is exaggerated owing to the fact that the Horizontal-Force scale is almost ten times as open as that of the Vertical Force, for which $1 \mathrm{~mm} . \equiv 13 \cdot 5 \gamma$.
$\S 50$. Plates VI to X show the Antaretic records during the remainder of the term hours. April 1 was the earliest term day for which a satisfactory record was obtained. The curves appear in chronological order, omitting, of course, the two days already considered, and likewise the two May term-day hours for which no record existed. The date is shown at the foot, and marks auswering to the beginning and end of the term hour appear on all the base lines. The hours (G.M.T.) are recorded on one only of the base lines. As before, the base lines are distinguished by the suffix 0 , and the letters $\mathrm{D}, \mathrm{H}, \mathrm{V}$, and T indicate Declination, Horizontal Force, Vertical Force, and temperature. The absolute values answering to the base lines and the scale values are shown on the left margin, and arrows assist in indicating the directions in which the elements increase (Vertical Force up the sheet, Declination and Horizontal Force down the sheet). To show better the general trend of the curve, the portions of curve reproduced represent in most cases about two hours, including half-an-hour before and half-an-hour after the true term hour. There are, however, exceptions to this on January 15, 1903, when the original record stopped before the end of the term hour, and on July 15, 1902. On the last-mentioned date the trace is restricted to the exact term hour-11 a.m. to noon, G.M.T. It consists of three portions. Starting at 11 a.m., we have a quick run covering about $I_{4}$ minutes, then a gap corresponding to a minute, due to the clamp intercepting the light,
then another quick-run portion covering about $5 \frac{1}{2}$ minutes, then a gap representing the time occupied in changing papers, and, finally, a slow-run portion answering to the period $11.18 \mathrm{a} . \mathrm{m}$. to noon. It is interesting to compare the quick- and slow-run portions of the $D$ and $H$ curves. At first sight one would hardly realise that the curves were really as quiet towards the end of the hour as at the begiming.

In all the curves the Declination has the same scale value, $1 \mathrm{~mm} . \equiv 1^{\prime} \cdot 5$. In the other elements there is considerable variation. In H the value of 1 mm . varies from $1.4 \gamma$ to $0.67 \gamma$, in V it varies from $16.5 \gamma$ to $3 \cdot 2 \gamma$. Full particulars will be found in Table I, or can be derived from the plates themselves. The curves in Plates VI to X were obtained as follows:-

The original Antarctic curves were photographed at Kew by Mr. W. J. Boxall, light being transmitted through the sheet on to photographic paper immediately in contact with it. From the negatives thus formed positives were made, and when unduly faint these were intensified with pen and ink. In some cases the base lines shown represent exactly the originals, but in other cases the original base lines were replaced by others occurring in greater juxtaposition to the curves to which they really belonged. This was done to minimise the risk of confusion between the different elements.

The latter part of the work was done by Mr. Brookes at Teddington, under the general supervision of Mr. F. J. Selby. Every care was taken to avoid introducing departures from the original curves.

In some cases, e.g. November 1 and 15, and December 15, 1902, there is, unfortunately, a good deal of intercrossing of the D and H curves. It is hoped, however, that by the aid of the letters attached it will be found possible to distinguish the traces.

The original V trace was lacking on October 1 and November 15, 1902, and on February 1, 1903, while the H trace was off the sheet during the term hours of January 1 and February 15, 1903.

The curves, being exact copies of the originals, are, of course, uncorrected for temperature, and, as will be seen from the TT curves, when these exist, temperature was far from constant. As already explained, the Vertical-Force magnet had a large temperature coefficient; thus a very appreciable temperature correction would in reality be necessary to bring the values of V derived from measurements of the V curves in some of the hours to a common basis. In other words, the base-line value in V was not really constant, and the values actually assigned it refer to only one instant of the hour.

The temperature variations seldom, however, showed rapid oscillations, and whilst the temperature change is, in some cases, responsible for a slight general drift up or down the sheet in the $V$ curve, it is practically without effect on the range of the oscillations shown in the curves.
§51. Owing to the difference in the type of the Antarctic magnetographs from those in use at Christchurch or Mauritius, the difficulty of identifying corresponding movements is materially increased. The difficulty is further enhanced by the fact that we have, in the case of the Antarctic, high sensitiveness in the instruments going along with large disturbance. But whatever the cause may be, I have not succeeded in tracing any close parallelism between individual Antarctic term-hour movements and those at either Mauritius or Christchurch. There is, however, a certain amount of parallelism between the Antarctic and Christchurch curves, as may be seen on reference to Table LXI, which gives the ranges of Declination and of Horizontal Force at the two stations during all the term hours for which information existed. The Christchurch ranges are derived from tabulations, at 20 -second intervals, made at Christchurch. The parallelism referred to consists in this, that on the term hours for which the range in the Antarctic is exceptionally low the range was also especially small at Christchurch. The cause, however, so far as Christchurch is concerned, has little to do with disturbance. The hours for which there were specially low ranges in the Antarctic occurred in Midwinter near midnight, and as midnight in the Antarctic and midnight at Christchurch differed by only about 23 minutes, the hours in question were naturally those for which the regular diurnal changes at Christchurch were least.

Table LXI.-Term-Hour Ranges.


The means at the foot of Table LXI are derived from all the term hours. Omitting the Declination for November 15, and the Horizontal Force for November 15, December 1, January 1, and February 15, so as to have only days for which both Antarctic and Christchurch records were complete, we obtain the following results for the hourly ranges:-

|  | Declination. |  | Horizontal Forse, in $\gamma$. |
| :---: | :---: | :---: | :---: |
|  | Arc. | Equivalent in $\gamma$. |  |
|  | , |  |  |
| Antarctic mean . . . . . | $28 \cdot 2$ | 54 | $20 \cdot 5$ |
| Christchurch mean . . . . . | 1.04 | 6.9 | $3 \cdot 1$ |
| $\frac{\text { Antarctic }}{\text { Christchurch }} \text { ratio }=$ | $27 \cdot 2$ | 79 | $6 \%$ |

The enormous disparity between the Antarctic and Christchurch results is partly due, of course, to the fact that the range of the regular diurnal inequality is so much larger at the former station. But this only accounts for part of the difference, as may readily be seen by reference to the hourly values on term days given in "Physical Observations," pp. 175-178. The range of the regular diurnal inequality derived from a number of days, even quiet days, is always materially less than the arithmetic mean of the ranges from the individual days, the difference being greater the more disturbed the days. Thus the ratio of the range of the regular diurnal inequality at Winter Quarters to that at Christchurch will be less than the ratio derived when instead of the diurnal inequality ranges one takes the arithmetic mean of the ranges on individual days.

In the Physical Observation tables there are 16 days for which Declination ranges are given at both stations and 12 for which Horizontal-Force ranges are given (the curve not being beyond the limit of registration). Deriving arithmetic means from these, wo find the following results for the duily ranges :-

|  | Declination. | IIorizontal Force. |
| :---: | :---: | :---: |
| Antarctic ${ }_{\text {Christchuroh }}$. . . . . . . | $\begin{array}{r} 84 \\ 6 \cdot 8 \end{array}$ | $\begin{aligned} & \gamma \\ & 62 \\ & 25 \cdot 8 \end{aligned}$ |
| $\frac{\text { Antarotic }}{\text { Christohurch }} \text { ratio }=$ | $12 \cdot 4$ | $2 \cdot 4$ |

The excess of the Antarctic figures is here very much less than for the hourly ranges.
§52. Plate XI gives a copy of a quick-run record obtained during a solar eclipse on September 21, 1903 (September 20, G.M.T.). According to Mr. Bernacchi's calculations the local times of the beginning, the central phase, and the end of the eclipse were respectively 3.40 p.m., $4.38 \frac{1}{2}$ p.m., and 5.37 p.m. The quick run was started at about 3.39 p.m. and continued until 5.46 p.m. The curves shown extend from $3.39 \mathrm{p} . \mathrm{m}$. to about $5.39 \mathrm{p} . \mathrm{m}$. They are copies taken with the Schmidt tracer, retaining the original scales for the ordinates, but reducing the time scale about 25 per cent.

The curves are shown in two portions, the upper panel representing the earlier half. The base lines $\mathrm{D}_{0} \mathrm{D}_{0}$ and $\mathrm{H}_{0} \mathrm{H}_{0}$ have been shifted from their positions in the originals so as to bring them within the range of the plate. The Declination, or D , curve and the Horizontal-Force, or H , curve cross one another in the second panel. The Declination scale is shown on the left, the Horizontal-Force scale on the right. The Vertical-Force curve to the eye was practically a straight line and so has been omitted along with the corresponding base line and the temperature curve to avoid confusing the trace.

The primary reason for reproducing the curves is that it is a question of interest whether an eclipse of the sun does exercise an appreciable influence on the variation of the magnetic elements. On various occasions there have been schemes of special observations during eclipses, and several authorities have been inclined to think that the results have shown a diminution in the changes as compared to those ordinarily encountered at the same hour of the day.

The curves in Plate XI show considerably more disturbance than those on June 15 and July 1, 1902, in Plates IV and V. They answer, however, to a time of year when the changes, both regular and irregular, were larger. It can be readily realised how useful open time-scale curves like these might prove for comparing disturbances of considerable magnitude at different stations.
§53. Plate XII gives an example of a series of regular waves, or "pulsations" as they have been very appropriately termed by Dr. van Bemmelen. The curves are exact copies of Antarctic quick-run curves of February 2, 1903. The letters D, H, V, T have their usual significance, the base lines being distinguished by the suffix 0 . The scale values appear at the left-hand side.

The pulsations are most apparent in the Horizontal-Force curve between 9.10 and 9.25 p.m. But they may be detected-at all events in the original-throughout most of the H curve between 9 and $10 \mathrm{p} . \mathrm{m}$., and can be recognised distinctly in the D curve about 9.15 p.m.

The special feature of pulsations is that they represent wave-like oscillations of force not varying much in amplitude and of at least approximately constant period during a considerable number of wave periods. In the present case there seem to be about 19 complete periods in 1 cm ., and, as the time scale is about 240 mm . to the hour, the period is accordingly about 8 seconds.

Very short-period pulsations cannot be seen except in open time-scale curves, but pulsations of 2 or 3 minutes' period are not infrequently seen in ordinary curves from Kew pattern magnetographs. The Eschenhagen type of magnetograph is better adapted for showing pulsations, but the Antarctic conditions were so disturbed that pulsations with periods of 2 or 3 minutes would be difficult to recognise
with certainty. There was, perhaps, only one slow-run trace on which they were distinctly recognisable, that of August 6, 1902. On that day pulsations were distinctly visible in both the D and H curves during most of the time between $8 \mathrm{a} . \mathrm{m}$. and $3.30 \mathrm{p} . \mathrm{m}$. They were distinct enough at times, the day being unusually free from large disturbances, to admit of an approximate determination of the period from both the D and H curves. The results obtained showed no certain variation in the period, the mean obtained being 1.9 minutes, i.e. there were about 32 complete pulsations in the hour.

Several quick-run curves showed pulsations of short period. On June 16, 1902, pulsations were visible in the D and H curves between 8.30 and 8.40 p.m., having a period of about 27 seconds, and again with about the same period from 10.30 to $10.40 \mathrm{p} . \mathrm{m}$. On August 15, 1902, pulsations appeared at intervals in both the $D$ and $H$ curves from $11.30 \mathrm{a} . \mathrm{m}$. to $2.10 \mathrm{p} . \mathrm{m}$. For 8 minutes, from 11.40 to 11.48 a.m., they had quite a considerable size, the maximum value of the double amplitude being at least 8 mm . (i.e. $12^{\prime}$ ) in D and 6 mm . $\left(8^{\circ} 4 \gamma\right)$ in H . During most of the time the amplitude was no larger than in Plate XII. The periods derived from the D and H curves were the same, about $7 \frac{1}{2}$ seconds, and appeared independent of whether the amplitude were large or small. This, of course, strongly supports the view that the largeness of the amplitudes between 11.40 and 11.48 a.m. was due to natural as opposed to instrumental causes.

On August 16, 1902, exceedingly small but very regular pulsations with a period of 7 or 8 seconds could be made out in the H curve from noon to 1 p.m. At one or two points the edge of the D curve seemed also serrated, but the movements were too small to be certain of.

Pulsations were also visible in the H curves of September 1, 1902, from 2.35 to $4.20 \mathrm{p} . \mathrm{m}$., and on those of September 16,1902 , from 10.5 to $11.0 \mathrm{p} . \mathrm{m}$. On both occasions the period was about 8 seconds. On September 1 the $D$ curve also showed a trace of pulsations. On September 16 there seemed a faint indication of minute pulsations in the H curve later in the day between 3.0 and $3.20 \mathrm{p} . \mathrm{m}$.

The above are the only cases in which pulsations were detected. As to whether they are more or less numerous in the Antarctic than elsewhere it would be impossible to say without a very minute investigation. It was really only when the magnetographs were quick run that pulsations were at all likely to be detected. The number of hours' trace in the Antarctic from quick-run curves was approximately 130, but this represents but a very small fraction of the time during which the magnetographs were in action.
§54. The curves in Plate XIII represent one of the quietest times, if not the very quietest time experienced in the Antarctic. The time represented is from 11 p.m. on June 27, 1903, to 9 a.m. on June 28. The slight movement that is apparent in the V curve is mainly due to temperature, there being a slight rise of temperature towards the end of the period. There were only 2 or 3 days' curves which even distantly approached those shown for quietness. Undisturbed as the D and H curves on June 27-28 are, when considered relatively to other Antarctic curves, they display a practically continual succession of small movements such as would not appear on an ordinary quiet day at Kew. The curves at Christchurch corresponding to those in Plate XIII were also very quiet. At Kew the time was also distinctly quiet, but not more so than on an average quiet day.

In the evening of the same day there was a large disturbance in the Antarctic, which is described later.

## CHAPTER VIII.

## Records of Disturbances from Co-operating Stations and tie Antarctic.

§55. What I had seen of the quick-run curves, obtained at Kow on the international term days of 1902 and 1903, made me anticipate that, so far as the inter-comparison of disturbances at different stations was concerned, the term-hour data were likely to prove of little service. Examination of the results obtained elsewhere, so far as available, and of the Antarctic records only served to confirm this. Accordingly I came to the conclusion that if any comparison of disturbances was to be carried out, additional data must be obtained. With this object, a list was made of days during 1902 and 1903 on which it appeared probable, from a consideration of the Kew and Antarctic curves, that comparative data from different stations would be valuable, and in January, 1905, a circular was issued, in the name of the Royal Society's Magnetic Committee, requesting that copies of the magnetic curves obtained on the days in question should be transmitted. This was sent to Dr. C. C. Farr, Director of Christchurch Observatory, New Zealand; Mr. C. T. F. Claxton, Director of the Royal Alfred Observatory, Mauritius; and to Mr. N. A. F. Moos, Director of the Colaba Observatory, Bombay. These gentlemen all gave favourable consideration to the request, and sent home a number of most valuable records. The copies received from Christchurch and Colaba were photographic, those from Mauritius were tracings of the original. After considering all the material, I made a selection of the disturbances which presented most points of interest, and these are dealt with here. Even in 1902 the Kew curves, especially those of Vertical Force, were sensibly influenced by electric-traction currents, and accordingly application was made to Mr. E. Kitto, Superintendent of Falmouth Observatory, for the loan of the Falmouth magnetic curves for the days selected. Mr. Kitro and the Committee of the Falmonth Observatory responded favourably to this request so far as was possible. The Falmouth Vertical-Force magnet was unfortunately out of action during part of 1902, so that the information for that element during this time was derived from Kew alone. So far as the disturbances here considered are concerned, the differences between the Kew and Falmouth curves are mostly infinitesimal. In general, I measured both sets of curves. In the discussion I chiefly use the Kew measurements, but the plates show the Falmouth curves. It will be most convenient to describe first the plates dealing with curves from Falmouth, Colaba, Mauritius, and Christchurch, and then the plates dealing with the corresponding Antarctic curves and some additional onee.

## Falmouth, Colaba, Mauritius, and Christchurch Disturbed Curves.

§56. The magnetographs in use at Falmouth, Colaba, Mauritius, and Christchurch are of what is known as the "Kew "pattern. The time scale of the present Kew magnetograph is 1 hour to 0.6 inch ( $15 \cdot 25, \mathrm{~cm}$.), and this, presumably, was what was aimed at in the case of the magnetographs at the four stations mentioned above. As a matter of fact, the time scales, though very nearly equal, are not absolutely so. Still, the difference is so small that it does not readily catch the eye when the curves are juxtaposed, unless one extends the comparison to a large number of hours. This enables an effective comparison to be readily made without modifying the scale values.

The curves in Plates XIV to XXI were actually derived from copies of the curves received from Falmouth, Colaba, Mauritius, and Christchurch, which were made with the Schmidt curve tracer. The tracing was done at Bushy House by Mr. Brookes, under the supervision of Mr. F. J. Selby. The scales, whether of ordinates or of abscisse, can be altered with the Schmidt tracer at the will of the operator, but, as a matter of fact, no intentional departure was made from the scales in the original curves. The tracings were made on separate sheets of paper, which were then arranged one below the other. In the originals the time is shown by breaks occurring every 2 hours in the time line, either at the even or at the odd local hours. The original time lines were dispensed with and replaced by others in which hours G.M.T. are indicated by short transverse lines. These lines were put on with considerable care, but they cannot claim to be more than approximately correct.

Further, ingenious as the Schmidt machine is, no ordinary operator can produce with it an absolutely perfect copy of the original. It is also not so well adapted for use with the broad trace of the Kew pattern magnetograph as the finer trace of the Fischenhagen instrument. It can deal, moreover, much less satisfactorily with rapid to-and-fro movements than with those of a rounded character. Thus Plates XIV to XXI are more appropriate for general descriptive purposes than for any minute investigation of details with a view to numerical calculations, and no use has been made of them for the latter purpose. Partly for this reason, and partly with as view to avoiding an excessive multiplication of lines, the ordinate scales are shown only at the left-hand margin of the curves. The Horizontal- and Vertical-Force scales varied slightly with the time, but the following values are sufficiently close considering the nature of the curves:-

|  | Value of 1 mm . of ordinate. |  |  |
| :---: | :---: | :---: | :---: |
|  | Declination. | Horizontal Force. | Vertical Force. |
| Filnouth <br> Colaba <br> Mauritius <br> ("histehureh | $\begin{aligned} & 1 \cdot 1 \\ & 1 \cdot 1 \\ & 1 \cdot 1 \\ & 1 \cdot 1 \end{aligned}$ | $\begin{aligned} & \gamma \\ & 5 \cdot 0 \\ & 5 \cdot 1 \\ & 2 \cdot 9 \\ & 4 \cdot 6 \end{aligned}$ | $\begin{gathered} \gamma \\ 5 \\ 16.5 \\ 5 \cdot 0 \\ 2 \cdot 7 \text { to } 3 \cdot 2 \end{gathered}$ |

At all four stations the Horizontal-Force curves were those which were most affected by disturbances, whilst the Vertical Force ones were the least affected. Thus the curves shown are mostly those of Horizontal Force.

As indicated by the scale figures and by the arrows, there is considerable variety in the procedure adopted at the different stations. At Falmouth and Colaba, as at Kew, movement up the sheet denotes increase in all three elements. At Mauritius, movement up the sheet means increase in Declination, but diminution in Horizontal or Vertical Force. At Christchurch, movement up the sheet means decrease in all three elements. In deciding, as I eventually did, not to invert any of the traces, I was influenced by several considerations.

In the Antarctic curves, which it was decided to reproduce photographically-the difficulty of obtaining reasonably accurate results with the Schmidt tracer appearing excessive-movement up the sheet meant increase only for the Vertical Force. The three elements were in this case all on one sheet, and to have retained one unaltered while inverting the others presented serious difficulties. Thus absolute uniformity of procedure was unattainable unless one was prepared to take a great deal of trouble, and to take much trouble appeared justifiable only if it were clear that the result was a decided gain. On the other hand, reflection showed that in any case only a semblance of uniformity was obtainable. Declination is westerly at Falmouth and Mauritius, but easterly at Colaba and Christchurch, thus an increase in the element means movement of the north end to the west at the two former stations, but to the east at the latter two. In the Antarctic Declination was easterly, but exceeded $90^{\circ}$; thus increase of the element meant increase of easterly Declination, but movement of the $\mathbf{N}$. pole to the west. At Falmouth and Colaba, where the N. pole dips, increase of Vertical Force means an additional force on the N. pole directed towards the Earth's centre; but at Mauritius and Christchurch, where the S. pole dips, it means a force urging the N. pole from the Earth's centre. As regards the Horizontal Force, the magnetic Meridians of the various stations varied from $18^{\circ} \mathrm{W}$. at Falmouth to $16^{\circ} \mathrm{E}$. at Christchurch and $152^{\circ}$ E. in the Antarctic. Thus the planes in which the Horizontal Force is measured were related in very different ways to the local astronomical Meridian.

Finally, there was the consideration that horizontal and vertical, east and west, are terms of purely local significance, and indicate directions whose relations to any one system of rectangular axes vary with the latitude and longitude of the station.
§57. Plates XIV to XXI follow the chronological order. Plate XIV (compare Plate XXII) shows a disturbance which seems to have been felt all over the world. It lasted for practically 8 hours, the termination being nearly as clearly indicated as the commencement. The ihree lower, Horizontal-Force, curves show
a sudden commencement at about 11.59 a.m., G.M.T., on May 8, 1902. At Kew, Falmouth, Colalia, Mauritius (curve not shown), and Christchurch the initial change in H was an increase, which continued for about 10 minutes. The Falmouth curve in Plate XIV shows a short stoppage of the upward motion at about 12.4. This stoppage actually exister, being as apparent in the Kew as in the Falmouth curve, but it has been somewhat exaggerated in the tracing. The original Colaba curve gives just a suggestion of a corresponding stoppage. It is not shown in the Mauritius curve (a tracing), nor on the Christchurch curve. The latter, however, is somewhat fuzzy, and a very short stoppage might not leave a visible effect. To the nearest $1 \gamma$ the total increase in $H$ from $11.59 \mathrm{a} . \mathrm{m}$. to 12.9 was $18 \gamma$ at Kew, $19 \gamma$ at Colaba, $9 \gamma$ at Mauritius, and $9 \gamma$ at Christchurch. The corresponding changes in Declination (W. to west, E. to east) were at Kew $1^{\prime} \cdot 3 \mathrm{~W}$., Colaba $0^{\prime} \cdot 7 \mathrm{E}$., Mauritius $0^{\prime} \cdot 7 \mathrm{~W}$., and Christchurch $0^{\prime} \cdot 2 \mathrm{E}$. The disturbance took place at an hour when the regular diurnal change at Kew and Falmonth was rapid and the elimination which was made of the regular diurnal change was, doubtless, only approximately correct.

The changes in Vertical Force were at Colaba $-7 \gamma$ and at Mauritius $+5 \gamma$. No Falmouth $V$ curve was available. At Kew no change was detected, but a small change might have been hidden by electrictram currents. At Christchureh the Vertical Force was, if anything, diminished, but if so, only to the extent of about $0.5 \gamma$. At Colaba the Vertical-Force scale was $16.2 \gamma$ to 1 mm ., so that an error of 0.2 mm . in the measurement of the ordinate means an error of $3 \gamma$ in force.

The initial change of force at $11.59 \mathrm{a} . \mathrm{m}$. is not the only one common to the different stations. A very similar movement, occurring from about 7.4 to $7.11 \mathrm{p} . \mathrm{m} .$, G.M.T., is prominent in the Kew, Falmouth, Colaba, and Mauritius curves, and seems recognisable also at Christchurch. This was followed by a decrease in H, which is very prominent at Kew, Falmouth, and Colaba, but which at Christchurch is interrupted by oscillations. We have here an excellent example of the difficulty of getting comparative results from different stations. To do this satisfactorily, one requires a movement practically continuous in one direction, occurring simultaneously in all the elements at all the stations. One can obtain a variety of such corresponding points-peaks and hollows as Dr. Balfour Stewart called them-on the Kew and Colaba curves, but when it comes to bringing in Mauritius and Christchurch it is comparatively seldom that turning-points common to all can be found. In the present instance only three movements could be found at all suitable for comparison. These answered to the intervals $11.59 \mathrm{a} . \mathrm{m}$. to 12.9 , 2.10 to 2.28 p.m., and 7.4 to 7.11 p.m., and even in their case one could not feel absolutely certain that the movements in all the elements at all the stations were really synchronous.

The magnetic storm illustrated in Plate XIV has received considerable prior attention owing to the fact that its commencement began within 4 or 5 minutes, possibly less, of the beginning of the eruption of Mont Pelée, which devastated St. Pierre in Martinique. This fact was pointed out by Dr. Bauer,* dealing with magnetic curves from American observatories.

The shortness of the interval between the commencement of the magnetic storm and the eruption-not to speak of the character of the movements-showed that it was not a case of mechanical shock due to seismic waves, and Dr. BaUer's remarks suggest the possibility of a true magnetic action at a distance due directly to the eruption. This was, I think, only a suggestion. Dr. Bauer's final opinion on the subject I have not seen. My own impression is that the coincidence was merely accidental. The reasons for this view will appear later ( $\$ 100$ ).
§58. Plate XV shows Horizontal-Force movements referring to three different disturbances, those of August 20, November 6, and November 24, 1902.

The storm of August 20 was ushered in, like that of May 8, by a sudden movement. This commenced at 9.6 p.m., G.M.T., and terminated practically in 4 minutes. During this time H increased by $14 \gamma$ at Kew, $10 \gamma$ at Colaba, $6 \gamma$ at Mauritius, and $4 \gamma$ at Christchurch. The changes in the other elements were mostly almost too small to measure, being in $\mathrm{D} 0^{\prime} \cdot 6 \mathrm{~W}$. at Kew, $0^{\prime} \cdot 2 \mathrm{E}$ 。 at Colaba, $0^{\prime} 0$ at Mauritius, and $0^{\prime} \cdot 6 \mathrm{~W}$. at Christchurch. In V there was an increase at Kew of about $1 \gamma$, and a fall at Colaba of under $2 \gamma$; no change could be detected at Mauritius or Christchurch. The movements between 9.6 and $9.10 \mathrm{p} . \mathrm{m}$. at the several observatories appeared synchronous, and there was no visiblo arrest of the movement corresponding to that seen in the Kew and Falmouth curves of May 8.

* 'Terrestrial Magnetism,' vol. 7, 1902, p. 57.

The disturbance on November 6 had also a sudden commencement at about 3.52 p.m., G. M.T., but in this case the commencing movement was of a distinctly double character, this being very clearly shown by the Falmouth curve which is reproduced. The first movement in $H$ was a decrease, lasting for about $1 \frac{1}{2}$ minutes, and it was followed by a notably larger increase. The upward movement will be seen to hesitate, and then diminish in steepness. The summit at Falmouth was reached about 3.58 p.m. The amplitude of the first change was about $1 \gamma$, that of the second about $9 \gamma$, but exact measurement of the first movement was difficult. The Falmouth D trace shows a corresponding oscillation, a movement of about $0^{\prime} .5 \mathrm{~W}$. being followed by a larger but slower movement of about $0^{\prime} \cdot 9 \mathrm{E}$. There was no Falmouth V curve; no measurable change appeared in that at Kew. The H and D curves at Kew appeared identical with those at Falmouth.

On November 24, as appears from the lowest curve in Plate XV, there was a disturbance of considerable magnitude at Christchurch. It did not, however, have a very definite commencement, and identification of correspouding movements at Christchurch and Falmouth could not be satisfactorily carried out.
§59. Plate XVI shows what appears to have been the most outstanding commencement exhibited by any of the magnetic storms experienced during 1902-3. The top and bottom curves are Declination ones, the others are of Horizontal Force. In the latter element, the initial movement was an increase at all the stations whose curves were reproduced. The copy of the Falmouth curve in Plate XVI suggests the presence of a small tremor before the upward movement commenced, but this tremor is very doubtfully shown in the original, either at Falmouth or at Kew. Neglecting it, the initial movement commenced at 11.25, and lasted until 11.29 p.m. on April 5. In the case of the Declination, the commencing movement at Kew and Falmouth was unquestionably double. For the first half minute or so, there was a small easterly movement, and then a much larger westerly movement, culminating at 11.29 p.m. Comparing the conditions at 11.25 and 11.29 p.m., we have for the changes in the three elements the following approximate values :-


Only Horizontal-Force curves were available for Mauritius.
Of the gaps apparent in the curves most are due to interruption of the trace by the time-shutter. The longer one at Mauritius represents the changing of the photographic paper.

The disturbance was most active between 2.30 and $5.30 \mathrm{a} . \mathrm{m}$. on April 6th, but it was by no means large at any time, the Declination range at Kew and Falmouth being only about 25'.
§60. Plate XVII shows some curves whose chief interest lies in connection with some Antarctic curves, which will be discussed presently.

The very symmetrical-looking movements shown by the $\mathrm{D}, \mathrm{H}$, and V traces at Christchureh, on June 19, 1903, occurred during a day which was otherwise very quiet.

The Christchurch curves of July 26, 1903, show what was for that station a very considerable disturbance. They illustrate what seems a general feature-the relatively small range of Vertical-Force disturbances at Christchurch, compared to those in the other elements.

The same phenomenon is also exhibited by the curves of Plate XVIII, which show the disturbances at Christchurch on August 22, 1903. This was one of the largest disturbances exhibited by the curves received from Christchurch. The storm showed no very definite beginning or end. The most distinctive movements are those occurring about $5 \mathrm{a} . \mathrm{m}$. and $9 \mathrm{a} . \mathrm{m}$. But even in their case the turningpoints for the different elements do not synchronise satisfactorily, and no satisfactory comparison proved possible with the curves at other stations.
§61. Plate XIX refers to a storm recorded on August 25-26, 1903. It shows the D, H, and V curves for Christchurch, and the H curves for Colaba and Falmouth. The largest movements did not occur until nearly 9 bours after the commencement, and to bring the principal part and the beginning within the range of one plate the trace is omitted from about 3.30 to 6.0 a.m.

The Falmouth and Colaba traces show a sudden commencement. At Falmouth and Kew the commencement, like that of April $\overline{\mathrm{j}}, 1906$, showed a distinct double movement, a very small diminution in H being followed by a large increase. The Declination curves at these two stations also show a double movement, a small easterly movement being followed by a somewhat larger but still small westerly one. The double movement occupied from 10.57 to $11.3 \mathrm{p} . \mathrm{m}$. on the 25 th , the to-and-fro movements occupying about equal times.

The Colaba H curve also shows at least a suggestion of a double movement, an increase of some size following a very small decrease, the two corresponding in time to the corresponding changes at Falmouth. The changes in the Colaba D and V curves are very small.

At Christchurch there was also a sudden commencement, though it is relatively inconspicuous. It is best seen in the D curve, where there was first a small movement to the east (i.e. down the sheet), from about 10.57 to 11.0 p.m., followed by a larger movement of similar duration to the west. Whether there was a double movement on the Christchurch V curve it is difficult to say, but there is at least an indication of a movement down the sheet, preceding the one distinctly visible up the sheet. There is also a distinct oscillation on the H curve, corresponding in time to that in the D curve. This was preceded by various small movements, which presumably were from a different source.

The results obtained from the measurements of the double movement were as follows :-


The increments $\Delta \mathrm{H}$, \&c., here, as elsewhere, measure the changes during the interval specified. Thus at Kew, on the occasion in question, $H$ fell $6 y$ between 10.57 and 11.0 p.m., and then fose $34 \gamma$ between 11.0 and 11.3 p.m. At 11.3 p.m. H was thus greater by $28 y$ than when the storm commenced. The change between 11.0 and $11.3 \mathrm{p} . \mathrm{m}$. was very probably in part a recovery from that experienced between 10.57 and 11.0 p.m., but that is purely a matter of surmise ; thus in the calculations subsequently made the movemént from 11.0 to 11.3 p.m. is treated as an absolutely independent one.
§62. Plates XX and XXI relate to a storm of considerable size which oecurred on December 13, 1903. At Falmouth, as Plate XX shows, a sudden commencement is distinctly visible in the $\mathbf{D}$ curve. There was a distinct double movement, first to the east from about 0.29 to 0.30 p.m., then to the west from 0.30 to 0.33 p.m., the westerly movement being considerably the larger. In the case of the H curve there were at the commencement several to-and-fro movements, of too small an amplitude to show details clearly. The principal part of the disturbance occurred from 6 to 10 hours after the commencement.

Plate XXI shows the changes in the Horizontal Force at Colaba and Mauritius corresponding to those at Falmouth. Both curves show a very distinct sudden commencement, the initial movement lasting from about 0.28 to $0.35 \mathrm{p} . \mathrm{m}$. H was falling at both stations at the time when the disturbance began (at Mauritius H increases down the sheet), and it is not absolutely clear that the first sudden change may not have been a further decrease, but if so it was too small and instantaneous to be clearly visible. The increase, on the other hand, is very decided. Following this there was a large but wonderfully uniform fall in H , continuing for about six hours, in the course of which a decrease of over $150 \gamma$ was experienced both at Colaba and Mauritius. At Falmouth the synchronous change in H was also on the whole a decrease,
but it was much less uniform and was interrupted by numerous oscillations of considerable size. Again, while the force at Falmouth continued to fall on the whole until 8.30 p.m., a marked rise commenced at about $6.30 \mathrm{p} . \mathrm{m}$. at the other two stations. At Mauritius and Colaba the most rapid and striking moments are those occurring from 8 to $10 \mathrm{p} . \mathrm{m}$.

In this case the correspondence is much closer between the Colaba and Mauritius curves than between cither of these and the Falmouth or Kew curves. At first sight the Colaba and Mauritius curves seemed to present a number of opportunities for comparing amplitudes at the two stations, but on further investigation this did not prove to be the case, the different elements not showing convenient synchronous turning-points.

## Antarctic Disturbed Curves.

§63. In the Antarctic, as already explained, all three elements and the temperature were recorded on one sheet. Thus at times of more than usual disturbance there was apt to be a good deal of intercrossing of curves, especially those of Declination and Horizontal Force. As a rule, however, the traces from the different elements differed in appearance to some extent, one being thicker or less exactly in focus than the others. Thus it was seldom that any serious doubt existed as to the identity of any particular portion of trace. The Vertical-Force trace was so much less open than the others, i.e., the sensitiveness was relatively so low, that it usually showed very little range of movement. Even at times of high disturbance changes of ordinate exceeding 2 cm . occurred very rarely. Thus the Vertical-Force trace, VV in the plates, was almost always easily distinguishable. It bore at times a resemblance to the temperature trace TT-which had the same base line $V_{0} V_{0}$ —but the movements in these two elements at times of magnetic disturbance were of such absolutely different types that confusion was impossible. One feature that often served to distinguish between the Declination trace DD and the Horizontal-Force trace HH was that the former was shown freely from one edge of the sheel to the other, whilst the latter became invisible at a considerable distance below the top edge of the sheet. This was due, as already explained, to the light being eclipsed by the Declination instrument. Exactly where the trace became invisible depended on the intensity of the light and the photographic conditions.

Whenever a trace died out gradually near the top of the sheet one knew it was that of the Horizontal Force. In the reproductions in the plates the letters $\mathrm{D}, \mathrm{H}$, and V appear whenever there seemed any risk of confusion. The suffixes 0 , as before, distinguish the base lines.

In the originals the base lines were interrupted for a short interval every hour. These breaks indicated hour intervals with the usual precision, but the eclipsing arrangement was not usually so set as to act exactly at the hour of local (or Greenwich) time. To determine the absolute time one had to refer to the times of starting or stopping the registration. Mr. Bernacchi always entered these times in a notebook, and he transferred them with his estimate of the error of the watch with which they were taken to the sheets themselves. Suppose, for instance, the exact local time of starting to be $10.30 \mathrm{a} . \mathrm{m}$., and 15 mm . trace to be recorded before the first break began, and suppose the interval between the left- (or between the right-) hand sides of the hour breaks to be 20 mm . (i.e. 1 mm . to 3 minutes). Then the time to which the left-hand side of the first break corresponded was $10 \mathrm{~h} .30 \mathrm{~m} .+\frac{1}{2} 5 \times 60$ minutes, or 11 h .15 m .

Supposing the magnetograph clock to have a negligible rate, the breaks in the above case would all commence at 15 minutes after the hour throughout the 24 hours, and the cross lines indicating the exact hour would all be drawn at $15 / 3$ or 5 mm . to the left of the left-hand side of each break.

Usually regard was had mainly to the time of commencement of registration. The clock rate was normally very small, but on a few occasions when it was large, or when there was some special cause of uncertainty, use was also made of the time of stopping. Of course the position of the last hour mark relatively to the end of the trace was always noted, but very exact measurements were not made unless there was some apparent inconsistency.

What accuracy can be claimed for the hour marks it is a little difficult to say. Whether owing to variations of temperature-which were unusually large-in the magnetograph room, or to differences in the shrinking or stretching of the photographic paper, the interval between the hour marks varied slightly from day to day, and was seldom exactly 20 mm ., usually somewhat less. There were even sometimes
visible variations in the course of the 24 hours. The marks were put on by an assistant using a millimetre scale and a fine pencil. If measurement showed the result of his first attempt to be unsatisfactory, the mark was rubbed out and a fresh one put on. Great care was used, but, of course, absolute accuracy cannot be claimed. The different elements were dealt with at different times, and to some extent by different people, and the time marks put on the three base lines were seldom absolutely in a straight line. There are several reasons for this quite apart from human liability to err. The dots of light were not all the same size, and none, of course, were mathematical points. They might not be all absolutely in a line, and the paper was doubtless at times not absolutely even on the cylinder, while the stretching or shrinking of the paper might not be the same at all parts of the sheet. Apart from the hour marks, measurements of time suffered from the fact that the curve and the corresponding base line were often very wide apart. If a measuring scale is imperfect, or if the observer does not get its base line to be absolutely colinear with the curve base line, an error comes into the measurement of the time which is proportional to the length of the ordinate. Uncertainties on this ground were very much greater for the Antarctic curves than for those at most of the co-operating observatories. The reason for dwelling on these points is that any uncertainty in the time increases the difficulty of being absolutely sure of the identification of corresponding rapid movements at different stations. As we shall presently see, there was sometimes a double to-and-fro movement in the Antarctic when only a single movement was clearly visible elsewhere, and a 2- or 3-minute mistake in the time might suffice to lead to a mistaken identification. All I can say is that in attempting these identifications I had regard to the original hour breaks and all the information available as to both stopping and starting. For the intercomparison of disturbances, if that had been the only object, it would have been more convenient to have used Greenwich time, as in the case of Falmouth, Colaba, Mauritius, and Christchurch. But to show Greenwich time would have entailed putting on a second set of hour marks, because those originally put on gave the local time which was employed in all the hourly measurements and tabulation tables, except those for the international term days and hours. Thus the times shown in all the Antarctic disturbed curves now to be discussed are local time (termed L.T.). The corresponding Greenwich times may be obtained by subtracting 11 h .7 m . In practice it is simpler to add 53 minutes, converting p.m. into a.m., and a.m. into p.m. of the previous day.

The original Antarctic curves were placed in contact with a photographic sheet, and a negative was obtained by transmitted light. Writing made on the original, and unnecessary details as to times, \&c., were removed from the negative. The positive from this was intensified, the faint parts being carefully inked in, and the hour gaps in the base lines filled in. The plates are reproductions from these positives. The reproduced curves cannot claim to be absolutely identical with the originals, but the differences between them are much less than if the originals had been traced in the same way as the term-hour curves were. Exact photographic copies of the originals were hardly feasible, because during the very rapid movements characteristic of the larger disturbances the trace was naturally faint, In fact, it was not infrequently so faint that it had to be traced over in the original to be sufficiently clear to be measured. These tracings were mainly done by Mr. Bernacche, and to distinguish the Declination and Horizontal Force traces at times he had been obliged to use coloured inks.

The base line values and the scales are shown at the margins, the information for one of the elements being on one margin, that for the two other elements on the other. To assist in distinguishing between the different curves all the traces stop just short of a scale-value line, except that of the base line for the element concerned. Arrows are added to mark the direction in which each element increases (V up the sheet, D and H down the sheet), though that can be inferred at once from the scale-value figures. The short intervals on the D scale represent $10^{\prime}$, only the $30^{\prime}$ intervals are numbered. On the H scale $10 \gamma$ intervals are shown, the $50 \gamma$ intervals being numbered. On the V scale $100 \gamma$ intervals are shown, but usually only the $200 \gamma$ intervals are numbered; the procedure varies according to the openness of the scale.
§64. Plate XXII gives the Antarctic record of the disturbance of May 8, 1902, corresponding to the records at the co-operating stations shown in Plate XIV. As elsewhere, there is an unusual definiteness about the end as well as the beginning of this storm, it being followed as well as preceded by an unusually quiet time. As already mentioned, it lasted almost exactly 8 hours. At the co-operating
stations, it will be rememhered, the notable feature at the commencement was the rise in the Horizontal Force. In the Antarctic the Horizontal-Force trace-whose sensitiveness is nearly four times that of the corresponding traces at Falmouth and Colaba-shows, at first, only the most trifling of movements. It is the D curve which is the disturbed one. This shows a very marked double movement. There is first an increase in I) (i.c. movement down the sheet), commencing at 11.6 p.m., L.T. ( 11.59 a.m., G.M.T.), and lasting about 4 minutes. This answered, apparently, to the increase in the Kew and Falmouth H curves from the commoncement up to the time when the curves show a temporary arrest in the movement. Following this, however, is a larger swing in the opposite direction. The second summit was not reached until 2 or 3 minutes after the conclusion of the increase in H at the co-operating stations, but the movemont during the last 3 minutes of the movement was small. The peculiar nature of the Antaretic phenomena is, perhaps, best brought out by contrasting the changes which occurred there during the two portions of the double movement, from $11.59 \mathrm{a} . \mathrm{m}$. to 12.9 , G.M.T., with those which occurred synchronously at Kew.


At the co-operating stations, movements from 2.10 to 2.28 p.m., G.M.T., and from 7.4 to 7.11 p.m., G.M.T., were fairly identifiable. In the Antarctic there are fairly recognisable turning- or stopping-points at 2.10 and 2.28 p.m., G.M.T., in all the elements, and a comparison was made between the changes of force at the various stations during these 18 minutes, the results of which are given afterwards. It should be noticed, however, that whilst 2.10 p.m. ( $1.17 \mathrm{a} . \mathrm{m} .$, L.T.) is a peak, and $2.28 \mathrm{p} . \mathrm{m}$. a prominent hollow, there is an intervening peak at 2.20 p.m. ( $1.27 \mathrm{a} . \mathrm{m} .$, L.T.), so that in the Antarctic the change from 2.10 to 2.28 p.m., G.M..T., was not persistent in one direction.

Owing, presumably, to the rapidity of the movements, the Antarctic $D$ trace became invisible during several minutes both before and after 7.11 p.m., G.M.T., so what happened from 7.4 to $7.11 \mathrm{p} . \mathrm{m}$. in that element is uncertain. The largest of the recorded D and H movements took place between 4.25 and 5.50 p.m., G.M.T. (3.32 and 4.57 a.m., L.T.), whilst the largest V movements took place between 6.51 and 7.30 p.m., G.M.T. ( 5.58 and $6.37 \mathrm{a} . \mathrm{m} .$, L.T.). These were also the times when the largest movements occurred at the co-operating stations.
§65. Plate XXIII shows what was, on the whole, a very quiet state of matters for the Antarctic. Its interest lies in the fact that the prominent oscillation in the D curve, which took place between 8 and 9 a.m., L.T., on August 21, synchronises apparently with the sudden change in $H$ at the co-operating stations shown in the top curves of Plate XV. Whilst there can be little doubt that the movement in the Antarctic is due to the same source as the commencing movement at the co-operating stations, there is a doubt as to what corresponds to what. At all the co-operating stations $H$ showed a sudden rise, commencing at from 9.6 to 9.7 p.m., G.M.T., on August 20, and continuing until 9.10 p.m. Before 9.6 p.m. the curves were very quiet. After 9.10 p.m. H diminished, but only very slowly. In the Antarctic the D curve shows a sudden movement up the sheet which lasts 4 minutes, but it seems to commence at $8.10 \mathrm{a} . \mathrm{m} .$, L.T. ( $9.3 \mathrm{p} . \mathrm{m} .$, G.M.T.), and it is immediately followed by a reverse swing of larger magnitude also lasting 4 minutes. This, in its turn, is followed after a minor oscillation by a movement up the sheet, bringing the D trace back to the position it occupied originally. The V curve shows an obvious double movement, but the first obvious movement (down the sheet, or decreasing force) sychronises with the second, not the first movement in D, and the second, or upward, movement follows thereafter.

However, on looking closely at the Antarctic V curve, it will be seen that its trend during the 4 minutes preceding the downward movement commencing at 8.13 or 8.14 a.m., L.T., is not quite in a line
with the immediately preceding portion. Prior to 8.10 a.m., L.T., the V trace was sloping gently down the sheet, representing most probably a gradual rise in temperature, which can, however, only be inferred, as the temperature trace got beyond the edge of the sheet at $2 \mathrm{a} . \mathrm{m} .$, L.T. But from 8.10 to 8.14 , L.T., the slope is slightly the other way. The measurements of tho original gave an increase of about 0.1 mm . of ordinate during this 4 minutes, answering to $1 \cdot 2 \gamma$. Thus we have apparently a real oscillation in all three elements, from approximately 8.10 to 8.18 a.m., L.T. ( 9.3 to 9.11 p.m., G.M.T.), for the magnitude of which I find:-


The time of commencing the second phase in the Antarctic is, if anything, before 9.7 p.m., G.M.T., whilst the times of the commencement of the increase in H at Kew, Falmouth, and Colaba are all, if anything, later than 9.6. Thus the evidence certainly points to the conclusion that it is the second phase in the Antarctic which answers to the commencement seen elsewhere. If so, it is certainly remarkable that what seems to a considerable extent a recovery from a sudden disturbance should exert an effect visible all over the world, whilst the effect of the preceding disturbance is visible only in the Antaretic.
§66. Plate XXIV shows a disturbance corresponding to the second of those appearing in Plate XV. In this instance the commencing movement was distinctly double at Falmouth and elsewhere, and in the Antarctic this oscillatory character is even more clearly shown in the H and V traces. Taking the H curve, we have first a sudden movement down the sheet (increase of H ) commencing at 2.59 a.m., L.T., on November 7 ( 3.52 p.m., G.M.T., on November 6) and lasting about $1 \frac{1}{2}$ minutes. This is followed by a. reverse motion of more than double the amplitude, occupying about $10 \frac{1}{2}$ minutes. The first movement, so far as can be judged, is exactly synchronous with that at Falmouth, but the second movement somewhat overlaps the corresponding Falmouth movement. The Falmouth trace is rounded at the top, but a slight fall in H is visible before the corresponding turning-point appears in the Antarctic curve. The Antarctic V trace shows a sudden rise synchronous with the first movement in H , followed by a larger fall. The second turning-point occurs, however, about two minutes earlier than that in H , and so corresponds more nearly to the summit of the Falmouth H curves. What happened to the Antarctic Declination is uncertain, as the trace is invisible during the most rapid part of the H movement. The fact that the D trace is invisible suggests that the movement was even larger than that in the $H$ trace; so presumably a to-and-fro movement of large amplitude occurred.

When the D trace becomes visible it shows two small peaks in rapid succession. It is the second of these that answers to the turning-point terminating the second movement in $H$. The first peak preceded this by three minutes.

The amplitudes of the first two movements in H and V are as follows:-

|  | $\Delta \mathrm{H}$. | $\Delta V$. |
| :---: | :---: | :---: |
| First movement <br> Second movement | $\begin{gathered} \gamma \\ +22 \\ -57 \end{gathered}$ | $\begin{array}{r} \gamma \\ +6 \\ -16 \end{array}$ |

The H and V magnets had at this time their highest sensitiveness, so that the large size of the apparent movements in these traces requires to be discounted when compared with that in other disturbances. The D trace, however, had the same sensitiveness as usual. Thus the fact that so much of the D trace is lost -partly through going off the sheet, partly from being too faint to be visible-is evidence that in the Antarctic the storm was a very considerable one, including some very rapid changes of force. At Kew
and Falmouth it was of a very trifling character, the range being under $6^{\prime}$ in D and under $20 \gamma$ in H . In the Antarctic the visille ranges, on the other hand, are $2^{\circ} 13^{\prime}$ in D and 80 y in H , and to all appearance the D range at least was considerably larger.
§67. Plate XXV shows the disturbance in the Antarctic on November 24, 1902, corresponding to the disturbance at Christchurch, shown in the lowest curve of Plate XV. This disturbance, though of considerahle size, had no very outstanding features either in the Antarctic or elsewhere. It occurred at a time when there was a good deal of general disturbance, and the day was not very exceptional as compared to adjacent days. The D curve, it will be seen, crossed from one edge of the sheet to the other -representing $4^{\circ} 50^{\circ}$-and trace was lost at both sides. The range in D is, however, partly due to the regular diurnal variation, which doubtless contributed considerably to the movement shown between noon and $5 \mathrm{p} . \mathrm{m}$. (L.T.). The H curve was off the sheet most of the time from noon to $4.30 \mathrm{p} . \mathrm{m}$. (L.T.) on the 24 th and the range probably largely exceeded that actually shown, $149 \gamma$.

The sensitiveness of the $V$ magnet was less than half of what it possessed on November 7. The movement shown on the trace represents about $171 \gamma$, but a portion of this is due to temperature. Between 11.30 a.m. (L.T.) on the 24th, when the minimum appears on the V trace, and 9 p.m. (L.T.) on the 24 th, when the maximum appears, temperature fell $0^{\circ} \cdot 4 \mathrm{C}$. Allowing for this change of temperature, the range is reduced to $154 \gamma$.

The identification of the D and H curves during the early part of the time covered by Plate XXV presented some difficultics. It answers to the earlier part of the photographic sheet for the day November 24-25 (L.T.). The interval between the putting on of this sheet and the taking off the previous one was unusually long, exceeding an hour. The disturbances during the early part of the 24th covered by the preceding sheet were very large, the to-and-fro oscillations in both D and H , if not so large as later in the day, being more numerous and rapid. At the time when the sheet was taken off, the D trace had been off the sheet for over an hour, while the $H$ trace, after being repeatedly beyond both limits of registration, was near the centre of the sheet.

On the sheet for November 24-25 the first absolutely clear discrimination between the D and H traces is afforded by the fact that the trace marked D did not become invisible at $5 \mathrm{p} . \mathrm{m}$. This enables one to identify the trace marked D as certainly Declination back to before 2 p.m. Prior to that one had to be guided by the appearance of the traces, and on that particular day there was no very marked difference as regards thickness of trace or definition. As to whether the small portion of curve which just came on the sheet at about $0.20 \mathrm{p} . \mathrm{m}$. (L.T.) represented D or H it was impossible to say.
§68. Plate XXVI shows disturbances in the Antarctic on April 6, 1903, the same date as those of Plate XVI.

Plate XXVI represents two portions of curves, the first extending from midnight on April 5 to $4.15 \mathrm{a} . \mathrm{m}$, (L.T.) on the 6 th, the second from $9.20 \mathrm{a} . \mathrm{m}$. to $3.10 \mathrm{p} . \mathrm{m}$. on the 6 th. The intermediate portion of curve was fairly quiet, and was omitted to bring the whole within the compass of one plate. The disturbances in the earlier portion of curve have a range of $63^{\prime}$ in $\mathrm{D}, 112 \gamma$ in H , and about $55 \gamma$ in V (the temperature effect in this instance is small). They answer to disturbances occurring from 10 to 6 hours earlier than those represented on Plate XVI. At Kew, the largest of these movements is represented by a bay on the H curve, the element first rising, and then falling about $30 \gamma$. This answered in time, at least very approximately, to the movement shown on the Antarctic $H$ curve between 3 and 4 a.m., L.T., but was of only about one-third the amplitude. The second portion of curve in Plate XXVI answers to the earlies portion of Plate XVI. What answers to the sudden commencing movement from 11.25 to 11.29 p.m., G.M.T., at Falmouth and the other co-operating stations, is apparently the sudden movement up the sheet in D and H , and movement down the sheet in V , shown in Plate XXVI as commeneing at $10.32 \mathrm{a} . \mathrm{m} ., \mathrm{L} . \mathrm{T}$, and lasting 4 minutes. But while the commencing movement at the co-operating stations was much more prominent than the movements immediately following it, the reverse is true of the Antarctic, where the commencing movement was immediately followed by an equally rapid and considerably larger movement in the opposite direction. Towards the end of the second movement the rapidity of change diminished notably in D and V . What happened in the case of H is uncertain, as the trace went off the sheet, but the movement terminated earlier than was the case in V or D . The double movement in the H trace has
many points of similarity to that shown at the commencement of the storm of November 7, 1902. It is, however, a much larger movement, and is exactly opposite in direction, the first movement on April 6, 1903, representing decrease in all the elements, whilst that on November 7, 1902, represented increase at least in H and V . The range shown in the second portion of Plate XXVI is $3^{\circ} 45^{\circ} \mathrm{in} \mathrm{I}$, $102 \gamma$ in H, and $105 \gamma$ in V, but as the D trace got off the sheet towards the end of the time, and the II trace was off for $2 \frac{1}{2}$ consecutive hours, the range in these two elements may have considerably exceeded that shown. The end of the traces in Plate XXVI answers to the end of the photographic shect of April 5-6. The commencing portion of the next sheet (not reproduced) shows a similarly disturbed set of conditions, lasting until nearly 2 a.m., L.T., on the 7th.
§69. Plates XXVII to XXXI afford illustrations of what I have called the "special typo of disturbance," which is discussed in detail in Chapter X. Plate XXVII is a good example of moderate size occurring on June 19, 1903. The disturbance referred to is that shown from about 6.40 to $7.20 \mathrm{p} . \mathrm{m}$., L.T. During what I have termed the "first phase"-extending from 6.40 to $7.1 \mathrm{p} . \mathrm{m}$.- the D and H traces go up, the V trace down the sheet, i.e. all three elements diminish. During the second phase the movements are in the reverse direction, i.e. the three elements increase. In this individual case the return D and $H$ movements are less than the original movements, something occurring to check them. The return movement in the V curve exceeds the first movement, as nearly always happened. The curves show the continual minor oscillations characteristic of practically all the Antarctic D and H records. Thus it is difficult to assign a very definite time for either the beginning of the first phase or the conclusion of the second. Most likely the duration of the phases is underestimated. The $H$ trace went beyond the limits of registration-apparently just exceeding them-about 7 p.m., L.T. The summit in the $H$ curve must have occurred a few minutes after the hour $7.1 \mathrm{p} . \mathrm{m}$. accepted above for the end of the first phase. The curve shows, however, a peak somewhat earlier, about the time of the turning-points in the 1 and V curves. The behaviour of the H trace during the special type of disturbance was less uniform than that of the other two elements-the first change in $H$ being sometimes an increase-and in fixing the end of the first phase more attention was given to the two other elements, especially to V.

The changes in H during the special disturbance of June 19, though looking much larger to the eye than those in D , represented in reality considerably less change of force. Thus, during the first phase, the change shown in H represents a decrease of $91 \gamma$, but the corresponding fall in D was $72^{\prime}$, answering to a force of about $138 \gamma$. The second phase movement in V , relatively insignificant as it looks, represents an increase of $53 \%$. After reaching the maximum shown at about 7.20 p.m., the $V$ curve shows only a very gradual diminution of ordinate, as if the cause of the special disturbance, whatever that may be, had an effect on the Vertical Force which required time to disappear.

Synchronously with the special disturbance in the Antarctic on June 19, 1903, there was a well-marked, though much smaller movement at Christchurch, shown in the upper curves of Plate XVII. The prominent crest in the Christchurch D curve, appearing a little before $8 \mathrm{a} . \mathrm{m}$. (G.M.T.), seems to answer to the end of the first phase in the Antarctic movement.
§70. Plates XXVIII and XXIX show movements of a very similar character to that shown by Plate XXVII, occurring on successive days, June 28 and June 29, 1903. The estirnated times of ending for the first phase were: for June $28,7.54$ p.m., L.T. ; for June 29, 8.33 p.m., L.T. The occurrence of closely similar movements in Declination on two successive days has been remarked on before. In the 9th edition of the 'Encyclopædia Britannica'* Dr. Balfour Stewart refers to a discussion by Senor Capello of a series of cases in which such changes appeared in the Lisbon magnetic curves, and mentions that he had observed correspondences of similar closeness in the Kew curves. It is not clear whether Dr. Stewart himself considered that the number of the instances of apparent repetitions and the closeness of the resemblance were such as to demonstrate that more than chance was involved. But provisionally, supposing some definite physical cause to be involved, he suggests that something might lead to the neighbourhood of some particular meridian of the Earth's surface becoming sensitive to radiations received from the Sun, and that the consequence might be the repetition of a special type of disturbance on

[^2]successive days at about the same hour until the sensitiveness had disappeared. I have seen in the Kew curves a considerable number of instances of the kind referred to by Senor Capello and Dr. Stewart, but in few, if any, has there been so close a resemblance-taking all the elements into account-as that presented by the Antaretic curves of June 28 and June 29, 1903. The interval between the two occurrences, as already stated, was only approximately 24 hours, but in this respect there is no departure from what has been observed of apparent repetitions at Kew and elsewhere. In considering whether more than chance enters into the occurrence regard must be had to the greater or less frequency of the particular type of disturbance, and as to whether it shows a preference for particular hours of the day. It is shown later, p. 189, that this special type of disturbance was in reality of somewhat frequent occurrence in the Midwinter months, and it happened much oftener during the four hours 6-10 p.m. than throughout the remaining 20. Thus the fact that two occurrences should follow one another at nearly a 24 -hour interval hardly possesses the significance one is disposed to give it at first sight.

On June 28 the second phase seems to have been hardly completed before the arrival of a second wave-like disturbance of much shorter period. After the passage of this short-period disturbance the V curve comes down the shect for a time more rapidly than was usual in the special type of disturbance. 'The disturbance of June 29 was more normal in this respect.
\$71. Plate XXX shows another example of the special disturbance occurring on July 26, 1903. The end of the first phase occurred about 5.55 p.m., L.T. In this instance the H trace indicates first an increase, then a decrease, the element being so to speak opposite in phase to the other two, in which the decrease in value comes first. As already stated, this opposition in phase in H sometimes happened, though in but a small minority of cases. The movement in the Antarctic synchronised with the first portions of the disturbance shown in the three lower Christchurch curves of Plate XVII. As Plate XXX shows, there were later disturbances on the same date in the Antarctic, and one or two of the movements show some resemblance to the principal one described above. The $V$ trace throughout its entire length shows a tendency to wave-like swelling forms, and it will be noticed that the hollows in these tend to synchronise more or less closely with peaks on the D and H curves.

The interval between the disturbances of June 29 and July 26 is approximately 26 days $21 \frac{1}{2}$ hours, and so is not far from the period $27 \frac{1}{4}$ days which Mr. Maunder believes his researches show in magnetic storms at Greenwich. It cannot be said that there is a closer resemblance between the storms of July 26 and June 29 than between those of the former date and the storms of June 28 and June 19. So that the evidence is equally strong for a considerable choice of periods.
872. Plate XXXI shows another example of the special type of disturbance occurring on August 17, 1903. The end of the first phase was about 7.54 p.m., L.T. The disturbance is not large, but it occurred during a very quiet period. The slowness with which V diminished after attaining its maximum is shown with especial clearness, because the temperature was unusually steady, so that no serious complications arise through it.
§73. Plate XXXII shows a somewhat highly disturbed state of matters in the Antarctic on August 22-23, 1903, synchronising with the disturbances at Christchurch shown in Plate XVIII. The H curve was off the sheet during a considerable portion of the time, first on the one side, then on the other. Probably the range in the element very largely exceeded that shown, $150 \gamma$. The D curve went also off the sheet, but only for a short time, a little after 4 p.m., L.T. ; the range shown is $3^{\circ} 45^{\prime}$. Even the V curve shows continual oscillations; the range shown by the trace, $200 \gamma$, is only slightly affected by the temperature variation, which was small. There is at least an approach to the special type of disturbance between 3.55 and 4.50 p.m., L.T., but much irregular disturbance existed at the time.
§74. Plates XXXIII and XXXIV show respectively the concluding portion of the trace of August $25-26,1903$, and the commencing portion of the trace of August 26-27. The days intermediate between Angust 22 and August 25 showed also a good deal of disturbance. During the latter part of August 26 the $V$ magnet was out of action and its trace does not appear in Plate XXXIV. Plates XXXIII and XXXIV cover the time shown by the disturbances at the co-operating stations which appear in Plate XIX. At the co-operating stations the storm had a sudden commencement, consisting of to-and-fro movements-the latter much the larger-occupying from 10.57 to 11.3 p.m., G.M.T., on August 25.

This answers to 10.4 to 10.10 a.m., L.'T., on the 26th ins the Antarctic. As Plate XXXIII shows, there were also very rapid oscillatory movements in the Antarctic during this time; but there is again some uncertainty about the exact identification of corresponding movements. Conspicuously the most rapid H movement in the Antarctic is the one of which the turning-point is lost owing to the trace going beyond the limits of registration. This consists of a to-and-fro movement, first up, then down the sheet. When the rapid upward movement began the curve was moving up the sheet, but at a comparatively leisurely rate. The time of the rapid acceleration of movement is almost, if not exactly, 10.4 a.m., L.T. Similarly the movement down the sheot continued for some time after a marked reduction appeared in the velocity. The reduction of velocity set in very approximately at $10.10 \mathrm{a} . \mathrm{m}$. , L.T. Thus, presumably, the very rapid part of this to-and-fro movement in H-representing a decrease of the element followed by a larger increase-answered to the commencing movement seen elsewhere. At the time when the H trace accelerated its movement up the sheet, the D trace began to move down; but after a minute or two the motion downwards diminished, and a small oscillatory movement intervened of which details are not clearly shown in the original. Thus during the latter part of the very rapid oscillation in H the value of D was nearly stationary. The V trace went through several oscillatory movements. The one which probably corresponds to the double movement in H is that of which the end is represented by a crest which appears nearly 1 mm . to the right of the rapid downward movement in H . The previous turning-point (a hollow) comes between the upward and downward H movements, and answered probably to the turning-point in H which was beyond the limits of registration. The previous turning-point on the V curve is not shown, because the curve was so near the D base line that it got eclipsed by the hour stop, which came on at about $10.3 \mathrm{a} . \mathrm{m} ., \mathrm{L} . \mathrm{T}$. The V double movement is somewhat larger than one which immediately preceded it, but is in no way outstanding.

The most conspicuous movements in Plate XXXIV are those appearing between 6 and 9 p.m., L.T. In particular, there is a bold movement of the D and H curves up the sheet, commencing at $6.45 \mathrm{p} . \mathrm{m} ., \mathrm{L} . \mathrm{T}$. Both curves went beyond the limit of registration, but the turning-point in each was apparently a few minutes after 7.0 p.m., L.'T. The return movements down the sheet were equally conspicuous and ended about 7.35 p.m., L.T. This double movement was in all probability of the "special type," but owing to the absence of the V curve our information on the point is incomplete. This movement corresponds in time to that shown by the Christchurch D and V curves in Plate XIX, culminating about 8 a.m., G.M.T., on the 26th. This double movement in the Antarctic was immediately followed by another lasting about 50 minutes, in which, however, the motion of the $H$ magnet was opposite in phase to that of the D magnet. There were large subsidiary oscillations in the H trace. Subsequent to these two large movements there were some rapid oscillations of considerable size during which the D and H traces are a little difficult to distinguish.
§75. Plate XXXV shows the first part of a highly disturbed Antarctic record which commenced a little before 6 p.m., L.T., on October 12, 1903, and ended between 5 and 6 p.m. on October 13. The portion reproduced ends at $4 \mathrm{a} . \mathrm{m} .$, L.T., on the 13 th. The $H$ trace was beyond the limits of registration continuously from $8 \mathrm{p} . \mathrm{m}$. on the 12 th to $8 \mathrm{a} . \mathrm{m}$. on the 13 th, whilst the D trace was off the sheet continuously from $4 \mathrm{a} . \mathrm{m}$. to $11.50 \mathrm{a} . \mathrm{m}$. on the 13 th . Subsequent to $8 \mathrm{a} . \mathrm{m}$. on the 13 th, the H trace came on the sheet at the top, executed a number of oscillations proceeding gradually down the sheet, and disappeared a little after 11 a.m. Subsequent to that, except for a few fugitive appearances, it was not seen until between 3 and $4 \mathrm{p} . \mathrm{m}$. on the 13th. After coming on the bottom of the sheet about $11.55 \mathrm{a} . \mathrm{m}$. on the 13 th, the D trace remained on, showing a number of rapid oscillations, until a little after 4 p.m., when it went off the sheet at the top. After remaining off for 10 or 15 minutes it came on again and the disturbance became less.

The range in D shown in Plate XXXV is $4^{\circ} 36^{\prime}$, and that in V $157 \gamma$, the latter being practically unaffected by temperature. Between $11.55 \mathrm{a} . \mathrm{m}$. and $4.15 \mathrm{p} . \mathrm{m} .$, L.T., on the 13 th the range in D exceeded $4^{\circ} 51^{\circ}$. The $\mathbf{V}$ trace throughout the whole time covered by the original sheet showed numerous short-period oscillations, but none of the subsequent movements were so large as that shown in Plate XXXV from 9.40 to 10.40 p.m., L.T.

There was a very considerable magnetic disturbance at Kew on October 12-13, but the principal
movements occurred between 6 p.m., G.M.T., on the 12 th and 3 a.m., G.M.T., on the 13 th ( 5 a.m. to 2 p.m. on the 13th, Antaretic L.T.). They are thus subsequent to the movements shown in Plate XXXV, and answered to a time when much of the Antarctic trace was beyond the limits of registration. During the time covered by Plate XXXV the range at Kew was $18^{\prime}$ in D and $75 \gamma$ in H , but later, during the two hours $10-12 \mathrm{p} . \mathrm{m}$. , G.M.T., on the 12th there was a range of $37^{\prime}$ in D and $180 \mathrm{\gamma}$ in H . During a portion of the two hours specified the Antarctic $\mathbf{D}$ and $H$ curves were off the sheet simultaneously; the $\mathbf{V}$ trace, however, was less disturbed then than at various other times.

This is a conspicuous example of the disadvantages attending the use of high sensitiveness in the Antaretic.
§76. Plate XXXVI shows disturbances in the Antarctic on December 13-14, 1903, corresponding to those at the co-operating stations shown in Plates XX and XXI. At the co-operating stations there was a sudden commencement, but it was not very clearly indicated in the H curves at Kew and Falmouth, and its duration in the D curves at these two stations appeared less than in the H curves at Colaba and Mauritius. The Antarctic curves also show a sudden movement, especially that for $D$. The trace for this element shows a peak at 11.28 p.m., L.T. ( 0.21 p.m., G.M.T.), followed by a downward movement. During the first two minutes of its occurrence the downward movement is slow, but it suddenly accelerates at 11.30 p.m., L.T., and the movement down the sheet during the next minute represents a rise of $36^{\prime}$ in D . This is followed by a less rapid, but larger, movement in the opposite direction, which continued until 11.40 p.m., L.T. ( 0.33 p.m., G.M.T.), and represented a fall of $77^{\prime}$ in D. The second movement, it will be observed, covers the time of the double oscillation seen at Kew and Falmouth. The first movement may possibly be represented by some movements indistinctly shown in the Kew and Falmouth H curves prior to 0.29 p.m., G.M.T.

The Antarctic V trace shows also a distinct double movement, first up, then down the sheet. The summit is distinctly shown at about 11.36 p.m., L.'T. ( 0.29 p.m., G.M.T.). The movements are small, and V was increasing when the first movement set in, thus it is difficult to assign an exact time for its commencement; it occupied apparently not more than one or two minutes, and was most likely synchronous with the rapid part of the commencing D movement. The second, or decreasing, movement in V lasted four or five minutes. It was about $18 \gamma$ and about double the first movement.

Plate XXXVI consists, it will be noticed, of two portions, separated by an interval of about $3 \frac{1}{2}$ hours. During this interval the D and H traces were both continuously beyond the limits of registration. The $H$ trace was, in fact, beyond the limits from 11 p.m., L.T., on the 13 th until $9.30 \mathrm{a} . \mathrm{m}$. on the 14 th. The mode of its reappearance, as shown in the end or right-hand portion of the plate, is such as to suggest that the limit of registration answering to low values of the element was very much exceeded. The limit answering to high values was also exceeded between 10 and 11 a.m., but only for about 12 minutes.

The left-hand portion of the plate shows D executing considerable oscillations, with a general drift down the sheet after midnight until 5 a.m., L.T., when the limit of registration was exceeded for a few minutes. The total range during these five hours was $4^{\circ} 12^{\prime}$. The right-hand portion of the plate shows the D trace reappearing after three hours' absence, and then alternately on and off the sheet until 10.30 a.m., L.T., when a very rapid movement up the sheet (diminution of the element) took place. The complete width of the sheet, answering to $4^{\circ} 52^{\prime}$, was crossed in about 18 minutes.

The final portion of Plate XXXVI answers to the end of the Antarctic sheet of December 13-14. The following Antarctic sheet (not reproduced) shows further large disturbances lasting until about 6 p.m., L.T., on the 14th. The most notable $V$ change occurred during the afternoon of the 14th. It consisted of a rise of about $240 \gamma$ between 1.10 and 1.50 p.m., L.T. D increased $4^{\circ} 8^{\prime}$ between 1.5 and 3.5 p.m., L.T., and then diminished $3^{\circ} 57^{\prime}$ between 3.5 and $5.15 \mathrm{p} . \mathrm{m}$. L.T. A rough estimate of the sum of the D movements, taken irrespective of sign, which were recorded between 10.30 a.m. and 5.15 p.m., L.T., on December 14, proved to amount to no less than $25^{\circ}$.

On October 31 and the early morning of November 1, 1903, there occurred the largest magnetic storm recorded at Kew during several years, but of this no record was obtained in the Antarctic, as the magnetographs were not in action on October 31, nor on the subsequent day until the late afternoon.
§77. The seven Plates XXXVII to XLIII contain copies of portions of a number of Antarctic curves
selected as representing various types of disturbance. They were first drawn by hand on tracing paper and then photographed. The disturbances included in any one of these plates have at least a general similarity of type. One or two hours (local time) are marked on the base lines. The curves are shown natural size, so that 20 mm . of abscissa represents very nearly an hour. The principal object being to illustrate the types of disturbances, numerical magnitudes were of minor importance, and scale values are thus not shown, to avoid complicating the figures. One could hardly have shown the scale values clearly without greatly extending the area of the plates. If, however, scale values are desired they can be obtained by reference to Table I, and they can be derived approximately from the information given as to the ranges of the elements. The range recorded ( D in minutes of arc, H and V in terms of $1 \gamma$ ) refers, in all cases, to the portion of the disturbance actually shown.
§78. Plate XXXVII contains examples of the special type of disturbance which is dealt with in detail in Chapter X, and some closely allied forms. The curve of May 22, 1903, represents a simple example of the normal kind. Shortly before 8 p.m the D and H curves start running up the sheet (i.e. elements both diminish), the V trace being nearly straight, but with a slight tendency downwards ( V diminishing). This constitutes the first phase. Superposed on the general drift in the D and H curves are minor oscillations. These render it difficult to say, to the minute, when the turning-point was reached ; but all the elements show one within a few minutes after 8 p.m. The second phase then begins. The D and H traces come down the sheet (elements increase) at a rate very similar to that of the first phase. V goes up the sheet (element increases) to a maximum, the rise being considerably in excess of the previous fail. This concludes the second phase, at the end of which $H$ has practically returned to its original value, while V retains an enhanced value for some time, no conspicuous fall setting in until half an hour after the maximum was reached.

On August 31, 1902, we have again a disturbance of the special type, but the fall in V during the first phase is unusually large, and the turning-point in the D curve appears delayed. The end of the special type of disturbance may be put at about 9.20 p.m., when the lowest point appears on the $H$ curve (maximum of force). Thereafter the $H$ curve rises sharply again, and the $V$ curve comes down the sheet, a second wave, as it were, rolling in. After the first 20 minutes of this second wave the $H$ curve hesitates and fluctuates, indicating presumably the action of some independent source of disturbance. This, though able only to check the H movement, sufficed to alter the direction of the D movement, which seemed at first to be about to follow the H . Thus we do not get a repetition of the special type of disturbance.

The limit of registration of the $H$ curve was exceeded-though apparently only very slightly-just before $9 \mathrm{p} . \mathrm{m}$.; the range recorded is measured from this limit. As previously explained, the light was really cut off through the D instrument coming in the way.

Unlike the two other disturbances, that of August 15, 1903, is an early morning one. There are a succession of waves, crests of D and H (i.e. minima of force) answering, at least approximately, to troughs (minima of force) in the V curve. The movements in V and D represent nearly equal force components, each fully double the visible change in H. Judging, however, by the appearance of the H trace, the limit of registration was a good deal exceeded. Towards $2 \mathrm{a} . \mathrm{m}$. a final wave will be noticed, which at least approached the special type of disturbance.
§79. Plate XXXVIII shows some transitional forms having a greater or less resemblance to the special type. The disturbance of June 4,1903 , is essentially of this type. H, it is true, is opposite in phase to D, but that happened in an appreciable minority of cases. There were clearly, however, subsidiary disturbances, $V$ showing a succession of small oscillations after the maximum was reached. Owing to subsidiary oscillations, the crest of the D wave is not very clearly indicated, and no assistance in fixing the turningpoint is derivable from the H curve, as the trace was off the sheet for fully an hour.

October 4-5 and October 11-12, 1902, are examples of midnight disturbances. In both, the H movement is at least strongly suggestive of the special type. The V magnet, however, was out of action, so that no corroboration is forthcoming from its trace. The D curve is of a somewhat different type. There is at least a suggestion of waves, but they seem of a shorter period than those in H . On both days the D and H curves appear opposite in phase when the large movement in H begins, but towards the end of the interval they seem nearly in the same phase.
§80. Plate XXXIX deals with two occasions in September, 1902, separated by 12 days, when the H trace shows in immediate suceession two well-marked oscillations of similar amplitude. On September 6-7 there are four well-marked crests in the H curve, but the first and the last of the oscillations are much smaller than the intermediate two. On this day there are also well-marked waves in the V curve, troughs (minima) answering to crests (minima) in the $H$ curve. The $D$ curve on this occasion is also of a distinctly undulatory type, but its phase seems on the whole a little in advance of that of the H curve.

On September 18-19, 1902, the V magnet was out of action. The D curve, though exhibiting a tendency to wave-like formation, is rather complex.

Both these occasions relate to hours near midnight, when the special type of disturbance tended to assume the more sharply oscillatory character usually associated with the morning hours.
§81. Plate XL shows four examples of early-morning disturbances. On June 21-22 and July 9, 1902, the movements in the V curve are still of a comparatively simple type. On the former occasion we have long wave-like movements in V , but towards $0.30 \mathrm{a} . \mathrm{m}$. minor oscillations are distinctly visible. The H movement in this case, though not quite symmetrical, closely resembles the special type; it representswhat was unusual-a considerably larger component of disturbing force than either the D or the V movements.
On July 9, 1902, the H range is again decidedly the largest. The disturbance on this day on the whole resembles the special type, but a second wave has apparently rolled in before the first was exhausted, so that the V trace appears double-headed. On the whole, the D and H traces are opposite in phase to one another.

January 10, 1903, and December 10, 1902, are examples of large disturbances. On January 10, 1903, the H trace was beyond the limits of registration during the whole interval represented, and even the D trace was off the sheet for a time. The range of $D$ measured to the edge of the sheet was almost $4^{\circ}$, and the time from the minimum to the maximum shown was only 12 minutes. The corresponding change in force is about $450 \%$. The V trace, though showing minor oscillations, is still of a rounded wave-like form, with crests (maxima) answering to troughs (maxima) in D. The next two days, January 11 and 12, 1903, it may be remarked in passing, exhibited movements of similar magnitude to those shown on January 10 throughout the same morning hours.

On December 10, 1902, the D trace was beyond-though apparently very little beyond-the lower edge of the sheet from 5.7 to 5.8 a.m., and it just reached the other edge at $6.3 \mathrm{a} . \mathrm{m}$. The most rapid changes were from 4.22 to $4.36 \mathrm{a} . \mathrm{m}$., when D increased $3^{\circ} 33^{\prime}$ in 14 minutes, and from 5.8 to $5.26 \mathrm{a} . \mathrm{m}$., when it fell $3^{\circ} 57^{\prime}$ in 18 minutes. In this instance there are numerous oscillations on the V trace, but, on the whole, it is clear that crests (maxima) on it answer to troughs (maxima) on the D curve. The H trace was beyond the limits of registration all the time.
§82. Plate XLI also illustrates morning disturbances. Only the H and V curves are shown, the D trace being largely off the sheet and highly oscillatory. All three examples exhibit numerous oscillatory movements, mostly very irregular. On February 8, 1903, the H curve, however, shows a marked tendency to form sharp-peaked waves, the time of the combined up-and-down movements in each occupying about 25 minutes. On the whole, crests (minima) in H seem to answer to crests (maxima) in V .

Even on February 10 and March 6, 1903, though there are numerous short-period oscillations, the movements in the $H$ curve suggest as a substratum a comparatively regular wave motion of longer period.
§83. Plate XLII contains examples of very sharp oscillations of considerable size. The disturbances on February 25 and 26, 1903, were recorded in the early afternoon; the others in the forenoon. Of the five examples that from April 1 is the most striking. On that occasion the three elements all exhibit one conspicuously large and rapid to-and-fro movement just after 9 a.m.

Within from 6 to 7 minutes D first diminishes $84^{\prime}$ and then increases $70^{\prime}, \mathrm{H}$ falls and rises $90 \gamma$ and V falls and rises $100 \%$. The turning-points for the three elements appear to be really identical in time.

Individually considered, the two specimens from June 10, 1903, are much less striking, but it is a little remarkable that two movements so similar in type and in magnitude should occur within an interval of five hours.
§84. Plate XLIII contains examples of morning disturbances. There are here no longer rounded wave-
like movements, possessing the regularity and considerable period of the special type of disturlance, but instead a constant succession of short-period oscillations or ripples in all three elements. The movements are so large, and the intervals between turning-points so short, that it is hardly possible to decide the exact relationship of the changes in the three elements. Judged by the ranges recorded, the disturbances are not very large; but this merely affords an illustration of how unsatisfactory a criterion of disturbance range alone may be. Thus, on August 15, 1903, during the $2 \frac{1}{2}$ hours considered, the D range was only $2^{\circ} 9^{\prime}$; but at least 80 turning-points can be recognised on the curve, and the aggregate of the changes in D , taken irrespective of sign, is about $12 \frac{1}{2}^{\circ}$.

Rapid as the oscillations are in the curves of Plate XLIII, they are, of course, infinitely slow compared to the magnetic changes that accompany the sudden creation or stoppage of an electric current under normal conditions. If they represent the waxing and waning, or the change in distance and direction of electric currents, these changes must presumably present nothing approaching to a discontinuity, in the mathematical sense.

It should be explained that there is no reason to suppose that the magnets under these rapid oscillations behaved in any but the most dead-beat way. There is no suggestion in the curves of vibrations possessing their period.

## CHAPTER IX.

## Discussion of Disturbances.

S45. In discussing the magnetic disturbances recorded in the Antaretic and the co-operating stations it is desirable to pay some regard to existing knowledge and theorios. It has long been known that large magnetic disturbances are usually felt over at least a large portion of the Earth's surface. In a considerable number of instances large storms have been ushered in by comparatively sudden changes in the magnetic dements, which the researches of Adams, Elis, Van Bemmelen, and others have shown to be in all probability simultaneous in their incidence wherever recorded. It has been observed that magnetic storms are more numerous near sun-spot maximum than near sun-spot minimum, and many persons have suspected a very intimate connection between the two phenomena. In temperate latitudes a display of aurora is practically always accompanied by a magnetic storm, and another invariable accompaniment of the larger and more rapid magnetic changes is the existence of earth currents.
Turning to theories, a considerable number have been advanced. It has been supposed that the Sun is a powerful magnet, and that the regular diurnal variation of terrestrial magnetism and the irregular changes arise from the direct magnetic action of the Sun acting, as we may say, at a distance. Another hypothesis is, that electric waves or Röntgen rays, arising in the Sun, travel to the Earth, and on their arrival set up aurora and magnetic disturbances. Others bave postulated the existence of negatively-charged particles (Arrhenius), or kathode rays (Birkeland), or some analogous form of ray discharge (Maunder). The discharge, whatever its nature, on reaching the Earth's atmosphere, occasions electric phenomena which create variations in the Earth's magnetic field. Arrhenius has calculated that his hypothetical negative particles will travel from the Sun to the Earth in about 45 hours. Birkeland has produced artificially, in the neighbourhood of a small magnetic sphere, or terella, in a high vacuum, electric discharges bearing a close resemblance to various types of aurora, and has attempted to establish a direct connection between individual auroras and the magnetic disturbances recorded during them. Adopting Birkeland's bypothesis of kathode rays, STÖrmer has carried out elaborate mathematical calculations as to the paths of these rays within the Earth's magnetic field, and has concluded that they can approach the surface only in the regions surrounding the magnetic Poles. Mr. Maunder has not attempted to assign any exact physical properties to the discharge which he postulates, beyond supposing that it emanates from a sunspot, and travels out not in all directions but in a compact mass, like a water jet from a hose. The commencement of a magnetic storm answers to the Earth commencing to cross the jet, and the storm continues until the transit of the jet has been accomplished. Supposing the sun-spot to continue active for a time in excess of one complete revolution of the Sun, the Earth will naturally cross the jet again after an interval of about $27 \frac{1}{4}$ days, and a second magnetic storm will be experienced.

If magnetic storms are due to electric carriers or ions of any type, the incidence of the magnetic storm will depend on the Earth's magnetic field, and until the properties of the ions are known the laws governing the intensity of disturbances at different parts of the Earth's surface cannot be inferred. If the carriers are kathode rays, then, assuming Stormer's calculations correct, magnetic disturbances would seem to be necessarily always greater near the magnetic Poles than towards the magnetic Equator, but further than this we cannot apparently go at present.
§86. Direct magnetic action from the Sun was considered long ago by Kelvin, who made a rough calculation of the amount of energy which on this hypothesis would be expended during a large magnetic storm. The result which he reached seemed to Kelvin inadmissably large. Physicists' ideas, however, as to solar energy have considerably altered of late, and I think it will be generally allowed that in our present state of knowledge it is safest to admit the possibility of expenditures of energy which 20 years ago would have appeared incredible. It must also be allowed that, whilst physical speculators in the past do not seem to have been deterred by the absence of outward and visible signs of magnetism in the Sun, still the discovery by Professor Hale of the Zeeman effect in light received from sun-spot areas affords a direct evidence which bas hitherto been lacking. Dr. Schuster, it is true, has already discounted Professor Hale's discovery to some extent, by publishing in 'Nature' an estimate of the
possible size of the effect on the Earth of the largest of sun-spot areas, endowed with what he considers the largest magnetic moment reasonably attributable. The resulting disturbance is only of the order of $\frac{1}{1000}$ of the range of large magnetic storms at Greenwich, and Dr. Schuster thus reaches much the same conclusion as Lord Kelvin. Just, however, as in Kelvin's case, I think it would be unwise at present to pass a final judgment. It must also be remembered that magnetic disturbances are of varied forms and sizes. Even if we were justified in concluding that large movements-of, say, $1000 \gamma$-in H camnot be due to direct magnetic action from the Sun, it is still conceivable that the comparatively small but outstanding movements which often precede by several hours the largest movements of a disturbance may represent such direct action. I have thus thought it worth while to make a few mathematical calculations as to what might be expected to happen if direct action took place. The Sun's distance being so great compared to the Earth's diameter, the field due to the Sun's action at any instant may be treated as constant throughout an element of space large enough to contain the Earth and her atmosphere. The presence of the Earth modifies the field, and thus leads to variations in the disturbing force experienced at different parts of her surface. The problem is not strictly a statical one, as the disturbing forces do not remain constant. Rapid, however, as magnetic changes may seem when recorded by ordinary slow-ruming magnetographs, they are usually infinitely slow as compared to the changes which occur in wireless telegraphy. Even the sudden commencements of storms, as we have seen, consist of changes persistent in one direction for several minutes, the rate of change being usually less than $10 \gamma$ per minute. There is thus reason to think that whilst the treatment of the problem as a statical one involves departures from actual conditions, the results ought not to suffer so much from this limitation as to be unworthy of consideration. A complete treatment ought of course to regard the Earth as a conductor of electricity. Earth currents, we know, have diurnal, annual and irregular variations, and there can be no doubt that a relationship exists between these currents and the phenomena of terrestrial magnetism. Our knowledge at present, however, as regards Earth currents, is exceedingly limited. Even close to the surface but few reliable observations have been made, and what depths the currents may extend to, or what the electrical resistance of the several strata is, we do not know. The fact that the greater portion of the Earth's surface is sea may be a most important consideration. We are thus almost perforce obliged to confine ourselves to the statical problem, which will now be considered.
§87. It must be admitted that our direct knowledge of the Earth's magnetic quality is very slight. There are few surface materials which are appreciably magnetic, and how the Earth comes to be a magnet is a mystery. A magnet however it must be unless there is something absolutely wrong in the application of the Gaussian analysis. According to that analysis, the source of the Earth's magnetism is almost entirely internal, and the potential, to a rough first approximation, is that due to a solid sphere or internal spherical shell uniformly magnetised. The natural inference is that the Earth's magnetie quality is, to a first approximation, a function only of the distance $r^{r}$ from the centre. We shall first consider the simplest case, viz., that of a sphere of uniform permeability $\mu$ and radius $a$. Let $\mathbf{F}$ denote the strength of the field in the absence of the sphere, and let $r$ and $\theta$ be polar co-ordinates, $\theta$ being measured from the diameter which coincides with the direction of the field F. On the introduction of the sphere the field in the medium outside it, supposed of unit permeability, is given by

$$
\begin{equation*}
\mathrm{V}=-\mathrm{F} \cdot \cos \theta+\frac{\mu-1}{\mu+2} \mathrm{~F} \frac{\ell^{3}}{r^{2}} \cos \theta \tag{1}
\end{equation*}
$$



In the figure, $\mathrm{APA}^{\prime}$ is a section through the centre 0 of the sphere, $\mathbf{A}^{\prime} \mathbf{A}$ being the diameter parallel to the direction of the field $\mathbf{F}$.

The force at any point $\mathrm{P}(a, \theta)$ just outside the sphere may be regarded as composed of a normal (i.e. vertical) component $3 \mu \mathrm{~F} \cos \theta /(\mu+2)$, and a tangential (i.e. horizontal) component $3 \mathrm{~F} \sin \theta /(\mu+2)$.

If $\psi$ be the inclination to 0 A of the resultant force R at P , then

$$
\begin{array}{r}
\tan \psi=(\mu-1) \sin 2 \theta \div\{(\mu+1)+(\mu-1) \cos 2 \theta\} \ldots \\
\mathrm{R}_{/} / \mathrm{F}=3(\mu+2)^{-1}\left\{\mu^{2}-\left(\mu^{2}-1\right) \sin ^{2} \theta\right\}^{1 / 2}=3(\mu+2)^{-1}\left\{1+\left(\mu^{2}-1\right) \cos ^{2} \theta\right\}^{1 / 2} . \tag{3}
\end{array}
$$

If $\mu-1$ be small, then $\psi$ is small and R/F differs but little from unity, as of course is obvious a priori.
Over the surface, $R$ obviously has its maximum value when $\theta=0$, its minimum when $\theta=\pi / 2$, and

Thus

$$
\left.\begin{array}{l}
\mathrm{R}_{\mathrm{max} .} / \mathrm{F}=3 \mu /(\mu+2)  \tag{4}\\
\mathrm{R}_{\mathrm{min} .} / \mathrm{F}=3 /(\mu+2)
\end{array}\right\}
$$

$$
\begin{equation*}
\mathrm{R}_{\max .} / \mathrm{R}_{\min .}=\mu \tag{5}
\end{equation*}
$$

The maximum occurs at the ends of the diameter $\theta=0$, where the force is entirely vertical, the minimum at all points on the perpendicular great circle; where the force is entirely horizontal.

Clearly the phenomena will depend mainly on whether $\mu$ is large. If it were possible to suppose $\mu$ large, we should have a wide range of values of $R$, and the disturbing forces-regarding $F$ as a disturbing field-would vary widely in direction at different parts of the surface. But it is difficult to suppose that $\mu$ is large. The Earth's own field is only of the order 0.5 C.G.S. units, and a disturbance as large as 0.01 C.G.S. is exceptional. Even in the best magnetic steel $\mu$ is low for such fields. Apart from more theoretical considerations there is the fact that, according to the above solution, if $\mu$ were large, the horizontal (or tangential) component $3 \mathrm{~F} \sin \theta /(\mu+2)$ would tend to be negligible compared to the vertical $3 \mu \mathrm{~F} \cos \theta /(\mu+2)$, except at places where $\theta$ is nearly $\pi / 2$. Now the tendency is not for the horizontal component of disturbances to be small compared to the vertical, but rather the opposite.

A point calling for special remark is that $R$ has the same value and the same absolute direction in space for places on the same great circle through $\mathrm{AA}^{\prime}$ whose angular co-ordinates are $\theta$ and $\pi+\theta$; in other words, the disturbing forces are equal and parallel at any two places diametrically situated with respect to one another.
§88. As the Earth's crust is on the whole non-magnetic, the above simple problem is obviously an imperfect representation of the facts. Thus it is worth glancing at the next simplest case, that presented by an earth composed of a nucleus of radius $a$ and permeability $\mu$, with a surrounding shell of permeability $\mu^{\prime}$ and radius $a^{\prime}$, the medium external to $a^{\prime}$ having unit permeability. The imposed field being F as before, I find for the potential external to the "earth"

$$
\begin{equation*}
\mathrm{V}=-\mathrm{Fr} \cos \theta+\mathrm{F} \frac{a^{\prime 3}}{r^{2}} \cos \theta \frac{\left(\mu+2 \mu^{\prime}\right)\left(\mu^{\prime}-1\right)+\left(\mu-\mu^{\prime}\right)\left(2 \mu^{\prime}+1\right)\left(a / a^{\prime}\right)^{3}}{\left(\mu+2 \mu^{\prime}\right)\left(\mu^{\prime}+2\right)+2\left(\mu-\mu^{\prime}\right)\left(\mu^{\prime}-1\right)\left(a / a^{\prime}\right)^{3}} \tag{6}
\end{equation*}
$$

In practice the only interesting case seems that in which $\mu^{\prime}-1$ is small, i.e. in which the material of the layer is only slightly magnetic. In this case, assuming $a / a^{\prime}$ no to be very small, the potential is approximately given by

$$
\begin{equation*}
\mathbf{V}=-\mathbf{F} r \cos \theta+\frac{\mu-1}{\mu+2} \mathbf{F} \frac{a^{3}}{r^{2}} \cos \theta \tag{7}
\end{equation*}
$$

This is identical in form with (1), the only difference being that in (1) a represents the radius of the "earth," whilst in (7) it represents the radius of the magnetic nucleus. If the slightly magnetic crust be thin, the phenomena are much the same as if the permeability were $\mu$ throughout. If the thickness of the crust be considerable, the variation in R over the surface is considerably reduced. Whether the crust be thick or thin, the value and the direction of $R$ are the same at diametral points.

In this investigation the only assumption made as to the disturbing field $F$ is that it may be regarded as uniform so far as the Earth is concerned. Thus it need not be due to the Sun, but might represent any stray field that happens to exist in any part of space traversed by the Earth, so long, of course, as the hypothesis of uniformity in strength and direction is approximately satisfied. Some of the phenomena
of terrestrial magnetism are a little suggestive of the existence of such fields, and it is well to bear the possibility of their existence in mind.
889. The elements which are usually recorded by magnetographs are the Vertical Force, the Horizontal Force, and the Declination. By an increase in Vertical Force is meant a force $\Delta V$ tending to pull towards the Earth's centre the dipping pole of the magnet. $\mathbf{A}+$ sign to $\Delta \mathbf{V}$ thus denotes a force urging the N-pole of a magnet towards the Earth's centre when the N-pole is, as at Kew, Falmouth, and Colaba, a dipping pole. But at Mauritius, Christchurch, and Winter Quarters, a + sign attached to $\Delta V$ means a force urging the $\mathbf{N}$-pole from the Earth's centre. An increase $\Delta H$ in the Horizontal Force means a force urging the N -pole in the direction of the magnetic Meridian drawn towards the magnetic Pole in the northern hemisphere. An increase $\Delta \mathrm{D}$ in the Declination means a force $\mathrm{H} \Delta \mathrm{D}$ urging the N -pole perpendicular to the magnetic Meridian, in the direction of D increasing. At Kew $\Delta \mathrm{H}$ when of + sign means a force on a $N$-pole inclined at $163^{\circ}$ to the west of geographical north, while $\Delta \mathrm{D}$ when of + sign means a force on a N -pole (tending to increase westerly Declination) inclined at $163^{\circ}$ to the south of geographical west. At Colaba, where the Declination is easterly but nearly zero, $\Delta \mathrm{D}$ when + means a force acting nearly due east. At Christchurch, where Declination is about $16 \frac{1}{3}^{\circ}$ east, $\Delta \mathrm{D}$ when positive means a force $\mathrm{H} \Delta \mathrm{D}$ inclined at about $16 \frac{1}{3}^{\circ}$ to the south of geographical east.

Thus, even as related to the local geographical directions, increments in $\mathrm{D}, \mathrm{H}$, and V have at the different stations widely different significance. As referred to fixed axes in space, the significance is even more complex. Christchurch and Falmouth, for instance, differ nearly $180^{\circ}$ in longitude, so that what is east at the one is west at the other.

The above considerations will show that in studying disturbances it is desirable not to confine our attention exclusively to the disturbances $\Delta \mathrm{D}, \Delta \mathrm{H}$, and $\Delta \mathrm{V}$, but to take account also of the disturbances $\Delta \mathrm{N}$ and $\Delta \mathrm{E}$ to geographical north and east, and, finally, to regard the disturbance at any place as a vector possessed of magnitude and having directions referred to three fixed axes at the Earth's centre.
For co-ordinate axes let us take the Earth's axis as axis of $z$, a perpendicular axis in the Meridian of Greenwich as that of $y$, and a second perpendicular axis in the Meridian $90^{\circ}$ east of Greenwich as axis of $x$. Let $\Delta \mathrm{D}, \Delta \mathrm{H}, \Delta \mathrm{V}$ denote the disturbing forces experienced at a place of latitude $\lambda$ and (easterly) longitude $l$, assuming $\Delta \mathrm{D}$ to be counted positively when easterly Declination increases, and $\Delta \mathrm{V}$ to be counted positively when it urges a N -pole towards the Earth's centre. (We are here departing, it should be noticed, somewhat from the common usage.) Then for the corresponding components $\Delta \mathbf{X}, \Delta \mathbf{Y}, \Delta \mathbf{Z}$ relative to the system of co-ordinates specified above, I find
$\Delta \mathrm{X}=\mathrm{H} \Delta \mathrm{D}(\cos l \cos \mathrm{D}+\sin l \sin \mathrm{D} \sin \lambda)+\Delta \mathrm{H}(\cos l \sin \mathrm{D}-\sin l \cos \mathrm{D} \sin \lambda)-\sin l \cos \lambda \Delta \mathrm{~V}$,
$\Delta \mathbf{Y}=\mathrm{H} \Delta \mathrm{D}(-\sin l \cos \mathrm{D}+\cos l \sin \mathrm{D} \sin \lambda)+\Delta \mathrm{H}(-\sin l \sin \mathrm{D}-\cos l \cos \mathrm{D} \sin \lambda)-\cos l \cos \lambda \Delta \mathrm{~V}$,
$\Delta \mathrm{Z}=-\mathrm{H} \Delta \mathrm{D} \cos \lambda \sin \mathrm{D}+\Delta \mathrm{H} \cos \lambda \cos \mathrm{D}-\Delta \mathrm{V} \sin \lambda$
The geographical co-ordinates of the several stations and the mean value of D for the epoch 1902-3 are given in the following table :-

| Station. | Latitude. | Longitude. | Deolination. |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |

In the formula, D is to be regarded as positive when Declination is easterly as at Christchurch, and $\lambda$ as positive when latitude is northerly as at Kew. Where the S-pole dips, $\Delta V$ is to be regarded as negative when the numerical value of V increases.

For illustration, take the case where numerical increases of $1^{\prime}$ in D and of $10 y$ in H and in V occur at each of the two stations Kew and Christchurch. We have for 1902-3 :-

|  | At Kew. | At Christehureh. |
| :---: | :---: | :---: |
| $\lambda$ | $+51^{\circ} 28^{\prime}$ | $-43^{\circ} 32^{\prime}$ |
| $\stackrel{l}{\text { l }}$ | $359^{\circ}$ $-11^{\prime}$ $-10^{\circ}$ $43^{\prime}$ | $\begin{array}{r}172^{\circ} \\ \\ +18^{\circ} \\ \hline 18 \\ \hline 18\end{array}$ |
| $\stackrel{\text { D }}{\text { D }}$ | $-16^{\circ}{ }^{\text {a }}{ }^{43^{\prime}}$ -1 | +16 ${ }^{\circ}{ }^{\circ} 18^{\prime}$ $+1^{\prime}$ |
| H | -185 | ${ }_{-22}$ |
| $\Delta \mathrm{H}$ | $+10 \gamma$ | +10\% |
| $\Delta \mathrm{V}$ | $+10 \%$ | $-10 y$ |
| $\mathrm{H} \Delta \mathrm{D}$ | $-5 \cdot 4 \gamma$ | +6.5 $\gamma$ |

As a matter of fact, in the calculations I took the Meridian of Kew, not that of Greenwich, as the $y z$ plane, measuring $l$ from it, but to the degree of accuracy aimed at that is immaterial.

If one desired to obtain the components along the fundamental $x, y, z$ axes of the components of the disturbing force $\Delta \mathrm{N}$ and $\Delta \mathrm{E}$ towards geographical north and east respectively, this would be readily effected by supposing $D=0$ in the formulæ, and writing $\Delta N$ for $\Delta H$ and $\Delta E$ for $H \Delta D$.
§90. A question calling for some consideration is: what is to be regarded as a disturbance, and how is its magnitude and direction to be determined ? Sabine, whose work on Terrestrial Magnetism still commands respect, regarded the value of an element at any particular instant as disturbed when it departed from the mean value of the element at that time of day-contributions from disturbed days having been removed-by more than a specified limiting value. This regards the departure from an undisturbed mean value as the measure of a disturbance. The difficulty, of course, is to arrive at the undisturbed mean value. Sabine's method, theoretically considered, was to do this by a process of sifting, rejecting first such individual values as departed notably from the mean derived from all, then forming a new mean from the individual values retained and repeating the process, and so on.

In practice the method would be very laborious if strictly followed. A serious difficulty is that the amplitude, and sometimes the type, of the regular diurnal inequality varies largely throughout the year, and that there is a large sun-spot influence on the amplitude at least, if not on the type.

From the point of view of Sabine's definition, the natural thing would be to take $\Delta \mathrm{D}, \Delta \mathrm{H}$, and $\Delta \mathrm{V}$ as given at any instant by the departures of $\mathrm{D}, \mathrm{H}$, and V respectively from mean undisturbed values appropriate to the hour. But to make even the pretence of doing this satisfactorily it would be necessary to have a knowledge of what might reasonably be regarded as normal or undisturbed values. To obtain such normal values for the Antarctic seemed wholly impracticable, and to have obtained them even for the co-operating stations would have required regular diurnal inequalities to have been formed appropriate for the several months of the years 1902 and 1903. One could, of course, have derived diurnal inequalities from the hourly observations on the international term days, but two days a month afford a very slender basis for diurnal inequalities.

In the case of the larger Antarctic disturbances comparatively little could be done, because one at least of the traces was pretty sure to be off the sheet, and even if all three traces happened to be on, the oscillatory movements were so large and rapid that it was difficult to assign the times with sufficient accuracy. Most of the measurements that proved practicable referred to disturbances of the special type dealt with in Chapter X.

On considering the situation, I decided that the most hopeful course to pursue was to focus attention on changes in the values of the magnetic elements occurring during comparatively short intervals, especially on the commencing movements introducing magnetic storms. In dealing with changes of short duration we are comparatively independent of anything but disturbance pure and simple, especially when the changes occur at hours when the regular diurnal changes are slow, or at seasons when the regular diurnal changes are small. This seens to be practically the same conclusion as that reached many years ago by Dr. Balfour Stewart, when he gave his attention to what he called "peaks" and "hollows" in the magnetic curves, i.e. turning-points separated by no long interval of time.

To give greater definiteness, take the most favourable case where three consccutive turning-points present themselves simultaneously in the D, H, and V records. Suppose the 1st and 3rd to represent minima, the 2nd a maximum, and suppose the values of the elements answering to the two minima to be identical.

Let $t_{1}, t_{2}, t_{3}$ be the times of the three turning-points, and let $\Delta \mathrm{D}, \Delta \mathrm{H}$, and $\Delta \mathrm{V}$ represent the excess of the values of $\mathrm{D}, \mathrm{H}$, and V at time $t_{2}$ over the values at time $t_{1}$ or at time $t_{9}$.
If we take the condition of matters existing at time $t_{1}$ as fundamental, we may regard the condition existing at time $t_{2}$ as arising from the action of a disturbing force $\Delta \mathrm{R}$, whose components $\Delta \mathrm{X}, \Delta \mathrm{Y}, \Delta \mathrm{Z}$ are obtained from (8) by assigning to $\Delta \mathrm{D}, \Delta \mathrm{H}$, and $\Delta \mathrm{V}$ their values with positive sign. If, however, we take the condition at time $t_{2}$ as fundamental, we regard the condition at time $t_{3}$ as due to a disturbing force $\Delta R^{\prime}$, whose components are obtained from (8) by assigning the same numerical values to $\Delta \mathrm{D}, \Delta \mathrm{H}$, and $\Delta \mathrm{V}$, but giving them the negative sign. The vectors $\Delta \mathrm{R}$ and $\Delta \mathrm{R}^{\prime}$-assuming the changes $\Delta \mathrm{D}, \Delta \mathrm{H}, \Delta \mathrm{V}$ to be small-are numerically equal, but are oppositely directed. Presumably $\Delta \mathrm{R}$ and $\Delta \mathbf{R}^{\prime}$ really represent, the one the application of a force, the other its removal. The difficulty in practice is to know which is which. As a rule, when an oscillatory movement occurs, the to-and-fro movements are not equal, and there are usually both preceding and succeeding movements. Even if we take the sudden movements ushering in storms there is room for some doubt. At Kew, for instance, when such a movement occurs, $H$ is usually found to be enhanced after the first rapid change has ceased, but at the very commencement of the movement there is at least sometimes a slight depression in the value. If no preliminary depression is seen, it is certainly the natural thing to regard the disturbance as simply an increase in H , but still there is the possibility that the rise represents in reality the removal of a disturbing force which has remained nearly constant for a considerable time. We are much in the same position as an observer who sees the length of a bar under test, but does not know whether it is loaded or unloaded, nor whether the load is a tension or a compression.
§91. An investigation was made of the disturbed curves received from the co-operating stations, and when the three elements presented synchronous changes and there appeared an agreement in time between the changes recorded at two or more of the different stations, measurements were made. The algebraical excess of the value of an element at the end of the interval considered over the value at the commencement was regarded as representing the action of a disturbing force. No attempt was made to determine whether the change was really due to the application of a force or to its removal. From the $\Delta \mathrm{D}$ and $\Delta \mathrm{H}$ thus found the values of $\Delta \mathrm{N}$ and $\Delta \mathrm{E}$ (the components to geographical north and east) were calculated, and also the values of $\Delta \mathbf{X}, \Delta \mathbf{X}, \Delta \mathrm{Z}$ as given by (8).

The resultant $\Delta \mathrm{R}$ of $\Delta \mathrm{X}, \Delta \mathbf{Y}$, and $\Delta \mathrm{Z}$ was then found, and finally the inclination $\theta$ of this resultant to the Earth's axis and the easterly longitude $\phi$ of the meridian plane containing it.

The positive direction of the Earth's axis was taken as given by the radius drawn to the North Pole. $\theta$ and $\phi$, are given, of course, by

$$
\begin{equation*}
\cos \theta=\Delta \mathrm{Z} / \Delta \mathrm{R}, \quad \tan \phi=\Delta \mathrm{X} / \Delta \mathrm{Y} \tag{9}
\end{equation*}
$$

The most uncertain measurements by far at the co-operating stations were those of $\Delta \boldsymbol{V}$. The changes in V were usually small. For 1902 no Falmouth V curves were available, so that there was nothing to check the Kew ones by, and at Colaba the sensitiveness was throughout so small that a small change was hardly visible.
§92. Tables LXII to LXVI at the end of this chapter give the results obtained from the measurement of corresponding movements at Kew, Colaba, Mauritius, Christchurch, and the Antarctic. The occasions dealt with are taken from the storms included in the plates. The times are all G.M.T.

Of the examples for Kew in Table LXII, p. 181, those from August 20, 1902, April 5 and August 25, 1903, refer to sudden commencements of storms. The same is true of the disturbance on May 8, 1902, between $11.59 \mathrm{a} . \mathrm{m}$. and $0.9 \mathrm{p} . \mathrm{m}$. This disturbance is treated as a whole for purposes of comparison with the other co-operating stations, but also as composed of two parts for comparison with the Antarctic, where the commencing movement is pronouncedly double. The rapid commencements were also treated as if the mean derived from them represented a disturbance.

The disturbances of June 19, 28, 29 and July 26, 1903, were simultaneous with the two phases of the special type of disturbance recorded in the Antarctic. In this case also an imaginary mean disturbance has its components calculated, and the resultant and its direction angles thence deduced.

Owing to electric tram disturbances, changes of the order $0.5 \gamma$ in $V$ could not be recognised with any approach to certainty. This is no doubt partly responsible for the fact that the value 0 is so frequently assigned to $\Delta V$. When the vector is small there is inevitably a good deal of uncertainty in the determination of both $\theta$ and $\phi$, and the values given must not be regarded as usually more than somewhat rough approximations to the truth.

On examining the results it will be seen that practically all the disturbances may be regarded as included under one or other of two types, distinguished by the letters $\mathbf{A}$ and $\mathbf{B}$. The disturbances of the type $\mathbf{A}$ have a value of $\theta$ in the neighbourhood of $60^{\circ}$ and a value of $\phi$ in the neighbourhood of $220^{\circ}$. They include those sudden commencements in which the increase of H is the most noticeable feature, and the movements corresponding in time with the second phase of the Antarctic disturbances of the special type. An exception is provided by the disturbance of June 19,1903 , in which the value of $\phi$ departs largely from $220^{\circ}$. On this occasion, however, the disturbance at Kew was so small that comparatively little weight attaches to the results deduced. The disturbance on May 8, 1902, from 7.4 to 7.11 p.m., is also included amongst the A's. The disturbances to which the letter B is attached have a value of $\theta$ which is in the neighbourhood of $120^{\circ}$. The values of $\phi$ are more variable, but are all included between $+63^{\circ}$ and $-63^{\circ}$. The movements corresponding in time to the first phase of the special type of Antarctic disturbances all come fairly under this category; so does the first phase of the sudden commencing movement of August 25, 1903.
§93. The disturbances at Colaba, in Table LXIII, p. 182, are also mainly of two types, distinguished as before by the letters $A$ and $B$. The characteristic feature at Colaba is the large size of $\Delta Z$, the component parallel to the Earth's axis, and the small size of $\Delta \mathbf{X}$ and $\Delta Y$. Owing to this latter fact the evaluation of $\phi$ is particularly uncertain, and comparatively little significance can be assigned it. We may thus regard the essential feature of the type $\mathbf{A}$ disturbances as the possession of a small value for $\theta$, and the essential feature of the type $B$ disturbances as the possession of a value for $\theta$ approaching $180^{\circ}$.

The sudden commencements in which H increases, and the disturbance synchronising with the second phase of the Antarctic disturbance of July 26, 1903, are included amongst the A's, and so is the movement between 7.4 and 7.11 p.m. on May 8, 1902. Thus the A's at Colaba and the A's at Kew correspond to one another.

Comparing the values of $\Delta \mathrm{R}$ at Kew and Colaba, it will be seen that the former are very decidedly the larger, except on May 8, 1902.
§94. For the Mauritius Table LXIV, p. 182, our information is less extensive, as the copies of disturbances in D and V were mostly confined to 1902.

As at Kew and Colaba, the sudden commencements of May 8 and August 20, 1902, and the movement from 7.4 to $7.11 \mathrm{p} . \mathrm{m}$. on the former date, are of similar type and are classed as $\mathbf{A}^{\prime} \mathrm{s}$. The values obtained from $\Delta \mathrm{R}$ are all smaller than the corresponding values at Kew or Colaba.
§95. The data for Christchurch, Table LXV, p. 182, include all the disturbances already considered for Kew. The results are more difficult of classification than those at the other co-operating stations. The commencing disturbances of May 8 and August 20, 1902, and of April 5, 1903, as well as the movements answering in time to the second phase of the Antaretic disturbances of June 19, June 28, and July 26, have been classed as A's. They present, however, a considerable range of values in $\theta$ as well as in $\phi$. This group also includes the first of the to-and-fro movements experienced on August 25, 1903. This last movement was, however, so minute that too much weight should not be ascribed to the fact.

As elsewhere, the movement from 2.10 to 2.28 p.m. on May 8, 1902, and that answering to the first phase on July 26, 1903, appear of similar type and have been classed as B's. But the first movement on July 26 is clearly of the same type as the second on June 29, so this also must be classed as a B. The first-phase movements on June 19, June 28, and June 29, 1903, are very similar in type, but differ from either the A or the B class, and have accordingly been classified as $\mathrm{C}^{\prime} \mathrm{s}$. This class seems also to include the second and larger movement on August 25, 1903.

If we consider the sudden commencements, we find that the values of $\Delta \mathrm{R}$ at Christchurch are decidediy
less than those at Colaba, and apparently a little less than even those at Mauritius. When, however, we consider the disturbances answering to those of the special type in the Antarctic, it is otherwise.
§96. Table LXVI, p. 183, gives results for the Antarctic corresponding to those given for the co-operating stations. The movements include those treated in Table LXII, except the movement on May 8, 1902, from 7.4 to 7.11 p.m., which was partly lost in the Antarctic, and they include, in addition, the second phase of the commencing movement on April 5, 1903, which was essentially peculiar to the Antarctic. The letter $i$ denotes that the Antarctic records were not absolutely complete. On April 5, 1903, the first phase was complete, but in the second phase the $H$ curve went off the sheet. The amplitude shown, $300 \gamma$, was probably not much exceeded. On June 19 and 28 , and July 26 , both phases suffered loss. On June 19 the loss was confined to H , and on June 28 to D, but on July 26 both D and H went beyond the limit of registration. In all these cases, so far as one could judge from the appearance of the curves, the major part of the movement was recorded, though on July 26 the loss was probably greater than on the other days. These losses, of course, introduce some uncertainty, and one would naturally have used other examples of the special type of Antarctic disturbance but for the fact that a comparison was desirable with other stations, and that days had thus to be chosen when records from the co-operating stations were available.

To reduce the uncertainty thus arising, a mean disturbance was derived from 51 examples of the special type of disturbance for which the records were complete. The values of $\Delta \mathrm{D}, \Delta \mathrm{H}$, and $\Delta \mathrm{V}$ were meaned for the two phases separately, and from the mean $\Delta \mathrm{D}$ and $\Delta \mathrm{H}$ corresponding mean values for $\Delta \mathrm{N}$ and $\Delta \mathrm{E}$ were found. From these and the mean $\Delta V$ there were calculated corresponding mean values of $\Delta \mathbf{X}, \Delta \mathbf{Y}$, and $\Delta Z$, and the force vector deduced therefrom. Comparing the results thus found with those obtained by meaning the disturbances of June 19, 28, 29, and July 26, 1903, in like fashion, we see that in the case of the first phase there is a very close agreement in the values of the angles $\theta$ and $\phi$. In the case of the second phase the agreement is not quite so good.

The sudden commencements were also grouped with a view to obtaining representative means. The operation was limited to May 8 and August 20, 1902, and August 25, 1903, these being the only three occasions on which both phases were completely recorded. In this case the phases in which the elements increased were grouped together, irrespective of whether they occurred first or second. The values of $\theta$ and $\phi$ obtained for the two phases in this case correspond fairly closely with those obtained for the two phases of the mean representative of the special type of disturbance.

All the examples in Table LXVI can be fairly included in two classes, distinguished as before by the letters $A$ and $B$. Class $B$ includes all the first phases of the special type of disturbance, and that phase of the sudden commencements during which the element decreased in value. Class A , on the other hand, includes all the second phases of the special type of disturbance, and that phase of the sudden commencements during which the element increased in value. In choosing which letter to apply, the guiding principle was, that at the co-operating stations $A$ represented a type of disturbance in which the element or elements of force chiefly affected during sudden commencements exhibited an increase.
§97. On July $24-25,1902$, there were recorded at Kew a number of oscillatory disturbances of no great magnitude. On examining these I found that the turning-points in the three elements appeared identical in a considerable number of cases, and also that corresponding movements could be traced at the other co-operating stations. Table LXVII, p. 184, deals with the measured changes of force on ten of these occasions, at Kew, Colaba, and Mauritius, and with a minor number of them at Christchurch and the Antarctic. As the afternoon of the 24th (G.M.T.) advanced, it became increasingly difficult to obtain movements that appeared to correspond, the difficulty appearing first in the Antarctic data, and then in the Christchurch ones.

The first four movements consisted, at Kew, of a rise in force from 2.23 to $2.30 \mathrm{p} . \mathrm{m}$., followed by a fall from 2.30 to $2.38 \mathrm{p} . \mathrm{m}$., then another rise from 2.38 to $2.43 \mathrm{p} . \mathrm{m}$., and a second fall from 2.43 to $2.54 \mathrm{p} . \mathrm{m}$. The two double movements appeared of the same type at all the stations; thus, instead of treating them separately I took a mean from the two, combining the two falls together, and the two rises together. These are numbered (1) and (2) in the table, as if the means represented each a single movement. The other cases, (3) to (8), represent actual single changes of force. If we examine the
eight cases, we see that at Kew Nos. (1), (5), and (7) may be regarded as of the Class A of Table LXII, whilst Nos. (2), (3), (4), (6), and (8) are of Class B.

Coming to Colaba we recognise No. (1) as of the Class A of Table LXIII, and Nos. (2), (3), and (4) as of Class B. So far as the angle $\theta$ is concerned, Nos. (5) and (7) approach Class A, and Nos. (6) and (8) approach Class B, but the $\phi$ angle differs rather notably from that characteristie of the respective classes.

At Mauritius we can recognise No. (1) as of the Class A of Table LXIV, and Nos. (2), (3), (4), and (8) as of the Class $\mathbf{B}$.

At Christchurch Nos. (1) and (5) are of the Class A of Table LXV, though No. (5) is rather outstanding, and Nos. (2), (3), and (4) are good examples of Class B. After 5. 42 p.m., G.M.T., the Christchurch curves seemed to lose their parallelism with those at Kew entirely.

At the Antarctic No. (1) is fairly of the type of Class A of Table LXVI, and No. (2) is fairly of Class B. The turning-points, however, in the Antarctic were not very clearly marked in the $H$ and $V$ curves, and whilst they were clearly marked in the D curve-in which alone the movements stood out from their neighbours-they appeared to be hardly absolutely coincident in time with the movements they were believed to correspond to at Kew.
§98. In the course of our discussion of Tables LXII to LXVII there have been references to the relative size of the corresponding disturbances at different stations. The information on this point is summarised in Table LXVIII, p. 185, which expresses the amplitude of the disturbance $\Delta \mathbf{R}$ at each station in terms of the amplitude of the synchronous disturbance at Kew. In addition to results from individual cases the table gives mean results, treating separately the sudden commencements, the special type of disturbance, and the disturbances which belong to neither of these categories. Considering the comparatively limited data, too much weight must not be attached to numerical resemblances which may be partly accidental.

We have already seen that, so far as type is concerned, sudden commencements do not appear to be in any way essentially different from other short-period movements such as those of July 24, 1902. The similarity seems to extend to the variation in amplitude with geographical position. Taking either of these types of disturbance, the amplitudes at Mauritius and Christchurch are decidedly less than those at Colaba, which in their turn are less than those at Kew and Falmouth. But the movements, even at Kew and Falmouth, are much exceeded by those in the Antarctic, and this seems the case habitually, irrespective of whether the time of occurrence is day or night at Greenwich.

With regard to the relatively small size of sudden movements at Mauritius and Christchurch there is one possible explanation which must be borne in mind. In selecting disturbed days for comparison I had before me only Kew and Antarctic curves, the latter being invariably disturbed. Thus, no doubt, it was the amplitude of the Kew disturbance, or some special feature in it, which mainly determined the choice. It is thus conceivable, if instead of the Kew curves I had had the curves from some station in the southern hemisphere to guide me, that a different selection might have been made, and that under these circumstances there might not have been the pre-eminence in the amplitude of the disturbances at Kew and Colaba as compared to those at Mauritius and Christchurch which the table shows. Some countenance to this view is supplied by the results for the special type of disturbance. Corresponding movements could indeed be traced at Kew, but in no case were they so outstanding as to have caught the eye if merely glancing through the curves generally. In their case it will be seen that the disturbance experienced at Christchurch was usually much larger than that at Kew. If, however, the selected sudden commencements represented magnetic effects whose seat was mainly in the northern hemisphere, the extraordinarily large size of these movements in the Antarctic would be even more remarkable than it appears to be. Thus I do not think that the suggested explanation suffices, though there may be some truth in it.

As regards the special type of disturbance, we must, I think, conclude that though the effects are felt in the northern hemisphere, yet the seat of the disturbance must be mainly at least in the southern hemisphere.
§99. There is one notable peculiarity about the Antarctic results that calls for comment. At Kew, Falmouth, Colaba, Mauritius, and Christchurch the result of a sudden commencing movement, whether single or visibly double, is normally-there may be exceptions-to leave the total force increased. At Kew the
change is usually greatest in $H$, though sometimes in I). Taking the $H$ change as the most notable, we can usually see with certainty only an increase, but sometimes a smaller decrease for a relatively short time is distinctly visible, and on other occasions, though not clearly visible, it is suggested by the appearance of the curve. In the Antarctic there were six occasions-May 8, August 20, and Novemher 6, 1902; and April 5, August 25, and December 13, 1903 (Greenwich dates)-when sudden movements were detected synchronous in time with sudden commencements at Kew and elsewhere. All these six were distinctly double or oscillatory movements, and in all the second movement was the larger. So far there is no certain difference from the phenomena at Kew, as we are not in a position to say with certainty that when only one movement was seen at Kew it was unaccompanied by a previous very small and short-lived diminution in H. But of the six Antarctic double movements, three had the first movement a decrease in the elements, while three had it an increase. The three occasions which showed the increase first were May 8 and November 6, 1902, and December 13, 1903.

On May 8, 1902, we have an Antarctic disturbance A followed by a larger disturbance B , the two synchronous with a Kew disturbance A. On August 20, 1902, we have an Antarctic B disturbance, unrepresented elsewhere, followed by an A disturbance which synchronised with an $\mathbf{A}$ disturbance at Kew. On November 6, 1902, we have an Antarctic A disturbance synchronising with a Kew B disturbance, and then an Antarctic B disturbance synchronising with a Kew A disturbance. On April 5, 1903, we have an Antarctic B disturbance synchronising with a Kew A disturbance, being followed by a larger A disturbance not represented at Kew. On August 25, 1903, we have an Antarctic B disturbance synchronising with a Kew B disturbance, there immediately following a larger A disturbance at both stations. On. December 13, 1903, we have an Antarctic A disturbance not apparently represented elsewhere, followed by a $B$ disturbance which covered the time occupied by a $B$ and $A$ disturbance at Kew, the latter the larger.

All of the special disturbances present the $\mathbf{B}$ type first in the Antarctic, and this is true in three of the four cases at Kew. Between 2.10 and 2.28 p.m., G.M.T., on May 8, 1902, we have a B disturbance at Kew, but an A disturbance in the Antarctic. On July 24, 1902, the A and the B disturbances at the two stations correspond.

It would thus appear that whilst A and B movements in the Antarctic are just as opposed to one another in type as they are elsewhere-the one representing an increase, the other a decrease in the elements which vary most-the order in which they occur shows a variability which is at least unusual, and either may synchronise with an $\mathbf{A}$ movement at the co-operating stations.
$\S 100$. The subject of the coincidence in time of the commencement of the magnetic storm of May 8, 1902, and the eruption of Mont Pelée has been already referred to ( 857 ), and a further discussion was promised. We see from Tables LXII to LXV that at Kew, Colaba, Mauritius and Christchurch-and the same is true of Falmouth-the movement was of the same type $\mathbf{A}$ as other sudden commencements of storms. The only peculiarity was that at Kew and Falmouth-and possibly at Colaba-there was a short suspension of the upward movement in $H$, only just recognisable in the curve. In the Antarctic the movement, it is true, was represented by an oscillatory $A$ and $B$ movement-instead of by a simple $A$ movement -but a similar phenomenon occurred on December 14, 1903. There seems thus to be nothing of an outstanding character in the type of the commencing disturbance of May 8. As regards the relative amplitudes of the disturbance at different stations, there is nothing at all outstanding in the ratio recorded on this occasion in Table LXVIII. The Colaba disturbance was certainly relatively larger than usual, though not so large relatively as two hours later in the same day, and the Mauritius disturbance was also a little above average; but there is nothing at all abnormal in the figures at these stations. Moreover, the fact, that relatively considered the Kew and Falmouth disturbances are somewhat less than usual compared to those at Colaba and Mauritius, is the reverse of favourable to the view that the disturbance was directly due to the Mont Pelée eruption. If an eruption, which consists of a vertical movement of material, causes a magnetic disturbance, one would certainly expect it to be of a more or less symmetrical character round the vertical at the place, and one would unquestionably expect the disturbance to fall off rapidly as the distance increases. Now, somewhat curiously, Colaba, Mauritius and Christchurch are not far from equidistant from Martinique, all being at an angular distance fully double that of Kew and

Falmouth; and Winter Quarters, though nearer than the first three of these stations, is comparatively little nearer. The angular distances are, in short, approximately as follows: Christchurch $126^{\circ}$, Colaba $123 \frac{1}{2}^{\circ}$, Mauritius $121 \frac{1}{2}^{\circ}$, Winter Quarters $113^{\circ}$, Kew $603^{\circ}$. Thus, if the disturbance had been a direct consequence of the eruption, what we would have expected to find would have been disturbances of nearly equal magnitude at Christchurch, Colaha, Mauritius, and Winter Quarters, that magnitude being much less than that of the disturbance at Kew and Falmouth. Whereas we see from Table LXVIII that the disturbance at Winter Quarters was of the order of ten times that at Christchurch, while the disturbance at Colaba exceeded that at Kew.

Again, if the commencement of the May 8 storm was directly due to the eruption, what are we to think of the remainder of the storm which lasted 8 hours and presented a remarkable unity of appearance; and how are we to explain the fact that the largest movements occurred 6 or 7 hours after the commencement, and that one of these movements, viz, that occurring between 7.4 and 7.11 p.m., G.M.T., presented at Kew, Colaba and Mauritius a remarkable similarity in type to the commencing movement? The conclusion we seem led to is that the coincidence in time with the eruption was purely accidental.
§101. The bearing of the results of Tables LXII to LXVIII on theory calls for a short comment. We have found that disturbances of comparatively short period-whether commencements of magnetic storms or not-show a general tendency to approximate to one or other of a small number of types. Except at Christchurch, in the case of the special type of disturbance, there was no very clear indication of more than two classes, confining ourselves, of course, to cases in which a distinct correspondence was visible between synchronous disturbances at distant stations. In the limited number of cases we have considered, there has also been-apart from the special type of disturbance-a general tendency for disturbances to be larger at some stations than others.

On what I have called the action-at-a-distance theory this would imply that the disturbing field supposed has a tendency to have a more or less fixed direction relative to axes fixed in the Earth. If the direction assumed by that field were largely variable-as it would be if it pointed to the Sun, of were at right angles to the line joining the Earth and Sun-then one would expect the maximum disturbance to be experienced at widely different places on different occasions. The sudden commencements we have had to do with occurred at different times, G.M.T., that of May 8, 1902, being near Greenwich noon, those of April 5 and August 25, 1903, near Greenwich midnight. Again, the sudden commencements were at widely different seasons of the year, yet we do not find any conspicuous difference between the phenomena experienced.

As already remarked, Christchurch is not very far from being a diametral point to Falmouth and Kew, and on the action-at-a-distance theory we should expect the disturbance vectors at diametral points to be parallel in direction and equal in magnitude. So far as the parallelism in direction is considered, there is unquestionably considerable support to the theory. If we take Table LXVII for instance, we have for the first five (or really seven) disturbances :-


There is here unquestionably a somewhat remarkable accordance. When, however, we come to the magnitude of the vectors, we find that the disturbance at Christchurch was invariably less than half that at Kew.

When considering the significance of the vector angles $\theta$ and $\phi$ there is one point that should be borne in mind. The assumption by a vector of a nearly constant direction unquestionably suggests the action of an external force having also a nearly fixed direction. It may mean, however, no more than that the action of any external impulse tends to influence the Earth's magnetism in a particular way, just as when
one hits a bar magnot in a neutral field one usually knocks out magnetism irrespective of where and how one bits it.

Suppose, for instance, that at cach station we had a change only in II, the horizontal component of the Earth's field, then the vectors we should derluce at the several stations would have the following values for $\theta$ and $\phi$, supposing H to increase at all :-


Under the circumstances supposed, these angles would be independent of the size of the change in $H$.
An examination of Tables LXII to LXVII will show that the values of $\theta$ and $\phi$ thus found are in no case very remote from the values actually calculated for these angles in the case of the disturbancos classed as A.

One phenomenon that unquestionably tells against the action-at-a-distance theory is that at all the stations considered $\Delta \mathrm{V}$ is invariably less than the resultant horizontal component $\sqrt{ } \Delta \mathrm{N}^{2}+\Delta \mathrm{E}^{2}$ of the disturbing vector. The horizontal component is usually much the larger, even in the Antarctic, where the disturbances are greatest.

Whatever other conclusion may be drawn from this investigation, it is, I think, clear that the application of the method to a variety of disturbances drawn from different hours of the day, G.M.T., and from different seasons of the year offers a promising field of research.

Table LXII.-Disturbance Components and Resultants at Kew. (Unit of force $1 \gamma$.)

| Date and time (G.M.T.). | $\Delta N$. | $\Delta \mathrm{E}$. | $\Delta V$. | $\Delta \mathrm{X}$. | $\Delta \mathbf{Y}$. | $\Delta Z$. | $\Delta \mathrm{R}$. | $\theta$. | $\phi$. | Type. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1902 |  |  |  |  |  |  |  | - | - |  |
| May 8, 11.59 a.m. to 0.4 p.m. | + 80 | - 46 | 0.0 | $-46$ | -6.3 | $+5.0$ | $9 \cdot 3$ | 57 | 216 | A |
| " 8, 0.4 p.m. " 0.9 " | + $7 \cdot 2$ | - $7 \cdot 2$ | $0 \cdot 0$ | $-7 \cdot 2$ | $-5 \cdot 7$ | $+4.5$ | $10 \cdot 2$ | 64 | 232 | A |
| " 8, 11.59 a.m. " 0.9 " | $+15 \cdot 1$ | -11.9 | $0 \cdot 0$ | -11.9 | -11* | $+2 \cdot$ | $19 \cdot 2$ | 61 | 225 | A |
| \% 8, 2.10 p.m. , 2.28 , | - 5 2 | + $7 \cdot 2$ | $0 \cdot 0$ | + 72 | + 4.0 | $-3 \cdot 2$ | $8 \cdot 9$ | 111 | 61. | B |
| , 8, 7.4 $\quad$, 7.11 " | +15.9 | $-16 \cdot 6$ | $+4.0$ | $-16 \cdot 6$ | $-14.9$ | + 5.8 | $23 \cdot 1$ | 75 | 228 | A |
| August 20, 9.6 " " 9.10 " | $+12 \cdot 5$ | $-7 \cdot 1$ | $+10$ | $-7 \cdot 1$ | $-10 \cdot 4$ | $+7$. | 14.4 | 61 | 214 | A |
| 1903 |  |  |  |  |  |  |  |  |  |  |
| April 5, 11.25 p.m. to 11.29 p.m. | $+34 \cdot 2$ | $-28 \cdot 8$ | +5.2 | -28 8 | $-30 \cdot 0$ | $+17 \cdot 2$ | $45 \cdot 0$ | 68 | 224 |  |
| June 19, 7.38 日.m. , 7.55 \&.m. | $-12 \cdot 4$ | $+2 \cdot 1$ | 00 | + $2 \cdot 1$ | + $9 \cdot 7$ | $-7.8$ | $12 \cdot 6$ | 128 | 12 | B |
| \% 19, 7.55 " $\quad 8.13$, | $+1.4$ | $\begin{array}{r} \\ +46 \\ \hline\end{array}$ | $0 \cdot 0$ | + 4.6 $+\quad 11.8$ | - $1 \cdot 1$ | + 0.9 $+\quad 4.8$ | 4.8 | 79 | 103 | A? |
| " $28,8.36$ " $\quad 8.47$, | - 77 | -11.8 | $0 \cdot 0$ | $-11.8$ | + 60 | - 4.8 | $14 \cdot 1$ | 110 | 297 | 13 |
| \% 28, 8.47 " $\quad$, 9.8 " | + 8.5 | - 4.9 | $0 \cdot 0$ | - $4 \cdot 9$ | $-6 \cdot 6$ | + $5 \cdot 3$ | $9 \cdot 8$ | 57 | 217 | A |
| " $29,8.59$ " ", 9.26 " | -9.3 | - $5 \cdot 1$ | $0 \cdot 0$ | $-5 \cdot 1$ | + 73 | - 5.8 | $10 \cdot 6$ | 123 | 325 | B |
| , 29, 9.26 ", 9.53 " | + $0 \cdot 3$ | $-14 \cdot 7$ | $0 \cdot 0$ | $-14 \cdot 7$ | -0.2 | + $0 \cdot 2$ | $14 \cdot 7$ | 89 | 269 | A |
| July 26, 6.23 , " 6.55 , | -39.9 | $-11 \cdot 1$ | $-1 \cdot 0$ | $-11 \cdot 1$ | +31.8 | $-24 \cdot 1$ | 41.4 | 126 | 341 | B |
| \% 26, 6.55 " " 7.30."3 | +22.8 | $-16 \cdot 4$ | +4.0 | $-16 \cdot 4$ | $-20 \cdot 3$ | +11.0 | $28 \cdot 3$ | 67 | 219 | A |
| August 25, 10.57 p.m. ", 11.0 p.m. | - 477 | + 5.3 | $0 \cdot 0$ | + $5 \cdot 3$ | + 36 | - 2.9 | $7{ }^{\circ} 0$ | 114 | 56 | B |
| " 25, 11.0 " "11.8 " | $+30 \cdot 7$ | $-16^{\circ}$ | $0 \cdot 0$ | -16 0 | $-24.0$ | $+19 \cdot 1$ | $34 \cdot 6$ | 56 | 214 | A |
| Mean from sudden commencements of May 8 and August 20, 1902, April 5 and August 25, 1903, taken until end of rise in H | $+23 \cdot 1$ | -16.0 | +1 $15_{5}$ | $-16 \cdot 0$ |  | +13 $2 \cdot 28 \cdot 1$ |  | 62 | - 220 |  |
|  |  |  |  |  | -19 0 |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |
| Mean from oscillations of June 19, |  |  |  |  |  |  |  |  |  |  |
| 28, and 29, and July 26, 1903- |  |  |  |  |  |  |  |  |  |  |
| First movement | - | - | - | - 6.5 | + 13 ' 7 | $-10 \cdot 6$ | 18.5 | 125 | 335 | B |
| Second " | - |  |  | $-78$ | $-7.0$ | $+4 \cdot 4$ | 11.4 | 68 | 228 |  |

Table LXIII.-Disturbance Components and Resultants at Colaba. (Unit of force $1 \gamma$.)

| Date and time (G.M.T.). | $\Delta \mathbf{N}$. | $\Delta \mathrm{E}$. | $\Delta \mathrm{V}$. | $\Delta \mathbf{X}$. | $\Delta \mathbf{Y}$. | $\Delta Z$. | $\Delta \mathrm{R}$. | $\theta$. | ¢. | Type. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1902 |  |  |  |  |  |  |  | - | - |  |
| May $8,11.59 \mathrm{a.m}$. to 0.9 p.m. | +19 0 | $+76$ | $-7.0$ | +2.7 | $-7 \cdot 2$ | $+20 \cdot 2$ | 21.6 | 21 | 159 | A |
| "8, 2.10 p.m. " 2.28 " | $-11{ }^{2}$ | - $3 \cdot 7$ | $+3 \cdot 2$ | -0.5 | + $3 \cdot 7$ | $-11 \cdot 6$ | $12 \cdot 2$ | 162 | 352 | B |
| , 8, 7.4 ", 7.11 " | $+18 \cdot 4$ | + 37 | $-6 \cdot 5$ | +1.2 | -3.5 | + 19 \% | $19 \cdot 8$ | 11 | 160 | A |
| August 20, 9.6 " "9.10 " | $+10 \%$ | + $2 \cdot 2$ | $-16$ | $-1 \cdot 1$ | $-2 \cdot 6$ | $+10 \cdot 2$ | $10^{\circ} 6$ | 16 | 203 | A |
| 1903 |  |  |  |  |  |  |  |  |  |  |
| April 5, 11.25 p.m. to 11.29 p.m. | + 25 . | + 4.4 | $-8 \cdot 3$ | $+1 \cdot 0$ | $-4 \cdot 3$ | $+26 \cdot 3$ | $26^{\circ} 8$ | 11 | 167 | A |
| July 26, 6.23 a.m. , 6.55 a.m. | $-25 \cdot 6$ | + $2 \cdot 2$ | $0 \cdot 0$ | $+8 \cdot 6$ | +0.3 | $-24.2$ | $25 \cdot 7$ | 160 | 88 | B |
| $" 26,6.55 \quad, \quad 7.30 \text {,. }$ | $+61$ | + 5.5 | $0 \cdot 0$ | $-0.3$ | -5.8 | + 5.8 | $8 \cdot 2$ | 45 | 183 | A |
| A.ugust 25, 10.57 p.m. " 11.0 p.m. | -175 | + $2 \cdot 2$ | $0 \cdot 0$ | $+1 \cdot 1$ | -1.9 | -1.4 | $2 \cdot 6$ | 123 | 150 | B ? |
| " 25, 11.0 " "11.3 " |  | $-1 \cdot 1$ | $-2 \cdot 5$ | $-3 \cdot 5$ | $+0 \cdot 1$ | +17 3 | 177 | 11 | 272 | A |
| Mean from sudden commencements (same times as for Kew) | +179 | + 3 3 | -4.8 | -0.2 | -3'5 | + 18 \% | $18 \cdot 8$ | 11 | 184 | A |

Table LXIV.-Disturbance Components and Resultants at Mauritius. (Unit of force $1 \gamma$.)


Table LXV.-Disturbance Components and Resultants at Christchurch. (Unit of force 1 $\mathbf{\gamma}$.)

| Date and time (G.M.T.). | $\Delta \mathrm{N}$. | $\Delta \mathbf{E}$. | $\Delta V$. | $\Delta \mathbf{X}$. | $\Delta Y$ | $\Delta Z$. | $\Delta \mathbf{R}$. | $\theta$. | $\phi$. | Type. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{array}{\|ccccc} 1902 & & & \\ \text { May 8, } & 11.59 \text { a.m. to } & 0.9 & \text { p.m. } \\ \text { 8, } & 2.10 & \text { p.m. } & , & 2.28 \\ \text { A } \end{array}$ | $\begin{array}{r} +83 \\ +\quad 40 \\ +\quad 51 \end{array}$ | $\begin{array}{r} +38 \\ -4.3 \\ -2.6 \end{array}$ | $\begin{array}{r} -0.5 \\ +0.5 \\ 0.0 \end{array}$ | $\begin{array}{r} 31 \\ +\quad 39 \\ +\quad 31 \end{array}$ | $\begin{array}{r} -58 \\ +\quad 29 \\ -\quad 3.1 \end{array}$ | $\begin{array}{r} +6.3 \\ -3.2 \end{array}$ | $\begin{aligned} & 9 \cdot 1 \\ & 5 \cdot 8 \end{aligned}$ | 46 | 208 | A |
|  |  |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  | 151 | 53 | II |
|  |  |  |  |  |  | + 37 | $5 \cdot 7$ | 49 | 135 | A |
| 1903 |  |  |  |  |  |  |  |  |  |  |
| April 5, 11.25 p.m. to 11.29 p.m. | +216 | $-6 \cdot 1$ | -2 2 | + 77 | -12.4 | $+17 \cdot 2$ | $22 \cdot 6$ | 40 | 148 | A |
| June 19, 7.38 a.m. \# 7.55 a.m. | + $2 \cdot 8$ | -41.9 | $-6 \cdot 4$ | + 41.2 | + 78 | +65 | $42 \cdot 4$ | 81 | 79 | C |
| " 19, 7.55 , \% 8.13 " | +22 1 | + $37 \cdot 4$ | $+6^{\circ} 0$ | -34 8 | -24.0 | $+119$ | $43 \cdot 9$ | 74 | 235 | A |
| " 28, 8.36 " „ 8.47 " | $-3 \cdot 4$ | -36.8 | -4.5 | + 358 | +10'1 | + 0.7 | $37-2$ | 89 | 74 | C |
| " 28, 8.47 " | $+16 \cdot$ | $+24 \cdot 1$ | +4.5 | $-22 \cdot 1$ | $-17 \cdot 5$ | + 8.8 | $29 \cdot 5$ | 73 | 232 | A |
| " 29, 8.59 , $\quad 9.26$ \% | $+1 \cdot 1$ | - 3.8 | $0{ }^{\circ} 0$ | $+3 \cdot 9$ | $-0.3$ | + 0.8 | $4{ }^{\circ} 0$ | 78 | 94 | C |
| , 29, 9.26 ", 9.53 " | -16*8 | $-16.6$ | $0 \cdot 0$ | $+15 \cdot 1$ | $+13 \cdot 5$ | $-12 \cdot 2$ | $23 \cdot 6$ | 121 | 48 | B |
| July 26, 6.23 " " 6.55 , | $-26.3$ | $-31 \cdot 1$ | $-3 \cdot 2$ | + 28.4 | $+24 \cdot 1$ | $-16.9$ | $40 \cdot 9$ | 114 | 50 | B |
| " 26, 6.55 ", 7.30 , | $+2 \cdot 1$ | +33 0 | +8:3 | -31.8 | -11.5 | - 4 "2 | $34 \cdot 1$ | 97 | 250 | A |
| August 25, 10.57 p.m. ${ }^{2} 11.0$ p.m. | + 10 | $+1 \%$ | $0 \cdot 0$ | $-16$ | $-0.9$ | + $0 \%$ | 2.0 | 69 | 241 | A |
| " 25, 11.0 , 11.3 , | + $0 \cdot 7$ | $-7 \cdot 4$ | $-1.9$ | + $7 \cdot 2$ | + 18 | + 18 | 7 '6 | 76 | 76 | C |

Table LXVI.-Disturbance Components and Resultants at Winter Quarters. (Unit of force 1\%)

Table LXVII.-Disturbance Components and Resultants, July 24, 1902. (Unit of Force 1\%)



Table LXVIII.-Ratio of Disturbances to those at Kew.


## CHAPTER X.

## Special Type of Disturbance.

§102. Reference has already been made, $\S 69$, to a "special type" of disturbance in the Antarctic, examples of which are afforded by the curves of June 19, 28 and 29, July 26, and August 17, 1903 (see Plates XXVII-XXXI). The close resemblance of the disturbances of June 28 and 29 first drew my attention, and, in order to judge of the real significance of the apparent repetition of a disturbance at a nearly 24 -hour interval, it appeared necessary to examine the curves in detail. Examination soon showed that the type of disturbance was of somewhat frequent occurrence during certain hours of the day.

The essential part of the phenomenon may be regarded as consisting of two phases. During the first there is normally a slight fall in V , during the second there is always a rise, usually considerably larger than the preceding fall.

During the first phase D practically always decreases, and H usually, though not always, does the same. During the second phase the D and H changes are in the opposite direction to those during the first phase

On the crests, so to speak, of the D and $H$ waves minor oscillations usually occur (cf. Plate XXIX), also the slope of the V curve near its lowest point is sometimes very slight. Thus the time when the first phase ends and the second begins is usually uncertain to a minute or two, and a greater uncertainty often attaches to the beginning of the first phase and the ending of the second. The main movement was usually accompanied by the minor variations which were of such persistent occurrence in the Antarctic. But these minor movements and the various uncertainties which enter into individual cases will, it is believed, have but little influence on the general conclusions reached below.

What happened after the end of phase 2 varied a good deal. June 29, 1903, represents the most usual order of events. Here V remains nearly uniform for some time. Not infrequently, however, after reaching a maximum $V$ made some further oscillations, and then diminished somewhat rapidly, as in the curve of June 28, 1903 (Plate XXVIII).

In some cases the to-and-fro movements in $D$ and $H$ were closely alike; in other cases there was a good deal of asymmetry. In some cases the return wave, so to speak, was checked by what seemed to be a second wave surging in before normal conditions had been restored. In a few instances there were two complete examples of the phenomenon in a single day, and once or twice in immediate sequence to one another.

As will be seen presently, the great majority of the occurrences took place between 6 and $10 \mathrm{p} . \mathrm{m}$. (local time), no occurrences being noted from $2 \mathrm{a} . \mathrm{m}$. to $1 \mathrm{p} . \mathrm{m}$. There is, however, more than one possible explanation of this fact. At most seasons of the year the type of the disturbed movements varied throughout the 24 hours. During the afternoon, up to 10 or 11 p.m., the curves exhibited a marked tendency to rounded swelling movements, like a succession of irregular waves with a considerable interval between the crests. There were incessant minor oscillations superposed on these, but they did not usually obliterate the long-period waves. In the early forenoon, on the other hand, slow-swelling movements were rarely recognisable, while short-period oscillations became dominant, the curves appearing tumultuously irregular and highly disturbed.

The hour marks in the $\mathrm{D}, \mathrm{H}$ and V curves were seldom exactly in a line transverse to the hour lines, and in the case of these short-period oscillations it was exceedingly difficult to decide whether crests on the different curves answered exactly to one another. It is by no means improbable that microscopic examination of the curves would have shown a number of cases in the morning where phenomena of the type now being discussed occurred, only with the phases much shortened. In many morning disturbances it was clear that a turning-point answering to the minimum value of $D$ during a rapid to-and-fro movement answered, at least very approximately, to the lowest point of a sharp V depression. There were, however, usually a succession of crests and depressions on the $V$ traces at such times, the successive waves tending to become confused,
§103. Table LXIX gives particulars of the 82 cases of the phenomenon actually found. It gives the date, the approximate hour answering to the end of the first phase and the beginning of the second, the duration of each phase, and the results of the curve measurements. A double entry, e.g. -5 , in the D column means that the turning-point on the D curve occurred distinctly after the time taken as the ending of the first phase, the fall experienced during the first phase continuing for a short time, and being then followed by a larger rise. The addition of $a+$ after a figure means that the trace went beyond the limit of registration to which the figure given answers. In such a case the amplitude assigned is, of course, an underestimate.

Table LXIX. -Special Type of Disturbance. Occurrences.


Table LXIX (continued).

§104. Table LXX summarises the results from the whole 82 occurrences as to the hour at which the first phase ended. An occurrence at an exact hour was assigned to the hour then commencing. It will be seen that more than half the occurrences happened between 7 and 9 p.m. There is no marked seasonal variation in the hour of occurrence. A larger proportion of occurrences took place after $8 \mathrm{p} . \mathrm{m}$. in April, May and June than in July, August and September, but that might be accidental. The great concentration of occurrences near 8 p.m. is obviously a vital point when considering the real significance of the sequence presented in the disturbances of June 28 and 29, 1903.

Table LXX.-Hour of Occurrence of End of First Phase.

§105. Table LXXI gives particulars of the average duration of the two phases, again from the whole 82 occurrences. There seems a distinct tendency for the second phase to last the longer, though instances in which the reverse is true are not rare. There seems no marked seasonal variation in the duration of the phases. There is an apparent difference between 1902 and 1903 which may be real. 1902 was a year of very small sun-spot frequency (Wolfer's relative value 5.0 ), while 1903 had a considerably larger frequency (Wolfer's value $24^{\circ 4}$ ) and showed larger magnetic ranges both for the regular and irregular variations.

Table LXXI.-Duration of Phases (Minutes).

|  | First phase. |  |  | Second phase. |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 1902. | 1903. | 1902 and 1903. | 1902. | 1903. | 1902 and 1903. |
| April and May | $13 \cdot 5$ | $15 * 6$ | $15 \cdot 2$ | $22 \cdot 0$ | $19^{\circ} 0$ | $19 \cdot 5$ |
| June i . . . . | $18 \cdot 0$ | $16 \cdot 3$ | $17 \cdot 1$ | $21 \cdot 6$ | $18 \cdot 6$ | $19 \cdot 9$ |
| July . . . . . . | $18 \cdot 1$ | $13 \cdot 3$ | $15 \cdot 4$ | $22 \cdot 1$ | $17 \cdot 5$ | $19 \cdot 6$ |
| August . . . . | $17 \cdot 9$ | $20 \cdot 0$ | $18 \cdot 3$ | $20 \cdot 2$ | $20 \cdot 0$ | $20 \cdot 1$ |
| September . . . | 22.5 | 21.0 | 21.7 | 25.0 | 18.0 | 21.5 |
| Other month. . | 18.8 | $14 \cdot 6$ | $16 \cdot 9$ | $26 \cdot 3$ | 21.8 | $24 \cdot 3$ |
| Means . . . | $18 \cdot 1$ | $15 \% 8$ | $17 \cdot 0$ | 21.9 | $18 \cdot 9$ | $20 \cdot 5$ |

§106. An attempt was made to arrive at a more exact idea of the type of disturbance by forming means based on a number of days. Only those occurrences were used for which the trace was complete. Incompleteness was naturally most common in the larger disturbances, so that the mean results obtained somewhat underestimate the average amplitude. Occurrences in which the phases in the different elements did not synchronise, or which were ill defined, were also omitted. The 51 occurrences which remained were grouped according to the season of the year as shown in Table LXXII, which gives the mean results obtained. The Declination results are expressed not in minutes of are, but in units of force, replacing $1^{\prime}$ by $1.92 \gamma$ as the force acting perpendicular to the magnetic Meridian necessary to alter the direction of the Declination needle by one minute of arc.

Table LXXII.-Amplitude and Direction of Disturbing Force.


To explain Table LXXII let us consider the mean result from all the observations, which may be assigned to a single representative occurrence of the phenomenon.

In the figure, NS, EW are the geographical north-south and east-west directions, while nos represents the position of the Declination needle, $n$ being the north pole. Taking
 as point of departure the position existing at the commencement of the first phase, the force required to produce the disturbance existing at the end of that phase called for the action on the pole $n$ of a force whose horizontal components were $60 \cdot 9 \gamma$ perpendicular to no, and $15 \cdot 3 \gamma$ along no, in the directions shown, the vertical component being $6 \cdot 1 \gamma$ downwards. The resultant of the two horizontal components amounted to $62.8 \gamma$, and its inclination $\psi$ to no was $75^{\circ} 53^{\circ}$. The total force $\Delta \mathrm{T}$, obtained by combining the horizontal resultant with the Vertical Force $6 \cdot 1 \gamma$, amounted to $63 \cdot 1 \gamma$, and its inclination $\chi$ to the horizontal plane was $5^{\circ} 35^{\prime}$. The direction of $\Delta T$ pointed below the surface of the ground.
For the second phase we take the position at the end of the first phase as point of departure. To produce the disturbance existing at the end of the second phase required the forces shown in the second half of the table. The accompanying figure illustrates the results in the horizontal plane answering to the mean or representative occurrence. The resultant of the Horizontal Forces, $55^{\circ} 2 \gamma$, is inclined at $70^{\circ} 48^{\prime}$ to on produced. It is thus not exactly opposite to the corresponding force experienced during the first phase, the two being inclined at approximately $175^{\circ}$. The Vertical Force in the second phase is nearly $3 \frac{1}{2}$ times that in the first, so that the angle $\chi$, which the total disturbing force $\Delta \mathrm{T}$ now makes with the horizon, is increased to $20^{\circ} 47^{\prime}$ and points above the horizon.

There were only four occurrences in September, thus little significance attaches to the large size of the average disturbance
 for that month.
§107. The data on which Table LXXII depends were all got out before any anticipation was made as to
the probable result, or any plan existed for combining the observations. Thus the remarkable similarity in the values obtained for $\psi$ and for $\chi$ from the several months owes nothing to any preconceived ideas.

Whatever may be the cause of the phenomenon, it is clear that so far as the forces in the borizontal plane are concerned the second phase may be regarded, to a first approximation, as simply a relaxation of the forces to which the first phase is due. On the average, D ends by being about $4^{\prime}$ (or $8 \gamma$ ) smaller, and H by being about $3 \gamma$ larger than at the start. In the case of V , however, there is something more than a mere relaxation, the "recovery" during the second phase being, on the average, $3 \frac{1}{2}$ times the drop during the first phase.

It must not, of course, be forgotten that during the disturbance the ordinary diurnal changes may naturally be expected to go on as usual. The hour of occurrence being variable, it is difficult to allow very exactly for this. If, however, we take the mean dimrnal inequalities for Midwinter as the most nearly applicable, we find that during the average occurrence of the phenomenon the regular change would be practically nil in D , about $-2.6 \gamma$ in H , and about $+0.6 \gamma$ in V . The subtraction of the effects of these regular changes makes but little difference to the results, especially as regards the amplitudes. For the angles in the case of the representative disturbance the results in Table LXXII are altered to

|  |  |  | $\psi \cdot$ |
| :--- | :--- | :---: | :---: |
| First phase | $\cdot$ | $\cdot$ | $77^{\circ} 3$ |
| Second phase | $3^{\prime}$ | $5^{\circ} 51^{\prime}$ |  |
| $69^{\circ} 31^{\prime}$ | $20^{\circ} 23^{\prime}$ |  |  |

§108. As to the possible cause of the special type of disturbance, the results derived from the co-operating stations suggest that its seat is, mainly at least, in the southern hemisphere, but it is clearly not a purely local phenomenon. The disturbances experienced at Christchurch, on the occasions for which Christchurch data existed, averaged in amplitude about a fifth of those experienced at Winter Quarters. If the cause is electric currents, the absence at Winter Quarters of any large vertical component during the first phase suggests a nearly uniform current sheet overhead, or else underground currents having similar direction and intensity over a considerable area, or a combination of the two sets of currents.

If we take the value $75^{\circ} 53^{\prime}$ for $\psi$ during the first phase, and assume the local magnetic Meridian to be $152^{\circ} 40^{\prime} \mathrm{E}$., then the direction of the hypothetical currents would be

> if overhead, from N.W. to S.E. (more exactly, from $48^{\circ} 33^{\prime}$ to north of west),
> if underground, from S.E. to N.W. (more exactly, from $48^{\circ} 33^{\prime}$ to south of east).

Judging by the solitary observation on the ice* on January 30, 1904, the undisturbed magnetic Meridian was about $148^{\circ} \mathrm{E}$., and the inclination of the hypothetical currents to this is roughly $9 \frac{1}{2}^{\circ}$. The small vertical component seen during the first phase might be explained by supposing that the intensity of the currents varied slightly with the distance to N.E. or S.W. from Winter Quarters.

The fact that after the horizontal movements had subsided there remained an enhanced value of the Vertical Force might be explained by supposing that the current system or systems did not really die out, but moved, if overhead towards the S.W., if underground towards the N.E. Or the explanation might be that the currents had in reality circular paths and tended to magnetise the Earth, producing a S-pole in the neighbourhood of Winter Quarters, which usually showed marked hysteresis in its disappearance.

In our present state of knowledge it would be pretty much pure accident if one happened to hit on the true explanation. But there are certain conclusions which may usefully be drawn.

The existence of such a phenomenon as the special type of disturbance emphasises the importance of simultaneous observations not merely from ordinary observatories, but from stations much less remote from the Antarctic station. If such stations existed, and were within reach of one another by wireless telegraphy, and if observations on earth currents and on the transmissibility of wireless signals were included in the programme, there would be a reasonable chance of a satisfactory explanation of the phenomena being reached. If the magnetograms were studied as soon as available, and striking phenomens noticed as they occurred, then, if they repeated themselves, it might be possible to recognise

* Physical Observations," p. 141.
their occurrence while still incident, and thus to make physical observations likely to discriminate between the different theories proposed.

It would also be of interest to have comparative records from some Arctic station. All the sudden commencements of storms which were recorded at the co-operating stations appeared of enhanced magnitude at Winter Quarters. Unless we suppose them to be due to some source of disturbance peculiar to the southern hemisphere-which seems unlikely considering the size of the disturbances experienced at Kew and Colaba as compared to those at Mauritius and Christchurch-it would appear not unlikely that a similar enhancement would appear at an Arctic station. If this should prove to be the case it would be a most interesting fact, in view especially of the views recently advanced by Villard as to the nature of aurora.

## CHAPTER XI.

## Comparison of Magnetic Disturbances and Aurora.

§109. Aurora in England seems always accompanied by more or less magnetic disturbance. Auroras and large magnetic disturbances are both rare events in the south of England, so that when a conspicuous aurora and a large magnetic storm occur there simultaneously, the inference that the coincidence is not a mere accident is almost inevitable. In the Arctic and Antarctic regions, however, magnetic disturbance is the rule rather than the exception, and if the same is not equally true of aurora, still aurora is so often visible that even if no physical connection existed between the two phenomena, accidental coincidences would naturally occur. During several Arctic expeditions-e.g. those in the polar year 1882-elaborate auroral observations have been carried out, and special attention has been given to the question of the relationship to magnetic storms. The general conclusion reached seems to be that many Arctic aturoras are unaccompanied by any noteworthy magnetic disturbance, but that auroras of a specially vivid and rapidly changing character are usually accompanied by marked magnetic disturbance.

In considering the incidence of auroras there is the serious complication that to be visible an aurora must occur when there is no other source of light sufficiently bright to render it invisible. Thus aurora is seldom if ever seen until the Sun is below the horizon, and even moonlight, when the Moon is near the full, suffices to render any but specially bright auroras invisible. The state of the sky as to clouds is also important. Faint auroras of limited extent have a better chance of being seen when clouds are few than when they are many.

Again, there is the fact that while the record of magnetic storms, thanks to magnetographs, goes on equally well in the absence of trained observers, this is not the case with auroras. At Winter Quarters the Meteorological Observers were on the outlook for aurora at the two-hourly observation hours, right through the 24 hours. Specially bright auroras would also naturally attract the attention of the watch on deck, whose instructions under these circumstances were to call Mr. Bernacchi. Still, during the night, a display of less than two hours' duration, if faint or only moderately bright, might fail to be noted. Thus the fact that at an observation hour when aurora was noted the magnetic curves were quieter than half an hour before or after, when no aurora was noted, may not possess any real significance.
§110. The magnetic curves, when existent, were examined at all the times at which auroras appear on Mr. Bernacchi's list.*

Before considering the results of this comparison, the following particulars of the auroral statistics may be mentioned. $\dagger$ The auroras recorded, with two exceptions occurring late in March, were limited to the six months April to September, the days on which they were observed being distributed as follows :-

|  | March. | April. | May. | June. | July. | August. | September. | Total days. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1902. . | - | 10 | 8 | 12 | 10 | 9 | 3 | 52 |
| 1903. . . . | 2 | 18 | 14 | 18 | 22 | 14 | 2 | 90 |
| Total. . . | 2 | 28 | 22 | 30 | 32 | 23 | 5 | 142 |

Thus the auroras seen were practically limited to the four months April to August, and so to the season of the year when magnetic movements, both regular and irregular, were smallest.

* "Physical Observations," p. 101.
+Cf. "Physical Observations," p. 126.

The absence of auroral records from the remainder of the year means nothing more than "too much daylight," to use Mr. Bernacchi's words.

Even in the Midwinter months there are gaps in the auroral record: in 1902, from April 16 to May 5 , from May 15 to 30, from June 16 to 29, from July 14 to August 4; in 1903, from April 12 to 18, from May 8 to 15, from June 7 to 13, from July 7 to 11, and from July 31 to August 10, all inchusive. Of the third of these gaps Mr. Bernacchi says: "From June 15 to June 30 bright moonlight or overcast skies prevented any aurora being seen," and, presumably, the explanation of the other gaps is similar, as the dates of full Moon were, in 1902, April 23, May 23, June 21, July 20, and, in 1903, April 13, May 12, June 11, July 10, and August 8.
§111. It will be observed that many more auroras were seen in 1903 than in 1902. I am informed by Mr. Bernacchi that this is not due, at least in any large measure, to difference in the observational methods or increased activity in the observers. It is thus, presumably, a true physical phenomenon. 1902 was near sun-spot minimum, with a Wolfer's frequency of only $5 \cdot 0$, while the frequency for 1903 was $24^{\circ} 4$. In northern temperate latitudes auroral frequency normally increases with sun-spot frequency [it is, however, doubtful whether the same is true to the north of the zone of maximum auroral frequency], thus a difference between 1902 and 1903 is not surprising. The difference between the two years seems due partly to the greater length in 1902 of the intervals in which no auroras were observed. This rather suggests that the cause was difference in intensity rather than anything else. In 1903 the average aurora may have heen more intense than in 1902, and so have suffered less from the causes tending to render it invisible. Some collateral evidence of this is afforded by a consideration of the number of separate entries of aurora in Mr. Bernacchi's list.

On some days aurora is noted at one hour only, but on most days, when it is recorded at all, at several hours.

There is a certain amount of overlapping, so that there may be a trifling error in my estimate of the number of separate entries. The figures show, however, a very large difference between the two years, there being 250 separate entries for 1903 as against 125 for 1902 . The ratio $20: 10$ between these two numbers is substantially larger than the ratio $17: 10$ between the number of days of aurora in the two years. This obviously supports the view that in 1903 the average aurora retained for a longer time the intensity necessary to render it visible, from which a higher maximum intensity would naturally be inferred.

If this is a main cause of the excess of visible auroras in 1903, then the preceding figures are, at least, strongly suggestive of the view that the phenomena whose visible side is aurora were seldom, if ever, wholly absent on Midwinter days at Winter Quarters. Thus the fact that the magnets were practically never undisturbed for five minutes at a time cannot safely be interpreted as evidence of the continued presence of some cause of irregular magnetic disturbance other than that associated with the seat of aurora.

Another consideration should be borne in mind. Assuming, what few people now doubt to be true, that aurora is the visible manifestation of electrical action in the atmosphere, the existence of a very bright auroral band or streamer may mean an electrical current of unusually high intensity reckoned per unit of cross-section, but one having a comparatively small section. Thus the existence of visible aurora may mean only local concentration and no great total quantity of current. Thus the ultimate causes of aurora and magnetic disturbance may be the same, without any close parallelism being exhibited between the apparent intensities of the two phenomena.
§112. Two attempts were made to trace the possible interconnection of aurora and magnetic disturbance. The character of the magnetic curves was considered at all the times when aurora was noted, and a rough judgment passed. The magnetic trace was characterised as "quiet," "normal," "moderately disturbed," and "somewhat highly disturbed." These terms may be thus interpreted:-"Normal" means that I regarded the amount of disturbance as about average; "quiet" means that distinctly less than the average amount of disturbance existed; "moderately disturbed" means somewhat above the average amount of disturbance; and "somewhat highly disturbed" that the existence of more than usual disturbance obtruded itself even on casual inspection. The average amount of magnetic disturbance
varied with the season, and 1903 was decidedly more disturbed than 1902 , so that nothing like an exact standard of disturbance could well be maintained.

The results were as follows:-
Table LXXIII.


On a good many of the occasions when the curves were quiet the aurora is described as faint, or very faint, but this was not always the case. Thus, on May 31, 1903, a species of corona is described as visible at 4 p.m. (see Plate 13, "Physical Observations") having "intensity bright as it rose to the zenith." The magnetic curves were, however, unusually quiet at 4 p.m. and for some time afterwards. There was a minor disturbance about 3.30 p.m., but it had subsided before $4 \mathrm{p} . \mathrm{m}$.
§113. The second attempt referred to above consisted in examining individual magnetic curves more minutely, to see whether anything special happened at the precise times when aurora was noted.

A judgment was passed as to whether a correspondence existed. Cases which, so to speak, seemed worth sending to a jury, were adjudicated as to whether the correspondence were "possible," "probable," or "apparent." These terms may be interpreted as follows :-
"Possible" correspondence means that whilst the average man would probably decide against a true correspondence, he would experience more or less hesitation in doing so. In the case of "probable" correspondence the verdict would naturally be favourable, but again with hesitation. In the case of "apparent" correspondence there could be little doubt that a marked magnetic movement occurred during the observed aurora.

The decisions reached were as follows, the numbers denoting days on some of which more than one correspondence was considered :-

| Correspondence. |  |  |
| :---: | :---: | :---: |
| 7 | Probable. | Apparent. |
| 7 | 7 | 11 |

§114. The cases of apparent correspondence are included in the following list. Details of the magnetic disturbances are given in parallel with the description of the auroras as given by Mr. Bernacchi.

Directions are geographical unless the contrary is explicitly stated.
. Jene 30, 1902.

At $9.27 \mathrm{p} . \mathrm{m}$. faint aurora are from Observation Hill to Crater Hill (i.e. its centre a little south of the east). Altitude $15^{\circ}$ to $18^{\circ}$.

It (aurora) had completely disappeared at 9.32 p.m.

V rose to a maximum at about $9.30 \mathrm{p} . \mathrm{m}$. ; the increase during the previous 3 minutes was specially rapid, amounting to about $20 \gamma$. In the course of about 5 minutes D increased $18^{\prime}$ and diminished $27^{\prime}$, the turning-point answering to aloout 9.27 p.m. Synehronous with this was a sharp, but not large movement in H .

After 9.32 p.m. curves normal.

The evening of June 30 was fairly quiet magnetically, but there were at about $9.15 \mathrm{p} . \mathrm{m}$. (when there is no record of aurora) sharp oscillations in D and H , very similar in character, though opposite in direction, to those recorded 12 minutes later, and these were accompanied by a small but sharp temporary depression in V . The general trend of V was upwards from 9.0 to 9.30 p.m., the total increment being about $65 \gamma$.

September 19, 1902.

Midnight ( 0 a.m.). Faint aurora extending from N.E. to S.W.

There was a large oscillation on the H curve. A decrease of $50 \gamma$ was followed by a rise of $60 \gamma$. The turning-point was at about $0.5 \mathrm{a} . \mathrm{m}$., the double movement occupying about 50 minutes. D diminished about $30^{\prime}$ between $11.55 \mathrm{p} . \mathrm{m}$. on 18 th and $0.7 \mathrm{a} . \mathrm{m}$. on the 19 th . The V magnet was out of action.

The night of September 18-19 was on the whole rather quiet, but there was another oscillation in H very similar to and not much less than the above, with the turning-point at about $10.50 \mathrm{p} . \mathrm{m}$. on the 18 th. The observer remarks under the date September 19 that daylight was getting too bright for aurora to be visible, so that the above display was presumably more than usually intense.

May 28, 1903.
$10.45 \mathrm{p} . \mathrm{m}$. Aurora semi-are emanating from below hill in N.W. by W. magnetic (i.e. about $5^{\circ}$ north of true east) and terminating abruptly over Observation Hill, N.N.W. magnetic (i.e. about $45^{\circ}$ east of true south), where altitude was $20^{\circ}$. Light very faint and diffused.

The H curve shows a marked oscillation, a decrease of $45 \gamma$ being followed by a slightly larger increase. The turning-point occurred about 10.47 p.m., the whole movement occupying about 50 minutes.

There was no noteworthy variation in either $D$ or V , but the latter element showed a slight rise after $10.55 \mathrm{p} . \mathrm{m}$.

May 29, 1903.
3.0 p.m. Continued display of aurora in the S . and S.E., low on horizon, from $0^{\circ}$ to $3^{\circ}$ in altitude. Low arcs rising close upon one another, sometimes as many as parts of four or five, the southern extremities only being complete. High rays occasionally shot towards the zenith. The darkness below the arcs was marked. Movement chicfly from $E$. to $S$. in the rays, but from 3.0 to 4.30 p.m. the whole display had shifted from S . to E .

A magnetic disturbance of the special type described in Chapter $\mathbf{X}$ was in progress. Between 2.45 and 3.5 p.m. D diminished $20^{\prime}$, while H increased a little over $40 \gamma$, $V$ remaining practically constant ; then in the course of the next 18 minutes D increased $36^{\prime}, \mathrm{H}$ diminished about $30 \gamma$, and V increased $15 \gamma$.

By 3.40 p.m. the elements had returned to about the values they originally possessed and appeared normally quiet.

In this instance the auroral display seems to have continued after the special magnetic disturbance had ceasel.

An especially brilliant aurora suddenly appeared a few minutes after $4.0 \mathrm{p} . \mathrm{m}$., in the shape of a curtain, or segment of an arc, extending from W. $20^{\circ} \mathrm{N}$. to N.E. magnetic $\left(8^{\circ} \mathrm{N}\right.$. of E. to $17^{\circ} \mathrm{W}$. of S . true). There was more movement, both vertical and horizontal, than has yet been observed. The vertical movement of the whole display en masse was fairly rapid from S. towards the zenith, and the horizontal motion of the huge shafts of light at one time too rapid for the eye to follow ... Altitude at first was about $10^{\circ}$ at the extremities E . and W ., and $20^{\circ}$ in the centre, but this gradually rose to $50^{\circ}$ and $60^{\circ}$ in the centre. The brightest display was at about 4.10 to 4.15 p.m. ; had almost entirely disappeared at 4.25 p.m. . . . During this special display a bright auroral glow showed up above the hills almost at right angles to the curtain. The display originated quite suddenly in the direction of Mount Discovery (i.e. nearly S.W.) and flashed across the sky towards Observation Hill (nearly S.E.) in a few seconds.

There was a prominent movement in V , consisting of an increase of about $240 \gamma$, commencing about 3.30 and culminating about 4.1 p p.m. During the next 40 minutes there was a decratm: of about $170 \gamma$. The increase in V was most rapid from 4.5 to $4.15 \mathrm{p} . \mathrm{m}$. (i.\% when the aurora was brightest), the rise during the 10 minutes being about $125 \gamma$.

The most notable change in I) consisted of tron very rapid oscillations between 4.0 and 4.20 p.m. During the first, which occupied about 9 minntes, D increased $60^{\prime}$ and returned to its original value. During the second oscillation, which followed immediately on the first, D increased $118^{\prime}$ and then diminished fully $130^{\prime}$. The H trace was off the sheet, in the direction of force increasing, most of the time from 3.30 to $5.0 \mathrm{p} . \mathrm{m}$. But some very rapid oscillations came on the shect about 4.5 to $4.15 \mathrm{p} . \mathrm{m}$.
4.0 p.m. Auroral display in E.

July 21, 1903.
A disturbance of the special type (though not very typical) was in progress. From 3.32 to 3.44 p.m. V fell about $5 \gamma$ and then rose about $87 \gamma$ in the course of the next 31 minutes. Between 4.15 and 4.55 p.m. V fell about $50 \gamma$. Between 3.32 and $4.15 \mathrm{p} . \mathrm{m}$. the D magnet executed a to-and-fro movement, first diminishing nearly $60^{\prime}$, and then returning to about its original value. The absolutely lowest value occurred about 3.57 p.m. H showed an increase of about $40 \gamma$, followed by a decrease; the maximum, which was off the sheet, occurred apparently about $4.10 \mathrm{p} . \mathrm{m}$. In the case of both D and H there were minor oscillations, which somewhat obscured the phenomena near the turning-points.
July 27, 1903.

0 a.m. Aurora just above hills from S . to S.W. magnetic (approx. N.N.W. to N.N.E. true), altitude about $5^{\circ}$.

V, which had diminished about $80 \gamma$ since $10.20 \mathrm{p} . \mathrm{m}$. on the 26 th, commenced to rise about 11.50 p.m., and during the next 45 minutes increased about $180 \%$. It then diminished rapidly for a few minutes, and then more slowly until nearly $2.0 \mathrm{a} . \mathrm{m}$. on the 27 th . H went beyond the limit of registration, in the direction of force diminishing, about $11.50 \mathrm{p} . \mathrm{m}$. on the 26 th , and continued so until $0.40 \mathrm{a} . \mathrm{m}$. on the 27 th . D went through a number of irregular oscillations, the extreme range between $11.45 \mathrm{p} . \mathrm{m}$. on the 26 th and $0.35 \mathrm{a} . \mathrm{m}$. on the 27 th exceeding $60^{\circ}$.

July 27, 1903 (continued).
$2.0 \mathrm{~g} . \mathrm{m}$. Isolated patches of diffused aurora from N. to S.E., altitude $10^{\circ}$ to $30^{\circ}$.
4.0 a.m. Streamers or rays from N. to S.E. ; various heights, mean $40^{\circ}$.

The aurora was more or less visible all night, and confined principally to N.E.; average altitude $20^{\circ}$.

At about $9.45 \mathrm{p} . \mathrm{m}$. an unusual form of aurora appeared. A band of light extended from due S . to due N, passing round through E. Breadth of band $5^{\circ}$ and averaged $12^{\circ}$ in altitude. Intensity fairly strong in N. . . . The display reached its maximum brilliancy at about 9.50 p.m., and had almost entirely disappeared at 10.10 p.m. A few arrow-like beams were visible here and there just above the band.

V , which had been falling rather markedly since $0.35 \mathrm{a} . \mathrm{m}$., rose about $20 \gamma$ between 1.55 and 2.5 a.m., and then diminished. There were approximately equal to-and-fro movements in both .
D and H , the turning-points occurring at about $1.55 \mathrm{a} . \mathrm{m}$. D first diminished about $30^{\prime}$ and then increased, while H diminished about $20 \gamma$ and then increased.

From about 3.35 to $4.25 \mathrm{a} . \mathrm{m}$. there was a deep bay on the H curve; the turning-point, which answered to a minimum in H , occurred within a few minutes of $4.0 \mathrm{a} . \mathrm{m}$., but was beyond the limits of registration.

The D curve was not specially disturbed at 4.0 a.m., but there were minor oscillations on the V curve.

Between 9.20 and $9.30 \mathrm{p} . \mathrm{m}$. V increased about $25 \gamma$, and then remained nearly stationary until about 9.42 p.m., when a further rapid rise ensued.

Between 9.42 and 9.57 p.m. the increase was about $80 \gamma$. During the next 5 minutes $V$ fell a little, and then was nearly steady until after 10.20 p.m.

There were sharp oscillations in D and H between 9.45 and $10.0 \mathrm{p} . \mathrm{m}$., but of no great magnitude.

July 28, 1903.
2.0 a.m. Faint aurora diffused over the E., forming a narrow curtain and a few streamers scattered irregularly.
3.0 a.m. Fine display of aurora, involving the whole heavens from N.W. by E. to S. Nothing ever visible in S.W. Three fine compact curtains in the E., one above the other, height of uppermost approximately $60^{\circ}$. Three more curtains, more diffused, but bright and much folded, extended from N.W. to zenith, where there were two large bright luminous clouds. The rest of the area was filled with more or less isolated streamers, small or fragmentary curtains or clouds. All was constantly changing ...
4.0 a.m. Only the remains of the above display visible ...

About 2.0 a.m. there were small oscillations on the $V$ trace. There was a deep bay on the $D$ curve, the turning-point, which answered to a maximum, occurring about $1.45 \mathrm{a} . \mathrm{m}$. Between that hour and 2.20 a.m. D diminished about $70^{\circ}$. There was a considerable movement in the $H$ trace, which remained beyond the limit of registration, in the direction of H diminishing, from 1.40 to $2.15 \mathrm{a} . \mathrm{m}$.

Between 2.55 and $3.20 \mathrm{a} . \mathrm{m}$. there were marked oscillations in $V$, there being on the whole a decrease of about $30 \gamma$. Between 3.20 and $3.38 \mathrm{a} . \mathrm{m}$. there was a very rapid increase, amounting to about $110 \gamma$. The H trace went beyond the limit of registration in the direction of $\mathbf{H}$ diminishing about 3.15 a.m., but only for a few minutes. The change preceding this was very rapid.

August 13, 1903.
10.0 p.m. Fairly brilliant display, consisting of a complete arc, extending from N . to S ., and two streamers. The highest point of the are was due E., with an altitude of $15^{\circ}$. At its E.S.E. point it was distorted by a relatively more brilliant and wider zone of light with a streamer rising out of it to $30^{\circ}$ altitude. An independent ray also rose to $30^{\circ}$ altitude to the S.E., but did not quite reach to the arc. The breadth of the arc was between $2^{\circ}$ and $3^{\circ}$, the lower edge more defined than the upper, but neither particularly definite. Very rapid movement and very rapidly changing in form.

By 10.10 p.m. the are had completely disappeared and was replaced by streamers of irregular altitude and interrupted in their lengths. The streamers rose at various points where the are had been, the extremes being at E.N.E. and S. by E. points, with two more in between. The altitude of the highest was $40^{\circ}$. This latter display was also very rapidly changing.
$10.20 \mathrm{p} . \mathrm{m}$. Diffused streamer from due N :, spreading out fan-like to about $30^{\circ}$ altitude, but one thin band from one side of the fan extending across the sky to the W.
$10.45 \mathrm{p} . \mathrm{m}$. No aurora visible.
At $11.10 \mathrm{p} . \mathrm{m}$. fine arc in S. , extending from N.W. magnetic to N.E. magnetic $\left(17^{\circ} \mathrm{S}\right.$. of E. to $17^{\circ} \mathrm{W}$. of S.), altitude of apex $25^{\circ}$, and exactly in magnetic Meridian.

The N.E. (magnetic) extremity much the brightest and formed of vertical rays, while N.W. (magnetic) and centre were rather faint and about $4^{\circ}$ in width. The whole display moved rapidly towards zenith and at the N.E. (magnetic) formed draped aurora of a greenish tint. . . . .

A few isolated rays in N., altitude $40^{\circ}$. Shortly after a bright draped curtain appeared a little to E. of zenith, altitude $80^{\circ}$, and arc became very faint.

At 11.20 p.m. only a few faint cloud-like patches here and there were visible.

After $9.40 \mathrm{p} . \mathrm{m}$. numerous oscillations appeared in the V trace. A little before $10.0 \mathrm{p} . \mathrm{m} . \mathrm{V}$ began to increase rapidly. The maximum was reached about $10.18 \mathrm{p} . \mathrm{m}$., the rise being most rapid during the last 6 minutes. The total increase since $9.30 \mathrm{p} . \mathrm{m}$. was about $70 \gamma$.

After 10.18 p.m. V remained nearly constant during 20 minutes.

There was a deep bay on the D curve from about 9.50 to $10.40 \mathrm{p} . \mathrm{m}$. A decrease of about $70^{\prime}$ was followed by an equal increase, the turningpoint answering to about $10.12 \mathrm{p} . \mathrm{m}$.

The H curve showed a somewhat similar bay, but it appeared somewhat earlier in time.

The H trace was, however, beyond the lower limit of registration from about 9.45 to $10.25 \mathrm{p} . \mathrm{m}$., so that the turning-point was not shown.

Between 11.0 and $11.30 \mathrm{p} . \mathrm{m}$. the V trace showed a sharp double oscillation, the amplitude in each being about $25 \gamma$. The sharpest turning-point, which answered to a maximum of $V$, occurred at about 11.14 p.m.

The adjacent minima, which were also clearly shown, were at about 11.9 and $11.25 \mathrm{p} . \mathrm{m}$. respectively.

There were sharp peaks on the $D$ and $H$ curves at about $11.10 \mathrm{p} . \mathrm{m}$., D increasing $27^{\prime}$ between 11.10 and 11.13 p.m. Another very rapid movement occurred between 11.16 and 11.25 p.m., when D fell about $57^{\prime}$.

It will be observed that both after $10.0 \mathrm{p} . \mathrm{m}$. and after $11.0 \mathrm{p} . \mathrm{m}$., whilst there were notable magnetic changes synchronous with the brightest phases of the aurora, these were not more notable than the magnetic changes recorded after the auroral display had become faint.

August 14, 1903.
$2.0 \mathrm{a} . \mathrm{m}$. Very faint, lut extensive, aurora. Patches scattered about asymmetrically from N. tos.s. W.

Between 1.57 and 2.12 a.m. V first fell $15 \gamma$ and then rose $25 \%$. This was accompanied by a sharp oscillation in the H curve, consisting of a fall of $38 \gamma$, followed by a rise of $30 \gamma$. There was also a to-and-fro movement of fully $15^{\prime}$ in D .

Subsequent to $1.30 \mathrm{a} . \mathrm{m}$. there was a good deal of oscillation in all the curves during the whole forenoon, and some of the D and H changes were decidedly more striking than those specified above.

August 26, 1903.

From 7.0 to $7.50 \mathrm{p} . \mathrm{m}$. brilliant aurora was onserved. Started with rays showing up above the hills from N. (magnetic) all the way round to S. (magnetic) ( $28^{\circ} \mathrm{E}$. of S . to $28^{\circ} \mathrm{W}$. of N.). Some of these rays were exceptionally long, extending, in some cases, to an arc of $50^{\circ}$ vertically. The display seemed to have no special form. All manner of sinhous evanescent streamers, ares, \&c., were observed.

At about $7.35 \mathrm{p} . \mathrm{m}$. one streamer, or ray, about $1^{\circ}$ in width, extended vertically above Observation Hill (a little to E. of S.E.) to about $83^{\circ}$ in altitude. This is the longest ray we have olserved. At 7.40 p.m. a winding streamer, or curtain, appeared in the "Gap" (or about E.S.E.) and extended to about $45^{\circ}$ in altitude. This was the most brilliant part of the display and was about equal to a star of the 2nd magnitude.

At 7.50 p.m. the display had almost dispersed, lut remained faint and very diffused, like a kind of light luminous mist for some time after.

There was a very large magnetic disturbance commencing about $6.46 \mathrm{p} . \mathrm{m}$. To all appearance the first part was of the special type (Chapter X), but as the V magnet was out of action, this is not absolutely certain.

Between 6.46 and 7.35 p.m. the $D$ and $H$ magnets each executed a large to-and-fro movement, on which were superposed numerous smaller and very rapid oscillations.

Both curves got beyond the limit of registration in the direction of element diminishing, so that the time of the turning-point cannot be fixed exactly, but it was within a few minutes of $7.15 \mathrm{p} . \mathrm{m}$. The to-and-fro movements in D exceeded $105^{\prime}$ and $120^{\prime}$ respectively. The to-and-fro movements in $H$ were equal and fully $100 \%$. The rise to the maximum at $7.35 \mathrm{p} . \mathrm{m}$. was very rapid, and after the maximum there were extremely rapid movements in the opposite direction, D falling $54^{\prime}$ and H falling $45 \gamma$ in about $1 \frac{1}{2}$ minutes of time.
Another way of putting the facts is that in the course of about 3 minutes-synchronous apparently with the existence of the very long auroral streamer--D rose and fell 54.

After $7.37 \mathrm{p} . \mathrm{m}$. there were oscillatory movements in D, but of a much less striking character: The H trace was highly oscillatory from 7.40 to 8.0 p.m.

## APPENDIX A.

Abstract of "Term Day" Observations at Christchurch Ohservatory, New Zealand, made and tahulated by' Dr. C. Coleridee Farr and Mr. H. F. Skey, and their discussion by Dr. C. Chree, F.R.S.

§1. The "Discovery," on her way to the Antarctic, called at New Zealand, and magnetic observations were made at the Observatory at Christchurch. Dr. C. Coleridge Farr, who was then Director of the Observatory, being anxious to utilise to the utmost the opportunities presented, arranged with Captain Scott and Mr. Bernacchi for an extension of the programme of simultaneous magnetic observations laid down before the Expedition left England. The original programme specified the 1st and the 15 th of each month as "term days," during which hourly readings should be taken of the magnetic elements at all co-operating stations. The programme further arranged that on each term day there should be a "term hour," during which the values of the magnetic elements should be determined at 20 -second intervals.

On the first regular term day-February 1, 1902-the term hour was to he $0-1$ a.m., G.M.T. On each successive term day the term hour was to be one hour later, so that on Jamuary 15, 1903, the twentyfourth regular term day, the term hour arranged was $11-12 \mathrm{p} . \mathrm{m}$. , G.M.T. This completed the regular year, but four additional term days were proposed, viz. February 1 and 15 and March 1 and 15, 1903; the term hours arranged for these were respectively $0-1,1-2,2-3$, and $3-4$ a.m., G.M.T.

At observatories provided with self-recording instruments the readings were to be derived from the curves, and, to admit of their being read at 20 -second intervals, it was intended that the drums carrying the photographic paper should be rotated more rapidly than usual, twelve times the usual rate being the speed commonly adopted. The extension of the programme thus arranged at Christchurch contemplated that the magnetograph drums should be "quick run" during the whole term day. By altering the position of the light after each revolution of the drum, six or eight hours' run were usually obtained on a single sheet, but the mere alteration of the light entailed the presence of an observer, and several times, in the absence of a second observer, Dr. Farr had to remain on duty during the whole 24 hours. Notwithstanding the difficulties encountered, the revised programme was actually carried out at Christchurch, quick runs being taken on 26 terms days from March 1, 1902, to March 15, 1903, inclusive, Some little trace was necessarily lost when the light was being moved and when fresh paper was being put on, and occasionally an hour or two's record is lacking, e.g. from 10 to 12 p.m., G.M.T., on September $15,1902$. Declination trace was lost for one whole day, November 15, 1902. Everything considered, however, the loss of trace is remarkably small.

The extended scheme proved impracticable in the Antarctic. Still, quick runs were taken during several hours of most term days. This was so far fortunate as a mistake had somehow crept into the Antarctic 'Manual,' which made each term hour 12 hours later than it should actually have been, and the observer, Mr. Bernacchi-in the absence of any information to the contrary-naturally supposed the 'Manual' to be correct. Thanks to the extended programme, there were three term days on which the quick run made in the Antarctio covered at least a part of the real term hour, though on one of these occasions the part thus covered is but short.
§2. The original programme-which originated in Germany, but was approved by the Advisory Committee of the British Expedition-specified certain forms on which term-hour observations should be entered. Thus, taking the case of the Horizontal Force, the actual curve readings were to ocenpy one column, the converted values a second, the temperature corrections a third, the last column giving the finally corrected values of the ordinates in absolute measure C.G.S. The absolute value answering to zero ordinate, with particulars of the scale values and of the formula for the temperature corrections, were to be given on a separate page. The form for each term hour provided for 180 entries under each of 10 separate columns. Dr. C. C. Farr and his successor at Christchurch, Mr. H. F. Skey, applied the
original scheme to the whole of the quick runs, with this exception that the curves were read at l-minute intervals only during a considerable part of the term day, readings at 20 -second intervals being limited to part of the day, including always the true term hour.

After the return of the German and British Expeditions the German authorities suggested that the curves should all be copied on squared paper and published in this form, so that anyone could read off for himself the absolute value answering to any specified instant of time. Photographic copies were thus taken at Christchurch of all the quick-run curves.

The Christchurch material for term days, due to the combined efforts of Dr. Farr and Mr. Skey, was thus of a most comprehensive and voluminous character.

In advising as to what should be published I was guided largely by the following considerations:It was obvious that to publish the tables in full and to reproduce all the curves would occupy an amount of space and entail an expenditure which it was unreasonable to expect the Royal Society to sanction unless the results were likely to prove of extreme value. Now, it so happened that the term hours were unusually badly adapted for securing the objects ordinarily aimed at in simultaneous observations at short intervals of time.

The scheme was presumably due in considerable measure to the success attending a similar scheme proposed some years earlier by the late Dr. Eschenhagen, of Potsdam Observatory. This earlier scheme applied, however, to only two or three hours, and, on one of these, as it happened, there was a well-marked magnetic disturbance. The simultaneous observations at the co-operating stations enabled this disturbance to be followed in minute detail by Dr. A. Schmidt, who succeeded Dr. Eschenhagen at Potsdam. But of the term days observed during 1902-3 at Christchurch there was not one which presented any noteworthy disturbance. This will be readily seen on consulting the tables of hourly values published in "Physical Observations," pp. 160-179. Referring to the Declination results given there for Christchurch, on p. 177, it will be seen that the daily range never exceeded $13^{\prime} \cdot 1$, and on eight days it was actually under $5^{\prime}$.

For the study of a disturbance of moderate size, whose phases develop smartly, quiek-run curves have some marked advantages, but when magnetic changes are small and develop slowly, the gradient in quick-run curves is so slight that they look uncommonly like straight lines, and really disclose less to the eye than ordinary slow-run curves. This will be readily realised after inspection of the curves in Plate III, which are fairly representative of term-hour conditions at Christchurch.

Taking these facts into consideration, I decided that the publication of the tabulations made at Christ-church-tabulations representing an immense amount of labour on the part of Dr. Farr and Mr. Skeywas likely to be more useful than reproduction of the curves.
§3. Coming now to the tabulations, it was obvious that what we may call the "raw material" in the international forms-consisting of the uncorrected curve readings and the corrections-might be omitted without seriously impairing the value of the results. The raw material is necessary if one's object is to check the accuracy of the results, but it is usual to assume such checking to have been adequately performed by the observers. Another abbreviation was obviously possible, which had indeed been to some extent adopted in the sheets as received from Christchurch.

The magnetic changes were usually so slow and regular that for an element to change appreciably occupied several minutes, sometimes even hours. Thus in one outstanding case the Declination showed no change as large as $0^{\prime} \cdot 1$, the unit adopted, between 13 h .41 m . and 16 h .3 m . In this instance the international form provided for the repetition of the same figure-even on the basis of 1 -minute observa-tions- 142 times. Such repetitions are avoided by the method adopted here. Put briefly, the method consists in recording the value of an element only at the time when a change took place. Supposing, for instance, the Declination to remain at $16^{\circ} 12^{\prime} \cdot 0$ from 1 h .30 m . to 1 h .39 m ., but to be $16^{\circ} 12^{\prime} \cdot 1$ at 1 h .40 m ., it is sufficient to record the values at 1 h .30 m . and 1 h .40 m ., it being laid down that the absence of an entry signifies no change. Even as thus reduced, the material appeared to require further reduction. Accordingly, selecting a few representative term days, I investigated from which of the three elements Declination, Horizontal Force, and Vertical Force-results of most value were likely to be derived. was at once obvious that the Vertical-Force data promised to be the least valuable.

Referring to the hourly values of this clement given on p. 179 of the "Physical Ohservations," it will be noticed that a considerable number of discontinuities presented themselves, while the daily ranges were exceedingly small. The former phenomenon throws some doubt on the smooth working of the instrument, and the latter indicates that, even if the data were as trustworthy as those for the other two elements, they would form a less promising field of investigation.

The choice between Declination and Horizontal Force was more difficult. The examination, however, of the results from the few representative days pointed to the conclusion-fully confirmed by subsequent investigation-that the diurnal variation of Declination is decidedly more regular at Christchurch than that of Horizontal Force, and, accordingly, that it is the latter element whose changes promise the greatest return to minute examination. Accordingly, I recommended that publication of complete details of the variation throughout the whole duration of term days should be confined to Horizontal Force (pp. 204-228), and that in the case of the Declination (pp. 229-230) full details should be given only for the term hours. The Vertical-Force changes are not individually shown at all. Though details of the variations of Declination and Vertical Force are published only partially, or not at all, these variations were carefully studied and tabulated and a number of conclusions have been deduced.

Changes of Christchurch Horizontal Force. March 1, 1902.


Changes of Christchurch Horizontal Force. March 15, 1902.


Changes of Christchurch Horizontal Force. April 1, 1902.


Changes of Christchurch Horizontal Force. April 15, 1902.

| Value of H. | $\begin{aligned} & \text { Begins } \\ & \text { at } \end{aligned}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | Begins $a t-$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | h. m. |  |  |  |  |  |  |  | h. m. |
| -22684 | h. 0 | -22694 | h. 48 | '22696 | \%. 8. | -22696 | ${ }_{13} 59$ | -22698 | 20. |
| 84 | 8 | 92 | 57 | 96 | 1 | 97 | 140 | 97 | 18 |
|  | - | 91 | 58 |  | - | 98 | 21 | 96 | 37 |
| 84 | 22 | 90 | 459 | 96 | 3 | 98 | 31 | 95 | 43 |
| 83 | 31 | 89 | 50 | 97 | 6 |  | - | 94 | 2059 |
| 84 | 43 | 90 | 4 | 96 | 26 | 98 | 38 | 94 | 21 4 |
| 85 | 050 | 91 | $4 \cdot 6$ | 95 | 837 | 699 | 41 |  | - |
| 86 | 18 | 92 | 9 | 96 | 94 | 700 | 49 | 93 | 9 |
| 87 | 15 | 93 | 23 | 97 | 14 | 699 | 54 | 92 | 27 |
| 88 | 19 | 92 | 23 \% | 97 | 34 | 98 | 1458 | 91 | 31 |
| 89 | 25 | 93 | $23 \cdot 8$ |  | - | 699 | 1516 | 90 | 33 |
| 89 | 35 | 94 | 34 -3 | 97 | 44 | 700 | 19 | 89 | 36 |
|  | - | 93 | $38 \cdot 6$ | 96 | 952 | 699 | 28 | 88 | 4148 |
| 88 | 38 | 92 | $39 \cdot 6$ | 95 | 100 | 98 | 1539 | 87 | 2211 |
| 89 | 159 | 91 | $44 \cdot 6$ | 96 | 11 | 99 | 163 | 86 | 17 |
| 90 | 210 | 92 | $45 \cdot 3$ | 47 | 1037 | 98 | 6 | 85 | 29 |
| 91 | 13 | 93 | 48 | 96 | 1112 | 99 | 9 | 84 | 40 |
| 92 | 17 | 94 | 5 50 -6 | 96 | 15 | 99 | 13 |  | - |
| 91 | 25 | 95 | 65 |  | - |  | - | 85 | 46 |
| 90 | 31 | 94 | 6 | 96 | 26 | 699 | 1615 | 84 | 2254 |
| 91 | 35 | 95 | 15 | 95 | 1149 | 700 | $17 \quad 2$ | 83 | 236 |
| 92 | 41 | 94 | 17 | 94 | 1212 | 01 | 14 | 82 | 15 |
| 93 | 245 |  | - | 95 | 14 | 00 | 1731 | 81 | 17 |
| 94 | 30 | 94 | 22 | 96 | 35 | 01 | 18 5 | 82 | 18 |
| 94 | 6 | 94 | 23 | 96 | 52 | 700 | 17 | 81 | 22 |
|  | - |  | - |  | - | 699 | 26 | 80 | 25 |
| 94 | 9 | 94 | 31 | 95 | 55 | 98 | 29 | 79 | 28 |
| 93 | 20 | 95 | 35 | 96 | 1256 | 99 | 32 | 80 | 31 |
| 94 | 344 | 93 | $44 \cdot 3$ | 95 | 134 | 98 | 1845 | 81 | 39 |
| 95 | 423 | 94 | 45 | 94 | 18 | 99 | 1910 | 80 |  |
| 94 | 27 | 95 | 54 | 95 | 24 | 99 | 20 | 79 | 2344 |
| 95 | 35 | 94 | 55 | 96 | 46 |  | - | '22679 | 240 |
| 94 | 37 | 95 | 656 | 95 | 47 | 99 | 31 |  |  |
| -22695 | 442 | 36 | 728 | 96 | 48 | 98 | 33 |  |  |
|  | - | -22697 | 734 | -22697 | 1352 | -22697 | 1954 |  |  |

Changes of Christchurch Horizontal Force. May 1, 1902.

| Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{aligned} & \text { Begins } \\ & \text { at_- } \end{aligned}$ | Value of H. | $\begin{aligned} & \text { Begins } \\ & \text { at- } \end{aligned}$ | Value of H. | $\underset{\substack{\text { Begins } \\ \text { at }}}{ }$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | h. m. |  | h. m. |  | h. m. |  | h. m. |  | h. m. |
| -22697 | 00 | - 22703 | 417 | $\cdot 22708$ | 738.5 | -22706 | 1057 | $\cdot 22709$ | 1940 |
| 96 | 1 | 04 | 19 | 09 | 39 | 05 | 1117 | 08 | 204 |
| 95 | 5 | 05 | 31 | 10 | 40 | 05 | 56 | 09 | 5 |
| 96 | 6 |  |  | 11 | 41 |  |  | 08 | 8 |
| 95 |  | 06 | 37 | 10 | 43 '8 | 05 | 1159 | 07 | 15 |
| 94 | 10 | 07 | 39 | 11 | 44 | 06 | 1212 | 03 | 21 |
| 93 | 17 | 06 | 43 | 11 | 45 | 05 | 1257 | 07 | 22 |
| 92 | 23 | 07 | 50 |  | - | 05 | 1318 | 06 | 43 |
| 91 | 34 | 06 | 55 | 11 | 47 |  | - | 06 | 55 |
| 90 | 47 | 07 | 456 | 10 | 53 | 06 | 1326 |  |  |
| 91 | 48 | 08 | 50 | 11 | 55 | 07 | 1411 | 06 | 2050 |
| 92 | 50 | 07 | $5 \cdot \hat{6}$ | 10 | 756 | 06 | 28 | 05 | 214 |
| 91 | 54 | 08 | 6 | 09 | 83 | 05 | 43 | 04 | 0 |
| 92 | 055 | 07 | $6 \cdot 6$ | 10 | 4 | 06 | 47 | 03 | 14 |
| 9 | 14 | 08 | 7 | 09 | 5 | 05 | 1448 | 02 | 15 |
| 92 | 20 | 06 | 9 | 08 | 11 | 04 | 150 | 01 | 37 |
| 92 | 24 | 07 | 10 | 09 | 13 |  |  | 00 | 48 |
|  | $\bar{\square}$ | 08 | 11 | 08 | 14 | 04 | 153 | 01 | 49 |
| 92 | 26 | 07 | 12 | 07 | 15 | 05 | 1618 | 700 | 50 |
| 91 | 34 | 06 | 15 | 08 | 16 | 04 | 19 | 699 | 2158 |
| 92 | 143 | 07 | 16 | 07 | 17 | 05 | 22 | 98 | 228 |
| 93 | 21 | 08 | 18 | 08 | 18 | 05 | 24 | 97 | 7 |
| 92 | 9 | 09 | 19 | 07 | 19 |  | $\overline{-}$ | 97 | 24 |
| 93 | 13 | 08 | 21 | 08 | 43 | 04 | 1629 |  | - |
| 92 | 16 | 09 | 27 | 08 | 48 | 03 | 173 | 95 | 32 |
| ${ }_{94}^{93}$ | 19 | 08 | 32 |  | 8 - | 02 | 4 | 94 | 40 |
| 94 | 46 | 09 | 40 | 07 | 854 | 01 | 8 | 93 | 43 |
| 93 | 47 | 10 | 543 | 06 | 93 | 02 | 18 | 92 | 55 |
| 94 | 48 | 09 | 627 \% | 07 | 20 | 03 | 21 | 91 | 2259 |
| 94 | 57 | 10 | 70 | 06 | 26 | 04 | 24 | 90 | 2303 |
|  | - | 09 | 1 | 05 | 946 | 05 | 40 | 91 | $0 \cdot 6$ |
| 94 | 259 | 08 | 7 | 07 | 1026 | 05 | 48 | 90 | 1 |
| 95 | 30 | 09 | 10 | 08 | 34 |  |  | 89 | 5 |
| 96 | 5 | 10 | 24 | 10 | $34 \cdot 8$ | 06 | 1759 | 88 | 8 |
| 97 | ${ }_{21} 0$ | 09 | ${ }^{25}$ | 08 | 35 | 07 | 1818 | 87 | 23 |
| 96 | 21 | 10 | 27 | 07 | 36 | 08 | 50 | 86 | 28 |
| 97 | ${ }^{23}$ | 09 | 33 | 08 | 38 | 09 | 1854 | 85 | 32 |
| 98 | 39 | 10 | 34 | 09 | 39 | 08 | 1921 | 84 | 34 |
| 699 | 44 | 09 | 35 | 08 | 43 | 08 | 26 | 83 | 2345 |
| 700 |  | 10 | ${ }^{\cdot 36}{ }_{37}{ }_{6}$ | 07 08 | 47 48 |  | 9 | -22683 | 240 |
| -22\%02 | 353 410 | $\begin{array}{r}11 \\ \hline 22709\end{array}$ | ${ }_{7}{ }^{378}$ | 08 $\cdot 22707$ | 48 1051 | -22708 | $19 \stackrel{29}{36}$ |  |  |

Changes of Christchurch Horizontal Force. May 15, 1902.

| Value of H. | $\begin{gathered} \text { Begins } \\ \text { at— } \end{gathered}$ | Value of H. | Begins | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { st- } \end{gathered}$ | Value of H. | Begins |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | h. m. |  | h. m. |  | h. m. |  | h. $m$ |  | h. $m$ |
| -22690 | 00 | -22693 | 236 | -2270コ | 449 | -22705 | 958 | -22709 | $20{ }^{25}$ |
| 89 | 13 | 92 | 37 | 04 | 51 | 04 | 1035 |  | - |
| 88 | 22 | 93 | 38 | 05 | 53 | 05 | 1041 | 08 | 27 |
| 87 | 29 | 94 | 40 | 04 | 459 | 06 | 110 | 07 | 30 |
| 86 | 34 | 93 | 41 | 03 | 55 | 06 | 31 | 06 | 37 |
| 85 | 040 | 94 | 42 | 04 | 6 |  | - | 05 | 2059 |
| 86 | 11 | 93 | 43 | 03 | 7 | 06 | 34 | 04 | 2116 |
| 85 | 5 | 94 | 44 | 04 | 9 | 05 | 48 | 03 | 47 |
| H8 | 15 | 95 | 47 | 03 | 18 | 06 | 52 | 02 | 2152 |
| 85 | 17 | 94 | 49 | 04 | 21 | 07 | 56 | 02 | 221 |
| 86 | 19 | 95 | 51 | 05 | 45 | 06 | 1159 |  | - |
| 85 | 20 | 95 | 252 | 05 | 52 | 06 | 1251 | 01 | 13 |
| 85 | 21 |  | - |  | - |  | - | 02 | 23 |
|  | - | 93 | 32 | 05 | 554 | 06 | 1259 | 01 | 28 |
| 86 | 31 | 92 | 3 | 04 | 624 | 06 | 1425 | 00 | 32 |
| 87 | 35 | 92 | 4 | 04 | 53 |  | - | 01 | 34 |
| 88 | 43 |  | - |  | - | 06 | 1428 | 00 | 35 |
| 89 | 45 | 92 | 10 | 04 | 656 | 06 | 1558 | 01 | 36 |
| 88 | 46 | 93 | 12 | 05 | $7 \quad 26$ |  | - | 00 | 37 |
| 89 | 48 | 92 | 13 | 04 | 3 | 06 | 160 | 01 | 45 |
| 90 | 152 | 91 | 19 | 05 | $8 \cdot 3$ | 07 | 41 | 700 | 2247 |
| 91 | 20 | 92 | 30 | 04 | $18 \cdot 6$ | 06 | 1647 | 699 | 233 |
| 92 | 3 | 91 | 33 | 05 | 21.3 | 05 | 170 | 98 | 24 |
| 91 | 7 | 92 | 34. | 04 | $25 \cdot 3$ | 06 | 6 | 97 | 33 |
| 92 | 14 | 93 | 36 | 05 | 35.3 | 06 | 18 | 97 | 37 |
| 91. | 15 | 94 | 37 | 04 | 40 \% |  | - |  | - |
| 92 | 23 | 95 | 42 | 05 | $53 \cdot 3$ | 05 | 29 | 97 | 39 |
| 91 | 26 | 96 | 49 | 04 | $754 \cdot 6$ | 06 | 43 | 96 | 46 |
| 92 | 27 | 97 | 55 | 04 | 823 | 05 | 51 | 95 | 52 |
| 93 | 30 | 98 | 359 |  | - | 06 | 1753 | 96 | 53 |
| 92 | 31 * ${ }^{\text {¢ }}$ | 699 | 44 | 05 | 31 | 07 | 1830 | 95 | 56 |
| 93 | $32 \cdot 5$ | 700 | 2 | 06 | 856 | 08 | 44 | 96 | 2357 |
| 92 | 33 | 00 | 13 | 05 | 92 | 08 | 55 | -22695 | 240 |
| $94^{\prime}$ | 34 |  | - | 06 | 28 |  | - |  |  |
| 92 | $34 \cdot 6$ | 02 | 23 | 05 | 39 | 07 | 1859 |  |  |
| 93 | 35 | 03 | 34 | -22705 | 955 | 08 | 1953 |  |  |
| -22694 | $235 \cdot 5$ | $\cdot 22704$ | 448 |  | - | -22709 | 205 |  |  |

Changes of Christchurch Horizontal Force．June 1， 1902.

| N |  |
| :---: | :---: |
|  | 禺荡 |
| No |  |
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|  <br>  <br> $\leftrightarrow$ © |  |
| 筑Q | 붕 |
|  | $\begin{gathered} \text { M. } \\ \text { 㵄 } \end{gathered}$ |
|  |  |
|  |  |

Changes of Christchurch Horizontal Force. June 15, 1902.

| Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | Begina at- | $\begin{aligned} & \text { Value of } \\ & \text { H. } \end{aligned}$ | Begins at- | Value of H. | Begins at- | Value of II. | Beging at- |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | h. m. |  | h. m. |  | h. 1 ll . |  | h. m. |  | h. $m$. |
| '22694 | 00 | -22703 | 73 | '22706 | 1220 | 22704 | 153 | 22709 | 1933 |
| 93 | 12 | 0.4 |  | 05 | 21 | 03 | 4 | 10 | 34 |
| 92 | 26 | 03 | 8 | 0.4 | 23 | 02 | 9 | ( $)^{4}$ | 39 |
| 91 | 36 | 02 | 13 | 702 | 24 | 01 | 11 | 18 | 44 |
| 90 | 43 | 03 | 15 | 6999 | 25 | 00 | 14 | 09 | 41 |
| 89 | 44 | 0.4 | 38 | 98 | 26 | 01 | 18 | 10 | 45 |
| 90 | 46 |  | - | 96 | 27 | 02 | 20 | (1) | 4f; |
| 89 | 51 | 03 | 40 | 9. | 28 | 01 | 26 | 10 | 48 |
| 90 | 53 | 04 | 41 | 95 | 30 | 02 | 29 | $(9)$ | 49 |
| 89 | 059 | 05 | 42 | 96 | 31 | 01 | 35 | 0 O | 50 |
| 88 | 123 | 04 | 43 | 698 | 32 | 700 | 42 | 09 | 52 |
| 88 | 31 | 03 | 744 | 702 | 33 | 699 | 43 | 08 | 1957 |
|  | - | 02 | 816 | 06 | 34 | 700 | 48 | 19 | $20 \quad 9$ |
| 89 | 33 | 03 | 19 | 07 | 36 | 01 | 52 | 10 | 17 |
| 88 | 35 | 02 | 22 | 10 | 37 | 00 | 53 | 08 | 19 |
| 89 | $13{ }^{\circ}$ | 03 | 30 | 08 | 39 | 701 | $53 \cdot 5$ | 09 | 20 |
| 90 | 214 | 02 | 36 | 10 | 40 | 699 | 54 | 10 | 21 |
| 89 | 18 | 03 | 38 | 08 | 41 | 700 | 54 \% | 09 | 24 |
| 90 | 22 | 03 | 49 | 07 | 42 | 01 | 1555 | 10 | 25 |
| 89 | 33 |  | - | 06 | 47 | 02 | 163 | 09 | 27 |
| 90 | 35 | 03 | 854 | 07 | 49 | 03 | 18 | 10 | 29 |
| 91 | 48 | 04 | $945 \%$ | 06 | 50 | 03 | 21 | 09 | 34 |
| 90 | 52 | 05 | $53 \cdot 6$ | 08 | 51 |  |  | 10 | 87 |
| 91 | 53 | 04 | $54 \cdot 6$ | 07 | 52 | 03 | 24 | 09 | 38 |
| 92 | 256 | 03 | 9593 | 08 | 54 | 04 | 41 | 10 | 39 |
| 91 | 30 | 04 | 1017 | 07 | 55 | 05 | 48 | 09 | 41 |
| 92 | 1 | 04 | 20 | 08 | 1259 | 06 | 1659 | 10 | 44 |
| 92 | 6 |  | - | 07 | 133 | 04 | 174 | 10 | 52 |
|  | - | 03 | 23 | 06 | 4 | 03 | 5 |  | - |
| 98 | 9 | 02 | 26 | 07 | 8 | 04 | 8 | 09 | 55 |
| 94 | 24 | 01 | 27 | 10 | 10 | 05 | 9 | 10 | 56 |
| 95 | 28 | 00 | 34 | 11 | 12 | 04 | 13 | 09 | 2057 |
| 96 | 30 | 01 | 38 | 12 | 13 | 03 | 15 | 08 | 2117 |
| 97 | 32 | 00 | 40 | 11 | 15 | 02 | 34 | 07 | 34 |
| 98 | 37 | 01 | 50 | 10 | 17 | 01 | 37 | 08 | 47 |
| 97 | 38 | 02 | 52 | 08 | 21 | 02 | 38 | 07 | 48 |
| 96 | 39 | 03 | 53 | 06 | 22 | 01 | 42 | 06 | 50 |
| 97 | 40 | 05 | 55 | 702 | 23 | 02 | 43 | 07 | 51 |
| 98 | 47 | 06 | 57 | 694 | 24 | 01 | 48 | 08 | 54 |
| 97 | 50 | 05 | 1059 | 92 | $24 \cdot 16$ | 01 | 50 | 09 | 57 |
| 98 | 51 | 04 | 111 |  | - |  | - | 08 | 2159 |
| 99 | 55 | 08 | 2 | 90 | 29 | 03 | 1757 | 07 | $22 \quad 2$ |
| 98 | 56 | 02 | 5 | 93 | 30 | 04 | 180 | 08 | 25 |
| 99 | 58 | 03 | 12 | 96 | 31 | 05 | 4 | 07 | 26 |
| 98 | 359 | 04 | 17 | 97 | 32 | 06 | 6 | C7 | 27 |
| 699 | 41 | 05 | 20 | 698 | 33 | 05 | 9 |  | - |
| 700 | 11 | 06 | 23 | 700 | 34 | 06 | 11 | 07 | 35 |
| 01 | 15 | 07 | 24 | 02 | 35 | 07 | 17 | 06 | 42 |
| 00 | 17 | 08 | 26 | 03 | 36 | 08 | 34 | 07 | 43 |
| 01 | 19 | 10 | 27 | 04 | 39 | 07 | 35 | 08 | 48 |
| 00 | 20 | 11 | 29 | 05 | 40 | 08 | 36 | 07 | 2259 |
| 01. | 21 | 10 | 30 | 04 | 45 | 07 | 39 | 06 | 230 |
| 02 | 25 | 09 | 33 | 03 | 48 | 08 | 43 | 07 | 3 |
| 02 | 40 | 07 | 34 | 02 | 51 | 09 | 44 | 06 | 5 |
|  | - | 06 | 35 | 01 | 53 | 10 | 47 | 05 | 7 |
| 02 | 449 | 04 | 36 | 02 | 1356 | 09 | 48 | 04 | 11 |
| 03 | 55 | 03 | 38 | 03 | 14.7 | 10 | 49 | 03 | 17 |
| 02 | 6 | 02 | 39 | 02 | 14 | 08 | 50 | 02 | 31 |
| 03 | 12 | 03 | 40 | 01 | 15 | 07 | 52 | 01 | 32 |
| 02 | 24 | 04 | 45 | 700 | 16 | 06 | 53 | 02 | 33 |
| 03 | 28 | 03 | 47 | 699 | 17 | 07 | 1856 | 01 | 38 |
| 02 | 31 | 04 | 49 | 700 | 21 | 08 | 190 | 00 | 40 |
| 03 | 39 | 03 | 51 | 01 | 22 | 09 | 1 | 01 | 41 |
| 02 | 557 |  | - | 700 | 28 | 10 | 3 | 00 | 42 |
| 02 | 68 | 04 | 53 | 699 | 30 | 09 | 7 | 01 | 43 |
|  | - | 03 | 1155 | 700 | 38 | 08 | 8 | 700 | 45 |
| 03 | 11 | 04 | 121 | 699 | 39 | 09 | 13 | 699 | 48 |
| 04 | 32 | 05 | 3 | 98 | 42 | 10 | 16 | 701 | 49 |
| 03 | 33 | 06 | 10 | 699 | 43 | 09 | 19 | 02 | 51 |
| 04 | 41 | 05 | 11 | 700 | 47 | 10 | 24 | 01 | 54 |
| 03 | 44 | 04 | 12 | 699 | 65 |  | - | 02 | 57 |
| 04 | 48 | 05 | 13 | 700 | 1458 | 10 | 26 | 01 | 2358 |
| 03 | 49 | 06 |  | -26701 | 150 | 09 | 27 | -22701 | 240 |
| -22704 | 659 | -22707 | 1216 |  | - | -22710 | 1928 |  |  |

2 E

Changes of Christchurch Horizontal Force. July 1, 1902.


Changes of Christchurch Horizontal Force. July 15, 1902.

| Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Talue of H. | Begins | Value of H. | Begins at - | $\begin{gathered} \text { Value of } \\ \text { H. } \end{gathered}$ | Begins at- | Value of 11. | Begins ut |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| -22697 | h. m. <br> 00 | -22704 | h. m. 3 555 | -22709 | $\begin{array}{cc}\text { h. } \\ 7 \\ 7 & 51\end{array}$ | '22704 | h. m. 10 E 2 | -22709 | h. m 1854 |
| 96 | 8 | 03 | 41 | 10 | 755 | 05 | 53 | 08 | 1856 |
| 95 | 17 | 04 | 13 | 09 | 80 | 07 | 54 | 09 | 191 |
| 94 | 27 | 05 | 16 | 08 | 1 | 08 | 56 | 10 | 6 |
| 93 | 046 | 06 | 22 | 07 | 2 | 07 | 1059 | 11 | - 10 |
| 92 | 112 | 05 | 28 | 06 | 3 | 06 | 11.7 | 10 | 29 |
| 92 | 16 | 04 | 37 | 05 | 8 | 05 | $13 \cdot 8$ | 11 | 1940 |
|  | - | 03 | 38 | 06 | 15 | 04 | $20 \%$ | 12 | $20 \quad 7$ |
| 92 | 18 | 04 | 43 | 06 | 21 | 03 | $28 \cdot 6$ | 13 | 12 |
| 93 | 40 | 05 | 48 |  | - | 06 | $34^{\circ} \mathrm{B}$ | 14 | 17 |
| 95 | 42 | 04 | 49 | 06 | 30 | 07 | 38 68 | 14 | 22 |
| 96 | 43 | 05 | 54 | 07 | 51 | 06 | 45 |  | - |
| 97 | 44 | 06 | 56 | 08 | 52 | 05 | $50 \% 3$ | 14 | 31 |
| 699 | 46 | 07 | 58 | 09 | 53 | 06 | $52 \cdot 3$ | 15 | 39 |
| 700 | 47 | 09 | 459 | 10 | 55 | 05 | 52 \% | 16 | 2954 |
| 699 | 52 | 11 | 50 | 11 | 859 | 06 | $1157{ }^{\prime} 6$ | 17 | 2149 |
| 98 | 53 | 12 | 1 | 10 | 96 | 05 | 1211 | 18 | 50 |
| 97 | 154 | 11 | 2 | 08 | 7 | 06 | 18 | 17 | 51 |
| 96 | 21.5 | 10 | 8 | 07 | 8 | 06 | 24 | 16 | 53 |
| 97 | 2 | 11 | 11 | 06 | 9 |  | - | 16 | 57 |
| 98 | 13 | 10 | 16 | 05 | 10 | 06 | 30 |  | - |
| 99 | 17 | 09 | 20 | 06 | 11 | 05 | 1255 | 16 | 2159 |
| 98 | 21 | 09 | 21 | 05 | 25 | 05 | 139 | 15 | $\because 211$ |
| 97 | 33 |  | - | 03 | 40 |  | - | 14 | 19 |
| 98 | 34 | 09 | 23 | 04 | 41 | 05 | 11 | 13 | 26 |
| 97 | 37 | 10 | 25 | 03 | 42 | 06 | 31 | 12 | 29 |
| 98 | 40 | 09 | 34 | 04 | 44 | 05 | 1357 | 11. | 31 |
| 99 | 41 | 07 | 37 | 05 | 50 | 04 | 144 | 10 | 40 |
| 98 | 42 | 09 | 42 |  | - | 05 | 12 | 09 | 45 |
| 699 | 43 | 11 | 543 | 05 | 52 | 05 | 21 | OS | 46 |
| 700 | 44 | 10 | 629 | 06 | 57 |  | - | 07 | 47 |
|  | - | 11 | 34 | 05 | 959 | 05 | 24 | 06 | 50 |
| 00 | 46 | 12 | 42 | 06 | 100 | 04 | 14.59 | 05 | 2256 |
| 01 | 47 | 11 | 45 | 05 | 1 | 05 | 1510 | 03 | 238 |
| 700 | 51 | 11 | 54 | 06 | 2 | 06 | 15 | 02 | 7 |
| 699 | 258 |  | - | 05 | 3 | 06 | 50 | 01 | 14 |
| 98 | 31 | 11 | 656 | 04 | 19 |  | - | 700 | 23 |
| 699 | 2 | 10 | 72 | 03 | 21 | 05 | 1554 | 699 | 27 |
| 700 | 3 | 11 | 4 | 04 | 22 | 06 | 169 | 98 | 28 |
| 699 | 4 | 10 | 19 | 03 | 24 | 07 | 33 | 97 | 30 |
| 700 | 8 | 07 | 20 | 02 | 26 | 08 | 1652 | 98 | 39 |
| 699 | 20 | 09 | 23 | 01 | 27 | 08 | 1717 | 97 | 47 |
| 700 | 29 | 10 | 24 | 700 | 28 |  | - | 98 | 51 |
| 699 | 30 | 11 | 25 | 698 | 29 | 08 | 20 | 97 | 52 |
| 98 | 32 | 12 | 27 | 699 | 31 | 07 | 33 | 96 | 53 |
| 97 | 33 | 11 | 32 | 700 | 32 | 06 | 45 | 95 | 54 |
| 98 | 35 | 10 | 33 | 01 | 35 | 07 | 52 | 96 | 56 |
| 699 | 37 | 09 | 34 | 02 | 39 | 08 | 1756 | 95 | 57 |
| 700 | 38 | 07 | 35 | 01 | 40 | 07 | 1814 | 96 | 2359 |
| 01 | 40 | 08 | 39 | 02 | 49 | 08 | 33 | -22694 | 240 |
| 02 | 46 | 09 | 40 | -22702 | 1050 | 09 | 40 |  |  |
| -22702 | 348 | -22710 | 743 |  | - | -22709 | 1851 |  |  |
|  | - |  |  |  |  |  | - |  |  |

Changes of Christchurch Horizontal Force. August 1, 1902.


Changes of Christchurch Horizontal Force. August 15, 1902.

| Value of H. | Begins at- | Value of H. | Begins at- | Value of $H$. | Begins at- | Value of H. | Begins at- | Value of H. | $\begin{gathered} \text { Begins } \\ \text { gt- } \end{gathered}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| -22697 | $\begin{array}{cc} \text { h. } & \text { m. } \\ 0 & 0 \end{array}$ | -22698 | h. m. 5 5 | 22696 | h. m. 8 | '22700 |  | -22698 | h. m. 18.52 |
| 96 | 10 | 97 | 13 | 95 | 7 | 698 | - 55 |  | - |
| 95 | 15 | 96 | 14 | 96 | 10 | 97 | 1059 | 98 | 1855 |
| 94 | 22 | 97 | 16 | 95 | 12 | 96 | 112 | 699 | 1940 |
| 93 | 38 | 97 | 41 | 96 | 33 | 97 | 24 | 700 | 2015 |
| 92 | 057 |  | - | 96 | 45 | 98 | 29 | 01 | 22 |
| 91 | 16 | 98 | 44 |  | - | 98 | 40 | 01 | 30 |
| 91 | 11 | 97 | 47 | 97 | 52 |  |  |  |  |
|  | - | 98 | 48 | 96 | 858 | 98 | 43 | 02 | 2032 |
| 92 | 13 | 97 | 551 | 97 | $9 \quad 0$ | 97 | 53 | 03 | 214 |
| 91 | 18 | 96 | 68.5 | 96 | 2 | 98 | 1154 | 04 | 41 |
| 90 | 22 | 97 | 10 | 95 | 13 | 97 | 129 | 04 | 53 |
| 89 | 132 | 96 | 12 | 94 | 21 | 98 | 29 |  | - |
| 89 | 246 | 97 | 13 | 95 | 25 | 97 | 35 | 05 | 55 |
|  | - | 96 | 15 | 94 | 26 | 97 | 50 | 05 | 2159 |
| 90 | 49 | 95 | 27 | 95 | 30 |  | - |  | - |
| 89 | 256 | 96 | 30 | 94 | 31 | 97 | 1255 | 05 | 2210 |
| 90 | 32 | 95 | 32 | 95 | 39 | 98 | $13 \quad 5 \cdot 3$ | 06 | 13 |
| 91 | 11 | 96 | 40 | 94 | 40 | 97 | $32 \cdot 6$ | 05 | 18 |
| 92 | 25 | 95 | 41 | 95 | 46 | 98 | 40 | 06 | 28 |
| 91 | 26 | 96 | 43 | 96 | 48 | 97 | $13 \quad 50 \cdot 3$ | 05 | 29 |
| 92 | 27 | 95 | 44 | 95 | 50 | 96 | 141 | 06 | 35 |
| 93 | 39 | 96 | 45 | 96 | 951 | 97 | 6 | 05 | 36 |
| 92 | 41 | 95 | 50 | 97 | 1011 | 98 | 21 | 04 | 48 |
| 93 | 342 | 96 | 658 | 96 | 14 | 98 | 25 | 03 | 2257 |
| 94 | 44 | 95 | 72 | 97 | 16 |  | - | 04 | 230 |
| 95 | 20 | 96 | 3 | 97 | 17 | 97 | 1427 | 03 | 1 |
| 95 | 23 | 93 | 4 |  | - | 96 | 152 | 02 | 3 |
|  | - | 96 | 31 | 97 | 20 | 97 | 35 | 03 | 4 |
| 96 | 31 | 95 | 32 | 98 | 34 | 97 | 50 | 02 | 8 |
| 97 | 37 | 94 | 37 | 97 | 35 |  |  | 01 | 21 |
| 96 | 41 | 95 | 49 | 98 | 36 | 97 | 1553 | 700 | 24 |
| 97 | 46 | 94 | 56 | 699 | 40 | 96 | 1613 | 699 | 32 |
| 96 | 48 | 95 | 57 | 700 | 43 | 97 | 1647 | 98 | 42 |
| 97 | 454 | 94 | 759 | 699 | 44 | 98 | 171 | 99 | 43 |
| 98 | 510 | 95 | 80 | 701 | 46 | 98 | 23 | 98 | 44 |
| 99 | $10 \cdot 3$ | 96 | 1 | 700 | 51 |  | - | 97 | 2355 |
| -22694 | 511 | -22695 | 83 | -22699 | 1053 | -22698 | 1728 | -22697 | 240 |

Changes of Christchurch Horizontal Force. September 1, 1902.

| Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | Begins at- | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { st- } \end{gathered}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\cdot 22680$ | $\begin{array}{rr}\text { h. } & \text { m. } \\ 0 & 0\end{array}$ | -22688 | h. m. | -2689 | h. 6 6 | $\cdot 22696$ | $\begin{array}{cc} \text { h. } \\ \hline 15 & 8 \end{array}$ | -22689 | h. m. <br> 2121 |
| 79 | 6 | 85 | 33 | 90 | 10 | 97 | 10 | 88 | 32 |
| 78 | 17 | 84 | 5 | 89 | 26 | 98 | 14 | 87 | 39 |
| 77 | 20 | 85 | 12 | 90 | 27 | 97 | 28 | 86 | 41 |
| 78 | 26 | 87 | 13 | 91 | 34 | 96 | 39 | 85 | 44 |
| 76 | 35 | 88 | 14 | 90 | 645 | 95 | 1558 | 84 | 51 |
| 77 | 50 | 89 | 15 | 89 | 7. 2 | 96 | 161 | 83 | 2157 |
| 78 | 51 | 88 | 24 | 90 | 12 | 95 | 3 | 82 | $22 \quad 2$ |
| 79 | 54 | 87 | 26 | 90 | 29 | 96 | 27 | 81 | 5 |
| 78 | 55 | 86 | 29 |  | - | 95 | 30 | 80 | 10 |
| 77 | 58 | 87 | 32 | 90 | 32 | 96 | 34 | 79 | 15 |
| 76 | 059 | 88 | 37 | 91 | 46 | $97^{\circ}$ | 38 | 78 | 20 |
| 75 | 14 | 87 | 44 | 90 | 50 | 96 | 40 | 77 | 27 |
| 74 | 6 | 88 | 47 | 89 | 754 | 97 | 43 | 76 | 28 |
| 73 | 7 | 89 | 354 | 90 | 817 | 96 | 45 | 75 | 37 |
| 72 | 8 | 90 | 411 | 91 | 40 | 97 | 47 | 75 | 44 |
| 71 | 9 | 89 | 18 | 92 | 45 | 97 | 48 |  | - |
| 72 | 10 | 89 | 22 | 91 | 46 |  | - | 74 | 53 |
| 73 | 11 |  | - | 91 | 48 | 97 | 51 | 73 | 54 |
| 74 | 12 | 89 | 27 |  | - | 98 | 1656 | 72 | 2257 |
| 75 | 13 | 88 | 33 | 92 | 53 | 97 | 1710 | 71 | 230 |
| 76 | 15 | 87 | 36 | 93 | 58 | 98 | 19 | 70 | 6 |
| 77 | 16 | 88 | 38 | 92 | 859 | 99 | 1741 | 71 | $10 \%$ |
| 78 | 17 | 87 | 41 | 91 | 99 | 99 | 1821 | 70 | 11 |
| 79 | 20 | 88 | 44 | 92 | 952 |  | - | 71 | 16 |
| 79 | 21 | 89 | 47 | 91 | 100 | 699 | 26 | 69 | $16 \cdot 5$ |
|  | - | 88 | 52 | 92 | 2 | 700 | 28 | 70 | 18 |
| 79 | 23 | 89 | 53 | 93 | 16 | 01 | 1831 | 69 | 20 |
| 80 | 32 | 88 | 54 | 92 | 40 | 00 | 1937 | 68 | 24 |
| 81 | 33 | 87 | 56 | 91 | 48 | 01 | 40 | 69 | 25 |
| 83 | 35 | 86 | 58 | 92 | 1054 | 700 | 43 | 68 | 28 |
| 84 | 42 | 87 | 459 | 91 | 115 | 699 | 45 | 69 | 35 |
| 83 | 45 | 89 | 55 | 92 | 8 | 700 | 49 | 68 | 36 |
| 81 | 48 | 87 | 6 | 93 | 17 | 700 | 56 | 67 | 38 |
| 80 | 51 | 88 | 7 | 94 | 24 '3 |  | - | 68 | 41 |
| 81 | 159 | 86 | 8 | 93 | 25 | 699 | 1959 | 69 | 42 |
| 82 | 21 | 87 | 9 | 94 | 1155 | 98 | 2017 | 68 | 43 |
| 83 | 2 | 88 | 20 | 94 | 120 | 97 | 23 | 69 | 44 |
| 82 | 11 | 87 | 21 |  | - | 96 | 30 | 70 | 46 |
| 81 | 13 | 88 | 22 | 94 | 3 | 95 | 33 | 71 | 49 |
| 82 | 22 | 87 | 30 | 93 | 23 | 94 | 38 | 70 | 52 |
| 83 | 24 | 88 | 31 | 94 | 42 | 93 | 2051 | 69 | 53 |
| 84 | 28 | 87 | 32 | 93 | 1256 | 92 | 214 | 70 | 2358 |
| 85 | 41 | 88 | 36 | 94 | 1319 | 91 | 10 | $\cdot 22670$ | 240 |
| 86 | 51 | 89 | 40 | 94 | 38 | 90 | 14 |  |  |
| 87 | 53 | 89 | 57 |  | - | 90 | 15 |  |  |
| $\cdot 22688$ | 254 |  | - | 94 | 1344 |  | - |  |  |
|  | - | -22690 | 559 | $\cdot 22695$ | 156 | -22690 | 2119 |  |  |

Changes of Christchurch Horizontal Force. September 15, 1902.

| Value of H. | $\begin{aligned} & \text { Begins } \\ & \text { at- } \end{aligned}$ | Value of H. | Begins at- | Value of H. | Bogins at- | $\begin{gathered} \text { Value of } \\ \mathrm{H} . \end{gathered}$ | Begins at- | Value of H. | Begins at |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| -22674, | h. m. 0 | -22693 | h. m. | -22695 | h. m. 5 5 | - 22687 | h. m. | -22693 | h. m. |
| - 73 | 14 | 94 | - 27 | 2205 | 52 | 2268 | 941 | 22093 |  |
| 74 | 16 | 93 | 28 | 96 | 27 | 88 | 45 | 94 | 24 |
| 75 | 22 | 94 | 37 | 97 | 35 | 89 | 46 | 95 | 27 |
| 76 | 23 | 95 | 42 | 96 | 39 | 90 | 48 | 94 | 39 |
| 75 | 24 | 94 | 47 | 98 | 40 | 91 | 951 | 93 | 40 |
| 74 | 25 | 93 | 48 | 99 | 41 | 90 | 1016 | 94 | 41 |
| 73 | 27 | 95 | 49 | 98 | 48 | 89 | 28 | 95 | 44 |
| 74. | 32 | 94 | 51 | 97 | 49 | 88 | 31 | 95 | 47 |
| 75 | 35 | 92 | 52 | 96 | 50 | 89 | 38 |  | - |
| 77 | 36 | 91 | 253 | 95 | 555 | 90 | 39 | 96 | 1451 |
| 77 | 51 | 92 | 31 | 96 | 65 | 91 | 40 | 95 | 152 ® |
|  | - | 93 | 3 | 95 | 13 | 92 | 41 | 95 | 165 |
| 77 | 55 | 92 | 5 | 94 | 15 | 93 | 42 |  | - |
| 78 | 58 | 95 | 6 | 93 | 17 | 94 | 43 | 95 | 11 |
| 79 | 059 | 96 | 8 | 94 | 19 | 93 | 50 | 96 | 1622 |
| 78 | 14 | 97 | 9 | 95 | 22 | 92 | 53 | 96 | 1741 |
| 77 | 6 | 98 | 11 | 96 | 25 | 91 | 56 |  | - |
| 78 | 7 | 699 | 12 | 95 | - 27 | 92 | 1059 | 96 | 43 |
| 79 | 10 | 700 | 13 | 96 | 28 | 93 | 1119 | 97 | 1746 |
| 78 | 11 | 699 | 15 | 95 | 40 |  | - | 98 | 1812 |
| 79 | 12 | 700 | 17 | 94 | 42 | 94 | 22 | 97 | 1828 |
| 80 | 15 | 699 | 20 | 93 | 44 | 93 | 23 | 98 | 190 |
| 81 | 16 | 700 | 22 | 92 | 46 | 92 | 45 | 97 | 1 |
| 82 | 21 | 699 | 24 | 93 | 47 | 91 | 51 | 98 | 5 |
| 81 | 26 | 98 | 25 | 92 | 49 | 92 | 1153 | 97 | 10 |
| 82 | 27 | 699 | 29 | 91 | 51 | 93 | 120 | 96 | 27 |
| 83 | 30 | 700 | 31 | 92 | 55 | 92 | 13 | 95 | 1953 |
| 84 | 32 | 01 | 33 | 91 | 659 | 93 | 19 | 94 | 205 |
| 85 | $32 \cdot 6$ | 00 | 38 | 91 | 72 | 92 | 37 | 93 | 13 |
| 84 | 34 | 00 | 42 |  | - | 91 | 39 | 92 | 38 |
| 85 | 35 |  | - | 91 | 76 | 91 | 45 | 91 | 43 |
| 86 | 36 | 01 | 50 | 91 | 831 |  | - | 89 | 53 |
| 87 | 40 | 00 | 52 |  | - | 91 | 51 | 88 | 64 |
| 88 | 43 | 01 | 53 | 92 | 42 | 92 | 53 | 89 | 55 |
| 87 | 45 | 00 | 358 | 91 | 849 | 93 | 1256 | 90 | 2057 |
| 88 | 46 | 01 | 49 | 92 | 911 | 92 | 1312 | 89 | 210 |
| 89 | 47 | 03 | 21 | 91 | 18 | 91 | 16 | 88 | 1 |
| 90 | 52 | 04 | 24 | 90 | 20 | 92 | 29 | 87 | 2 |
| 89 | 153 | 03 | 29 | 89 | 21 | 91 | 40 | 86 | 4 |
| 90 | 20 | 02 | 40 | 88 | 22 | 92 | 41 | 85 | 11 |
| 91 | 3 | 01 | 43 | 89 | 28 | 93 | 43 | 84 | 15 |
| 90 | 7 | 00 | 457 | 90 | 30 | 94 | 44 | 83 | 19 |
| 91 | 8 | 01 | 55 | 91 | 31 | 95 | 1345 | 82 | 22 |
| 92 | 15 | 700 | 12 | 89 | 33 | 94 | 145 | 83 | 24 |
| 93 | 17 | 697 | 18 | 88 | 34 | 95 | 7 | -22682 | 2125 |
| 94 | 19 | 96 | 19 | 87 | 35 | 94 | 13 |  |  |
| -22694 | 223 | -22695 | 524 | -22686 | 936 | -22693 | 1417 | Record | stops. |

Changes of Christchurch Horizontal Force. October 1, 1902.

| Value of II. | Begins at- | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | Begins at- | Value of H. | Begins at- | $\begin{gathered} \text { Value of } \\ \mathbf{H} \end{gathered}$ | Begins at- |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 22675 | h. m. 0 | -22680 | h. m. 288 | -2681 | $\begin{aligned} & \text { h. } \\ & 4 . \text { m } \\ & \text { 4, }\end{aligned}$ |  | h. m. | '22686 | $\begin{array}{ll}\text { h. } & \text { m. } \\ 19 & 40\end{array}$ |
| 74 | 9 | 81 | 25 | 82 | 53 | $\cdot 22681$ | 854 | 85 | 48 |
| 75 | 10 | 80 | 26 | 83 | 56 | 85 | 96 | 84 | 1959 |
| 74 | 12 | 79 | 38 | 84 | 457 | 84 | 26 | 83 | 2010 |
| 73 | 14 | 80 | 39 | 83 | 51 | 85 | 38 | 82 | 13 |
| 74 | 15 | 81 | 41 | 84 | 2 | 84 | 49 | 81 | 16 |
| 73 | 17 | 82 | 42 | N3 | 5 | 83 | 957 | 80 | 24 |
| 72 | 26 | 83 | 46 | 82 | 8 | 84 | 107 | 80 | 32 |
| 73 | 33 | 84 | 48 | 81 | 10 | 85 | 1057 |  | - |
| 74 | 40 | 84 | 49 | 80 | 12 | 86 | 1128 | 77 | 40 |
| 73 | 42 |  | - | 81 | 19 | 86 | 36 | 76 | 47 |
| 72 | 45 | 84 | 56 | 82 | 23 |  | - | 75 | 2054 |
| 73 | 47 | 85 | 57 | 81 | 24 | 86 | 40 | 74 | 21.2 |
| 74 | 49 | 84 | 58 | 82 | 32 | 85 | 42 | 73 | 4 |
| 73 | 51 | 83 | 259 | 82 | 36 | 86 | 1153 | 72 | 12 |
| 74 | 058 | 82 | 30 |  | - | 87 | 120 | 71 | 21 |
| 73 | 11 | 81 | 1 | 82 | 39 | 88 | 2 | 70 | 34 |
| 74 | 2 | 80 | 2 | 83 | 545 | 87 | 21 | 69 | 37 |
| 75 | 6 | 79 | 7 | 82 | 60 | 86 | 27 | 68 | 39 |
| 76 | 8 | 80 | 10 | N1 | 4 | 85 | 34 | 68 | 2152 |
| 75 | 13 | 81 | 16 | 82 | 6 | 86 | 1242 |  | - |
|  | - | 82 | 18 | *3 | 10 | 87 | 134 | 67 | 220 |
| 76 | 16 | 83 | 20 | 84 | 14 | 87 | 12 | 68 | 16 |
| 77 | 19 | 84 | 25 | 83 | 18 |  | - | 67 | 19 |
| 76 | 20 | 85 | 33 | N 2 | 25 | 86 | 18 | 68 | 26 |
| 75 | 24 | 84 | 34 | 83 | 35 | 85 | 34 | 69 | 37 |
| 76 | 28 | 85 | 35 | 82 | 36 | 86 | 41 | 70 | 41 |
| 77 | 36 | 84 | 36 | 83 | 38 | 85 | 1359 | 71 | 46 |
| 76 | 38 | 85 | 39 | 82 | 39 | 86 | 1417 | 70 | $47 \cdot 5$ |
| 77 | 43 | 86 | 44 | 83 | 41 | 87 | 31 | 71 | 48 |
| 75 | 44 | 85 | 53 | 82 | 44 | $8{ }^{6}$ | 40 | 73 | 2255 |
| 76 | 45 | 86 | 355 | 83 | 45 | 86 | 41 | 73 | 235 |
| 75 | 46 | 85 | 44 | 82 | 46 |  | - | 74 | 9 |
| 76 | 48 | 86 | 5 | 81 | 51 | 86 | 1443 | 75 | 11 |
| 77 | 53 | 87 | 7 | 82 | 652 | 87 | 152 | 76 | 15 |
| 79 | 58 | 88 | 10 | 81 | 72 | 87 | 52 | 77 | 19 |
| 80 | 159 | 87 | 15 | 82 | 4 |  | - | 77 | 21 |
| 81 | 20 | 86 | 16 | 83 | 7 | 86 | 1557 |  | - |
| 82 |  |  | - | 82 | 16 | 87 | 161 | 76 | 29 |
| 83 | 2 | 87 | 24 | 81 | 27 | 88 | 6 安 | 77 | 32 |
| 82 | 6 | 88 | 25 | 82 | 29 | 89 | $20 \cdot 6$ | 76 | 33 |
| 81 | 8 | 87 | 26 | 81 | 33 | 90 | 34 3 | 77 | 40 |
| 80 | 9 | 86 | 28 | 82 | 55 | 89 | 1637 | 78 | 47 |
| 81 | 10 | 85. | 29 | 81 | 57 | 89 | 1712 | 79 | 2358 |
| 80 | 12 | 84 | 31 | 82 | 758 |  | - | -22679 | 240 |
| 79 | 17 | 83 | 32 | 81 | 88 | 89 | $17 \quad 20$ |  |  |
| 77 | 18 | 82 | 35 | 82 | 12 | 89 | 1912 |  |  |
| 78 | 21 | 81 | 43 | 83 | 24 |  | - |  |  |
| -22679 | 222 | -22680 | 445 | -22683 | 848 | -22687 | 1920 |  |  |

Changes of Christchurch Horizontal Force. October 15, 1902.

| Value of H. | Begins at- | Yalue of $H$. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | Begins at- | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | h. m. |  | h. m. |  | h. m. |  | h. m. |  | h. m. |
| -22672 | 00 | -22690 | 428 | -22684 | 644 | 22678 | 109 | -22680 | 19 4 |
| 73 | 13 | 89 | 31 | 83 | 47 | 77 | 13 | 80 | 12 |
| 72 | 19 | 90 | 32 | 84 | 57 | 76 | 20 |  | - |
| 73 | 22 | 88 | 37 | 83 | 658 | 75 | 22 | 81 | 10 |
| 74 | 39 | 89 | 38 | 84 | 73 | 74 | 27 | 80 | 32 |
| 75 | 42 | 90 | 39 | 83 | 4 | 75 | 28 | 79 | 46 |
| 74 | 46 | 89 | 43 | 84 | 9 | 76 | 32 | 80 | 50 |
| 75 | 48 | 88 | 44 | 83 | 14 | 77 | 38 | 79 | 54 |
| 76 | 49 | 89 | 45 | 83 | 16 | 76 | 40 | 78 | 1959 |
| 75 | 53 | 90 | 46 |  |  | 77 | 41 | 77 | 209 |
| 76 | 56 | 89 | 48 | 82 | 19 | 76 | 1046 | 76 | 14 |
| 77 | 57 | 88 | 453 | 83 | 47 | 77 | 1111 | 75 | 20 |
| 76 | 059 | 89 | 55 | 84 | 49 | 78 | 14 | 74 | 24 |
| 77 | 16 | 88 | 11 | 83 | 56 | 79 | 20 | 73 | 30 |
| 78 | 8 | 87 | 13 | 82 | 758 | 78 | 25 | 72 | 37 |
| 78 | 10 | 86 | 17 | 81 | 80 | 79 | 26 | 72 | 41 |
|  | - | 85 | 21 | 80 | 1 | 80 | 27 |  | - |
| 79 | 13 | 84 | 23 | 79 | 2 | 79 | 30 | 70 | 50 |
| 80 | 18 | 83 | 24 | 78 | 5 | 78 | 39 | 69 | 53 |
| 81 | 20 | 84 | 25 | 79 | 9 | 78 | 42 | 68 | 2059 |
| 82 | 24 | 85 | 26 | 80 | 10 |  | - | 67 | 217 |
| 83 | 36 | 86 | 28 | 81 | 12 | 80 | 45 | 66 | 12 |
| 84 | 45 | 85 | 34 | 82 | 18 | 81 | 55 | 65 | 16 |
| 85 | 47 | 86 | 39 | 83 | 26 | 82 | 1159 | 64 | 19 |
| 86 | 49 | 87 | 50 | 84 | 34 | 81 | 126 | 63 | 25 |
| 87 | 156 | 86 | 54 | 83 | 39 | 82 | 1229 | 62 | 33 |
| 88 | 22 | 85 | 65 | 82 | 41 | 82 | 1316 | 61 | 37 |
| 89 | 4 | 84 | 558 | 82 | 44 |  | - | 60 | 41 |
| 88 | 12 | 85 | 60 |  | - | 82 | 22 | 59 | 2148 |
| 89 | 15 | 86 | 1 | 81 | 50 | 83 | 44 | 58 | 2215 |
| 90 | 16 | 85 | 4 | 82 | 852 | 82 | 1348 | 57 | 21 |
| 91 | 22 | 84 | 0 | 81 | 94 | 82 | 1651 | 56 | 28 |
| 92 | 40 | 83 | 11 | 80 | 6 |  | - | 55 | 31 |
| 93 | 56 | 84 | 12 | 79 | 14 | 82 | 1658 | 56 | 2254 |
| 94 | 259 | 85 | 14 | 78 | 17 | 83 | 1714 | 57 | 2315 |
| 95. | 33 | 84 | 16 | 79 | 18 | 82 | 184 | 58 | 18 |
| 94 | 8 | 85 | 18 | 80 | 22 | 82 | 12 | 59 | 28 |
| 98 | 12 | 84 | 19 | 81 | 25 |  | - | 59 | 39 |
| 92 | 14 | 83 | 21 | 80 | 28 | 82 | 17 |  | - |
| 91 | 332 | 84 | 28 | 79 | 30 | 81 | 18 | 59 | 42 |
| 91 | 412 | 85 | 33 | 78 | 936 | 80 | 29 | 60 | 2351 |
| -22691 | 4-19 | 84 -22885 | 38 642 | -22677 | 10 5 | -22681 | 1848 | -22661 | 240 |

Changes of Christchurch Horizontal Force. November 1, 1902.


Changes of Christchurch Horizontal Force. November 15, 1902.


Changes of Christchurch Horizontal Force. November 15, 1902-continued.

| Value of 11. | $\begin{gathered} \text { Beqins } \\ \text { at- } \end{gathered}$ | Value of 11. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | Begins | Value of H. | Begine at- |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| '2268) | $\begin{array}{cc} \mathrm{h} . & \mathrm{m} . \\ 15 & 8 \end{array}$ | -22f86 | $\begin{array}{ll}\text { h. m. } \\ 17 & 23\end{array}$ | -22691 | h. m. $1934{ }^{\circ} \mathrm{B}$ | -22669 | h. m. 2057 | -22656 | h. m. <br> 2228 |
| 81 | 10 | 87 | 26 | 90 | $35 \cdot 6$ | 68 | 213 | 55 | 29 |
| 82 | 11 | 87 | 41 | 89 | $37 \cdot 6$ | 66 | 5 | 56 | 30 |
| 81 | 16 |  | - | 88 | $40 \cdot 6$ | 65 | 13 | 55 | 36 |
| 80 | 18 | 87 | 47 | 89 | 45 | $6{ }^{6}$ | 17 | 56 | 37 |
| 79 | 20 | n6 | 50 | 88 | $47 \cdot 3$ | 65 | 21 | 55 | 38 |
| 78 | 24 | 87 | 53 | 87 | 1954 | 64 | 22 | 56 | 40 |
| 77 | 29 | 86 | 54 | 87 | 2013 | 63 | 28 | 55 | 43 |
| 76 | 41 | 87 | 1755 |  | - | 62 | 29 | 54 | $43 \%$ |
| 77 | 42 | 88 | 189 | 87 | 15 | 63 | 30 | 55 | 43.5 |
| 78 | 1557 | 89 | 12 | 86 | 18 | 62 | 31 | 53 | $43 \cdot 6$ |
| 79 | 166 | 91 | 14 | 85 | 19 | 61 | 33 | 55 | 44 |
| 79 | 14 | 90 | 16 | 84 | 20 | 60 | 38 | 54 | 45 |
|  | - | 89 | 17 | 83 | 23 | 59 | 41 | 55 | 49 |
| 80 | 16 | 91. | 18 | 82 | 25 | 59 | 42 | 54 | 50 |
| 81 | 20 | 92 | 20 | 81 | 28 |  | - | 54 | 54 |
| 82 | 23 | 93 | 22 | 80 | 29 | 59 | 48 |  | - |
| 81 | 30 | 91 | 24 | 79 | 30 | 58 | 52 | 56 | 2258 |
| 82 | 31 | 90 | 34 | 78 | 33 | 57 | 55 | 57 | 2314 |
| 81 | 35 | 91 | 40 | 77 | 35 | 56 | 2158 | 56 | 30 |
| 82 | \$9 | 90 | 45 | 76 | 37 | 55 | 2212 | 57 | 31 |
| 81 | 40 | 89 | 51 | 75 | 39 | 56 | 13 | 56 | 32 |
| 82 | 47 | 89 | 54 | 74 | 45 | 55 | 15 | 57 | 47 |
| 83 | 1658 |  | - | 73 | 47 | 56 | 17 | 59 | 52 |
| 84 | 1711 | 90 | 1857 | 72 | 49 | 55 | 18 | 60 | 2356 |
| 85 | 13 | 89 | 190 | 71 | 51 | 56 | 19 | -22660 | 240 |
| -22684 | 1715 | $\cdot 22690$ | $1933 \cdot 6$ | -22670 | 2052 | -22655 | 2224 |  |  |

Changes of Christchurch Horizontal Force. December 1, 1902.


Changes of Christchurch Horizontal Force. December 15, 1902.

| Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ | Value of H. | Begins at- | Value of H. | Begins at- | Value of H. | Begins at- | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at- } \end{gathered}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | h. m. |  | h. m. |  | h. m. |  | h. m. |  | h. m. |
| -22647 | 00 | -22672 | 239 | '22666 | $7 \quad 2$ | -22667 | 1311 | -22647 | 2024 26 |
| 46 |  |  | - | 65 | 6 | 68 | 13 | 46 | 26 |
| 47 | 10 | 71 | 41 | 66 | 24 | 67 | 18 | 45 | 29 |
| 48 | 12 | 70 | 44 | 65 | 26 | 67 | 19 | 45 | 35 |
| 49 | 15 | 68 | 45 | 64 | 35 |  | - |  | - |
| 48 | 19 | 67 | $45 \cdot 3$ | 65 | 41 | 66 | 24 | 44 | 37 |
| 49 | 24 | 69 | 46 | 66 | 758 | 67 | 33 | 43 | 42 |
| 50 | 27 | 70 | 47 | 65 | 86 | 68 | 1346 | 42 | 48 |
| 51 | 30 | 71 | 49 | 64 | 10 | 67 | 1411 | 41 | 49 |
| 52 | 35 | 72 | 51 | 65 | 11 | 66 | 25 | 40 | 2054 |
| 53 | 37 | 73 | 52 | 66 | 19 | 67 | 28 | 41 | 210 |
| 52 | 39 | 72 | 53 | 65 | 21 | 67 | 40 | 40 | 1 \% |
| 53 | 40 | 73 | 54 | 66 | 32 |  |  | 39 | 7 \% |
| 52 | 42 | 72 | 258 | 67 | 39 | 67 | 43 | 38 | $15 \cdot 6$ |
| 51 | 46 | 71 | 30 |  | - | 66 | 48 | 37 | 16 '3 |
| 50 | 48 | 70 | 1 | 68 | 45 | 67 | 14.53 | 36 | 20 |
| 51 | 52 | 72 | 4 | 67 | 51 | 68 | 1518 | 35 | 33 |
| 52 | 53 | 71 | 14 | (38 | 55 | 67 | 26 | 84 | 44 |
| 53 | 54 | 70 | 16 | 67 | 858 | 66 | 33 | 35 | 50 * 6 |
| 54 | 055 | 68 | 18 | 68 | 95 | 67 | 49 | 34 | $2159 \cdot 6$ |
| 54 | 13 | 67 | 20 | 66 | 7 | 68 | 1555 | 35 | 221 |
|  | - | 66 | 21 | 67 | 8 | 68 | 168 | 36 | 5 |
|  |  | 67 | 25 | 68 | 12 |  | - | 35 | 8 |
| 58 | ${ }_{7}^{6}$ | 68 | 26 | 67 | 14 |  |  |  | 13 |
| 58 | - | 69 | 36 | 68 | 18 | 68 | 17 | 33 | 13 |
|  | - | 70 | 37 | 67 | 19 | 70 | 25 | 32 | 16 |
| 58 | 12 | 71 | 39 | 68 | 27 | 71 | 32 | 33 | 20 |
| 59 | 16 | 72 | 43 | 67 | 46 | 72 | 1644 | 32 | 22 |
| 58 | 18 | 73 | 46 | 68 | 954 | 73 | 1734 | 33 | 23 |
| 59 | 19 | 74 | 47 | 68 | 1015 | 73 | 1739 | 32 | 30 |
| 58 | 23 | 75 | 49 |  | - | 73 | 38 | 33 | 32 |
| 57 | 24 | 76 | 53 |  |  |  | 46 | 32 | 33 |
| 58 | 25 | 75 | 55 | 68 | 17 | 72 | ${ }^{46}$ | 33 | 35 |
| 57 | 26 | 76 | 359 | 70 | 24 | 71 | 1751 | 32 | 37 |
| 58 | 32 | 77 | 42 | 69 | 27 | 70 | $18 \quad 4$ | 33 | 38 |
| 59 | 34 | 77 | 7 | 68 | 33 | 69 | 10 | 34 | 42 |
| 60 | 37 |  | $\overline{14}$ | 68 69 | 33 <br> 34 | 68 | 13 | 33 | 43 |
| 61 | 39 | 77 | 14 | 68 | 34 39 | 67 | 22 | 34 | 47 |
| 62 | 40 | 78 | 19 | 68 69 | 1040 | 66 | 29 | 35 | 49 |
| 61 | 41 | 77 | 22 | 69 68 | 111 | 65 | 33 | 36 | 53 |
| 62 | 43 | 76 | 55 | 69 | 118 | 64 | 40 | 37 | 56 |
| 63 | 46 | 75 | 458 | 70 | 8 12 | 63 | 1849 | 38 | 2259 |
| 64 | 55 | 74 | 51 | 69 | 13 | 62 | 195 | 39 | 236 |
| 63 | 158 | 75 | ${ }^{6}$ | 68 | 13 | 62 | 8 | 40 | 9 |
| 64 | 26 | 74 | 19 | 69 | 28 |  | - | 41 | 15 |
| 65 | 8 | 73 | 31 | 69 | 35 | 61 | 10 | 42 | 16 |
| 66 | 11 | 72 | 37 | 69 | 35 | 60 | 13 | 41 | 23 |
| 67 | 14 | 73 | 39 |  | - | 59 | 18 | 42 | 27 |
| 68 | 17 | 73 | 44 | 68 | 1137 | 58 | 27 | 43 | 28 |
| 70 | 22 |  | - | 67 | 127 | 57 | $33 \cdot 3$ | 42 | 32 |
| 69 | 24 | 73 | 47 | 66 | 11 | 56 | 34 | 43 | 36 |
| 70 | 25 | 72 | 557 | 65 | 26 | 55 | 41 | 43 | 39 |
| 71 | 26 | 71 | 65 | 66 | 29 | 54 | 47 |  | - |
| 70 | 28 | 70 | 6 | 67 | 32 | 53 | 1958 | 43 | 41 |
| 69 | 30 | 69 | 26 | 66 | 36 | 52 | $20 \quad 2$ | 44 | 46 |
| 68 | 31 | 68 | 28 | 67 | 39 | 51 | 6 | 45 | 2350 |
| 69 | 35 | 67 | 642 | 66 | 46 | 50 | 8 | -22645 | 240 |
| 70 | 37 | -22666 | 70 | 67 | 1257 | 49 | 18 |  |  |
| -22671 | 238 |  | - | -22666 | 130 | -22648 | 2018 |  |  |

Changes of Christchurch Horizontal Force Jamary 1, 1903.

| Value of H. | $\begin{aligned} & \text { Regins } \\ & \text { at—— } \end{aligned}$ | Value of H. | $\begin{aligned} & \text { Begins } \\ & \text { at- } \end{aligned}$ | Value of H. | $\begin{aligned} & \text { Begins } \\ & \text { at- } \end{aligned}$ | Value of H. | $\begin{gathered} \text { Begins } \\ \text { at } \end{gathered}$ | Value of II. | $\begin{aligned} & \text { Begins } \\ & \text { at- } \end{aligned}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | h. m. |  | h. m. |  | h. m. |  | l. m. |  | h. m. |
| -22647 | 00 |  | - | -22680 | 755 | -22684 | 1355 | 22645 | 2041 |
| 48 | 9 | '22697 | 418 | 81 | 57 | 85 | 1357 | $8 \downarrow$ | 46 |
| 49 | 11 | 96 | 21 | 80 | 759 | 84 | 142 | 83 | 50 |
| 48 | 12 | 97 | 25 | 82 | 81 | 8.1 | 37 | 8 | 51 |
| 49 | 16 | 96 | 30 | 83 | 1.5 |  |  | 81 | 53 |
| 50 | 18 | 95 | 45 | 82 | 2 | 84 | 1441 | 80 | 56 |
| 51 | 22 | 44 | 48 | 80 | 2.5 | 86 | 150 | 79 | 2059 |
| 52 | 24 | 95 | 53 | 81 | 3 | 85 | 2 | 78 | 211 |
| 53 | 26 | 94 | 457 | 82 | $3 \cdot 5$ | 84 | 5 | 77 | $\stackrel{2}{7}$ |
| 52 | 29 | 93 | 50 | 83 | 6 | 85 | 53 | 76 | 7 |
| 53 | 31 | 92 | 10 | 82 | 7 | 86 | 56 | 75 | 8 |
| 54 | 44 | 93 | 11 | 81 | $8 \stackrel{3}{3}$ | 85 | 57 | 74 | 13 |
| 53 | 45 | 92 | 16 | 82 | 9 | 86 | 1559 | 73 | 16 |
| 54 | 056 | 91 | 21 | 84 | 10 | 88 | 160 | 72 | 18 |
| 54 | 10 | 90 | 23 | 83 | 11 | 87 | 8 | 71 | 20 |
|  | - | 91 | 26 | 84 | 13 | 87 | 11 | 70 | 23 |
| 55 | 3 | 90 | 28 | 83 | 15 |  |  | 69 | 26 |
| 56 | 8 | 91 | 33 | 84 | 16 | 87 | 13 | 68 | 29 |
| 57 | 18 | 92 | 35 | 83 | 17 | 86 | 14 | 67 | 32 |
| 58 | 20 | 93 | 37 | 84 | 18 | 85 | 15 | 66 | 35 |
| 59 | 24 | 92 | 40 | 83 | 20 | 86 | 16 | 66 | 39 |
| 60 | 30 | 92 | 44 | 84 | 21 | 87 | 20 |  | - |
| 61 | 37 |  | - | 84 | 43 | 88 | 21 | 64 | 45 |
| 62 | 41 | 91 | 46 |  | - | 89 | 22 | 63 | 46 |
| 63 | 45 | 90 | 48 | 84 | 48 | 88 | 23 | 62 | 50 |
| 64 | 48 | 89 | 52 | 83 | 52 | 87 | 28 | 61 | 53 |
| 63 | 49 | 90 | 54 | 84 | 53 | 86 | 30 | 60 | 54 |
| 64 | 50 | 91 | 557 | 82 | 55 | 87 | 35 | 59 | 2156 |
| ${ }^{65}$ | 51 | 89 | 62 | 83 | 56 | 86 | 37 | 58 | 220 |
| 66 | 54 | 88 | 4 | 84 | 857 | 85 | 38 | 57 | 5 |
| 67 | 157 | 89 | 7 | 85 | 98 | 86 | 44 | 56 | 10 名 |
| 68 | 20 | 90 | 9 | 84 | 10 | 87 | 48 | 55 | 14 |
| 70 | 2 | 91 | 10 | 83 | 13 | 86 | 50 | 54 | $17 \cdot 6$ |
| 71 | 4 | 90 | 11 | 84 | 14 | 87 | 51 | 55 | 19 |
| 72 | 7 | 89 | 12 | 83 | 27 | 86 | 52 | 54 | 20 |
| 73 | , | 88 | 19 | 84 | 940 | 87 | 1653 | 55 | 21 |
| 74 | 13 | 87 | 21 | 84 | 104 | 88 | 171 | 54 |  |
| 75 | 15 | 86 | 24 |  | - | 89 | 19 | 53 | 26.3 |
| 76 | 16 | 85 | 25 | 84 | 6 | 90 | 30 | 52 | 38.3 |
| 77 | 18 | 86 | 26 | 83 | 7 | 90 | 39 | 53 | 39 |
| 78 | 19 | 85 | 28 | 84 | 29 |  | - | 52 | 41 |
| 79 | 21 | 84 | 31 | 83 | 1056 | 91 | 46 | 51 | 46 |
| 80 81 | ${ }_{28}^{26}$ | 83 | ${ }_{38} 34$ | 84 | $11{ }^{1}$ | 92 | 1749 | 50 | 51. |
| 81 | ${ }_{29}^{28}$ | 84 <br> 83 <br> 8 | 45 | 86 85 | ${ }_{27}^{23}$ | 91 | 18 3 | 49 | 51.8 |
| 83 | 30 | 82 | 52 | 86 | 28 | 93 | 12 | 49 | 54 |
| 84 | 32 | 81 | 656 | 85 | 31 | 92 | 13 | 48 | 2259 '3 |
| 85 | 35 | 82 | 71 | 84 | 34 | 93 | 21 | 47 | 2314 |
|  | - |  | - | 85 | 38 | 92 | 22 | 47 | 16 |
| 87 | 38 | 81 | 4 | 85 | 39 | 93 | 30 |  | - |
| 88 | 42 | 80 | 5 |  | - | 93 | 48 | 46 | 19 |
| 89 90 | 48 68 | 79 80 | ${ }_{8}^{6}$ | 84 83 | 1141 | 94 | 1852 | 45 | 23 |
| 91 | 57 | 81 | 9 | 82 | ${ }^{12} 4$ | 95 | 1923 | 43 | 31 |
| 92 | 259 | 82 | 14 | 83 | 49 | 94 | 35 | 42 | 34 |
| 93 | 31 | 83 | 19 | 84 | 1255 | 93 | 1957 | 43 | 37 |
| 94 | 16 | 82 | 26 | 85 | 138 | 92 | 206 | 44 | 41 |
| 93 | 19 | 81 | 29 | 86 | 10 | 92 | 7 | 45 | 44 |
| 94 | 22 | 82 | 30 | 85 | 13 |  |  | 46 | 49 |
| 93 | 27 | 81 | 32 | 86 | 15 | 92 | 9 | 47 | 53 |
| 94 | 31 | 82 | 35 | 86 | 16 | 91 | 11 | 46 | 54 |
| 95 | 38 43 | 81 82 | 37 40 | 84 | 21 | 90 89 | 20 24 | 47 | - 56 |
| 97 | 345 | 81 | 41 | 85 | 39 | 88 | 29 | -22648 | 23 24 |
| 98 | 43 | 82 | 47 | 86 | 42 | 87 | 32 |  |  |
| -22698 | 411 | -22681 | 752 | 22685 | 1344 | '22686 | 2038 |  |  |

Chavges of Christchurch Horizontal Force. January 15, 1903.

| Value of H. | Begins | Value of H. | Begins <br> at- | Value of H. | Begins <br> at- | Value of H. | Begins | Value of II. | $\begin{aligned} & \text { Begins } \\ & \text { at- } \end{aligned}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 4. m. |  | h. m. |  | h. m. |  | h. m. |  | h. m. |
| -22673 | 00 | 22690 | 346 | -22692 | 627 | -22689 | 1018 | -22693 | 1744 |
| 74 |  | 89 | 49 | 93 | 29 |  | - | 92 | 1753 |
| 75 | 7 | 90 | 52 | 94 | 33 | 86 | 1147 | 91 | 1813 |
| 76 | 13 | 91 | 356 | 93 | 35 | 87 | 49 | 90 | 19 |
| 77 | 15 | 90 | 40 | 94 | 37 | 86 | 51 | 89 | 28 |
| 78 | 18 | 89 | 1 | 95 | 43 | 87 | 1158 | 89 | 55 |
| 79 | 22 | 91 | 5 | 94 | 644 | 86 | 124 |  | - |
| 80 | 24 | 90 | 11 | 93 | 7. 1 | 87 | 10 | 89 | 1857 |
| 81 | 27 | 90 | 15 | 94 | - 4 | 88 | 20 | 88 | 192 |
| 82 | 29 |  | - | 94 | 6 | 87 | 22 | 87 | 5 |
| 83 | 38 | 89 | 20 |  | - | 89 | 23 | 86 | 9 |
| 84 | 44 | 90 | 23 | 94 | 9 | 90 | 30 | 85 | 19 |
| 85 | 49 | 89 | 25 | 93 | 20 | 91 | 46 | 84 | 28 |
| 86 | 53 | 90 | 30 | 92 | 26 | 90 | 51 | 83 | 38 |
| 87 | 55 | 89 | 32 | 91 | 28 | 91 | 52 | 82 | 47 |
| 88 | 56 | 90 | 34 | 90 | 31 | 90 | 54 | 81 | 49 |
| 87 | 57 | 91 | 40 | 89 | 36 | 91 | 125 | 80 | 53 |
| 86 | 58 | 90 | 44 | 88 | 758 | 92 | 131 | 79 | 1958 |
| 87 | 059 | 91 | 46 | 87 | 84 | 91 | 6 | 78 | 203 |
| 88 | 10 | 90 | 47 | 88 | 26 | 92 | 7 | 77 | 8 |
| 89 | 1 | 92 | 48 | 89 | 28 | 93 | 8 | 76 | 14 |
| 88 | 4 | 91 | 51 | 90 | 32 | 94 | 13 | 75 | 17 |
| 89 | 7 | 92 | 55 | 89 | 36 | 93 | 14 | 74 | 25 |
| 89 | 8 | 91 | 56 | 88 | 43 |  | - | 74 | 30 |
|  | - | 93 | 458 | 89 | 44 | 92 | 20 |  | - |
| 89 | 10 | 92 | 50 | 88 | 45 | 91 | 35 | 73 | 32 |
| 88 | 14 | 94 | 4 | 88 | 47 | 90 | 41 | 72 | 34 |
| 89 | 17 | 93 | 10 |  | - | 89 | 56 | 71 | 38 |
| 88 | 20 | 92 | 11 | 89 | 50 | 90 | 1358 | 70 | 43 |
| 87 | 22 | 91 | 12 | 88 | 51 | 91 | 14.5 | 69 | 48 |
| 86 | 29 | 92 | 20 | 87 | 53 | 92 | 10 | 68 | 51 |
| 87 | 31 | 93 | 23 | 89 | 859 | 93 | 14 | 67 | 53 |
| 88 | 40 | 92 | 25 | 87 | 90 | 92 | 15 | 66 | 2055 |
| 89 | 47 | 94 | 27 | 88 | $0 \cdot 5$ | 91 | 19 | 65 | 217 |
| 88 | 51 | 93 | 30 | 86 | 1 | 90 | 22 | 65 | 219 |
| 87 | 52 | 92 | 31 | 88 | 2 | 89 | 26 |  | - |
| 88 | 54 | 93 | 34 | 89 | 3 | 90 | 36 | 54 | 2218 |
| 89 | 56 | 92 | 35 | 88 | 4 | 91 | 45 | 53 | 21 |
| 88 | 159 | 91 | 36 | 89 | 5 | 92 | 47 | 52 | 27 |
| 87 | 22 | 92 | 89 | 88 | 17 | 91 | 1451 | 51 | 31 |
| 86 | 16 | 93 | 48 | 89 | 19 | 90 | 152 | 50 | 40 |
| 85 | 24 | 94 | 49 | 86 | 20 \% | 89 | 7 | 51 | 47 |
| 86 | 31 | 94 | 51 | 88 | 22 | 88 | 1526 | 51 | 50 |
| 85 | 34 |  | - | 87 | 26 | 87 | 160 |  |  |
| 86 | 36 | 93 | 54 | 88 | 29 | 88 | $\underline{\square}$ | 51 | 53 |
| 87 | 40 | 94 | 559 | 87 | 33 | 89 | 10 | 52 | 2259 |
| 86 | 42 | 95 | 60 | 88 | 34 | 88 | 12 | 53 | 23 4 ${ }^{3}$ |
|  | - | 96 | 1 | 89 | 36 |  | - | 54 | 13 '6 |
| 87 | 48 | 95 | 3 | 88 | 43 | 88 | 14 | 55 | 17 * |
| 86 | 50 | 94 | 5 | 89 | 48 | 89 | 20 | 56 | $20 \cdot 6$ |
| 87 | 259 | 93 | 7 | 88 | 49 | 90 | 40 | 57 | 24 -3 |
| 88 | 32 | 92 | 9 | 89 | 53 | 91 | 43 | 58 | $54 \cdot 3$ |
| 89 | 7 | 91 | 10 | 88 | 956 | 90 | 48 | 59 | 2357 |
| 88 | 8 | 92 | 14 | 89 | 101 | 91 | 1659 | '22659 | 240 |
| 89 | 0 | 91 | 15 | 89 | 4 | 92 | $17 \quad 7$ |  |  |
| 90 | 44 | 90 | 19 |  | - | -22692 | 1738 |  |  |
| -22689 | 345 | -22691 | 620 | -22689 | 107 |  | - |  |  |

Changes of Christchurch Horizontal Force. February 1, 1903.


Changes of Christchurch Horizontal Force. February 15, 1903.


Changes of Christchurch Declination during International Term Hours, 1902-3.


Changes of Christchurch Declination during International Term Hours, 1902-3 (continued).

§4. To make sure that the tables will be correctly interpreted by the reader, let us consider the Horizontal-Force data for the first term day, March 1, 1902. At 0h. Om., G.M.T., when the term day commenced, the value of $H$ was $\cdot 22672$, and no departure from this of as much as 00001 (the unit adopted) was recorded until 0 h .3 m ., when the value had fallen to 22671 . During the next minute it rose again to its initial value, and then, after the lapse of 3 minutes, fell a second time to 22671 . So it went on, now up, now down, with on the whole an upward tendency until 1 h . 15 m ., when the value recorded was 22680 . This persisted for the next 16 minutes-i.e. until 1 h .31 m .-when a break of 6 minutes occurred in the observations, indicated by a dash in the time column. Readings were resumed at lh .37 m ., when the force had risen by $\cdot 00004$. Details of what occurred during the 6 -minute interval are necessarily unknown.
The term hour on March 1, 1902, was from 2 h .0 m . to 3 h .0 m ., and a considerable number of the changes recorded during this hour were at 20 seconds ( 0.3 minutes) or 40 seconds ( $0 \cdot 6$ minutes) after the exact minute. Observations at 20 -second intervals, as has been already explained, were as a rule made during several hours of each term day; but in addition to the sheets in which these observations appeared there were others which recorded only the results answering to 1 -minute observations, and except for the actual term hours I have confined myself to the latter. Occasionally observations of turning-points taken at fractions of a minute were inserted in the 1-minute tables, and when these existed they were made use of.

In the case of H the first three significant figures were always either • 226 or $\cdot 227$, and the difference between two consecutive readings was never large. It was thus seldom necessary to print more than the
two last significant figures. Three were necessary only when, as sometimes happened, $H$ was rising or falling in the immediate neighbourhood of the value $\cdot 22700$. In the case of D the value always lay between $16^{\circ}$ and $17^{\circ}$ east, thus minutes only are recorded except at the leginning and end of each hour.

Before drawing conclusions from the term-day data it will be convenient to consider the nature of the regular diurnal variation.
§5. Christchurch is a comparatively new magnetic station, and particulars of the regular diurnal variation there do not seem as yet to have been published. There is thus no direct evidence that the hourly readings on term days, pp. 177-179 of the "Physical Observations," can be regarded as representative of undisturbed conditions at Christchurch. This is, however, the conclusion to which the character of the results points, and it seems sufficiently confirmed by the character of the synchronous results from Greenwich, Kew, Falmouth, Pola, Bombay, and Mauritius, stations whose general characteristics are known. Considering the limited number of days, perfectly smooth diurnal inequalities could not be expected from the Christchurch data, even for the year as a whole. There being only two days a month, diurnal inequalities for individual months would clearly have been too uncertain. Examination of the data showed, however, that, as is usual in corresponding northern latitudes, the variation in the type of the diurnal inequality throughout the year is not large, and that sub-division of the year into Midwinter (May to August), Equinox (March, April, September, October), and Midsummer (November to February) would combine months in which the type was closely similar and the amplitude not widely different. Diurnal inequalities for D and H were thus derived for the three seasons and the year. In the case of V the apparent range was so small that, taking the discontinuities into account, the results appeared too uncertain for publication.

Referring to "Physical Observations," pp. 177 and 178 , it will be seen that data are lacking for the second midnight of November 1, and for hours 10 p.m. to midnight on September 15. These missing values were replaced by interpolation, in such a way as not to influence the conclusions one would have derived from the remaining days as to the diurnal variation during the hours of the day concerned. This seemed preferable to omitting two entire days from so small a number. As is generally the case when diurnal inequalities are derived from a limited number of selected days, there was an appreciable non-cyclic element. This was eliminated in the usual way.

The diurnal inequalities resulting for D and H are given in Table $\mathrm{IA}, \mathrm{p} .238$. It will be noticed that the hours are numbered 1 to 24 . The reason for this is as follows: The usual practice is to refer diurnal inequalities either to local mean time or to the standard mean time of the country where the observatory is situated. Thus the natural thing would have been to have employed standard New Zealand time, counting hours 1 to 12, a.m. and p.m. But standard New Zealand time is $11 \frac{1}{2}$ hours fast on Greenwich, so that what is a.m. in the one is p.m. in the other. And as G.M.T.-with hours 1 to 12, a.m. and p.m.-was employed in the Christchurch tables in the "Physical Observations," a similar nomenclature with New Zealand time substituted would have been extremely apt to create confusion. I accordingly decided to employ Greenwich time, but to count the hours 1 to 24 , which was really the course proposed in the international term-hour forms. Thus in the tables of this Appendix, hours 6 and 18, for example, are equivalent, respectively, to $5.30 \mathrm{p} . \mathrm{m}$. and $5.30 \mathrm{a} . \mathrm{m}$. New Zealand time. In the discussion a.m. and p.m. mean forenoon and afternoon in New Zealand.
§6. The principal features disclosed by Table IA in the case of Declination are as follows: At all seasons there are two distinct maxima and two distinct minima-distinguished by figures in heavy type. There is thus a distinct double daily oscillation, but the oscillation during the night (say from hours 7 to 19 in Midwinter, and 8 to 17 in Midsummer) is trifling compared to that during the day.

To bring out minute details of the night movements, records would be necessary from a much larger number of days. The extreme westerly position of the needle is reached from 2 to 3 hours before noon. As is usual in the southern hemisphere, the most conspicuous movement during the day is the swing over to the east. The extreme easterly position is reached 2 or 3 hours after noon, and so later than is usual for the corresponding westerly maximum in Britain. As at Kew, the inequality derived from the four equinoctial months closely resembles both in type and range that derived from the whole year. The
ratio borne by the range in Midwinter to that in Midsummer is considerably less than at Kew. The only seasonal variation of type that appears at all marked is that the extreme westerly position is reached earlier in Summer than in Winter.

Horizontal-Force changes, as already stated, were less regular than those of Declination, thus the inequality figures for the former element in Table Ia are less smooth than those for the latter. Still, they suffice to bring out the general character clearly. As with D, there is a distinct double daily period, even in Midwinter. The principal minimum is, as at Kew, more conspicuons than the maximum. This minimum occurs near noon-later in the day than at Kew-and it seoms some two hours later in Midwinter than in Midsummer. There is a morning maximum some 5 hours before the principal minimum, and a second maximum about $5 \mathrm{p} . \mathrm{m}$., but while the former maximum considerably predominates in Midwinter, the latter is the larger in Midsummer.

The variation in the hour of occurrence of the principal minimum reduces the range in the diurnal inequality for the year, which falls appreciably short of that for the equinoctial months. The difference between the ranges during Midsummer and Midwinter, though well marked, is less conspicuous than in the case of the Declination.

This brief summary will suffice to show the main features of the diurnal variation.
§7. Looking at inequalities such as those in Table $\mathrm{I}_{\mathrm{A}}$, and still more so when looking at inequalities based on a very large number of days, one sees a magnetic element rising persistently to a maximum, and falling as steadily to a minimum. But if one takes the more familiar case of a tide in the sea, one knows that while the combination of the observations of a large number of days produces the semblance of continuous rise for hours, followed by continuous fall, on any individual day either rise or fall consists of numerous ebbs and flows in the height, dependent partly on the waves. The number and amplitude of these subsidiary movements may, for all I know, differ widely even on the same day at two spots on the same coast no great distance apart. Very probably the phenomena on any given day are largely dependent on the wind and on the distance of storm centres. Still, I am disposed to think it is probable that the minor phenomena in the rise and fall of the tides in different oceans, or different parts of the same large ocean, possess characteristics of a distinctive character, though these may not be apparent in tidal data calculated in the usual way. Now it is probably the same with the magnetic diurnal variation, with this difference that purely local features have little if any effect. These considerations encouraged the hope that a systematic study of the changes shown by the Christchurch data might disclose phenomena whose comparison with corresponding phenomena at other stations might some day throw a valued light on the nature of the diurnal variation. If similar data should be got out on a future occasion for Christchurch from a year of large sun-spot frequency the comparison could hardly fail to be instructive.
§8. The magnetic changes recorded by ordinary magnetographs rarely if ever occur quite suddenly. Even with slow-run curves the movements usually appear continuous. The fact that the value of $H$ is given on March 1 as 22672 at 0 h .0 m . and as $\cdot 22671$ at 0 h . 3 m . does not mean that the value was exactly .22672 at 0 h .0 m ., 0 h . 1 m . and 0 h .2 m ., but exactly $\cdot 22671$ at 0 h .3 m . Even supposing the photographic trace perfection and the tabulator unerring, it may mean a number of different things. The values at four successive minutes might for instance be

$$
\begin{array}{llll}
\cdot 22672_{3}, & \cdot 22672_{0}, & \cdot 22671_{7,}, & \cdot 22671_{4} \\
\cdot 22672_{0}, & 22671_{8}, & \cdot 22671_{6,} & \cdot 22671_{4}
\end{array}
$$

or
and there might have been numerous oscillations in value too small to be detected.
The conclusions one would draw as to the number of changes of force occurring in a given time would certainly depend somewhat on the sensitiveness of the instrument, even when one required a given minimum amplitude to count as a change. The number of changes recorded would inevitably vary with the size of the minimum accepted and with the interval between successive readings. On the other hand, I do not think the number of changes in Declination or Horizontal Force (I cannot say the same of Vertical Force) would depend much on the type of the magnetographs (if reasonable damping exists) as distinct from their sensitiveness. In the absence of artificial sources of disturbance, such as persistent air
currents or electric-tram currents, magnets which are properly damped do not get into vibration, and their position at any instant-excopt during specially violent magnetio storms-may be regarded as one of equilibrium, answering to the instantaneous value of the magnetic field.
§9. Returning now to the term-day data wo find very different types of change. Taking, for instance, the Declination changes during the term hours of January 15, February 1, and February 15, 1903 (see p. 230), we have a persistent steady rise in Declination from beginning to end of the hour. On the other hand, during the term hours of July 15 and September 15, 1902, p. 229, the Declination went on oscillating through a narrow range without any marked drift in one direction. Then, on August 1, 1902, we have the magnet not budging by as much as $0^{\prime} \cdot 1$ from the position it occupied at the beginning of the hour. A consideration of these varied phenomena suggested that account should be taken both of the number of changes and of the range for each hour.

Tables IIA to IVA, pp. 239-241, summarise the results from the readings at 1-minute intervals as to the number of changes in each hour of each term day as derived from the corrected values of the elements, $0^{\prime} \cdot 1$ being the minimum change for D , and $\mathrm{l} \gamma$ the minimum for H and V . In preparing these tables, allowance had to be made for the gaps in the record. This was done on the assumption that changes took place during an interval for which data were lacking at the same average rate as during that part of the same hour for which observations existed. It is these corrected totals that appear in the tables. To show the comparatively trifling character of the uncertainties due to lack of record, the total number of additional changes thus introduced during each day is given in brackets in the last column of each table. With a view to obtaining the diurnal variation in the hourly number of changes, it was also necessary to interpolate values for a few hours-e.g. the last two hours of September 15 -for which no data existed. These interpolated values-enclosed in brackets-were arrived at by assuming that the ratio which they bore to the total from the remaining hours of the day was the same as for the average day whose records were complete.
§10. Tables Va and VIa, pp. 242-243, summarise the results for the range of Declination and Horizontal Force throughout each hour of each term day. As a rule, no allowance was made for gaps in the record. This was partly because no one guiding principle could be applied, but mainly because there was usually reason to believe that the loss of record had had little, if any, influence on the result. An allowance was, however, made when a gap commencing towards the end of an hour extended into the next hour, and the earliest recorded reading of the second hour lay outside the range covered by the readings during the complete part of the first hour. In such an instance the element was assumed to have increased or diminished, as the case might be, at a uniform rate whilst the trace was lacking. Values were also interpolated for a few hours-e.g. the last three hours of September 15 -for which readings were totally or largely wanting. These interpolated values are enclosed in brackets; they were arrived at in a similar fashion to the interpolated values in Tables IIA to IVA. The two last columns in Tables VA and VIA give respectively the arithmetical sum of the 24 hourly ranges and the absolute range, i.e. the difference between the algebraically greatest and least of the readings obtained at 1 -minute intervals throughout the day. These absolute ranges are necessarily never less and are usually larger than the ranges on pp. 177 and 178 of the "Physical Observations," which are derived from hourly readings, but the difference is in no case large. On September 15, 1902, and January 15, 1903, there was reason to believe that one of the extreme readings for the day was lost through a break in the readings, so $a$ " + " is put after the range as deduced from the readings recorded.
§11. If we examine Tables IIA and IIIA in detail, we observe a great variability from day to day in the number of changes. On March 15,1903 , there were 426 changes in the Declination reading, or an averago of about 18 an hour, though the total range for the day was only $9^{\prime} \cdot 2$. On December 1,1902 , and February 1, 1903, with the closely similar absolute ranges of $9^{\prime} \cdot 7$ and $8^{\prime} \cdot 9$, the numbers of changes were respectively only 244 and 181. The least number of daily changes-averaging about $5 \frac{1}{2}$ per hour-was recorded on April 15, 1902, although the absolute range for the day exceeded that on the nine immediately subsequent term days. The greatest number of Declination changes recorded in an hour, viz. 41 on March 1, 1903, took place in an early morning hour, for which the normal change due to the regular diurnal inequality was under $0^{\prime} \cdot 2$. There are 13 hours in Table IIA during which no change in D was
recorded-all included between $6 \mathrm{p} . \mathrm{m}$. and $4 \mathrm{a} . \mathrm{m}$. Two consecutive hours during which no change was recorded are met with on March 1, 1902, August 1, 1902, and Eebruary 1, 1903. The longest interval during which no Declination change as large as $0^{\prime} \cdot 1$ was recorded occurred on August 1, 1902, and amounted to no less than 2 hours 22 minutes. On this day the Declination showed a range of only $0^{\prime} \cdot 1$ between 11 h .45 m . and 16 h .3 m . (i.e. between $11.15 \mathrm{p} . \mathrm{m}$. and $3.33 \mathrm{a} . \mathrm{m}$. ). The intervals on March 1 and February 1 during which no change occurred were but little shorter, being on the former occasion 138 minutes (from 1.30 to 3.48 a.m.), and on the latter 133 minutes (from 0.25 to $2.38 \mathrm{a} . \mathrm{m}$.). On March 1 and August 1 there were a few minutes' loss of trace during the intervals specified, but none on February 1.

In the case of H the greatest number of changes in one day took place on November 15, 1902, the day for which Declination data were lacking; the changes averaged nearly 19 an hour, though the absolute range was only $47 \gamma$. November 15 is, perhaps, the outstanding example of an almost incessant oscillation, the tide of magnetic change never setting in one direction for any length of time. The immediately previous term day, November 1, with exactly the same absolute range, had a number of changes less by nearly 200. The smallest number of changes, 136, occurred on April 15, 1902, the same day as gave the minimum number for D . The greatest number of changes of H in any one hour was 47 on March 15, 1903 , the hour representing 4.30 to 5.30 p.m. The range encountered during this hour was only $8 \gamma$. There are 18 hours in Table IIIA during which no change of H as large as $1 \gamma$ was recorded, and of these seven were included between hours 14 and 16 (i.e. 1.30 to 3.30 a.m.).

On August 1 and October 1, 1902, no change in H was recorded during two consecutive hours; on October 15, 1902, during three consecutive hours; and on May 15, 1902, during four consecutive hours. The complete interval without change on this last occasion extended really to 4 hours 41 minutes (from $11.29 \mathrm{p} . \mathrm{m}$. to 4.10 am m .). Trace, it is true, was lacking for three short periods, amounting in all to 10 minutes, and it is, of course, possible, though unlikely, that a measurable change may thus have failed to be recorded.

In the case of the Vertical Force the number of changes (the minimum change being $1 \gamma$ as with $H$ ) showed a maximum of 61 on March 15, 1902, and a minimum of 21 on May 1, 1902. The largest number of changes in one bour, 12, occurred on March 1, 1903, shortly after midnight. No change was recorded during four consecutive hours on June 15, 1902, and February 15, 1903, while on December 15, 1902, no change was recorded during $5 \frac{3}{4}$ hours (from $10.30 \mathrm{p} . \mathrm{m}$. to 4.15 a.m.). The sensitiveness of the H and V magnetographs was for $H 1 \mathrm{~cm} .=46 \gamma$, for $\mathrm{V} 1 \mathrm{~cm} .=27 \gamma$. There is thus nothing in the sensitiveness to account for the extraordinary excess in the number of changes in $H$. There is, however, some room for doubt whether in the case of $V$ absence of apparent change necessarily meant absolute uniformity in the field. The discontinuities presented on several days are rather suggestive of friction between knife-edge and plane, which might suffice to prevent any response on the part of the magnet to minute oscillations in the field.

If the term days are arranged according to the number of changes of reading, the order is by no means the same for the different elements. This is even more strikingly true as regards individual hours. For example, during hour $5-6$ on March 15,1903 , whilst the changes in $H$ numbered 47 , those in $D$ and $V$ did not exceed the average. Thus what may fairly be described as "restlessness" by no means always affects the different elements simultaneously. The fluctuations may thus answer for many consecutive minutes, sometimes even for hours, to disturbing forces acting persistently in one general direction.
§12. When comparing the number of changes in D and H it is necessary to remember that the minimum changes, viz. $0^{\prime} \cdot 1$ in $D$ and $1 \gamma$ in $H$, are arbitrary quantities and do not represent equal disturbances. The force perpendicular to the magnetic Meridian required to produce a change of $0^{\prime} \cdot 1$ in Declination varies as the local value of $H$. For Christchurch it amounts to $0 \cdot 66 \gamma$. Thus the unit for changes in $H$ was really 50 per cent. larger than that for changes in $D$. The number of changes must decrease as the size of the unit is enlarged, but the exact relationship between the two quantities is uncertain.* If we

- Taking June 1 and December 1, the effect was tried of taking 0 ' 2 as the unit change in $D$ and $2 \gamma$ as the unit change in $H$. The number of changes in $D$ was reduced only about 50 per cent., but the number in $H$ was reduced by 70 per cent. The mean of these two days represented fairly average conditions so far as the total number of changes is concerned.
divide the number of observed changes in H by the corresponding number in D multiplied by 0.66 we shall probably overestimate the relative variability in H , while we shall certainly underestimate it if we divide the number for $H$ by the number for $D$. The results of these two operations are as follows:-

|  | Year. | Midwinter. | Equinox. | Midsummer. |
| :---: | :---: | :---: | :---: | :---: |
| Number of changes in $\mathbf{H}$ $0.6 \overline{6} \times \overline{n u m b e r}$ of changes in D | $1 \times 67$ | 177 | 1.55 | 167 |
| Number of changes in If <br> Number of changes in 1) | $1 \cdot 10$ | $1 \cdot 17$ | 102 | $1 \cdot 10$ |

If we put the variability in H as 40 per cent. greater than the variability in $D$, we shall probably not be very far wrong. The figures suggest a seasonal variation in the ratio, but it does not seem to be large A variation in the ratio would suggest, of course, a seasonal variation in the dominant direction of the disturbing forces.

The tables of hourly ranges call for less comment. If the diurnal variation consisted of a regular single oscillation, the numbers in the penultimate columns of Tables VA and VIA would be approximately double the corresponding numbers in the last columns. The ratio in reality varies within wide limits. In Table VA the largest value of the ratio, $3 \cdot 6$, is met with on the two June term days; in Table VIa the largest value also occurs on June 15 and is no less than $5 \cdot 4$. The largest hourly ranges in $D$ and $H$ occurred respectively on December 15, 1902, and January 1, 1903, being $3^{\prime} \cdot 9$ for D and $24 \gamma$ for $H$. The smallness of these maxima is eloquent testimony to the remarkable quietness of the term days.
§13. To ascertain more exactly the character of the diumal variation in the frequency of the changes of reading given by Tables IIA to IVA means were derived for each hour from the first 24 term days, and, in the case of $D$ and $H$, also from combinations of these days forming the same three seasons as in Table IA. The two term days of March, 1903, were not taken into account.

The results are given in Table VIIA, p. 244, the absolutely largest and least of the hourly values being in heavy type. The penultimate column gives the mean of the 24 hourly values, and so represents the average number of changes of reading per hour for the whole year and the seasons. The last column gives the average number of changes per diem. The number of changes in the case of both D and H is considerably larger at Midsummer than at Midwinter, but the difference between the seasons is much less conspicuous than it is in the case of the range of the diurnal inequality.

In the case of $D$ changes of reading are most numerous close to noon, whilst the magnet is swinging over to the early afternoon maximum of easterly Deelination. In the case of H changes of reading are most numerous in the afternoon, either an hour or two before or else an hour or two after the afternoon maximum. In all three elements changes appear to be least numerous in the early morning in the neighboufhood of 4 or 5 a.m. Whilst the day changes are decidedly more numerous than the night, the difference is much less than one would have anticipated from a consideration of the diurnal inequality. Thus if we define "day "as the twelve hours commencing $6 \mathrm{a} . \mathrm{m}$., and assume that half of the entries under hours 7 ( $6.30 \mathrm{p} . \mathrm{m}$.) and 19 ( $6.30 \mathrm{a} . \mathrm{m}$.) in Table VIIA belong to the day, while the inequality figures appropriate to $6 \mathrm{a} . \mathrm{m}$. and $6 \mathrm{p} . \mathrm{m}$. are the arithmetic means of those for 18 and 19 and for 6 and 7 respectively, we obtain for the ratios of the day to the night number of shanges, and the day to the night inequality ranges, the following values:-

| Day/Night. |  | Year | Midwinter. | Equinox. | Midsummer. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Number of change | D | $1 \cdot 62$ | 1.22 | 1.86 | 1.75 |
| Number of changea . . $\cdot$. | H | $1 \cdot 31$ | 1.08 | $1 \cdot 63$ | 1.25 |
| Inequality ranges | D | 5.5 | $5 \cdot 7$ | 5.8 | $3 \cdot 2$ |
| Inequity range | H | 5.0 | 4.5 | $3 \cdot 8$ | $8 \cdot 7$ |

\$14. The diurnal variation in the amplitude of the hourly rauges $(\mathrm{R})$ in D and $\mathrm{H}-\mathrm{as}$ deduced from Tables VA and VIA-is given in Tables VIIIA and IXi, p. 24 , for the year and the three seasons; maxima and minima are in heavy type. For comparison, the tables also show the corresponding hourly increments (1), as derived from the mean hourly readings for the year and the seasons, the largest positive and negative increments being in heavy type. These last data differ from the increments derivable from the diurnal inequalities in Table IA only in containing the non-cyclic change which was eliminated in forming the diurnal inequalitios. The presence of the non-cyclic element makes only a triffing difference. For example, Table VIIIA gives as the increment in $D$, for the year as a whole, for the hour ending at 2 the value $+1^{\prime} \cdot 22$, whereas from the diurnal inequality in Table IA we have $+3^{\prime} \cdot 02-1^{\prime} \cdot 81$, or $+\mathrm{l}^{\prime} \cdot 21$.

Before discussing Tables VIITA and IXA, two anomalous results should be explained, viz, the excess in the I over the R figure for hour 24 for Year and Equinox in Table VIIIA. The excess obviously cannot represent a physical fact, but only an imperfection in the data. It arises from the absence of observational data during one or both of the two last hours of the 24 . The hourly readings and the hourly ranges missing were replaced independently by interpolation methods already described. The inconsistency could no doubt have been removed by other methods of interpolation, but it did not seem worth while to take the trouble, as the results would not necessarily have made any nearer approach to real accuracy.

The $R$ and I data in Tables VIIIA and IXA would agree if there were nothing but a regular diurnal inequality with maxima and minima occurring at exact (Greenwich), hours. As a matter of fact, the difference between the two sets of figures in Table VIIIA is very small near the middle of the day, when the Declination inequality change is most rapid. This is due, I think, to more than one cause. During an hour when the inequality change is very rapid any small irregular disturbance has no chance of influencing the hourly range, unless it occur near one end of the hour. If, for instance, the normal D change during the hour is a rise of $2^{\prime}$, the minimum value will occur very near the commencement, and the maximum very near the end of the hour, unless there is an irregular disturbance amounting to at least several tenths of a minute. It is thus obvious that the hourly range in D during one of the mid-day hours when the needle is moving rapidly to the east cannot on an ordinary undisturbed day be much in excess of the range derived from the hourly readings. This is unquestionably accountable in part for the observed phenomenon, but it cannot, I think, wholly explain it, especially in the Midsummer months, and the conclusion seems warranted that the tendency to an oscillatory, or ebb-and-flow, type of variation in D was at its minimum during the day hours when the regular inequality changes were most rapid.

In the night hours the $\mathbf{R}$ and I values differ greatly in both the Tables VIIIA and IXA, in fact there is little relationship between them. It would, however, be incorrect to assume that the $R$ changes owe nothing to the diurnal inequality even during an hour when the I change vanishes. The difference between solar and mean time is not wholly negligible in any one of the seasons, still less so in the course of the year, also the times of maxima and minima in the diurnal magnetic inequality vary throughout the year whether we refer them to solar or to mean time. A low I value may mean that the diurnal inequality changes are very slow throughout the whole hour. It may signify, however, that the mean time of occurrence of a maximum or minimum falls about the middle of the hour, the actual time of occurrence varying considerably in different months comprised in the same season. The contributions from different portions of the season to I will then differ in sign, and may largely neutralise one another, but this of course is not the case with their contributions to $R$. This is unquestionably very largely the cause of the great difference between the R and I figures for some of the day hours, e.g. the hours ending at 22 and 3 (i.e. at 9.30 a.m. and 2.30 p.m.) in Equinox and Year in Table VIIIA. In most of the night hours, however, the difference between the $R$ and I figures must be otherwise explained. In Table IXA there are some day hours, e.g. in Midsummer those ending at 1, 21, and 22, when the R and I figures are very close, but this is exceptional, the differences being on the whole much more conspicuous for H than for D . This is in strict accordance with the greater variability already observed in $H$, but is not, I think, due exclusively to this cause. If we divide the arithmetic mean of the $R$ by
the arithmetic mean of the I values, as given in the last columns of Tables VIIIA and IXA, we find the following values for the ratio :-


In the case of D the value thus obtained for the year is less than that for Midwinter, and only a little larger than that for Equinox; but in the case of $H$ the value for the year is larger than even that for Midwinter. This can only be ascribed to variability in the phase of the diurnal inequality, and this greater variability of phase no doubt is partly accountable for the larger excess of the R over the I values in the case of H .

The phenomena presented by 24 days from 12 consecutive months at a particular station may owe a good deal to "accident." Thus it would be a mistake to make sweeping deductions from the Christchurch data discussed here. I trust, however, that the sidelights that have been thrown on the nature of the diurnal variation will at least suggest lines of profitable investigation.
Table Ia.-Diurnal Inequalities.

Table Ma.-Number of Changes in Declination.

Table IIIA.-Number of Changes in Horizontal Force.

Table IVA.-Number of Changes in Vertical Force.

Table Va．－Hourly Ranges in Declination．（Unit $0^{\prime} \cdot 1$ ．）

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Table VIIA．－Diurnal Variation in Hourly Number of Changes．

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| $\stackrel{\circ}{\square}$ |  | $\stackrel{\square}{5}$ | $?$ | ： | $\bigcirc$ | $\stackrel{-}{\square}$ | $\begin{aligned} & \underline{0} \\ & \dot{\theta} \end{aligned}$ | $\begin{aligned} & 4 \\ & 0 \\ & 0 \end{aligned}$ | $\begin{aligned} & \infty \\ & 0 \\ & 0 \end{aligned}$ |  | $\stackrel{\square}{-}$ |
| 0 |  | $\stackrel{7}{\square}$ | $\cdots$ | $\stackrel{¢}{8}$ | $\cdots$ | $\stackrel{\dot{8}}{\circ}$ | $\begin{aligned} & \stackrel{0}{2} \\ & \stackrel{y}{2} \end{aligned}$ | $\begin{aligned} & 0 \\ & 0 \\ & 0 \end{aligned}$ | $\begin{aligned} & \text { N } \\ & \underset{\sim}{\sim} \end{aligned}$ |  | $\stackrel{\infty}{-}$ |
| $\infty$ |  | $\infty$ | $\stackrel{\square}{6}$ | $\stackrel{\text { ® }}{\substack{\text { d }}}$ | $\stackrel{3}{0}$ | $\begin{aligned} & \infty \\ & \vdots \end{aligned}$ | $\begin{aligned} & \stackrel{P}{9} \\ & \text { in } \end{aligned}$ | $\stackrel{\square}{\square}$ | $\stackrel{\varrho}{\rightrightarrows}$ |  | $\stackrel{\infty}{\sim}$ |
| ＊ |  | $\bigcirc$ | ？ | － | $\begin{aligned} & \approx \\ & \# \end{aligned}$ | $\stackrel{\infty}{\text { ¢ }}$ | $\square$ | $\underset{\sim}{\cong}$ | $\stackrel{H}{i 0}$ |  | $\stackrel{\square}{-}$ |
| $\pm$ |  | $\stackrel{7}{5}$ | $\cdots$ | ¢ | $\vec{~} \overrightarrow{=}$ | $\stackrel{\oplus}{\oplus}$ | $\overbrace{\infty}^{\infty}$ | $\underset{F}{5}$ | $\underset{\sim}{ \pm}$ |  | $\stackrel{-}{-}$ |
| $๙$ |  | $\stackrel{8}{=}$ | $\overrightarrow{\dot{\phi}}$ | $\begin{aligned} & \stackrel{8}{2} \end{aligned}$ | © | $\begin{aligned} & 8 \\ & \underset{\sim}{4} \end{aligned}$ | $\begin{aligned} & \because \\ & = \end{aligned}$ | $\stackrel{\square}{2}$ | $\stackrel{4}{0}$ - |  | $\stackrel{\infty}{-}$ |
| $\dot{\square}$ |  | 8 | 4 | － | $\stackrel{-}{\circ}$ | $\stackrel{7}{9}$ | $\begin{aligned} & \Gamma \\ & \text { N } \end{aligned}$ | \％ | ¢ |  | $\stackrel{-}{-}$ |
| $\cdots$ |  | $\cdots$ | $\stackrel{0}{6}$ | $\cdots$ | $\stackrel{\square}{6}$ | \％ | $\begin{aligned} & \stackrel{̣}{1} \\ & \underset{\sim}{1} \end{aligned}$ | $\overrightarrow{~ ت}$ | ت |  | － |
| ब่ |  | ＋ | $\stackrel{\cong}{\vdots}$ | $\begin{aligned} & \infty \\ & \stackrel{\infty}{\sim} \end{aligned}$ | $\begin{aligned} & \vec{~} \\ & \underset{\sim}{n} \end{aligned}$ | $\stackrel{0}{\underset{\sim}{\circ}}$ | $\stackrel{\infty}{\infty}$ | $\stackrel{7}{-9}$ | $\underset{\sim}{\infty}$ |  | $\stackrel{-}{-}$ |
| $\cdots$ |  | － | $\stackrel{\oplus}{-1}$ | $\begin{aligned} & 0 \\ & i \\ & i \end{aligned}$ | $\begin{aligned} & \dot{0} \\ & \dot{-} \end{aligned}$ | － | $\cdots$ | $\stackrel{\rightharpoonup}{\text { in }}$ | $\begin{aligned} & -1 \\ & \text { in } \end{aligned}$ |  | $\stackrel{4}{4}$ |
|  | Dechination． | $\stackrel{\text { L゙ }}{\substack{\text { E }}}$ |  |  |  | $\begin{aligned} & \stackrel{\circ}{*} \\ & \text { * } \\ & \text { * } \end{aligned}$ |  | $\begin{aligned} & \text { K } \\ & \text { 荡 } \\ & \text { 苗 } \end{aligned}$ |  | 'mozoly tyonitax | $\begin{aligned} & \text { d } \\ & \text { d } \end{aligned}$ |

Table VIIIA．－Hourly Ranges（R）and Hourly Increments（I）of Declination（Diurnal Variation）．

| Hour ending |  | 1. | 2. | 3. | 4. | 5．， | 6. | 7. | 8. | 9. | 10. | 11. | 12. | 13. | 14. | 15. | 18. | 17. | 18. | 19. | 20. | 21. | 22. | ${ }^{23 .}$ | 24. | Mean． |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Year ．．．．\｛ | R | （ $\begin{gathered}1.74 \\ +1.74\end{gathered}$ | 1 131 | $\left.\begin{array}{\|c\|} 1 \\ 0.61 \\ +0.27 \end{array} \right\rvert\,$ | $\left(\left.\begin{array}{c} 1 \\ 0 \cdot 76 \\ -0.87 \end{array} \right\rvert\,\right.$ | $\begin{aligned} & 1 \cdot 07 \end{aligned}$ | $\int_{-0.71}^{1}$ | $\begin{array}{c\|c\|c} 0.62 \\ 0 & 0.62 \\ \hline-0.45 \end{array}$ | 10.47-0.24 | ${ }_{0 \cdot 46}$ | $c_{1}^{0.41} \begin{gathered} 0.20 \end{gathered}$ | $\begin{aligned} & 1 \\ & 0.55 \\ & \hline \end{aligned}$ | $0_{0}^{\prime} 50$ | $0 \cdot 44$ | ${ }_{0.35}^{\prime}$ | ${ }_{0}^{\prime}$ | $\sigma_{0.27}$ | $0.37$ | ${ }^{\circ} \mathrm{O} 50$ | ¢ $0 \cdot 79$ | ${ }_{0.93}^{\prime}$ |  | $\begin{gathered} \prime \\ 0 \cdot 80 \end{gathered}$ | $\left\|\begin{array}{c} 1 \cdot 3 \\ 1 \cdot 33 \end{array}\right\|$ | 1－ts | $\overline{1}$ |
|  |  |  | ＋1．22 |  |  | －1．08 |  |  |  | －0．14 |  | －0．27 | $+0 \cdot 17$ | ＋0．04 | ＋0．10 | ＋0．13 |  |  | －0．38 | －0．70 | －0．85 | －0．76 | ＋0．08 | ＋1．06 | ＋1．87 | 0.546 |
|  | R | 20 | 1－14 | 0.53 | $0 \cdot 55$ | 0.96 | 0.48 | \％ | 0.39 | 0.51 | 0.39 | 0.35 | 0.29 | 0.40 | 0.46 | － 21 | 0.21 | 0.23 | 0.26 | 0.29 | 0.23 | 0.54 | 0.80 | 0.50 | 0.90 | $0 \cdot 498$ |
|  | 1 | ＋1／19 | $+1.01$ | $+0.34$ | －0．40 | －0．92 | －0．34 | －0．18 | －0．21 | ＋0．0．3 | －0．16 | －0．16 | ＋0．03 | ＋0．26 | ＋0．05 | ＋0．03 | ＋0．15 | －0．02 | －0．01 | －0．11 | －0．09 | －0 | －0．32 | －－ 30 | ＋0．86 | 0.328 |
|  | R | － 10 | $1 \cdot 24$ |  | 90 | 1.07 | 0.79 | 0.48 | $0 \cdot 49$ | 0.39 | 0.45 | 0.61 | 0.60 | $0 \cdot 45$ | $0 \cdot 29$ | 0.24 | 0.21 | 0.23 | 0.26 | 62 | 111 | 1.08 | 0.76 | 1.32 | $2 \cdot 09$ | 0.764 |
|  | 1 | 9 | 16 | －0．03 | －0．80 | －1．04 | －0．69 | $-0.16$ | －0．15 | 23 | －0．34 | －0．27 | ＋0．29 | －0．07 | ＋0．16 | －0．04 | ＋0．05 | ＋0．01 | －0．17 | －0．80 | －1．05 | －0 | －0．16 | ＋1．31 | ＋2．24 | 0.5 |
|  | ${ }^{\text {R }}$ | 4 | 59 | 0.74 | $0 \cdot 84$ | 1.19 | － 19 | 1114 |  | $0 \cdot 49$ | － |  | 0.63 | $0 \cdot 47$ | 0 | 0.67 | 41 | 0.70 | 1.06 | 1.56 | 1.54 | 0.89 | 1.06 | $2 \cdot 27$ | 8.69 | 1.037 |
|  | 1 | ＋1．94 | $+1 \cdot 56$ | ＋0．51 | －0．83 | －1．16 | $-1 \cdot 16$ | －1．07 | －0．37 | －24 | －0．11 | －0．39 | ＋0．20 | －0．09 | ＋0 | ＋0．41 | －0．21 |  | －1．04 | －1．47 | －1．50 | －0．81 | ＋1．06 | ＋2．31 | ＋2．57 | 0.883 |

Table IXA．－Hourly Ranges（R）and Hourly Increments（I）of Horizontal Force（Diurnal Variation）．（Unit $1 \gamma$ ．）

| 盶 | \％\％ | $\cdots$ | \％ | $\overbrace{0}^{3}$ |
| :---: | :---: | :---: | :---: | :---: |
| $\stackrel{3}{8}$ | \％ | 응 | － | $\stackrel{3}{\square}$ |
| 約 | － | $\begin{array}{ll} 0 \\ i x \\ i \end{array}$ | $\cdots$ | $\because$ |
| ชั่ | $\begin{array}{r} \because \\ \because i \\ i \\ \hline \end{array}$ | $\stackrel{x}{6}$ |  | $\begin{aligned} & \circ \stackrel{\circ}{\dot{B}} \underset{\vdots}{\vdots} \\ & \hline \end{aligned}$ |
| ล่ | io | $\stackrel{\circ}{\dot{\circ}}$ | $\cdots$ | ¢ ¢ |
| ¢่ | －\％ | $\stackrel{\circ}{-}$ | $\stackrel{\square}{9}$ | － |
| $\stackrel{\text { ¢ }}{ }$ | $\cdots$ | 을 | $\cdots$ | 为 |
| $\stackrel{\sim}{\sim}$ | \％$\%$ |  | $\stackrel{\square}{\square}$ | $\cdots$ |
| $\stackrel{\square}{\square}$ | $\bigcirc$ | － | 皆 | － |
| $\stackrel{\oplus}{9}$ | ¢ | $\stackrel{\oplus}{\stackrel{\oplus}{+}}$ | 范 | \％ |
| $\stackrel{\square}{\square}$ |  | $\vec{\therefore} \underset{i}{0}$ | 范 | 的 |
| $\pm$ | － | $\stackrel{\square}{\square}$ | － |  |
| ¢ | $\cdots$ | $\bigcirc$ | 它 | 号 |
| ® | － | $\overrightarrow{\mathrm{n}} \mathrm{i}$ | $\because$ | ¢ |
| $\cdots$ | $\stackrel{\text { ¢ }}{\substack{\text { ¢ } \\++++ \\ \hline}}$ |  | 为 | $\cdots{ }_{\square}^{4}$ |
| $\stackrel{\square}{\square}$ | ¢ | 范 | ¢ | $\stackrel{\square}{\text { ¢ }}$ |
| $\cdots$ | － | $\stackrel{\square}{\circ} \mathrm{B}$ | － | \％ |
| $\infty$ | $\stackrel{\circ}{\circ}$ | ¢ | \％ | ¢ |
| $\therefore$ | \％\％ | $\stackrel{¢}{6}$ | 简 | － |
| $\dot{\sim}$ | $\cdots$ | － | $\bigcirc$ | $\stackrel{7}{\square}$ |
| $\stackrel{\circ}{\circ}$ | ¢ | ¢ | 管 | ¢ |
| $\div$ | 芯笭 | － | ※\％ | ¢ |
| $\cdots$ | $\begin{array}{r}\text { ¢ } \\ \stackrel{\circ}{\circ} \\ + \\ + \\ \\ \hline\end{array}$ |  | $\circ$ <br>  <br> 0 <br> + <br> + | $\stackrel{\text { \％}}{\text { ¢ }}$ |
| ค |  | $\begin{gathered} 40 \\ 0 \\ 0 \\ 0 \\ + \end{gathered}$ | － $\begin{array}{r}\text { ¢ } \\ \stackrel{\circ}{\dot{\circ}}+ \\ + \\ \hline\end{array}$ |  |
| $\cdots$ |  | $\stackrel{3}{i}$ | $\stackrel{8}{3}+$ + + + | － |
| － | $\pm \square$ | $\pm$ | ＊－ | $\stackrel{\square}{*}$ |
| 最 量 点 | $\underbrace{\prime}$ |  |  |  |

## APPENDIX B.

An Examination of Antarctic disturbances from October, 1902, to March, 1903, which are simultaneous with Arctic disturbances discussed by Prof. Kr. Birkeland.
§1. After the discussion of the Antarctic magnetic data on pages 73 to 200 had been completed, there appeared a large volume by Prof. Kr. Birkeland* which contains a great mass of information respecting magnetic disturbances in the Arctic, from October, 1902, to March, 1903, which were contemporaneous with the observations discussed in this volume.
Prof. Birkeland had the following four Aretic stations, all provided with self-recording magnetographs:-

| Station, | Latitude $\mathbf{N}$. |  | Longitude. |  |
| :---: | :---: | :---: | :---: | :---: |
|  | - | , | - | , |
| Axelöen (Spitzbergen). | 77 |  | 14 | 50 E |
| Matotehkin Schar (Nova Zembla) . |  | 17 |  | 57 E |
| Karjord (Finmark) . . . . . . . . |  |  |  | 58 E |
| Dyrafjord (Iceland) . . . . . . . . |  |  |  | 30 W |

Guided by the records from these stations and from Potsdam, Prof. Birkeland made a list of disturbances and issued a circular to magnetic observatories requesting copies of the records obtained on the days specified on his list. He thus became possessed of records of disturbances at 25 stations, including the four mentioned above and the following:-

In Europe-Bossekop, Pawlowsk, Stonyhurst, Wilhelmshaven, Potsdam, Kew, Val Joyeux, Munich, Pola, San Fernando ; in Asia-Tiflis, Zi-ka-wei, Dehra Dun, Bombay, Batavia; in N. America-Sitka, Toronto, Baldwin, Cheltenham; in the Pacific-Honolulu and Christchurch. The names in each group are in order of latitude from north to south.
Birkeland reproduces the disturbed curves in 21 plates, each dealing with a disturbed period varying in length from 2 to 20 hours. Some of the observatories supplied no Vertical-Force curves, and only a few supplied material for all the magnetic storms. Still, the plates represent what is probably the most extensive series of contemporaneous magnetic data that has yet been published. In addition to the plates, the volume contains over 160 charts representing the results which Birkeland has deduced for the disturbing forces at the different stations, at different stages of the 21 storm periods. These charts are based on elaborate measurements of the magnetic curves and represent a large amount of work. Besides discussing the charts, Birkeland deals with experiments which he has made with a miniature Earth, or terrella, magnetised and exposed to kathode-ray discharges in a high vacuum. The experiments are intended to serve as an auxiliary to the elucidation of the causes that produce magnetic disturbances. The discussion of observations, experiments and theory occupies more than 300 large quarto pages.
§2. As already mentioned, the discussion originally contemplated of Antarctic magnetic storms had been entirely completed when Prof. Birkeland's volume came into my hands in February, 1909, and I was exceedingly reluctant to re-open the subject, considering the long time that had already elapsed since the work was begun. It was accordingly decided that no change should be made in what had been already written, and pages 73 to 200 are thus absolutely unaffected by any results or views in Birkeland's $\dagger$ volume. When Prof. Birkeland's conclusions and mine harmonize, the harmony thus merits increased consideration; when they differ, the difference at least owes nothing to prejudice.

* 'The Norwegian Aurora Polaris Expedition, 1902-1903,' Volume I.
† The reference in Chapter IX is to Prof. Biregland's earlier work, 'Expédition Norvégienne de 1899-1900.'

After having read Prof. Birkeland's volume, I came, somewhat reluctantly I must confess, to the conclusion that those responsible for the production of the present volume would be open to criticism if the opportunity afforded for contrasting Arctic and Antarctic records were not utilised. This appendix has accordingly been written with the object of supplying information as to what was happening in the Antarctic during the disturbed times selected by Prof. Birkeland. Its primary object is to inform, not to criticise, and if it contains anything that savours of criticism, this is mainly in order to explain why Prof. Birkeland's exact procedure has not been followed.
§3. The method adopted by Birkeland in dealing with magnetic disturbances is fundamentally different from that adopted in Chapters IX and X. Practically following Sabine, he defines disturbance at any given hour as the difference from the normal for that hour ; and what his tables give, and his charts illustrate, are the values of disturbances so defined at different stages of each disturbed period. He believes that the canse of the disturbance is an electrical discharge, or a series of electrical discharges, in the Earth's atmosphere, such as he has succeeded in producing in the vicinity of his terrella, and his ultimate object apparently is to calculate the position and intensity of these discharges.

One great difficulty, as I remarked in $\$ 90$, Chapter IX, is the fixing of a normal value. In doing this Brrkeland seems to have been materially helped by the fact that at some observatories, including those provided with Kew-pattern magnetographs, it is usual to have two days' curves on the same photographic sheet. If the one day's curve happens to be undisturbed, its form greatly assists the eye in deciding as to the nature of the disturbance in the other. This is an advantage which I have often had occasion to sppreciate myself. The Arctic stations, however, and some of the others had magnetographs of the Eschenhagen pattern, like those of the "Discovery," and there must have been considerable difficulty at times, as Birkeland himself allows, in deciding what the departure from the normal really was. This difficulty had probably a considerable indirect influence on Brakeland's choice of disturbance periods. Those he has selected are largely represented at Kew and other non-polar stations by "bays" of comparatively short duration.

What is meant by a "bay" will be readily grasped by reference to the accompanying figure. The

continuous line ABCDEFGH represents an imaginary magnetic curve having two bays, one, BCD , occurring at a time when the regular diurnal variation is slow, the other, EFG, occurring at a time when it is rapid. The ordinates CM, FN drawn on to the base line represent the excess above the constant base-line value of the values of the magnetic element answering to the times $M$ and $N$. The broken lines $\mathrm{BC}^{\prime} \mathrm{D}, \mathrm{EF}^{\prime} \mathrm{G}$ are intended to represent the imaginary undisturbed curve, and the intercepts $\mathrm{CC}^{\prime}, \mathrm{FF}^{\prime}$ on the ordinates represent from Birkeland's point of view the disturbances.

In such a case as that represented by the figure the method appears simple, especially when the disturbance occurs at a time when the diurnal change is slow. In practice, however, there is usually a difficulty in deciding where the "bay" begins or ends, and the relative position of its two extremities is usually not quite what one would anticipate from the trend of the curve prior to its commencement. This latter difficulty naturally increases the longer the duration of the bay, and the more rapid the regular diurnal inequality changes at the hour.

As a general rule, during really active magnetic disturbances, whilst bays of a kind are not infrequent, the curve adjacent to them is itself disturbed and sinuous, and affords very imperfect guidance as to where the normal curve would come. On days of real disturbance, one is usually obliged to have recourse to the
curves of adjacent days, if quiet, or to a regular diurnal variation derived from a number of days. The latter alternative is, of course, the more satisfactory theoretically ; but there is this difficulty, which I have dealt with elsewhere, that the regular diurnal.inequality which one gets depends on the nature of the days from which it is derived. For instance, taking means from 11-year results at Kew, the departure from the mean Declination for the day at $2 \mathrm{p} . \mathrm{m}$. on a representative January day is :-

$$
\begin{aligned}
& +2^{\prime} \cdot 21 \text { on the Astronomer Royal's selected quiet days, } \\
& +2^{\prime} \cdot 66 \text { on ordinary undisturbed days, } \\
& +4^{\prime} \cdot 88 \text { on the average highly disturbed day. }
\end{aligned}
$$

For simplicity, we may suppose the secular change and sun-spot influence non-existent, though in reality these are complications which have to be reckoned with. Let us suppose that on a certain disturbed day in January the Declination at 2 p.m. departs from the mean for the month by $+3^{\prime} \cdot 66$. Is the "disturbance" $+1^{\prime} \cdot 45,+1^{\prime} \cdot 00$, or $-1^{\prime} \cdot 22$ ? The $+4^{\prime} \cdot 88$ departure from the daily mean on the representative disturbed day in January, be it noted, represents a regular diurnal inequality, or at all events something which we have no present means of distinguishing from a regular inequality. Its excess over the ordinary day value may, of course, indicate a tendency for a particular phase of disturbance to occur at a particular hour of the day, but it may mean that the causes operative in producing the regular diurnal inequality are for some reason-e.g., increased conductivity in the upper atmosphere-more effective on disturbed days than others.

This source of uncertainty is equally present when the normal curve is derived, as Birkeland seems to have derived it, by reference to a day or days adjacent to the disturbed one. The adjacent days may be of the very quietest type, or may themselves be considerably disturbed.

There is a further complication in that disturbance has sometimes a tendency to be associated with a temporary alteration in the value of magnetic elements, which disappears gradually like after-strain in a metal. Horizontal Force, for instance, is sometimes very considerably depressed for days after a really large disturbance. Thus it may make all the difference in the world to one's decision as to whether at a particular hour the element is above or below the normal value, if one happens to take for comparison the day after a disturbance rather than the day before. In one instance I observe that Prof. Birkeland noticed the occurrence of this precise source of uncertainty.

In the case of the Antarctic curves, the uncertainties which exist elsewhere are mostly much enhanced. We have seen in Chapter III, Tables XII and XIV, that disturbance exerts an unusually marked influence on the amplitude of the regular diurnal inequality. The Antarctic D and H curves were always sensibly disturbed, and the disturbance was usually sufficient to obscure, if not to totally conceal from the eye, the trend of the natural diurnal variation. The diurnal variation of V would very likely have been readily recognisable if the scale-value had been as low as $5 \gamma$ throughout, and the temperature coefficient had been small; but, as matters stand, mere optical inspection of the curve tells merely whether disturbance has been specially active or not. The days of Prof. Birkeland's choice lay entirely in the months October, 1902 , to March, 1903, and so came in the Antarctic Summer, when the magnetic variations and the temperature changes in the hut were both at their maximum.
§4. There is a rather different kind of uncertainty to which Birkeland refers more than once, e.g. p. 64, viz, the uncertainty in the estimate of the hour at which a particular movement took place. When one is dealing with stations not too far apart, especially when furnished with magnetographs of the same type -such, for instance, as Stonyhurst, Falmouth and Kew--this source of uncertainty may usually be largely eliminated, provided there are any rapid movements which are shown at all the stations. The curves from the different stations are sufficiently alike to render identification of corresponding points easy. The accuracy of one's identification of any one peak can be verified by reference to other peaks. Thus while one cannot decide what the error actually is in the times shown on the curves of any one of the stations, one can usually determine approximately the difference between the errors at the different stations, and so make sure that the comparisons made refer to one and the same absolute time within a minute or so.

When we have rounded curves like that on p. 247 representing a slowly varying disturbance, a mistake even of several minutes is unlikely to make much difference in one's estimate of a disturbance, provided
one selects a time when the phenomenon is near its maximum. This consideration most likely influenced Birkeland's choice of disturbances. If a bay or bays such as those on p. 247 occurs in the curves of one European station, it usually occurs simultaneously at other European stations, and in the cases selected by Birkeland this was true not merely of the European stations, but also to a greater or less extent of the other non-polar stations. Thus when the bay answered to a deep depression on the Aretic curves -and not infrequently trace was lacking from one or more of the Arctic stations, so that only two or three had to be considered-uncertainties of time may have been without any very serious influence on Birkeland's conclusions. Another circumstance favourable to his measurements was that all relate to what was Winter at nearly all his stations, while a large majority relate to what were night hours at the Arctic and European stations. His 21 plates deal with a total of 207 hours, and of these 104, or one half, fall between $6 \mathrm{p} . \mathrm{m}$. and $2 \mathrm{a} . \mathrm{m}$. , G.M.T. Uncertainties from the diurnal inequality were thus much reduced. When, however, one turns to the Antarctic curves which correspond to Birkeland's, one is met by the converse of all this. The season is the Antarctic Summer, when changes, regular and irregular, are largest. The time in a large majority of cases falls in the Antarctic morning, often during the hours when, as already explained, oscillatory movements were especially numerous and rapid. A mistake of a few minutes in the time may make a huge difference to the result, and the curves are so unlike those at any of Birkeland's stations that no help is usually forthcoming from the identification of peaks. If Arctic records exist synchronous with the occurrences of the "special type of disturbance" dealt with in Chapter X—which took place in the Antarctic Winter-anyone who attempts to apply Birkeland's method to them will appreciate my difficulties.

After carefully comparing the Antarctic curves with Birkeland's plates-in which the time-scale of the originals is usually much reduced-I decided that the application of Birkeland's method was absolutely impossible in a number of cases, and that, in general, the uncertainties attending it were too great to justify the labour necessary. It will, however, I think, be found that the comparison which it proved possible to make has led to results of no small interest.
§5. Professor Birkeland believes that he has succeeded in recognising several distinet types of disturbance, to which he attaches specific names. The following table gives particulars as to the date and duration of the disturbed periods dealt with in his 21 plates. It also gives the total range in D and H at Kew during the time covered by each plate. Hours are counted from 1 to $24-0$ or 24 signifying Greenwich midnight-that being the plan adopted by Bireeland. An entry such as October 11-12, hours 12-2, means that the time extended from 12 (i.e. noon) on October 11 to $2 \mathrm{a} . \mathrm{m}$. on October 12. The range recorded for Kew represents the difference between the highest and lowest values of the element during the hours covered by Birkeland's plate. It generally owes a good deal to the regular diurnal variation natural to the period of the day, except in cases where only night hours are included.

§6. The principal object in mentioning the ranges at Kew is to bring out an important point which I hardly think Birkeland himself quite realised, and which I am confident will not be realised by readers of his volume who are not themselves experts in Terrestrial Magnetism.

The mean value of the absolute daily range in D at Kew derived from all days of the eleven years 1890 to 1900 was $13^{\prime} \cdot 57$. This value, it will be observed, was exceeded on only six occasions in the table. During the eleven years the absolute daily range at Kew exceeded 20'-a value not once attained during Birkeland's disturbances-on no less than 12 per cent. of the total number of days.

It must not, of course, be forgotten that the average length of the period covered by one of Birkeland's plates is slightly under 10 hours, and that the majority of the periods do not include the hours at which the daily maximum and minimum most frequently occur. Still, taking everything into account, the fact remains that the great majority of the days selected by Birkeland were not what are ordinarily called disturbed days. In the Aretic, it is true, there were movements which would rank at Kew or any other non-polar station as magnetic storms, but there is not a single one of the occasions on which the phenomena at Kew would ordinarily be dignified with that name. On perhaps three occasions, October 31, to November 1, 1902, November 23-24, 1902, and February 8, 1903, one would have little hesitation in describing the day as disturbed, but on the other hand there is quite a considerable proportion of the days which one would be likely to describe as quiet. It is not merely that the movements on Birkeland's selected days were small, but that they were few in number, and in many cases represented slow changes. In the case of an ordinary magnetic storm at Kew, not only would the range be much larger than in any of the days selected by Birkeland-a range of $30^{\prime}$ in D and $300 \gamma$ in H represents what may be called a second class storm-but the large movements would be much more numerous and some of them much more rapid. If we take a really first-class storm, like that recorded at Kew on October 31, 1903, it represents an altogether different order of conditions. Not merely is the range five or six times larger in D , more in H , but there are dozens of rapid oscillations, altogether without parallel in the most disturbed cases selected by Birkeland. The expenditure of energy during a first-class storm may, for all we know to the contrary, be 10,100 , or even 1000 times greater than that during Birkeland's most disturbed day, and we cannot even say with certainty that the ultimate source of the energy, or the way in which it is expended, is the same in the two cases. What I have called a first-class storm is apparently experienced as a large storm over all the world-or at least over a very large part of it-and is invariably, or almost invariably,
accompanied by aurora visible even in the south of England, whereas the disturbances treated by Birkeland attained tho development of a first- or a second-class storm if anywhere only in a portion of the polar regions. In the stations in temperate latitudes they were mostly of the size one meets with every other day, and if there were any auroral discharge accompanying them, it was not visible outside the region where auroral frequency is high.
§7. Returning to the table particularising Birkeland's plates, we have now to consider the significance attached to the descriptive torms applied to the disturbances.

Explaining the term "equatorial," on p. 62, Birkeland says:
". . . it is not unusual to find perturbations that are best developed and most powerful at the Equator. It has even been found that these perturbations in the regions about the Equator act principally upon the horizontal intensity . . . Such perturbations we . . . call equatorial . . . Of these there are . . . two kinds . . . , such as produce an increase in the horizontal intensity, and such as produce a diminution . . . The first . . . we have called positive . . . , the second . . . negative equatorial perturbations."

Of polar elementary storms Birkeland says, pp. 84-85:
(1) "They are comparatively strong at the poles (meaning the north polar regions). The simultaneously perturbing forces, even as far north as the 60th parallel, have already sunk to about a tenth of their strength in the auroral zone.
(2) "They are of short duration, frequently lasting not more than 2 or 3 hours.
(3) "The conditions before and after are comparatively quiet.
(4) "The oscillations at the (north) polar stations, especially the more southern ones, run a simple course. At the poles, they are often characterised by a simple increase to a maximum, and decrease to zero. We may sometimes, even at the northern stations, have to some extent an undulating form, answering to a slow turning of the perturbing force."
Of compound perturbations no general definition seems to be given. Judging by individual cases they are a combination of phenomena, "equatorial perturbations" predominating at one stage, and "polar elementary storms" at another.

Birkeland's discussion of "cyclo-median" storms on p. 144 is somewhat lacking in clearness. After expressing his belief that "electric cyclones, wandering over the Earth's surface," according to a suggestion of Dr. Ad. Schmidt, are a very rare phenomenon at least in large storms, he adds: "It appears, however, that there is a class of perturbations that are due to current-systems which appear in lower latitudes at a height above the Earth that is small in proportion to the Earth's dimensions . . . In the whole of our material, we have not found more than one considerable perturbation that in its entirety must be due to systems that come near to the Earth in lower latitudes."

It is this single occurrence (October 6, 1902) that is characterised as "cyclo-median." Judging by a remark op p. 150, the term was intended to signify that the disturbance was "as great in medium as in high latitudes," and that the electrical currents to which it was due were "vortical in form."
§8. It is difficult to discriminate between Prof. Birkeland's observations and his theories, as the two are so interwoven in his pages. Thus some reference to his theoretical views may tend to clearness. He believes that the "equatorial" perturbations are due to electric currents encircling the Earth near the plane of the magnetic Equator, at a distance above the Earth which is similar to, possibly greater than, the Earth's radius. A current thus situated in a plane perpendicular to the Earth's magnetic axis would naturally give a force which in the equatorial regions would be roughly in the magnetic Meridian-thus affecting H almost exclusively-and the intensity would be greater in the Equator. Unless the height of the current were large, the disturbance would fall off rapidly as we departed from the Equator, and in even low latitudes there would be a large vertical component.

There does not seem to be in the volume any close comparison of the amplitude of the "equatorial" perturbations experienced on the same occasion at different stations; but the numerical data as to the disturbances during individual equatorial perturbations do not show such predominance in the equatorial regions as the definition leads one to expect.

So far as I can judge, a considerable number of the movements discussed in our Chapter IX, including
the "sudden commencements," would be classified by Bireeland as "equatorial" perturbations. In their case, however, as will be remembered, the amplitude at Colaba and Mauritius was usually less than that at Kew, and much less than in the synchronous movements seen in the Antarctic.

Bhreland believes an "elementary polar"* storm to be due to what he calls "precipitation" in a comparatively limited Arctic area. By this he apparently means an influx (and efflux) of charged ions. In some calculations in the volume this is treated as equivalent to an electric current, approaching and receding from the Earth in lines which, if produced, would intersect at the Earth's centre. The current is regarded as stopping short of the Earth, then travelling in a straight line to a point equidistant from the surface, and finally receding. In some special cases on p. 103 -which seem to be supposed to represent probable actual conditions-the height of the connecting (so-called "horizontal") portion is put at 200 or 300 kms , its length being taken as 1600 kms . and upwards. In the case of the "elementary polar" storms the disturbance at any given instant is large over only a very limited polar area. In temperate latitudes the disturbance is small and diminishes rapidly as the distance from the area of "precipitation" increases. This area usually keeps shifting its position, so that the disturbance travels across the polar regions.

If Prof. Birkeland had seen what has been called in Chapter $\mathbf{X}$ the "special type of disturbance" he would not unlikely have called it a south polar elementary storm. It seems, however, to present a much greater definiteness of type than Biriecand's "polar" storms, and its duration is usually much less than the two or three hours which Birkeland speaks of.
§9. In his volume Birkeland does not take the disturbances in chronological order when discussing them, but treats first the "equatorial," secondly the "clementary polar," thirdly the "cyclo-median," and finally the "compound." One finds, however, that there are few of the occasions on which Birkeland failed to detect at one stage or another the presence of both "polar" and "equatorial" disturbances. I have thus thought it simplest to follow the chronological order, as Birkeland himself has done in the case of the 21 plates at the end of his volume. In what follows, the references are to Birkeland's plates unless the contrary is explicitly stated, and the reader is strongly advised to have these plates before him while consulting the details given here.
§10. October 6, 1902 (hours 13 $\frac{1}{2}-15 \frac{1}{2}$, Plate I).
Of this "cyclo-median" disturbance, Birkeland, p. 150, says: "Its chief characteristics are that it is as great in medium as in high latitudes," also "the effect over the district Wilhelmshaven, San Fernando, Stonyhurst, Pola is of about the same magnitude." The difference from the "equatorial" disturbances, which otherwise it closely resembles, is that there appears hardly any movement in Asia or the Tropics. Bibkeland infers that it must be due to electrical currents at a height "small in proportion to the Earth's dimensions."

The phenomenon, as presented at Kew and the other stations where it was best developed, had for its most prominent feature a sudden change in D , commencing at about 14 h . 14 m . At Kew , in the course of 5 or 6 minutes, the Declination needle moved about $4^{\prime} .5$ to the west and then returned much more slowly towards its normal position, taking nearly 30 minutes to reach it. $H$ fell as $D$ increased, but the total fall was only about $6 \gamma$; the return to the normal position was slow as in the case of $D$. The curves were exceedingly quiet for some hours before and after this movement. While the amplitude of the movement was very trifling, the isolation of the D movement and its nature are certainly musual.

Whether there was or was not a corresponding movement in the Antarctic is open to doubt. In the Antarctic, as elsewhere, the day as a whole was exceptionally quiet; but, as was invariably the case in the Antarctic, both the D and the H curves show numerous small oscillations. The V curve was certainly not more disturbed between 14 h .10 m . and 14 h .50 m . than it was earlier in the day, and decidedly less disturbed than it was a few hours later. The same appears true of the D curve. There was, howeverit may be a merely chance coincidence-a distinct bay on the $H$ curve, whose inception was at least very nearly simultaneous with the commencement of the disturbance in Europe. Between 14h. 10 m . and $14 \mathrm{~h} .21 \mathrm{~m} . \mathrm{H}$ fell $9 \gamma$, two-thirds of the fall taking place between 14 h .13 m . and $14 \mathrm{~h} .18 \mathrm{~m} . \quad \mathrm{H}$ remained below its previous value until 15 h .23 m ., the return movement being much the slower.

* The terms "polar elementary" and "elementary polar" are applied indifferently.

This case is a good example of the uncertainties attending the application of Birkeland's method to the Antarctic curves, even under the most favourable conditions. From inspection of the adjacent portions of trace one would conclude that D and V were normal, but H distinctly below the normal value from 14 h .13 m . to 15 h .0 m . If, however, we compare the values of the elements at 14 h . 53 m . (otherwise 2 a.m., L.T., on October 7) with the means for the same hour from the four nearest days, this is what we find:-


These figures point to exactly the opposite conclusion to that suggested by the form of the curves themselves.
811. October 11-12, 1902 (hours 12-2, Plate II).

This "compound" perturbation is regarded by Birkeland, p. 251, as divisible into three so-called "sections," (i) from 11 h . to 17 h .20 m , and (ii) from 17 h .20 m . to 18 h .30 m . on October 11, (iii) from the last-mentioned hour to 0 h .30 m . on the 12 th .

The disturbance in section (i) is regarded as "mainly a positive equatorial perturbation," accompanied, however, from 12 h .25 m . to 13 h .15 m . by a "considerable polar perturbation." "The farther we go," Birkeland says, p. 252, "from the (N.) polar regions, the less perceptible does this brief polar perturbation become . . . . At Zi-ka-wei and Dehra Dun it is distinctly noticed, at Batavia it is almost imperceptible. At Christchurch, on the other hand, there is a rather violent perturbation . . . . (which) cannot have been produced by the same system .... for the effect of the latter is imperceptible (?) even at Honolulu and Batavia. The explanation of this seems to be that simultaneously with the descent (of ions or corpuscles) in the north, a similar phenomenon appears near the South Pole, and it is the effects of the latter that we olserve at Christchurch."

This quotation has been given at length because it constitutes one of the very few references which I have observed to the possible existence of disturbance centres near the S-pole. The "explanation" was presumably purely a hypothesis on Birkeland's part, as he had no records from south of Christchurch.

Section (ii), 17 h .20 m . to 18 h . 30 m ., was characterised by "violent storms in the Arctic," especially at Matotchkin Schar, but "the effect of the equatorial storm is still perceptible."

Section (iii), 18 h .30 m. to 0 h .30 m ., "is characterised by a long polar storm," in the course of which, however, there appeared three short "intermediate" polar storms, the first with maximum about 18h. 34m., the second lasting from 20 h .45 m . to 21 h .20 m ., the third from 23 h .10 m . to 0 h .25 m .

At Kew and all the other non-polar stations whose curves appear in Plate II there was a fairly sudden commencing movement, which does not seem to be mentioned by Birkeland. The H movement at Kew commenced about 12 h .18 m , a fall of $2 \gamma$ and a rise of $10 \gamma$ occurring in about 6 minutes. A peak, representing a movement of about $l^{\prime}$ to the west, appeared also in the Kew D curve at about 12 h .24 m ., i.e. simultaneously with the maximum in H. Plate II shows this commencing movement distinctly at Colaba, Batavia, and Christchurch. It forms in fact the commencement of the movement at Christchurch, which Birkeland suggests may be due to currents near the S-pole.

At. Kew and the other non-polar stations the disturbance consists mainly of bays in the H and D curves, with two or three small but fairly sharp peaks, the greatest departures from the normal appearing between 17 h . and 22 h . Christchurch differs a little from the other non-polar stations in that there is a bay on the H curve between 12 h .18 m . and 14 h .10 m ., which commences less suddenly, but is larger than those encountered elsewhere; the later movements at Christchurch, on the other hand, are exceptionally small. The division into three sections seems somewhat arbitrary, especially the line of demarcation drawn at 18 h .30 m . between Sections (ii) and (iii). This comes in the middle of a very rapid rise
in H at Matotchkin Schar, which according to Plate II extended from abont 18 h .10 m . to 18 h .35 m , and which was immediately followed by a movement in the opposite direction, of so closely similar a character as to suggest its being an essential part of the same phenomenon. The significance of this will appear presently.

The Antarctic curves were failly quiet, according to the Antarctic standard, until about 17h., but the II curve shows a prominent bay from about 12 h .13 m . to 13 h .57 m ., the element being depressed and the minimum coming at about 13 h .18 m . The fall and rise were each about $60 \gamma$.

During part of the time covered by the H bay there was also a bay on the D curve, a rise of $58^{\prime}$ being followed by a fall of 45 , and the maximum coming at about 12 h .53 m .

As $1^{\prime}$ in D represents a force of about $1.9 \gamma$, the D movement was really the larger as well as the more rapid, but owing to the high sensitiveness of the $H$ magnetograph the $H$ movement appeals more to the eyc. During the occurrence of the bays on the D and H curves the V trace shows numerous small oscillations of an irregular character, bearing no obvious relationship to the D and H changes, and not so suggestive of the "special type of disturbance" as the H trace is.

These D and H movements are synchronous with the "polar" storm of Section (i). Also, whilst there is no conspicuously rapid initial movement, the time of commencement is at least very approximately the same as that of the small sudden movement seen at Kew and elsewhere.

After 17h. the Antarctic curves certainly deserve to be called disturbed. The H trace showed the following changes, superposed on which were the usual short-period smaller oscillations:-

| h. 11 m. to 17 h .16 m . fall 28h .16 m. |  |
| :---: | :---: |
|  |  |
|  |  |
|  |  |

D showed a number of small oscillations, but none conspicuously large; it increased, however, $90^{\prime}$ between 16 h .53 m . and 18 h .3 m ., going off the sheet at 18 h .3 m . for a few minutes. During this time V rose and fell only about $10 \gamma$, though there were numerous small oscillations. The above movements seem to be associated. They are synchronous with the earlier part of Birkeland's Section (ii), but also with the end of his Section (i). They are followed by a larger H movement.

Commencing to rise at about 18 h .21 m ., H, after increasing $74 \gamma$, got off the sheet at 18 h .27 m . Reappearing at 18 h .31 m ., in the next 29 minutes it fell $108 \gamma$, going beyond the limit of registration on the negative side. D commenced to fall at 18 h . 21 m ., when H began to rise, and in the course of 26 minutes fell $61^{\prime}$ and rose 70'. During this time V oscillations, though numerous, were small. The form of the Antarctic curves suggests that the phenomena from 18 h .21 m . to 19 h .0 m . were associated together, and the most natural inference seems to be that they form part of the disturbance which Birkeland regards as the first "intermediate" storm of his Section (iii). This was the time, it may be added, of one of the most prominent movements seen at Birkeland's co-operating stations.

The next movements in the Antarctic worth mentioning were fairly synchronous with Birkeland's second "intermediate" storm ( 20 h .45 m. to 21 h .20 m .) They are somewhat imperfectly shown, owing to lack of trace. Between 20 h .40 m . and $21 \mathrm{~h} .3 \mathrm{~m} . \mathrm{H}$ fell $45 \gamma$ and rose $50 \gamma$ (possibly more, as the trace is very faint and part may be invisible). The D trace, which had gradually got off the sheet on the positive side, suddenly re-appeared at 20 h .51 m . and fell $58^{\prime}$ in 6 minutes. After a minor oscillation it began to rise rapidly at 21 h .6 m ., rising $55^{\prime}$ before it again went off the sheet at 21 h .15 m . During this time there were some very rapid movements in $V$; the total range between 20 h .46 m . and 21 h .3 m . was $22 \gamma$.

There was no Antarctic trace from 23 h .20 m . to 23 h .45 m . on the 11 th. Between 23 h .48 m . and 0 h .31 m . on the 12 th a bay appeared in the V trace, the maximum depression being about $12 \gamma$. There is a synchronous bay in the D curve, a fall of $43^{\prime}$ being followed by a rise of $33^{\prime}$. The H trace was off the sheot practically all the time. The above D and V movements occur simultaneously with Birkeland's third "intermediate" storm. They are by no means of an outstanding character and do not appeal much to the eye. They are followed, however, by a relatively quiet time.
§12. October 23-24, 1902 (hours 17-5, Plate III).
This is described, p. 76, as "a positive equatorial perturbation. It commences suddenly at 19 h .11 m ., simultaneously all over the Earth." It is added, however, "About $1 \frac{1}{2}$ hours later, a polar storm... characteristic, simple, and well defined, appears around the Norwegian (i.e. Arctic) stations . . especially distinct at Matotchkin Schar." The date was not in Bireeland's list, so he obtained curves from only a few stations, including, however, Bombay and Dehra Dun.

The sudden commencement is shown clearly at all the stations, including the Arctic ones. At Turonto, and at Axelöen in the D curve, the commencing movement appears distinctly double, the principal movement being preceded by a smaller movement in the opposite direction. At Kew there is only a suggestion of an initial fall in H , but nothing certain prior to a sudden rise, amounting to $17 \%$ in the course of 3 or 4 minutes. After the summit was reached at 19 h . 14 m ., there was a gradual return to an undisturbed value at about 19 h .35 m . Simultaneously with the change in H there was a very small change in D , westerly Declination rising and falling about $0^{\prime} \cdot 7$.

In the Antarctic the curves had been unusually quiet for some hours when there suddenly began, at 19h. 10 m ., an exceedingly rapid rise in H . The movement is too rapid to be distinctly shown, but the trace seems to have gone off the sheet, remained off for 2 minutes, and returned to near its original position in 5 or 6 minutes. How much the oscillation exceeded that shown, $+40 \gamma$, then $-35 \gamma$, it is of course impossible to say. After slackening for a minute or two about $19 \mathrm{~h} .16 \mathrm{~m} ., \mathrm{H}$ continued moving in the same direction as before, to a peak at about 19 h .22 m . The fall in H since the curve came on the sheet was $77 \gamma$. Simultaneously with the commencing movement in $H$ there was an oscillation in $D$, a fall and rise each of $15^{\prime}$ taking place in about 4 minutes; this was followed by a second smaller oscillation.
The V trace, which had been exceedingly quiet, showed also a marked commencing movement, consisting of a rise of $6 \gamma$ in 3 or 4 minutes to a sharp peak at about 19 h .14 m ., followed by a fall of $26 \gamma$, occupying about 8 minutes. Halfway during the fall there was a nearly stationary position during about 2 minutes.
The natural conclusion unquestionably is that these large sudden movements in the Antarctic correspond to the smaller commencing movements which appeared simultaneously elsewhere.

The principal disturbance in the Arctic occurs between 21 h , and 23 h ., the maximum coming about 22 h .20 m ., but the Axelöen curves appear considerably disturbed until $3 \mathrm{a} . \mathrm{m}$. on the 24 th . The nonpolar curves in Plate III show only very trifling disturbances, the largest, between 21 h . and 23 h . on the 23 rd , being only of the same order as the sudden commencement at 19 h .11 m .

In the Antarctic the conditions remained highly disturbed from 19 h .10 m . until the traces got on the clamp at about 23 h .6 m . The D and H traces show incessant large oscillations. The largest movements recorded in H were a rise of $84 \gamma$ between 19 h .56 m . and 20 h .13 m ., and a rise and fall each of $67 \gamma$ between 22 h .51 m . and 23 h .0 m . The curve came on the sheet at 22 h .51 m . and went off 9 minutes later. Coming on again immediately, H rose $44 \gamma$ in a few minutes, the trace coming on to the clamp and so being lost.

In D there was a rise of $94^{\prime}$ between 22 h .18 m . and 22 h .25 m ., followed by a fall of $100^{\circ}$ during the next 10 minutes. Between 22 h .51 m . and 23 h ., synchronous with the large changes in H , there was a fall of $54^{\prime}$ and a rise of $87^{\prime}$, the latter immediately followed by a rapid fall of $135^{\prime}$, the trace coming on to the clamp before the movement was completed.

Amongst the larger V movements were a fall of $26 \gamma$ and a rise of $22 \gamma$ between 19 h .54 m . and 20 h .26 m ., a fall of 20 y between 22 h .26 m . and 22 h .33 m ., and a rise of $51 \gamma$ between 22 h .33 m . and 22 h .53 m . Between 22 h .53 m . and 23 h .6 m ., when the trace got on the clamp, V fell $24 \gamma$, rose $37 \gamma$, and fell $27 \gamma$.

The larger D and V movements, it will be noticed, occurred during the time when the Arctic stations were most disturbed; but the disturbance in the Antarctic was continuously large between 19 h .10 m . and 23 h .

The conditions in the Antarctic had become distinctly quieter by 0 h .11 m . on the 244 , when the next sheet was put on, but might fairly be described as disturbed until $3 \mathrm{~h}, 30 \mathrm{~m}$.

There was a small bay from 0 h .17 m . to 0 h .53 m ., the changes in the three elements being at least approximately simultaneous. D rose $25^{\prime}$ and fell $51^{\prime}$, just going off the sheet on the positive side; while V rose $16 \gamma$ and fell $13 \gamma$. The $\mathbf{H}$ trace was off the sheet on the positive side for 10 minutes and the rise,
$21 \gamma$, and fall, $14 \gamma$, actually shown were probably a good deal exceeded. After 0 h .53 m . the oscillatory movemonts in the V curve were much reduced. The D trace, however, showed two moderate bays. During the first-from 0 h .53 m . to 1 h .16 m .-D rose $33^{\prime}$ and fell $69^{\prime}$. During the second-from 1 h .16 m . to 3 h .23 m .-D rose $42^{\prime}$ and fell $70^{\prime}$. Between 1 h .39 m . and $2 \mathrm{~h} .17 \mathrm{~m} . \mathrm{H}$ fell and rose $39 \gamma$. The trace went off the sheet at 2 h .17 m . and did not reappear until 3 h .27 m . It went off and came on steeply, so there may have been a considerable movement in the interval.
\$13. October 27-28, 1902 (hours 14-1, Plate IV).
This "compound" disturbance is divided by Birkeland, p. 209, into two sections. The first section, from 14 h . to 20 h .30 m ., is regarded as composed of a long storm, largest on the whole in the Equator, during which there is an "intermediate" storm, most powerful in the Arctic, especially at Axelöen and Matotchkin Schar, which lasted from about 15 h .30 m . to 16 h .45 m .

At Kew, which was fairly representative of non-Arctic Europe, the most prominent phenomenon of Section (i) was a bay in the D curve from about 15 h .30 m. to 16 h .55 m ., the element being depressed. The greatest departure from the normal value was about $7^{\prime}$ and occurred about 16 h .20 m . The corresponding H movement was a fall of $22 \gamma$ from 15 h .15 m . to 16 h .5 m ., interrupted by two small recoveries, and a rise of $19 \gamma$ from 16 h .5 m . to 16 h .30 m ., followed by a smaller fall.

Birkeland's Section (ii), from 21h. 40 m . to about midnight, consisted of a "polar" storm, largest at Axelöen. A table on p. 212 gives particulars of its beginning and end and also as to the time of occurrence and the value $P_{1}$ of the largest disturbance in the horizontal plane. The commencement is about 21 h .40 m . and the hour of maximum about 22 h .50 m . at most stations; the end varies from 23 h .20 m . on the 27 th to 0 h .20 m . on the 28 th . $\mathrm{P}_{1}$ varies from $265 \gamma$ at Axelöen to $4 \gamma$ at Dehra Dun, the value at Kew, $29 \gamma$, being slightly above the average for non-Arctic European stations.

The general nature of the disturbance during Section (ii) outside the Aretic is fairly represented by the phenomena observed at Kew. D there was distinctly depressed from 21 h .40 m . on the 27 th until about 0 h .10 m . on the 28 th . The most rapid change was a fall of $3^{\prime}$ between 22 h .20 m . and $22 \mathrm{~h} .45 \mathrm{~m} . \mathrm{H}$, on the other hand, was distinctly above the normal value from 21 h .40 m. to 23 h .20 m . The maximum occurred about 22 h .54 m ., and the most rapid change was a fall of $20 \gamma$ between that hour and 23 h .34 m .

The Antarctic curves during the time covered by Plate IV show a moderate amount of disturbance, but nothing, perhaps, that would naturally attract attention. During the "intermediate" storm of Bireeland's Section (i) there was a bay on the D curve between 15 h .29 m . and 16 h .53 m ., the element rising $47^{\prime}$ and falling $33^{\prime}$. The V trace showed numerous oscillations, but the largest only $3 \gamma$ or $4 \gamma$ in amplitude. The H trace was beyond the limit of registration in the negative direction from 15 h .43 m. to 16 h .53 m ., and may of course have been considerably disturbed.

The largest movements recorded during Section (i) took place later, between 18 h .53 m . and 19 h .53 m . During this hour D rose and fell $66^{\prime}, \mathrm{H}$ rose $99 \gamma$, while V fell $14 \gamma$ and rose $35 \gamma$.

The D trace was off the sheet on the positive side for some time after 21 h . and there was no trace from 22 h .16 m . to 22 h .36 m . There is thus rather a lack of information as to what was happening during Birkeland's Section (ii). After 22 h .36 m ., when registration was resumed, there was no really striking D movement. There was, however, a small bay between 23 h .16 m . and 23 h .56 m ., the value rising $18^{\prime}$ and falling $30^{\prime}$. The H trace, which had just got off the sheet on the positive side at 22 h .46 m ., came on again at 22 h .47 m , and between that hour and $23 \mathrm{~h} .33 \mathrm{~m} . \mathrm{H}$ fell $53 \gamma$. After 23 h .33 m . H rose gradually with minor oscillations until 0 h .53 m . on the 28 th , when the trace again went off the sheet on the positive side and remained off for two hours.

V fell $32 \gamma$ between 22 h .36 m . and 23 h .17 m ., and rose $40 \gamma$ between 23 h .17 m . and 23 h .43 m . After a slight halt it continued to rise, but much more slowly, until 0 h .17 m . on the 28 th. Between 0 h .17 m . and 0 h .45 m . it fell $19 \gamma$. Thereafter the V trace was relatively quiet for some hours.
§14. October 28-29, 1902 (hours 14-1, Plate V).
This was another "compound" storm, which resembled that of the previous day in containing two "intermediate" elementary polar storms. The interval between these was, however, much less than on the previous day, and there was, according to Birkeland, p. 222, in lower latitudes, "no trace on the 28 th of the long storm that occurred on the 27th, and was especially powerful at the Equator." According to Brrkeland's
table, p. 222, the first "intermediate" storm commenced about 18 h .10 m . and ended about 19 h .30 m . at most European stations, the maximum coming about 18 h .45 m . The value of $\mathrm{P}_{1}$-the maximum horizontal disturbing force-varied from $248 \gamma$ at Axelöen to $3 \gamma$ at Toronto, the value at Kew, $16 \gamma$, being slightly under the average for non-polar European stations.

The time of commencement of the second "intermediate" storm is given as about 21 h .30 m . at most stations, but somewhat earlier in the Arctic. The end is given as usually somewhat after 23h. The maximum is said to have occurred about 22 h .10 m ., the value of $\mathrm{P}_{1}$ varying from $266 \gamma$ at Axelön to $2.5 \gamma$ at Batavia. The value given for Kew, $16.5 \gamma$, is again slightly below the mean for non-Arctic Europe.

The disturbances outside the Arctic were really very trifling. The most notable change at Kew was a rise of $20 \gamma \mathrm{in} \mathrm{H}$ between 21 h .30 m . and 21 h .48 m .

In the Antarctic there was no loss of V-trace during the time covered by Plate V, except from 22 h .43 m . to 23 h .13 m ., when there was no sheet on the drum, and the H trace was off the sheet only for a short time before the end. The D trace, however, was off the sheet a good deal between 20 h . and 22h. 43 m .

The largest D movements recorded took place between 19 h .10 m . and 19 h .58 m ., and so synchronise with or overlap the latter part of Birkeland's first "intermediate" storm. During this time D rose $41^{\prime}$, fell $28^{\prime}$, rose $45^{\prime}+$ (going off the sheet), fell $52^{\prime}+$, and rose $50^{\prime}$.

Some rather notable oscillations also occurred in H. The element rose $37 \%$ and fell $53 y$ between 18 h .38 m . and 19 h .0 m ., the turning point (a maximum) being at 18 h .50 m . and so practically simultaneous with the maximum in Birkeland's first "intermediate" storm. Another considerable oscillation took place between 19 h .40 m . and 20 h .10 m ., H first rising $55 \gamma$ and then falling $52 \gamma$. The V trace showed numerous small oscillations. There was one rather sharp oscillation between 20 h .5 m . and 20 h .16 m ., a fall of $19 \gamma$ being followed by a rise of $18 \gamma$. The intervening minimum occurred at 20 h .9 m . Between this hour and 21 h .20 m . there was a total rise of $36 \gamma$ in V , which was followed during the next 20 minutes by a fall of $27 \gamma$.

The D trace went off the sheet rather steeply at 21 h .45 m ., and came on rather steeply at 22 h .25 m ., so there may have been a considerable oscillation in this element during the time of Birkeland's second "intermediate" storm. The H and V traces, however, after 21 h .40 m . were quieter than they had been for some hours previously. Thus whilst there was decidedly more than the average amount of disturbance in the Antarctic during Birkeland's first "intermediate" storm, it is at least doubtful whether the same was true of the second "intermediate" storm.
§15. October 29-30, 1902 (hours 16-4, Plate VI).
This "compound" perturbation is said, p. 161, to consist of an "equatorial" perturbation-which commenced suddenly on the 29 th at 16 h .52 m ., and whose most active phase in the southern stations appeared at about 1 h .30 m . on the 30 th -and of "polar" storms. Whether Birkeland supposed the same equatoriad perturbation to last continuously all the time is not clear. As to the nature of the coincidence of the equatorial and polar disturbances, p. 161 says: "The positive equatorial perturbations observed by us are always accompanied by polar storms. As a rule, the polar storms do not begin until a little while after the equatorial ; but on this occasion they begin almost simultaneously. . . ."

In discussing his Chart I , which includes results for hours 18 h .52 .5 m . and 20 h .30 m . on the 29 th , Birkeland concludes, p. 164, that "it is the polar systems that give the field its character," and he puts the "centre" of the polar system near Matotchkin Schar. In discussing Chart II for 1 h . 0 m . on the 30th, he regards the field as "now mainly conditioned by the equatorial perturbation."

During the major part of the polar storm Birkeland had records from only two polar stations, Axelöen and Matotchkin Schar. The largest movements shown at either occur between 18 h . and 21 h .

At the non-polar stations there was a distinct sudden commencement-not clearly apparent at the polar stations-whose time of occurrence Birkeland puts at 16 h .52 m . The original Kew $H$ trace shows a small fall, about $1 \gamma$, between 16 h .52 m . and 16 h . 54 m ., followed by a rise of about $4 \gamma$ during the next 4 minutes.

The most noteworthy movement at Kew and the other non-polar stations took place between 1 and
$2 \mathrm{a} . \mathrm{m}$. on the 30 th, the maximum displacement being that already referred to as occurring about 1 h .30 m . The disturbance at this hour appears to have been a good deal larger at Dehra Dun, Batavia, and Christchurch than at the non-polar European stations, Birkeland's estimate of the horizontal disturbing force being $43 \gamma$ at Batavia and $40 \gamma$ at Christchurch, as against $14 \gamma$ at Kew.

In the Antarctic there was an outstandingly rapid rise of H from about 16 h .55 m . to 16 h .59 m ., followed by an equally rapid and larger fall. The trace was very faint near the time of the turning-point, and got beyond the range of registration at 17 h .5 m ., so that all one can be sure of is that between 16 h .55 m . and 17 h .5 m . H rose at letst $27 \gamma$, and fell at least $37 \%$. A slow rise in H commenced about 16 h .53 m ., but this was checked for a few seconds at 16 h .55 m ., and the movement did not attain its highest rapidity until perhaps 16 h .56 m . This H movement occurred at a time when the trace had been rather quieter than usual for an hour or more, and there can be little doubt that it represents the sudden commencement seen at Birkeland's non-polar stations. In the Antarctic, synchronous apparently with the commencing movement in H , there was a sharp oscillation in V , a rise of $10 \gamma$ and fall of $6 \gamma$ taking place in about 6 minutes. D , which had been rising on the whole fairly steadily with minor oscillations, began a more decisive though not conspicuously rapid rise about 16 h .56 m . During the next 27 minutes it rose $73^{\prime}$ and the trace then got off the sheet. It came on the sheet 3 minutes later, but, after being on for about 18 minutes, got off once more, and thereafter was seen only at rare and short intervals during the remainder of the time covered by Plate VI.

The H trace remained beyond the limits of registration until about 19 h .18 m ., and was again lost sight of about 19 h .45 m . The light in the Antarctic magnetographs evidently became very faint towards the end of the sheet, as even the base lines are but faintly indicated after 21 h . The V trace, which suffered less from weak illumination than the $D$ and $H$ traces, had become invisible by this hour. It is thus possible that faintness of light may have been partly responsible for the non-appearance of the $\mathbf{D}$ and $\mathbf{H}$ traces after 20 h .

The persistence, however, of active disturbances until the time when the V trace became invisible may be safely inferred from the following list of observed changes in V :-

| From 17h. 6 m . to 17 h .48 m . rise $57 \gamma$, |  |
| :---: | :---: |
| ," | 17h. 48 m . , 17 h . 54 m . fall $36 \gamma$, |
| ", | 17h. 54 m. , 18 h . 5m. rise 33 y , |
| " | 18 h .57 m . , 19 h . 8 m . rise $38 \mathrm{\gamma}$, |
| , | 19h. 8 m. , 19 h .30 m . fall $58 \gamma$, |
| " | 19h. 30 m . " 19 h .45 m . rise $33 \gamma$, |
| " | 19h. 45 m. " 20 h . 1 m . fall $36 \gamma$, |
| , | 20h. 1m. "20h. 8 m . rise $30 \gamma$, |
|  | 20h. 8 m . " 20 h .23 m . fall $35 \%$. |

The major part of these $V$ disturbances synchronise with Birkeland's "polar" storm, but some precede it. Shortly after the last movement recorded above the trace became invisible.

After the next sheet was put on at 23 h .20 m , distinctly disturbed conditions existed until after $3 \mathrm{a} . \mathrm{m}$. on the 30 th. The changes in $V$ were especially noteworthy. Between 0 h .0 m . and 1 h .4 m . on the 30 th there was a rise of $92 \gamma$, between 1 h .4 m . and 1 h .46 m . a fall of $74 \gamma$, and between 1 h .46 m . and 2 h .46 m . a rise of $93 \gamma$. H rose $52 \gamma$ between 0 h .0 m . and 0 h .40 m ., the trace then going off the sheet on the positive side. After being off for 9 minutes it came on, but 20 minutes later it went off again on the positive side. Coming on once more at 1 h .36 m ., it showed a fall of $59 \gamma$ between 1 h .36 m . and 2 h .17 m ., and a rise of $59 \gamma$ between 2 h .17 m . and 2 h .35 m . The trace went off the sheet at the latter hour, and except for a short appearance of about 10 minutes re-appeared no more until after 8 h .

During the above changes in V and H the chief movements in D were a fall of $129^{\prime}$ between 1 h .10 m . and 1 h .43 m ., and a rise-interrupted for 25 minutes by minor oscillations-of $87^{\prime}$ between 1 h .43 m . and 2 h .45 m . After 3 h .30 m . the D and V traces were specially quiet during the next 5 hours. The H trace being off the sheet, one cannot be certain that it was equally quiet.

It will doubtless have been noticed that the largest movements recorded in the Antarctic occurred
during the time when Birkeland's "polar" storm was largest, and during the time of his largest "equatorial" disturbance. There is, however, no sign of intermission in the disturbance, though the evidence is not complete owing to failure of the trace.

Later in the 30th, it may be added, after the time covered by Plate VI, there was a very prominent bay in the Antarctic $D$ curve, extending from about 8 h .20 m . to 9 h .33 m . It included a fall of $142^{\prime}$ and rise of $115^{\circ}$. This disturbance was at least an approach to the "special type," $V$ oscillating about a mean position during the fall of D , and rising about $30 \gamma$ during the rise of D . There was a synchronous bay on the $H$ curve, but details are lacking, as the trace was off the sheet most of the time.
§16. October 31 to November 1, 1902 (hours 6-2, Plate VII).
Of this "compound" storm Birkeland says, p. 230, "It appears at the poles with tremendous violence, although perhaps its strength is even more unusual at the equatorial stations. Considering its long duration and its universal distribution, we may say that it is the greatest storm that has been observed by us." He regards the disturbance as consisting of a long storm lasting from about 9 h . on the 31 st to 3 h . on Nov. 1, with two, if not more, "intermediate" storms.

Referring to his first eight charts, which answer to times from 9 h .0 m . to 12 h .30 m . on October 31, he says, p. 232, that the equatorial stations show "powerful perturbing forces directed southwards," the forces at Dehra Dun and Batavia being almost double those in central and southern Europe. During this time Birkeland's Arctic stations showed no very large disturbances, but Sitka was highly disturbed.

Charts IX, X, and XI, for $13 \mathrm{~h} .30 \mathrm{~m} ., 13 \mathrm{~h} .42 \mathrm{~m}$., and 14 h .0 m. , represent the conditions during the "first powerful intermediate storm," whose maximum is put at 13 h .42 m . This includes the time of largest movements at the polar and equatorial stations. There are also movements at all the non-polar European stations, but these are on the whole smaller than the movements later in the day.

After 14 h .0 m . conditions were everywhere less disturbed for some hours. But from 17 h .45 m . to 1 h .0 m . on Nov. 1 there were further large disturbances in the Aretic and the European stations, which are dealt with in Birkeland's Charts XII to XIX.
The second "intermediate" storm is regarded as extending from 23 h .12 m . to 0 h .42 m . on Nov. 1 , with maximum about 23 h .45 m ., and after its conclusion the conditions became much quieter.
As usual, Kew seems to be fairly representative of non-polar European stations. It is very difficult there to assign even an approximate time for the commencement of the disturbance. One has to go back to 20 h . on the 30 th to get a time really free from the small undulatory movements which represent the disturbance up to noon on the 31st. The end of the disturbance between 3 and $4 \mathrm{a} . \mathrm{m}$. on Nov. 1 is more definite. The Kew D curve shows two slow wave-like movements in immediate sequence, extending from 7 h .30 m . to 10 h .0 m . on the 31 st , the rise and fall in each being from $1^{\prime}$ to $2^{\prime}$. From 11 h .40 m . to 14 h .30 m . there was another group of movements of a more irregular character, which included of fall of $4^{\prime}$ and rise of $3^{\prime}$ between 13 h .20 m . and 13 h .50 m . This corresponds to Birkeland's first "intermediate" storm.

From 17 h . on the 31 st to 2 h . on Nov. 1 there was considerably more disturbance at Kew than earlier. Between 17 h .0 m , and $17 \mathrm{~h} .48 \mathrm{~m} . \mathrm{D}$ fell and rose $3^{\prime}$, reaching a sharp peak at the latter hour. After 17 h .48 m . D continued to fall generally, with minor oscillations, until 22 h .10 m ., the fall in this time amounting to $13^{\prime}, D$ then rose $4^{\prime} \cdot 3$ in two steps to a rounded peak at 23 h .10 m . Between this hour and Oh. 45 m . on Nov. 1 it fell $6^{\prime}$ and rose $7^{\prime}$; the turning-point, which was the minimum during the disturbance, was at about 23 h .42 m .

In the Kew $H$ curve the most rapid changes were a fall of $26 \gamma$ and rise of $23 y$ between 13 h .15 m . and 13h. 42 m .-corresponding to Birkeland's first "intermediate" storm-and a fall of $25 \gamma$ between 17 h .45 m . and 17 h .53 m . There was a comparatively quiet time from 14 h .40 m . to 17 h .45 m . After the latter hour there was no cessation of disturbance until about 2 h . on Nov. 1.

In the Antarctic a highly disturbed state of matters existed from about $8 \mathrm{~h}, 50 \mathrm{~m}$. to 14 h . 0 m . on the 31st. The phenomena resembled four disturbances of the "special type," following one after the other without any interlude; but D and V were not quite in phase, and most of the turning-points on the H trace were beyond the limits of registration, so one can only see that this element was approximately in phase with V.

We may regard the V movements during this time as composed of four "waves," whose times and amplitudes were as follows ( + denotes a rise, - a fall):-

| Wave | From | To | Change (in $\gamma$ ). |
| :---: | :---: | :---: | :---: |
| 1 | $\left\{\begin{array}{c} \mathrm{h} . \mathrm{m} . \\ \mathrm{P}^{?} \mathrm{si} \end{array}\right.$ | $\begin{array}{ll} \text { h. } & \text { m. } \\ 8 & 51 \\ 9 & 23 \end{array}$ | $\begin{gathered} ? \\ +26 \end{gathered}$ |
| 2 | $\left\{\begin{array}{rr}9 & 23 \\ 10 & 13\end{array}\right.$ | 1013 1059 | -16 +43 |
| 3 | $\begin{cases}10 & 59 \\ 11 & 50\end{cases}$ | 1150 1230 | $\begin{aligned} & -29 \\ & +21 \end{aligned}$ |
| 4 | $\begin{cases}12 & 30 \\ 13 & 33\end{cases}$ | $\begin{array}{lr}13 & 33 \\ 14 & 1\end{array}$ | $\begin{aligned} & -34 \\ & +49 \end{aligned}$ |

Each "wave" except the third left V enhanced, so that the final value exceeded the original by $60 \%$. There were four "waves" in D roughly corresponding to those in V.

| During the first D fell | $75^{\prime}$ | and rose | $72^{\prime}$. |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| $"$ | "second ", | $102^{\prime}$ | $"$ | $"$ | $133^{\prime}$. |
| $"$ | $"$ third " | $"$ | $81^{\prime}$ | $"$ | $"$ |
| $195^{\prime}$. |  |  |  |  |  |
| $"$ | , fourth ", | $173^{\prime}$ | $"$ | $"$ | $80^{\prime}$. |

H fell $108 \gamma$ between 9 h .32 m . and 10 h .23 m ., rose $56 \gamma$ between 10 h .23 m . and 11 h . 5 m ., and fell $57 \gamma$ between 11 h .5 m . and 11 h .30 m . The trace went off the sheet on the negative side at 11 h .30 m ., remaining off until 12 h .46 m . It came on again, but only for a few minutes, showing a double peak at 12 h .50 m ., and was thereafter no more seen until 20 h .53 m .
The last of the four waves above mentioned in the Antarctic synchronises with Birkeland's first "intermediate" polar storm. But it seems impossible to draw any line of demarcation, such as Birkeland draws, between it and what precedes. The four waves are of the same type and follow in immediate sequence, and it is difficult to believe that the first three can be due to "equatorial" perturbation if the last is due to "polar." The storm in the Antarctic probably existed for some hours before 8 h .51 m ., as the D trace contains two bays similar to, though smaller than, the four we have described. The H trace, however, was off the sheet after 1 h .20 m ., and the V magnetograph not working until nearly 8 h . 51 m ., so our information is very limited before this hour.

After 14h. 1m. there was probably an absence of large rapid movements in the Antarctic until nearly 21 h . The H trace was off the sheet, and the D trace also after 17 h .40 m ., but the V trace remained on until the sheet was taken off at 22 h .7 m ., and its course was unbroken by any oscillations at all comparable with the four we have described. There was, however, a persistent fall in $V$, as if the cumulative effect of the four waves were gradually disappearing. The total fall between 14 h .1 m . and 20 h .41 m . amounted to fully $70 \gamma$.

After 20 h .50 m . there were some more wave-like movements in the V trace, though not so large as the earlier ones. The two principal consisted, the first of a fall of $21 \gamma$ and rise of $12 \gamma$ between 20 h .50 m . and 21 h .28 m ., the second of a fall of $30 \gamma$ and rise of $16 \gamma$ between 21 h .28 m . and 22 h .9 m .

- The H trace, after having been beyond the limits of registration on the negative side for nearly 8 hours, came on the sheet at 20 h .58 m ., and H rose $89 \gamma$ between that hour and 21 h .30 m . Rising further, with minor oscillations, it got off the sheet on the positive side at 22 h .2 m ., having risen $110 \gamma$ since 20 h .58 m .

There was no record from 22 h .7 m . to 23 h .46 m ., which covers half the time of Birkeland's second "intermediate" storm. When the next sheet was put on at 23 h .46 m ., the H trace was off the sheet on the positive side, and it remained off until 4 h .30 m . on November 1. The D and V traces on the second sheet show, however, no signs of special disturbance. The largest movements in the V trace are two bays, the first, from 1 h .33 m . to 2 h .45 m ., having a rise and fall of only $16 \gamma$, the second, from 2 h .45 m . to 3 h .56 m ., showing a rise of $43 \gamma$ and fall of $27 \gamma$.

## §17. November 23-24, 1902 (hours 15-7, Plate VIII).

This "compound" storm, divided by Brakeland into three sections, is said, p. 272, to be the most powerful of a series which developed daily in the Arctic from the 19 th to the 26 th of November.

Section (i), from 15 h .20 m . to about 16 h . 0 m ., is described, p. 273 , as a "typical positive equatorial storm," "strongest at the equatorial stations in the South of Asia." Brekeland speaks of a sudden rise occurring in H at 15 h .30 m ., except at the American stations, where the rise is slow. The time is not, I think, intended to be exact. For at Kow H begins to rise at about 15 h .22 m ., rising $10 \gamma$ to a rounded maximum at about 15 h .32 m ., and then diminishing about $6 \gamma$ during the next 7 minutes. There was a corresponding small movement in D , whose maximum was, however, a few minutes later. These triffing movements-which can, however, be recognised at Axelöen, Kaafjord, and all the European and Southern stations-constitute the largest disturbance of Section (i).

During Section (ii), from 16 h . to about 22 h ., the disturbing forces are generally small, but from 17 h .30 m . to 18 h .20 m . the disturbance is somewhat greater, "especially in southern latitudes."

In the Arctic Birkeland sees indications of a "polar" storm of minor intensity. During this time the most noteworthy feature at the non-polar European stations is a bay on the $\mathbf{D}$ curve, extending from about 17 h . 0 m . to 18 h .30 m . at Kew , where the element fell and rose about $3^{\prime}$. The most prominent H changes at Kew were a fall of $12 \gamma$ between 17 h .23 m . and 17 h .35 m ., and a fall of $11 \gamma$ between 19 h .52 m . and 20h. 3 m .

Birkeland's discussion of his Section (iii), 22h. to 7h., appears involved. He attempted apparently to recognise a series of "elementary polar" storms in it, but was unable to disentangle them. Speaking generally, the Arctic curves show numerous rapid oscillatory movements going on pretty continuously all the time.

At the non-polar stations the most disturbed time was from 22 h . on the 23 rd to 3 h . on the 24 th . The movements during this time were really of some size. At Kew, for instance, D fell $10^{\prime}$ between 22 h . 5 m . and 22 h .42 m ., rose $5^{\prime}$ between 0 h .30 m . and 0 h .50 m ., fell $6^{\prime} .8$ between 0 h .50 m . and 1 h .20 m ., rose $10^{\prime} .4$ between 1 h .20 m . and 2 h .10 m ., fell $9^{\prime}$ between 2 h .10 m . and 2 h .43 m ., and rose $7^{\prime}$ between 2 h .43 m . and 3 h .12 m . H fell $37 \gamma$ in three steps between 21 h .48 m . and 23 h . 1 m ., rose $35 \gamma$ in two steps between 23 h .1 m . and 23 h .42 m ., fell $29 \gamma$ in two steps between 23 h .42 m . on the 23 rd and 1 h .35 m . on the 24 th , rose $42 \gamma$ between 2 h .2 m . and 2 h .24 m , and fell $30 \gamma$ between 2 h .48 m . and 3 h .52 m .

The conditions existing at the southern stations during the early morning of the 24 th will be best realised by consulting the Christchurch H curve given in Plate XV of the present volume.

In the Antarctic there were some very rapid oscillatory movements synchronous with the commencing movements seen at Kew and elsewhere which Birkeland characterised as "equatorial." The D movements were so rapid that the turning-points are not very clearly shown, and the following measurements may be slightly under-estimates. Commencing suddenly, during a fairly quiet time, we have between 15 h .21 m . and 15 h .37 m . a rise of $81^{\prime}$, a fall of $78^{\prime}$, a second rise of $60^{\prime}$ and a second fall of $85^{\prime}$, with minor oscillations during the larger movements. There were presumably corresponding movements in $H$, as the curve which had been off the sheet for some time on the negative side came on for a couple of minutes about 15 h .24 m ., forming a very sharp peak (maximum). During the second oscillatory movement in D there was a marked depression and recovery in V , of amplitude $44 \gamma$. After this commencement the Antarctic curves remained highly disturbed until 10 h . on the 24 th . They are reproduced from 0 h .18 m . to 10 h .31 m . (or $11.25 \mathrm{a} . \mathrm{m}$. to $9.38 \mathrm{p} . \mathrm{m} . \mathrm{g}$ L.T.) in Plate XXV of the present volume. The most conspicuous movements prior to the time covered by this plate were as follows:-

D after falling $226^{\prime}$, with numerous oscillations, between 15 h .28 m . and 16 h .25 m ., rose again, oscillating vigorously, about $197^{\prime}$ between 16 h .25 m . and 18 h .8 m ., and then fell $127^{\prime}$ between 18 h .8 m . and 18 h .58 m . A specially rapid movement about 19 h .19 m . is referred to presently in conjunction with a corresponding $H$ change. Between 19h. 19 m . and $21 \mathrm{~h} .3 \mathrm{~m} . \mathrm{D}$ rose, oscillating largely, $208^{\prime}$, and the trace went off the sheet on the positive side for a few minutes at the latter hour. After coming on the sheet, the D trace in the course of about half an hour fell $97^{\prime}$, rose $72^{\prime}$, fell $75^{\prime}$, and rose $100^{\prime}$, going off the sheet again about 21 h .40 m . Its reappearance a few minutes later is not very clearly shown, and at 22 h .8 m . it disappeared, to be seen no more before the sheet was removed at 23 h .13 m .

The movements just recorded extend over Birgeland's Sections (i) and (ii) and show no interlude.
Until 21 h . the H trace was mostly off the sheet on the negative side, appearing at intervals for a few minutes at a time. The most striking movement seen during this time answered to a peak at about 19 h .19 m . The trace was visible for about 8 minutes, during which H rose and fell $68 \gamma$. This happened synchronously with an exceedingly rapid oscillation in D , that element in the course of 11 minutes falling $97^{\prime}$ and rising $88^{\prime}$.

Later the H trace was more in evidence. Between 20 h .53 m . and $21 \mathrm{~h} .8 \mathrm{~m} . \mathrm{H}$ rose $66 \gamma$ and fell $49 \gamma$ with minor oscillations. Between 21 h .8 m . and 21 h .20 m . it rose $140 \gamma$, the trace then going off the sheet on the positive side and remaining off for 20 minutes.

On reappearing at 21 h .40 m . the H trace fell $136 \gamma$ in 7 minutes, then rose $54 \gamma$ and fell $70 \gamma$ in the course of the next 11 minutes, and went off the sheet on the negative side about 21 h .58 m . It was not again seen, except for a few minutes, until about 22 h .37 m , when it appeared and rose $103 \gamma$ during the next 10 minutes.
There were numerous minor oscillations all this time in the $V$ trace, the most striking being a fall of $65 \gamma$ and rise of $50 y$ between 19 h .17 m . and 19 h .28 m ., a time during which there were also rapid oscillations in D and H . But perhaps the most outstanding feature is the continuous downward tendency of the curve from 16 h .15 m . to 21 h .40 m ., the total fall during this time being about $300 \%$. After some small oscillations V then commenced to rise rapidly, the rise between 21 h .57 m . and 22 h .42 m . representing $182 \gamma$.

From 23 h .13 m . on the 23 rd to 0 h .19 m . on the 24 th there was no paper on the drum, and though substantial alterations took place in the values of the elements during this interval it may have been quieter than the times before or after. But there is no direct evidence of any marked lull from the commencement, about 15 h .20 m . on the 23 rd , until 10 h . on the 24 th . No subdivision into sections nor difference in source is at all suggested by the Antarctic $D$ and $H$ curves.

The case of $V$ is rather different, as that element, on the whole, rose between 15 h .35 m . and 16 h .15 m ., and again between 21 h .40 m . and 22 h .42 m ., while it fell during the intervening time.

In the Antarctic there can be no question that the disturbance of October 31 to November 1 was smaller than that of November 23-24. The range of the elements appears greater at Kew on the former occasion, but that is due to the fact that the period was longer and included the ordinary hours of the daily maximum and minimum. The appearance of the Kew curves would indicate that the later disturbance was the more intense, and this is really borne out by Birkeland's figures on pp. 240 and 280 for the amplitude of the disturbing force.
§18. December 9, 1902 (hours 5-18, Plate IX).
This is entered amongst the "equatorial" storms. When discussing it on p. 70 Birkeland describes the disturbance as illustrating at its commencement "all the properties that characterise the positive equatorial perturbations. It commences quite suddenly, simultaneously all over the Earth, at $5 \mathrm{~h} .40 \cdot 6 \mathrm{~m}$." This movement is seen in the Arctic as well as the non-polar regions. It is, however, by no means large at most stations. At Kew , for instance, the initial rise in H was only about $5 y$ and it was not very rapid. Birkeland regards "equatorial" conditions as persisting until nearly 15h. During this time there are no large movements even in the Arctic.
"Between 15 h . and 18 h ., the character of the perturbation conditions is essentially changed. It is this feature that we continually find repeated, namely, that when the equatorial storm has lasted for some hours, polar systems appear." p. 70. During the time stated there was a disturbance of some size at all the Arctic stations, and even at the non-polar stations there were appreciable bays on both the H and D curves. Thus at Kew the H trace shows a bay from 16 h .25 m . to 17 h .40 m ., the greatest depression representing about $15 y$; corresponding to this was a hump on the D curve, the maximum representing a rise of about $2^{\prime} \cdot 5$.
In the Antarctic there were of course numerous movements larger than any at Kew, but there is no very decided trace of disturbance until towards the end of the period covered by Plate IX. On this occasion it is very doubtful whether there is anything in the Antarctic curves corresponding to the sudden movement seen elsewhere about 5 h .41 m . There is, it is true, apparently at this exact time a trifling but sharp peak in the V trace-which previously was very quiet-representing a very rapid rise and fall of
$1 \gamma$ or $2 \gamma$, and this is followed by a rise of about $7 \gamma$ in the course of the next 6 or 7 minutes. There were peaks on the D curve at 5 h .41 m . and 5 h .45 m . Between these two times D rose $7^{\prime}$, and during the next 9 minutes it fell $14^{\prime}$. There was also a peak on the $H$ trace at about 5 h .41 m ., the next turning. point being about 5 h .51 m . ; in the interval H rose $20 \gamma$. These movements in the D and H curves are, however, not conspicuously different either in size or rapidity from a good many others, and the apparent coincidence in time may be accidental.

One would hardly describe the Antarctic curves as disturbed until after 16h. From then until 23h. the disturbance was continuous and very considerable. During the time covered by Plate IX the largest movements recorded in D were a fall of $156^{\prime}$ between 16 h .44 m . and 17 h .13 m , followed by a rise of $213^{\prime}$ between 17 h .13 m . and 17 h .33 m . The HI trace was mostly off the sheet after 15 h . V, however, showed some considerable movements, rising $79 \gamma$ between 17 h .10 m . and 17 h .35 m ., falling $62 \gamma$ between 17 h .37 m . and 17 h .46 m ., then rising $38 \gamma$ between 17 h .46 m . and 17 h .57 m ., and falling $80 \gamma$ between 17 h .57 m . and 18 h .17 m . This last movement, however, extends beyond Birkeland's period.

The Antarctic movements after 16 h .47 m . are shown in the right-hand figure of our Plate XL. In it 4 a.m., December 10, answers to 16 h .53 m. ., G.M.T., on the 9 th. This figure shows the Antarctic disturbance for more than an bour subsequent to the time covered by Birkeland's Plate IX. The movements which it shows, though large, are if anything inferior in size to some recorded between hours 20 and 22, G.M.T., i.e. about two hours later. The disturbed conditions continued until nearly 23 h . One might thus be inclined to infer that the Antarctic storm, while so far synchronous with that observed in the Arctic and elsewhere after 16 h ., continued long after the disturbance elsewhere had ceased. The D and H Kew curves show, however, between 19h. and 22 h ., some movements which, though less than those between 16 h . and 18 h ., are larger than those occurring between 5 h . and 7 h . It would be of interest to know what was happening in the Arctic between 18h. and 22 h .
§19. December 14-15, 1902 (hours 23-5, Chart X).
This "elementary polar" storm is described by Birkeland, p. 87, as appearing "upon an otherwise very calm day . . . . without any preceding equatorial perturbation," and as consisting of "a great storm in the north, about Dyrafjord and Axelöen . . . . accompanied by a perturbation, small indeed, but well defined, . . . . observed in Northern America and Europe."

The effect is described as "just perceptible" at Dehra Dun, but not visible at Batavia. At Dyrafjord, where the movement was largest, the storm is said to have lasted from 0 h .10 m . to 3 h .15 m ., the maximum value of the disturbing force, $386 \gamma$, being met with about 1 h .8 m . At Axelöen, where the maximum disturbing force was about half that at Dyrafjord, the times were somewhat later, the maximum not appearing until 1 h .46 m . In temperate Europe the disturbance is said to begin "rather suddenly at 0 h .45 m ." and to last about 3 hours. Birkeland adds, "This perturbation . . . has its origin in the northern regions. Its sphere of action . . . . is concentrated about the neighbourhood of Dyrafjord and Axelöen. The shortness of its duration, as also the comparatively calm character of the curves . . . . seems to indicate that this is a polar elementary storm of the most typical nature; it appears to be produced by a coherent impulse, which increases to a certain size, and then again decreases to $0 \ldots$. . At the same time, as the perturbation does not make its appearance at all places simultaneously, the perturbing cause must be supposed to move with a somewhat continuous motion," p. 87. This remark has been quoted at length, because in several respects it is so suggestive of the Antarctic "special type" of disturbance, the principal difference being that the value of V in the Antarctic usually remained elevated for some time after the apparent end of the disturbance in D and H .

In Europe, as Birkeland says, the disturbance was small outside the Arctic. At Kew, for instance, H rose about $10 \gamma$ between 0 h .45 m . and 1 h .5 m . and then fell very gradually until about 2 h .40 m ., the total fall being about $15 \%$. D rose about $2^{\prime} \cdot 9$ between 0 h .45 m . and 1 h . 5 m . and then fell about $3^{\prime} \cdot 5$ to a badly defined minimum about 2 h .5 m .

In the Antarctic there were some rather striking movements about three hours before the earliest time on Plate X, and one would put the commencement of the disturbances there at about 18 h .30 m . on the 14th. There were, however, very sudden movements commencing about 23 h .5 m . in both D and H , consisting of a rise and fall occupying in all some six minutes. H rose $22 \gamma$ and fell $34 \gamma$. The D oscillation
was apparently larger，but owing to its rapidity the trace is too faint to follow to the turning－point． Simultancously there was a very sudden rise of $20 \gamma$ in $V$ ，which was preceded and followed by slower movements in the opposite direction．On examining Plate X one observes a small sudden movement at least approximately synchronous with these Antarctic movements at Matotchkin Schar，Dyrafjord，and the American stations，but it cannot be identified with certainty in the Kew curves．There was no record obtained in the Antarctic from 23 h .56 m ．on the 14 th until 0 h .19 m ．on the 15 th ．After the latter hour，however，there was a deep bay on the D curve，the element rising $99^{\prime}$ between 0 h .33 m ．and 1 h .11 m ，and falling $106^{\prime}$ between 1 h .11 m ．and 2 h .12 m ．H rose $48 \gamma$ between 0 h .53 m ．and 1 h .23 m ．； the trace then went off the sheet on the positive side，remaining off until 2 h .33 m ．After coming on for a few minutes，it was again off until 3 h ．V rose $65 \gamma$ between 0 h .47 m ．and 1 h .23 m ．，and then fell $36 y$ between 1 h .23 m ．and 2h． 8 m ．

These movements，it will be observed，occur about the time of the principal movements in Birkeland＇s Arctic stations．On their conclusion the Antarctic curves were relatively quiet during the next 12 hours．
§๊0．December 24－25， 1902 （hours 23－5，Plate XI）．
This disturbance is included amongst the＂compound．＂In temperate Europe Birkeland says，p．165， ＂the conditions are slightly disturbed from 23 h ．on the 24 th to 5 h ．on the 25 th．There are especially distinct perturbations about midnight，and from 2 h .30 m ．to 4 h ．＂In Toronto and at Baldwin and Cheltenham，U．S．，the perturbation is practically confined，Birkeland says，to the short interval 3 h .14 m ． to 3 h .57 m ．on the 25 th ，with maximum at 3 h ． 21 m ．At Dehra Dun，Batavia，and Christchurch 3 h ．to 4 h ．is also decidedly the most disturbed time．The Arctic stations are as much disturbed between 23 h ．on the 24 th and 0 h .15 m ．on the 25 th as they were later，while temperate European stations are also sensibly disturbed at the earlier hour．Birkeland concludes，p．165，that in Europe，as a whole，the conditions ＂are in the main connected with the polar storms at the Norwegian（i．e．Aretic）stations．＂

The disturbance at the non－polar European and Asiatic stations was really very trifling．At Kew the only changes in D at all conspicuous were a rise of $1^{\prime} \cdot 2$ between 23 h .5 m ．and 23 h .18 m ．，followed by a fall of $1^{\prime} \cdot 6$ between 23 h .24 m ，and 23 h ． 35 m ．on the 24 th ，and a rise of $2^{\prime} \cdot 6$ between 3 h .14 m ．and 3 h .21 m ．，followed by a fall of $2^{\prime}$ ending about 4 h .5 m ．on the 25 th ．In H there was a rise of $9 \gamma$ between 3 h .34 m ．and 3 h .55 m ．on the 25 th ．The other changes hardly catch the eye．

In the Antarctic it was rather quieter than usual for 2 or 3 hours prior to 23 h ．on the 24 th ． There then ensued a decidedly more disturbed time，extending from about 23 h ． 15 m ．on the 24 th to 0 h .13 m ．on the 25 th ．Between 23 h .15 m ．and 23 h .54 m ．D rose $28^{\prime}$ ，fell $84^{\prime}$ and rose $57^{\prime}$ ．H，during this time，had a total range of only $45 \gamma$ ，but there was rather a prominent double oscillation composed of a rise of $38 \gamma$ in 4 minutes，a fall of $45 \gamma$ in 6 minutes，a rise of $42 \gamma$ in 9 minutes，and a fall of $24 \gamma$ in 8 minutes．Between 23 h .19 m ．and 23 h .56 m ．V fell $40 \gamma$ and rose $53 \gamma$ ．Owing to the trace coming on the clamp，there was no Antarctic record from 0 h .18 m ．until the new sheet was put on at 0 h .55 m ．on the 2⿹勹口th．Conditions were distinctly quiet for over an hour after this．The H trace was off the sheet on the positive side from 2 h .23 m ．to 4 h .48 m ．，so there may have been high values in that element．Between 2 h .15 m ．and 4 h .19 m ．D fell gradually $130^{\prime}$ ，but a very appreciable fraction of this must be ascribed to the regular diarnal variation．The fall was interrupted as usual by a good many minor oscillations，the largest retrograde movement being a rise of $26^{\prime}$ between 3 h .12 m ．and 3 h .21 m ．The V trace showed no large oscillations，but there were a number of minor oscillations between 3 h .20 m ．and 4 h .35 m ．

The Antarctic movements are synchronous with and larger than those seen at the American and non－polar European stations．

The reader should，however，be warned that one would not naturally regard any portion of the time covered by Plate XI as more than usually disturbed in the Antarctic，with the exception of the last hour of the 24th．At the same time，the end of Plate XI answers to the early afternoon in the Antarctic，and the diurnal changes at Midsummer were then so rapid that irregular disturbances appeal less to the eye than at most hours of the day．
§21．December 26－27， 1902 （hours 18－2，Plate XII）．
This is included amongst the＂polar elementary＂storms．From the discussion on p． 137 we learn that it comprised two distinct＂elementary＂storms，the first especially powerful at Matotchkin Schar，having
a maximum at about 20 h .45 m . to 21 h ., the second especially powerful at Dyrafjord, being most developed between 22 h .30 m . and 24 h . There is also a somewhat vague reference to a "more lengthy perturbation" as covering the time of the two polar storms, and to the possibility of "cyclo-median" perturbations being felt in lower latitudes.

At the time of the first "polar elementary" storm the most prominent feature at Kew was a bay on the D curve. The element fell $4^{\prime} \cdot 2$ between 20 h .36 m . and 20 h .46 m ., and then rose gradually to about its original value at about 21 h .45 m ., the rate of recovery slackening after 21 h .10 m . H rose about $7 \gamma$ between 20 h .45 m . and 20 h .57 m ., having been falling slowly for some time previously.

At the time of the second "polar elementary" storm the most prominent feature at Kew was a bay on the H curve, the value of the element rising and falling about $14 \gamma$ between 23 h .8 m . and 23 h .48 m . In D there was a rise of about $1^{\prime} .5$ between 23 h .0 m . and 23 h .15 m , followed by an equal fall and a further small rise. After midnight on the 26 th one would ordinarily describe the curves as very quiet. The conditions at $\mathrm{K}_{e w}$ appear fairly representative of the European non-polar stations.

The Antarctic curves were unmistakably disturbed after 19 h . on the 26 th until the sheet was removed at 23 h .31 m . The disturbance commenced apparently with a very rapid rise in D , the value increasing $141^{\prime}$ in 6 minutes from 18 h .59 m . to 19 h . 5 m . After some small rapid oscillations D continued to rise until 19 h .15 m ., the total rise since 18 h .59 m . being $158^{\prime}$. This was followed by a fall of $196^{\prime}$ between 19 h .15 m . and 19 h .53 m ., and a second rise of $224^{\prime}$ between 19 h .53 m . and 20 h .55 m ., the trace then going off the sheet on the positive side for a few minutes.
H commenced to fall at 18 h .59 m ., when D began to rise, and fell $79 \gamma$ in about 9 minutes, shortly thereafter getting beyond the limits of registration on the negative side. The largest change shown in H was between 20 h .27 m . and 21 h .53 m , when the element rose 160 y in several steps interrupted by minor oscillations.

The $V$ trace during this time showed numerous oscillations, including the following ( + denotes a rise, - a fall):-

|  | h. | m. |  | h. | m, | $\gamma$. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| From | 18 | 59 | to | 19 | 4 | +22, |
| " | 19 | 4 | " | 19 | 16 | -68, |
| " | 19 | 16 | " | 19 | 24 | +40, |
| " | 19 | 24 | " | 19 | 31 | -65, |
| " | 19 | 31 | " | 19 | 36 | +36, |
| " | 19 | 36 | " | 19 | 46 | -29, |
| " | 20 | 7 | " | 20 | 33 | +47, |
| " | 20 | 33 | " | 20 | 48 | -33, |
| " | 20 | 57 | " | 21 | 13 | -49. |

After 21 h .45 m . there were further considerable movements in the Antarctic, the most notable being on two occasions when the traces from the three elements show deep bays, occurring synchronously or very nearly so. On the first occasion, lasting from about 21 h .55 m . to 22 h .30 m .,

$$
\begin{array}{lllll}
\text { D fell } & 136^{\prime} & \text { and rose } & 130^{\prime}, \\
\text { H } & " & 89 \gamma & \# & 69 \gamma, \\
\text { V } & \prime & 120 \gamma & " & 112 \gamma .
\end{array}
$$

The turning-points (minima) occurred within a few minutes of one another.
On the second occasion the oscillation was apparently not quite completed when the paper was changed. So far as recorded, it lasted from about 22 h .56 m . to 23 h . 31 m . During it

| D rose | $33^{\prime}$ | and fell | $39^{\prime}$, |  |
| :--- | :--- | :--- | :--- | :--- |
| H | $"$ | $52 \gamma$ | $"$ | $66 \gamma$, |
| V | $"$ | $23 \gamma$ | $"$ | $37 \gamma$ |

The second movement is comparatively small, and is mentioned chiefly because the three elements appear approximately in phase and the time synchronises with that of Brkeland's second "polar. elementary" storm.

For some time after 23 h .50 m . on the 26 th , when the next sheet was put on, there were some fair movements, notably some rapid nearly synchronous oscillations in the traces of the three elements between 0 h .20 m . and 1 h .0 m . on the 27 th . After this the conditions became distinctly quieter.

The temperature trace was lacking or imperfectly visible during most of the time, and the V changes recorded for times prior to 21 h .30 m . are not corrected for temperature. The oscillations were, however, so rapid that the uncertainty thus arising is small.
§22. December 28, 1902 (hours 3-8, Plate XIII).
This date was not given in Prof. Birkeland's circular, and he had few data for it from non-polar regions except from America. He has classified the storm, p. 169, as "compound," though the phenomena at Dyrafjord, where it was largest, suggested a "polar elementary" storm, whose centre was originally somewhere to the south of Greenland, and which moved to the westward. The classification seems to have been partly determined by the fact that the time of the chiof disturbance-from 4 h .40 m . to 6 h .was an unusual one for the occurrence of "elementary polar" storms.
The curves reproduced from the European co-operating stations include only D and H from San Fernando and D from Stonyhurst. The Kew D curves between 4 h .30 m . and 6 h .30 m . showed two wavelike movements; in each a rise of about $1^{\prime} \cdot 5$ was followed by a more gradual fall, the crests coming at about 4 h .50 m . and 6 h .0 m . The $H$ curve also showed two waves, $H$ falling about $6 \gamma$ between 4 h .30 m . and 4 h .50 m ., and rising $7 \gamma$ between 4 h .50 m . and 5 h .15 m ., then falling $5 \gamma$ between 5 h .15 m . and 5 h .30 m ., and rising very slowly about $5 \gamma$ to a maximum near 6 h .20 m .

There was a somewhat larger disturbance at Kew about midnight on the 27th, but it precedes the time covered by Plate XIII. The movements lasted from about 21 h . on the 27 th to 0 h .30 m . on the 28 th.

In the Antarctic the most notable phenomenon was a deep bay on the $D$ trace, of the same character as appeared during the "special type" of disturbance. The element began to fall suddenly about 5 h .13 m . and in 13 minutes fell $90^{\prime}$, going off the sheet on the negative side and remaining off for 11 minutes. During the next 12 minutes it was off and on the sheet once or twice, and then continued to rise with minor oscillations. The rise was not so rapid as the fall, and owing to the minor oscillations it is difficult to assign a definite time for the conclusion. D had, however, returned to its original value by about 6 h .20 m .

H was off the sheet on the positive side from 3 h .35 m . to 6 h .25 m ., and so far as that element was concerned the disturbance might have been of the "special type." The V trace, however, was not of that type. There was a rapid fall of $18 \gamma$ between 5 h .14 m . and 5 h .19 m ., followed by a rise of $29 \gamma$ between 5 h .19 m . and 5 h .37 m . But during the rise there were numerous small oscillations which continued until 6h. 20 m .

The Antarctic curves, it may be added, were somewhat highly disturbed on the 27 th , from 16 h .15 m . until 23 h .55 m ., when the sheet was taken off.
§23. January 26, 1903 (hours 7-15, Plate XIV).
During this time the Antarctic magnetographs were not in action.
§24. January 26-27, 1903 (hours 18-7, Plate XV).
This "compound" storm appears in Europe outside the Arctic as "a long perturbation . . . lasting from about 18 h .0 m . on the $26 \mathrm{th} \ldots$. . to 7 h .0 m . on the 27 th . . . We have . . . three . . . sharply defined intermediate storms," p. 287. These "intermediate" storms are said to coincide, on the whole, in time with three storms recorded in the Arctic, especially powerful at Axelöen.

The last, however, of these three "intermediate" storms attained its maximum about 0 h .35 m . on the 27 th, and the Antarctic records do not commence until 1 h .5 m . By this time there was comparatively little disturbance except in the Arctic, and even there it was much reduced.

In the Antarctic after 1 h .5 m . there was a fall in progress in D until 4 h .8 m ., when the curve got off the sheet for a few minutes. A fall is what we should expect in the ordinary course of events, but the fall shown between the hours mentioned, $198^{\prime}$, is notably in excess of the average. The fall in D was interrupted as usual by minor oscillations. The most conspicuous of the retrograde movements, one of $57^{\prime}$, took place between 2 h .23 m . and 2 h .46 m . The D trace was off the sheet between 5 h .25 m . and 6 h .0 m ., and the form of the curve when going off and coming on is not inconsistent with the existence of a bay of considerable depth.

On the H trace there were some fairly large oscillations, but only partly shown. H rose $56 y$ hetween 1 h .13 m . and 1 h .26 m . ; the trace was then off the sheet for 21 minutes on the positive side. Between 1 h .47 m . and 2 h .31 m . H fell and rose $53 \gamma$, the trace again going off the sheet. Between 4 h .16 m . and 5 h. 13 m . H rose $105 \%$.
The V trace was, on the whole, quiet. It contained, however, a small bay from about $1 \mathrm{~h}, 13 \mathrm{~m}$. to 1 h .48 m ., V rising and falling about $19 \gamma$.
§25. February 8, 1903 (hours 8-12, Plate XVI ; and hours 13-24, Plate XVII).
Before discussing this disturbance Birkeland mentions, p. 187, that conditions in the Arctic had been on the whole very quiet since the end of November, 1902, until February 7, 1903. On that day, he says, a fairly powerful storm was experienced in the Arctic from 21 h .5 m . on the 7 th until about 1 h . on the 28 th. There was, I may add, also marked disturbance in the Antarctic at the same time, but it commenced earlier, about 19 h .30 m . on the 7 th .

The disturbance of February" 8 is classed amongst the "compound." It is regarded as divisible into three sections, the first of which covers the time to which Plate XVI refers. As this section is allowed a separate plate, and there is a gap of an hour between the times represented by the two plates, I shall treat the two parts separately. During his Section (i) Brakeland remarks, p. 187, that "The perturbation is particularly powerful at Sitka, and is (there) especially violent from 9 h . to 9 h .35 m ." It continued to be considerable at Sitka until 11h., but ran a rather irregular course there and at the Arctic stations. Birkeland adds, p. 188, "The simple conditions found between San Fernando in the west and Zi-ka-wei in the east, and between Kew in the north and Batavia in the south, form a strong contrast . . the perturbation is throughout chiefly in H . It is well defined, and, as far as we can determine, commences everywhere simultaneously at about 8 h .35 m . . . It terminates simultaneously at about 10 h .50 m ." The time of maximum at the co-operating stations outside North America is given as from 10 h .0 m , to 10 h .10 m .

Birkeland, p. 189, considers that the phenomena at the American stations suggest "a polar elementary storm, at first not very far north-east of Sitka," but this could not, he says, account for the phenomena "over the district between Kew and Batavia," which suggest a "negative equatorial storm," i.e. a disturbance due "to a current round the Earth from east to west . . . at a distance from the Earth of at least a magnitude equal to the radius of the Earth." If I rightly follow him, the view he finally inclines to, p. 189, is "that at first the perturbation partakes most of the nature of a cyclo-median storm, and subsequently changes into a more purely polar one."

At the non-polar European stations the phenomena were similar to those at Kew. There $H$ fell gradually about $40 \gamma$ between 8 h .35 m . and 10 h .0 m ., and then rose $30 \gamma$ between 10 h .0 m . and 10 h .50 m . For some reason which I do not understand, Birkeland, p. 188, regards this as the end of the storm at Kew, and elsewhere, during his Section (i). But as a matter of fact the Kew H curve began to fall again slightly immediately after 10 h .50 m ., and between 11 h .20 m . and 11 h .50 m . it rose $17 \%$, falling $14 \gamma$ during the next 30 minutes. The trace was quieter for a short time after 12 h .20 m . The D trace at Kew showed a rise of from $2^{\prime}$ to $2^{\prime} .5$ above the normal between 9 h .30 m . and 10 h .40 m . D was, if anything, falling at 10 h .50 m ., the normal change at that hour being a rise. Decidedly the most conspicuous movement was a rise of $4^{\prime} \cdot 8$ between 11 h .35 m . and 11 h . 55 m . There was a fall of $3^{\prime} \cdot 5$ between 12 h .0 m . and 12 h .40 m .

In the Antarctic the working of the V magnet appears doubtful. The H trace was off the sheet on the positive side, except during a few minutes, until 8 h .56 m . Between 9 h .38 m . and 10 h .39 m . the trace shows a prominent to-and-fro movement-a fall followed by a rise-strongly suggestive of the special type of disturbance. The full extent of the movement is not shown, as the trace was beyond the limits of registration on the negative side for some 15 minutes, but the amplitude exceeded $68 \%$. The turningpoint, a minimum, must have occurred within a few minutes of 10 h . 0 m . The D trace was off the sheet at this time on the negative side, having been off since about 8 h .10 m ., and did not appear until 10 h .53 m . It went off and came on steeply, suggesting a deep bay or bays.

Between 10 h .53 m . and 11 h .33 m . D rose $82^{\prime}$. Between 11 h .33 m . and 12 h .25 m . it executed a to-and-: fro movement, somewhat suggestive of the special type of disturbance, a fall exceeding $82^{\prime}$ being followed by a rise of over $105^{\prime}$. The trace was off the şheet only for a minute or two, the turning-point coming at
about 11 h .50 m . The H trace shows a bay synchronous with that on the D trace, H rising as D fell and conversely, but the H movements were relatively small, the rise being $17 \gamma$, the fall $25 \gamma$.

It will be noticed that the large bay in the $H$ curve occurs during the time of the first movements at Kew and elsewhere, while the hay in the D curve synchronises closely with the disturbance recorded at Kew near noon.

Birkeland, p. 190, says that his Section (ii) extends from 14h. to 18 h ., but Plate XVII begins at 13 h ., and the disturbance is said to commence at more than one place at 13 h .45 m . The conclusion drawn, p. 191, is that the perturbations represent "a series of short, principally polar impulses with somewhat changing centre."

The disturbances in the Arctic are not very striking, but there is a considerable movement at Sitka, Kiafjord, and Matotchkin Schar with a maximum about 14 h .40 m . This is also about the time when the departure from the normal is largest at the non-polar stations. At Kew H fell $17 \%$ between 14 h .28 m . and 14 h .44 m ., and then returned gradually to the original value at about 15 h .50 m .

Birkeland's Section (iii), 18 h . to 23 h ., had for its principal feature in the Arctic a "violent storm . . . at all the . . stations simultaneously, most powerful at Axelöen and Matotchkin Schar," p. 192. The time of commencement of this powerful storm is said to vary from 18 h .33 m . at Dyrafjord to 19 h .7 m . at Axelöen, the time of ending being about 22 h .30 m . The intensity appeared to be greatest between 19 h .15 m . and 20 h .15 m .

At the non-polar stations the disturbance was mainly from 19 h .5 m . to 20 h .30 m ., the maximum being placed by Birkeland at 19 h .18 m . At Kew the range between 19 h . and 20 h . was 48 y in H and $14^{\prime} .5$ in D ; the most prominent movements were a rise of $44 \gamma$ in H from 19 h .18 m . to 19 h .30 m ., and a fall of $14^{\prime}$ in D from 19 h .5 m . to 19 h .24 m .

In the Antarctic it does not seem possible to draw any line corresponding to Birkrland's division into Sections (ii) and (iii). There was a comparatively quiet time for about half an hour prior to 13 h ., but thereafter there was a constant succession of large movements until after the time included in Plate XVII. The largest of several considerable oscillations in D during Birkeland's Section (ii) consisted of a fall and rise each about $84^{\prime}$ between 13 h .33 m . and 14 h .28 m . During Birkeland's Section (iii) there were larger, but not more rapid movements. Between 18 h .47 m . and 20 h .40 m . there was a sort of bay in the D trace, a rise of $189^{\prime}$ being followed by a fall of $151^{\prime}$. The decreasing movement which began about 20 h .5 m . was interrupted by a rise of $78^{\prime}$ between 20 h .41 m . and 21 h .13 m . When resumed, it continued until 22 h .21 m ., the value of D at that hour being $261^{\prime}$ below that at 20 h .5 m . Between 22 h .37 m . and 23 h .34 m . D rose $110^{\prime}$.

The $H$ trace was off the sheet on the negative side, most of the time from 14 h .10 m . until 21 h .20 m . Its longest appearance was from 16 h .28 m . to 17 h .42 m ., when it rose and fell $68 \gamma$, with numerous minor oscillations. The chief turning-point (a maximum) was at 16 h .53 m . After 21 h .20 m . the H trace remained on the sheet until the end of the 8 th. There was rather a rapid rise of $52 \gamma$ between 21 h .20 m . and 21 h .33 m ., followed by a slower motion in the same direction interrupted by numerous oscillations. The total range in H between 21 h .20 m . and 24 h .0 m . was $101 \gamma$.
§26. February 10-11, 1903 (hours 20-3, Plate XVIII).
This is classed amongst the "polar elementary" storms. Of it Bireeland says, p. 106, "This magnetic disturbance is brief, and commences without any previous equatorial perturbation on an otherwise very quiet day. First a small disturbance appears rather suddenly at about $21 \mathrm{~h} .6 \mathrm{~m} . \ldots$... most powerful at the northern stations . . . . but is also perceptible in (temperate) Europe and North America. After about 30 minutes, the conditions are once more almost normal . .. . The powerful perturbation .... does not commence until 23 h . . . . After about an hour and three-quarters the storm is over . . . . At 2 h .30 m . on the 11th February another short, slight perturbation appears. . . ."

On p. 107 Birkeland gives a table containing his estimate of the times of beginning and ending and of the maximum for the principal storm, and also the value of $\mathrm{P}_{1}$, the maximum disturbing force in the horizontal plane. At the non-polar stations the beginning is put at about 23 h . 0 m ., the maximum at about 23 h .20 m . The end is given for the non-polar European stations as 1 h .0 m . on the 11 th. In the Arctic the times of beginning and ending are more variable, but do not depart much from those
already given; the maximum appears about half an hour later than elsewhere. $P_{1}$ is said to vary from $373 \gamma$ at Matotchkin Schar to less than $5 \gamma$ at Batavia. It is given as $35 \cdot 8 \gamma$ at Kew, a value similar to that assigned to the other non-polar European stations.

Commenting on the fact that the value of $\mathrm{P}_{1}$ at Christchurch, $12 \gamma$, is larger than the values at Honolulu, Zi-ka-wei, and Batavia, Birkeland says, p. 108, "This may be explained by the fact that the perturbation in the Arctic regions is often accompanied by simultaneons perturbations in the Antarctic regions, and it is the effect of these latter that is noticed in Christchurch. Our material does not, however, allow of certain conclusions being drawn in this matter."

As Birkeland does not seem to have had any records from south of Christchurch, the above is presumably pure surmise.

At Kew the short commencing movement was represented by a fall of 5 y in H between times which I make 21 h .8 m . and 21 h .16 m ., and by a bay lasting from 21 h .8 m . to 21 h .35 m . in the D curve, the element falling $1^{\prime} \cdot 9$ and rising $1^{\prime} \cdot 0$.

The principal disturbance at Kew was represented in H by a rise of 29 from 23 h .4 m . to 23 h .19 m ., a fall of 41 y from 23 h .19 m . on the 10 th to 0 h .16 m . on the 11 th , and a rise of $13 \gamma$ from 0 h .16 m . to 0 h .45 m . In the D curve there was a bay from about 23 h .0 m . on the 10 th to 0 h .40 m . on the 11 th, the element being depressed. The maximum depression occurred about 23 h .40 m . and amounted to $5^{\prime} \cdot 2$.

In the Antarctic, after being comparatively undisturbed for over 12 hours, the traces became distinctly disturbed about 18 h .36 m . on the 10 th , or $2 \frac{1}{2}$ hours before the commencing disturbance noted by Birkeland, and the disturbance continued without appreciable intermission until 23 h . 0 m ., when the lamp went out. When registration was resumed at 0 h .11 m . on the 11th the disturbance appeared much reduced. The period for which trace is lacking includes unfortunately the major part of Birkeland's principal storm.

The commencing movement in the Antaretic was a rapid fall of $72 \gamma$ in H between 18 h .36 m . and 18h. 53m. The most striking movement, however, prior to Plate XVIII, was an oscillation which appeared simultaneously in $\mathrm{D}, \mathrm{H}$, and V . The turning-point, a maximum, occurred at about 19 h .46 m . In the course of about 14 minutes D rose $70^{\prime}$ and fell $54^{\prime}$, H rose and fell $39 \gamma$, while V rose $25 \%$ and fell $33 \%$. There is, however, no distinct movement at this time in the Kew curves.

During the time covered by Plate XVIII the largest Antarctic movements were as follows, + denoting a rise and - a fall in the element :-

| D. | H. | V. |
| :---: | :---: | :---: |
| $$ |  |  h. m.  h. m.   <br> From 19 46 to 19 53 <br>  19 53  23 29 <br>  20 29  20 49 |

The H trace was off the sheet on the negative side for a few minutes after 21 h .6 m ., and got off on the positive side at 2 h .28 m , remaining off for over $4 \frac{1}{2}$ hours.

After the new sheet was put on at 0 h .11 m . on the 11th, the D and V traces appeared no more disturbed than usual.
Owing to the lack of trace our information is mainly confined to the fact that the disturbances near the commencement of Plate XVIII, which Birkeland aseribes to equatorial perturbations, were synchronous with large oscillations in the Antarctic, which formed part of a series which commenced about 18h. 36 m ., and which persisted without intermission until at least 23 h .0 m .
§27. February 15, 1903 (hours 13-20, Plate XIX).
This is classified as a "compound" storm on the following grounds, p. 174,... "It must thus be assumed that these are in the main polar perturbations ; but the conditions are not simple, indicating, as
they do, both in the Arctic regions and in lower latitudes, that there are a number of systems acting to some extent simultaneously."

According to Brrkeland, one noteworthy peculiarity is that the disturbance at most stations commenced nearly 2 hours later, while ending nearly half an hour earlier in D than in H .

In Europe, outside the Arctic, the H disturbance is given as lasting from about 14 h .15 m . to 18 h .15 m ., and the D disturbance as lasting from about 16 h .10 m . to 17 h .45 m . The time of the maximum disturbance is given at most stations as near 16 h .40 m . The maximum value of the disturbing force in the horizontal plane varied, according to Birkeland, from $392 \gamma$ at Axelöen to $10 \cdot 6 \gamma$ at Honolulu, the value at Kew, $38 \gamma$, being similar to that at most non-polar European stations.

There were no Christchurch results.
In Europe generally, outside the Arctic, the disturbance consisted of a slight depression in $H$ from about 14 h .15 m . to 18 h .15 m .-the maximum depression at Kew below the normal being about $20 \gamma$ and occurring about 16 h .30 m . -and of a more conspicuous depression in D. At Kew there was a fall of $6^{\prime} \cdot 5$ in D between 16 h .10 m . and 16 h . 45 m ., followed by a rise of $4^{\prime} .8$ between 16 h .45 m . and 17 h .20 m . After a stoppage of nearly 20 minutes there was a further rise of about 2 ' going on until nearly 18 h .30 m . This final rise, however, is so slow that opinions might well differ as to its duration and significance.

In the Antarctic the curves would hardly be described as disturbed before 16 h . Between 13 h .41 m . and 14 h .53 m ., however, D rose $66^{\prime}$, while H fell $64 \gamma$ between 13 h .17 m . and 14 h .53 m . There were two successive bays on the D curve, the first from 16 h .3 m . to 17 h .10 m . with a rise of $51^{\prime}$ and fall of $31^{\prime}$, the second from 17 h .10 m . to 18 h .10 m . with a fall of $52^{\prime}$ and rise of $57^{\prime \prime}$. The H trace, which had been off the sheet on the negative side for about 45 minutes, came on the sheet at 17 h .23 m ., and H rose $63 \gamma$ by 17 h .50 m . Following this were a number of oscillations of some size, $H$ rising on the whole until 19 h .33 m ., the total rise since the trace came on the sheet at 17 h .23 m . being about $105 \mathrm{\gamma}$. During this time the V trace shows two fairly prominent bays, the first lasting from 15 h .53 m . to 17 h .5 m , the second from 17 h .5 m . to 18 h .13 m . The first bay represents a rise and fall of about $25 \dot{\gamma}$, the second a rise and fall of about $48 \gamma$.

Up to 18 h .40 m . there were no very rapid movements. Between 18 h .40 m . and 18 h . 55 m ., however, there were simultaneous rapid oscillations in D and H ; D rose $52^{\prime}$ and fell $48^{\prime}$, while H fell $62 \gamma$ and rose $75 \%$. During part of this time, from 18 h .43 m . to 18 h .48 m ., there was a smart rise of $25 \gamma$ in V . This was followed by a larger but slower fall of $32 \gamma$ from 18 h .48 m . to 19 h .3 m .
Conspicuous movements in D and H were just concluding at 20 h .0 m ., the hour answering to the end of Plate XIX. D had risen $42^{\prime}$ in the course of the previous 11 minutes, whilst $H$ had fallen $64 \gamma$ since 19 h .46 m . These movements seem, however, to be related rather to what follows than to what precedes them. For after a small retrograde movement $H$ continued to fall to a very sharp peak at 20 h .3 m ., the total fall since 19 h .46 m . amounting to $81 \%$. The peak at 20 h .3 m . answers to a depression (minimum) in V , rather suggesting that the preceding fall answered to the first phase of a disturbance of the "special type." If this view is correct, the second phase is represented by a rise of $78 \gamma$ in H and of $23 \gamma$ in V , which took place between 20 h .3 m . and 20 h .16 m . This was immediately followed, between 20 h .16 m . and 20 h .38 m ., by another double movement, also suggestive of the "special type" of disturbance. The H change in this second oscillation was a fall of $72 \gamma$ and rise of $56 \gamma$, the V change a fall of $36 \gamma$ and rise of $43 \gamma$. This movement was followed by yet a third sharp oscillation in H and V , terminating about 20 h .54 m . The range of H , however, in the third oscillation was only $32 \gamma$, the fall and rise being equal, while the V movements were under $20 \gamma$. During these three oscillations there were also oscillatory movements of some size in D , the total range in D between 20 h .3 m . and 20 h .54 m . amounting to $50^{\circ}$. The oscillations in D were, however, much interrupted by minor oscillations, and it is difficult to decide on their exact relationship to the H and V movements.

One would put the commencement of the quiet time in the Antarctic at the end of the third oscillation, i.e. at 20 h .54 m ., and one would unquestionably regard the disturbances from 19 h .46 m . to 20 h .54 m . as forming part of a common system.
§28. March 22-23, 1903 (hours 12-1, Plate XX).
This is classed amongst the "elementary polar" storms, but Birkeland, p. 127, explains that "The
perturbation . . . is in reality . . . composed of two principal phenomena, an equatorial perturbation and a short, well-defined, comparatively powerful elementary polar storm." Of the "equatorial" perturbation he writes: "This ...begins quite suddenly, at 12 h .58 m ., with an oscillation that is noticed simultaneously all over the world. In the equatorial regions, this sharp deflection . . . appears principally in H . About the auroral zone the curve oscillates, and the perturbation is noticeable both in D and H. ." The time at which this commencing disturbance reached a maximum is given as 13 h .4 m . Birkeland concludes, apparently, that the small enhancement of $H$ caused by the commencing disturbance, or at least a fraction of it, persists in low latitudes right through the subsequent polar storm until midnight. At Dyrafjord and Axelöen, however, he notes "a peculiar circumstance," p. 128, viz., that the commencing movement is distinctly oscillatory. The H curve, in fact, at the two stations shows a fall, a rise and then a second fall, and possibly a second, but much smaller, rise.

According to Birkeland's table, p. 130, the time of commencement of the "polar" storm varied from 20 h .30 m . to 21 h .30 m ., the time at non-polar European stations being about 21 h .10 m . At most stations the termination is put at about 23 h .45 m . At most non-polar European stations the maximum occurred about 22 h .10 m. ; in North America it was about half an hour later. The maximum value of the horizontal disturbing force is given as varying from about $370 \gamma$ at Axeloen to $12.4 \gamma$ at Batavia, the value at Christchurch being too small to measure satisfactorily. The value at Kew, $44 \gamma$, is about a mean for non-polar Europe.

My examination of the original Kew curves makes the sudden commencement 4 or 5 minutes earlier than the hour, 12 h .58 m ., given by Birkeland. But the movement appears distinctly oscillatory in both H and D , a small fall, lasting 2 or 3 minutes, preceding the rise. In the case of $H$ the fall seems about $1 \gamma$, the rise about $13 \gamma$; while in D the fall is $1^{\prime} \cdot 0$, the rise $2^{\prime} \cdot 4$. There are a number of small movements at Kew in both $H$ and $D$ throughout the remainder of the 22 nd, but much the most conspicuous phenomenon is a bay on the D curve from about 21 h .6 m . to 23 h . 50 m ., the element being depressed in value. The largest value of the depression is about $8^{\prime} \cdot 8$ and it occurred about 22 h .5 m .

Looking at the curves in Plate XX, it is easy to recognise in most, if not all, a second movement which somewhat resembles the commencing movement both in size and character, and which occurred about an hour and fifty minutes later. At Kew this second movement occupied some 14 minutes. During it H fell $4 \gamma$ and rose $10 \gamma$, while D fell $0^{\prime} \cdot 8$ and rose $1^{\prime} \cdot 0$, the turning-point in each case coming at about 14 h .46 m . The reason for mentioning this is because on looking at the Antarctic curves one's eye is caught by two oscillatory D movements larger than their neighbours, occurring during an otherwise somewhat unusually quiet time. So far as I can judge, these movements correspond in time to the commencing movement and the second movement just referred to. The following are the results of the measurement made of the two groups of movement (+ denotes a rise, - a fall):-


In each case it is a little doubtful whether the final rise in D forms part of the same system as the preceding rise and fall. The $H$ movements do not stand out very conspicuously, and it is mainly their approximate coincidence in time with the others that has led to their enumeration. The V movements, though small, do stand out.

Unless we assume an accidental coincidence of a truly remarkable kind, the conclusion seems inevitable that the above two groups of movements are due to the same cause as the movements seen at Kow and elsewhere at the same times. If this be so, then the fact that the Antarctic movements represent
disturbing forces much larger than those shown by the curves in equatorial and temperate latitudes is a fact which requires a good deal of explanation on the hypothesis of equatorial electric currents.

From 15 h .5 m . to 18 h .15 m . the Antarctic curves, like those elsewhere, were, on the whole, distinctly quiet. From 18 h .15 m . to 18 h .25 m . there was a more prominent oscillation in H , a fall of $14 \gamma$ being followed by a rise of $16 \gamma$, while V rose $16 \gamma$ between 18 h .22 m . and 18 h .26 m .

Between 19 h .13 m . and 19 h .22 m . H fell $20 \gamma$ and rose $24 \gamma$, while between 19 h .32 m . and 19 h .35 m . it rose $16 \gamma$ and fell $17 \gamma$, the latter an excoptionally rapid though not large oscillation.

After 20 h .15 m . conditions became more disturbed in all the elements. The oscillations, though not really large, were unusually rapid between 20 h . 15 m . and 22 h . 15 m .

The $\mathbf{D}$ and H traces got rather mixed up and are difficult to distinguish. Amongst the movements in the V trace after 20 h . were-

$$
\begin{array}{ccccc}
\begin{array}{c}
\text { a fall of } 16 \gamma \\
\text { a between } 20 \mathrm{~h} . ~ \\
\hline
\end{array} 14 \mathrm{~m} . \text { and } 20 \mathrm{~h} .19 \mathrm{~m} ., \\
\text { a rise of } 24 \gamma & " & 20 \mathrm{~h} .19 \mathrm{~m} . & 20 \mathrm{~h} .24 \mathrm{~m} ., \\
\text { and a fall of } 39 \gamma & " & 20 \mathrm{~h} .24 \mathrm{~m} . & , 20 \mathrm{~h} .31 \mathrm{~m} .
\end{array}
$$

After this there were numerous small but rapid oscillations, during whose incidence V rose to a maximum at about 21 h .13 m . There then ensued a fall, interrupted by minor oscillations, which continued until about 22 h .40 m , in the course of which V fell $56 \gamma$. This was followed by a rise of $97 \gamma$ in two steps between 22 h .40 m . and 23 h . 25 m ., the rise being conspicuously most rapid during the first 20 minutes. From 22 h .20 m . to 23 h .3 m ., simultaneously with the most rapid part of the change in V , there was a prominent bay on the H curve, the element falling $61 \gamma$ and rising $36 \gamma$.

After 23 h . the D and H traces were somewhat uncertain owing to operations connected with the clearing away of snow near the Magnet Hut, and the record was suspended altogether from 23 h .37 m . to 1h. 5 m . on the 23 rd .

It will have been noticed that the large rise in V and the prominent bay on the H curve took place during Birkeland's "elementary polar" storm.

So far as mere amplitude is concerned, the large rise in V after 22 h .40 m . is, perhaps, the only phenomenon which merits the title of disturbance as judged by the Antarctic standard. The number, however, of minor oscillations in V and the rapidity of the oscillations in D and H , especially H , are certainly a little outstanding.
§29. March 30-31, 1903 (hours 19-3, Plate XXI).
This is classed amongst the "elementary polar" storms. Birkeland explains, however, pp. 115-117, that this "elementary" storm was preceded by an "equatorial" perturbation. He adds, p. 116: "As early as 19 h ., those little, sudden, very variable perturbations are noticed, which occur simultaneously all over the Earth. . . The ("equatorial") perturbation appears to be over at about 23 h .12 m. . . ."

The "polar" perturbation is said to be recognisable at Kaafjord about 23 h ., being earlier there than elsewhere. It did not commence at Dyrafjord until about 0h. 24 m . At non-polar stations the commencement was, in general, about 23 h .50 m ., the end about 2 h .10 m . on the 31 st . The hour of occurrence of the maximum is given as 0 h .30 m . for temperate Europe, but 0 h .58 m . for Dyrafjord. The maximum value of the horizontal component of the disturbing force varied from $546 \gamma$ at Dyrafjord to $10.5 \gamma$ at Batavia. The value for Kew, $41 \cdot 5 y$, is about a mean for non-Arctic Europe.

The earliest decided movement shown in Plate XXI at non-Arctic stations is a small oscillation apparent in almost all the H traces and in some of the D traces. At Kew the double movement lasts from about 19 h .23 m . to 19 h .31 m ., the turning-point coming about 19 h .27 m . There was a rise of $7 \gamma$ and then a fall of $6 \gamma$ in H , while D fell and then rose $1^{\prime} \cdot 3$.

While examining the original Kew curves I noticed a somewhat conspicuous movement of similar size and character about an hour and fifty minutes earlier. This consisted in the case of H of a rise of $6 \gamma$ between 17 h .32 m . and 17 h .36 m ., followed by a fall of $9 \gamma$ between 17 h .36 m . and 17 h .40 m . The reason for mentioning this will appear presently.

During the whole time covered by Plate XXI the most conspicuous phenomenon at Kew (and the same is true generally of the other non-polar stations) took place between 0 h .10 m . and 1 h .10 m . of the 31 st, i.e during the "elementary polar" storm. D, which had fallen 1 ' 8 between 23 h .45 m . and 0 h .10 m . on
the 31 st , rose $9^{\prime} \cdot 6$ between 0 h .10 m . and 0 h .32 m ., and then fell $11^{\prime} \cdot 5$ between 0 h .32 m . and 1 h .15 m . Between 1h. 15 m . and 1 h . 40 m . there was a rise of 2 , and a further very gradual rise continued until nearly 4 a.m. H fell $10 \gamma$ between 23 h .45 m . and 0 h .6 m . on the 31 st , rose $25 \gamma$ between 0 h .15 m . and 0 h .42 m ., and fell $25 \gamma$ between 0 h .42 m . and 1 h .27 m . The only other movements on the Kew curves likely to attract attention occurred shortly after 21 h ., i.e. $2 \frac{1}{2}$ hours before the commencement of the "elementary polar" storm. Between 21 h .1 m . and 21 h . 35 m . H rose $5 \gamma$ and fell $12 \gamma$, then rose $19 y$ in two steps and fell $6 \gamma$. The $D$ trace shows only a very small movement about 21 h . 3 m ., consisting of a fall of $0^{\prime} \cdot 8$, followed by an equal rise.

The Antarctic curves had been rather quiet for some hours prior to 19 h . on the 30 th , but they exhilit a somewhat prominent oscillation occurring apparently simultaneously in the $\mathrm{D}, \mathrm{H}$, and V curves, all showing a sharp peak at 17 h .35 m . The oscillation lasts from about 17 h .31 m . to 17 h .39 m . In D there is a rise and fall, each 12 ', in H a rise of $16 \gamma$ and fall of $12 \gamma$, in V a rise and fall, each $7 \gamma$. The V movement, though small, is clearly shown, the curve being unusually quiet for some time before and after. The fall in D and $H$ only paused for a minute at 17 h .39 m . and then proceeded for some minutes, but at a slower rate. The oscillation is synchronous with the earlier movement at Kew, which occurred prior to the time considered by Birkeland.

The next movement in the Antarctic which catches the eye occurs between 19 h .25 m . and 19 h .31 m ., with turning-point at about 19 h .28 m . in all the elements. The D movement is small and not very distinct. It consists apparently of a fall of about $2^{\prime}$ and a rise of $10^{\prime}$. H falls and rises $18 \gamma$, while V falls and rises $6 \gamma$. These movements synchronise with the second of the two movements at Kew, i.e. with Birkeland's "equatorial "perturbation. One is rather reminded of the occurrence of the two oscillatory movements on March 22 ; the interval between the two movements is nearly the same on the two occasions. The movements of March 30 do not appeal so much to the eye as those of March 22.
The fact that the disturbances in the Antarctic are again larger than the corresponding ones at stations nearer the Equator will doubtless have been noted.

Though not seriously disturbed, the Antarctic curves all show oscillations of increased amplitude between 19 h .25 m . and 21 h .40 m . The largest movements during this time consisted of a bay, apparently simultaneous in the three traces, from about 21 h .6 m . to 21 h .26 m . During it D fell $35^{\prime}$ and rose $52^{\prime}$, H rose and fell $26 \gamma$, while V fell $25 \gamma$ and rose $38 \gamma$. This movement, it should be noticed, occurred at a time when the curves at Kew and the other non-polar stations showed increased activity of disturbance.

From 21 h .40 m . to 22 h .25 m . there was a relatively quiet time in the Antarctic. There then ensued from 22 h .25 m . to 23 h .30 m . an interval during which the V trace showed a number of rather sharp oscillations, which were accompanied by more or less simultaneous oscillations in D and H , the H oscillations being of enhanced size. Several of these oscillations rather suggest the "special type" of disturbance, but their period is short and the rises in V do not exceed the falls. The two following cases gave the largest changes in V :-


The oscillations in V continued for some hours, but with diminished size. In the case of D and H , however, movements of increased size and duration took place.
The chief movements in D were a fall of $99^{\prime}$ between 23 h .30 m . on the 30 th and 0 h .20 m . on the 31 st , a rise of $120^{\prime}$ between 0 h .20 m . and 0 h .56 m ., and a fall of $111^{\prime}$ between 0 h .56 m . and 1 h .57 m . These movements were accompanied by shorter oscillations of comparatively small size.

During this time there was a large bay on the H curve, but its nature is imperfectly shown, as the trace was off the sheet on the positive side from 0 h .16 m . until 1 h .7 m . Between 23 h .40 m ., when the bay may
be said to commence, and 0 h .16 m ., when the trace went off the sheet, H rose $54 \gamma$; between 1 h .7 m . and 1 h .42 m . it fell $42 \gamma$. Between 1 h .42 m . and 2 h .12 m . H rose $42 \gamma$, again going off the sheet, but only for about a minute.

After 2 h .10 m . the Antarctic curves were quiet, according to the Antarctic standard.
The large D and H movements observed in the Antarctic after 23 h .30 m . on the 30 th synchronise with Birkeland's "elementary polar" storm. It is impossible to decide from the form of the curve alone at what instant $D$ was really most remote from the normal, but this would seem to have been at the turning-point at 0 h .56 m ., when the maximum for the day occurred. This aunswers practically to noon, L.T., thus rendering possible a comparison with mean results for that hour from adjacent days. Taking a mean from the five previous days we deduce an excess on the 31st of $110^{\circ}$, answering to about $200 \gamma$ in force. The simultaneous value of H was also substantially in excess of the normal, but how much it is impossible to say, as the trace was off the sheet and very steep both when going off and coming on. Probably the maximum value of the horizontal component of the disturbing force in the Antarctic, measured after Birkeland's method, was fairly similar to that at Axelöen, $280 \gamma$, or, say, half that at his most disturbed station, Dyrafjord. The disturbance of March 30-31, regarded from an Antarctic standpoint; was, of course, by no means a large one.
§30. The intercomparison of the Antarctic curves with Birketand's records points to the following conclusions:-

1. At the times of the small sudden movements which Birkeland assigns to "equatorial perturbations" there seem to be almost always (no certain exception has been noted) corresponding movements in the Antarctic. The Antarctic movements are larger, usually much larger, than those recorded at the equatorial or non-polar stations, and are usually, if not always, oscillatory in type.
2. At the times of Brakeland's "elementary polar" storms in the Aretic, the conditions in the Antarctic are generally more than usually disturbed. The times of the largest movements in the Antarctic usually occur not far from the times of maximum disturbance in the Arctic. Not infrequently, however, the disturbances in the Antarctic continue without any marked subsidence of intensity throughout times which Birkeland regards as including two or more elementary polar storms.
3. In the Antarctic, the sudden commencing movements which Birkeland assigns to equatorial perturbations are sometimes immediately followed by disturbances which are not separated by any markedly quiet interlude from subsequent disturbances which synchronise with Arctic disturbances which Birkeland believes of the "elementary polar" type.
4. The "elementary polar" disturbances of Birkeland seem practically all confined to the hours 13 to 2, G.M.T. (or 0 a.m. to 1 p.m., Antarctic L.T.), and so occur at a time when the Antarctic movements are usually of a rapidly oscillatory character, and avoid the hours when the rounded wave-like movements of the "special type" are usually found.

Further, Birkeland's results are limited to the Antarctic Summer, i.e. to the season at which few if any disturbances of the "special type" were recorded.

It would clearly be of great interest to know what was the nature of the phenomena in the Arctic during the times of occurrence of the Antarctic "special type" of disturbance; also whether the Arctic stations show a marked diurnal variation in the type of disturbance corresponding to that seen in the Antarctic.

One of the results of the comparison has been to make me realise even more clearly than before the desirability of much reduced sensitiveness in magnetographs intended for use in polar regions. It is clear that most fundamentally important results might be hoped for if simultaneous complete records were obtainable from a series of stations in the Aretic and Antarctic regions so situated that the effects of day and night could be adequately brought out. This comparison should extend over a complete year, so as to bring out seasonal effects.

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Term Hours, slow run, at Mauritius.

Plate III.

June 15, 1902, $\quad 9$-ı0 G.M.T.

July I, 1902, $10-$ IIG.M.T.
Term Hours, June 15, and July 1, 1902, at Christchurch.

Ilate IV.








June I. I902.


April $15,1902$.
Term Hours, April 1 and 15, and June 1, 1902, at Winter Quarters.


$.06660 \frac{T}{H_{0}+\quad H_{0}}$


August 15, 1902.



\}
1
\%
${ }^{8}$


$152^{\circ} 37^{\prime}$







September 15, 1902. October I, 1902.
Term Hours, September 1 and 15, and October 1 and 15, 1902, at Winter Quarter:

J


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-



Term Hours, January 1 and 15 , and February 1 and 15, 1903, at Winter Quarters.


- COGI'I אnenuer
- 





Harizontal force August 20, 1902.

Falmouth


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- Horizontal



Vertical force
July $26,1903$.


Plate XVIII.



Plate $X X$

$\square$



1)isturbance of August 21, 1902, at Winter Quarters.

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Plate XXXII.


B, 1903, at Winter Quarters.





May 22, 1903.


August 3r, 1902.

N

$\mathrm{Dr}^{\circ} 39^{\prime} ; \mathrm{H} 71 \gamma ; \mathrm{V}_{74 \gamma}$.

August 15, 1903.


D $\mathrm{I}^{\circ} 2 \mathrm{I}^{\prime} ; \mathrm{H} 77 \%$; Vi62 $\%$.


June 4, 1903.

$$
\begin{array}{ll}
\mathrm{D} & \mathrm{I}^{\bullet} 5 \mathrm{I}^{\prime} \\
\mathrm{H} & 45 \gamma \\
\mathrm{~V} & 113 y
\end{array}
$$



Oct 4-5, 1902.


Oct.11-12, 1902.

-

September 6-\%, 1902.


September 18-19, 1902.


D $50^{\prime}$
H 68y



February 10, 1903.


March 6, 1903.



$H_{H_{0}}^{H_{0} 14: 1161 \gamma: V_{4} 5 \gamma}$

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[^0]:    * Applied also during March. May base-line value used after Noon on April 22.
    + After August 20. September base-line value applied.

[^1]:    * The results obtained for the secular change by Commender Chetwynd, R.N., from the absolute observations alone, disregarding diurnal variation, were for Declination $-26^{\prime} 4$, for Horizontal Force $+130 \gamma$ ("Physical Observations," pp. 137, 140).

[^2]:    *Terrestrial Magnetism (under Meteorology), Art. 87.

