



TECHNICAL REPORT

A PRACTICAL METHOD OF PREDICTING
SEA ICE FORMATION AND GROWTH

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ABSTRACT

A technique for predicting ice formation and growth is presented. The ice potential calculations originally discussed by Zubov and Defant form part of the method, as do the standard heat budget equations developed by various authors. In addition, new formulas are derived for computing ice growth in terms of known or predicted oceanographic and meteorological data. The method is applied to the problem of predicting the general features of the ice distribution in the Baffin Bay-Davis Strait area for the season 1952-53. It is pointed out that use of this method must be limited to open-sea areas where winter and not polar ice is the dominating feature. However, the technique has important applications to estimating the general ice features of such areas several months in advance.

FOREWORD

The increasing importance of defense installations in northern areas has increased greatly the responsibilities of the U. S. Navy in supplying bases in Arctic waters, where sea ice is often an operating obstacle. The Hydrographic Office is charged with the responsibility of developing and testing techniques for observing and forecasting sea ice conditions. Standardized techniques for observing, charting, and reporting sea ice are now in operational use by the Navy, as described in publications issued by the Hydrographic Office. Heretofore, techniques for forecasting the formation, growth, and movement of sea ice have not been published by this Office. This publication describes a method of long-range forecasting of ice formation and growth. Since this technique is still in the developmental stage, the Hydrographic Office welcomes comments as to its operational value.



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A. INTRODUCTION

The expanding program of ice observation and forecasting at the Hydrographic Office has emphasized the desirability of long-range ice forecasting. There are many phases of the ice program for which long range forecasts giving advance estimates of ice conditions of as much as 150 days are required. This paper deals with long-range forecasting of ice thickness in open seas whose salinity and density remain relatively constant, such as Baffin Bay and the Labrador Sea.

An operational method of forecasting for a long period must be easy to use, require a minimum number of involved calculations, and yield forecasts of reasonable accuracy. Most formulas previously developed have been complicated expressions which are not suited for operational use. In developing the following method, the goal has been to perfect a technique from which quantitative results can be derived with a minimum of computation and observation.

B. THE ICE POTENTIAL OF A WATER COLUMN

The severity of an ice season depends, among other things, upon the amount of thermal energy stored in the water mass and upon the rate at which convective mixing takes place. It is necessary to select a model of convective mixing which will explain the variation in the thermohaline conditions caused by heat removal and ice accretion. The method proposed by Defant (1949) and by Zubov (1938) of computing the ice potential and potential heat loss has been adopted as a basis for the present forecasting technique.

Before a forecast can be made, an analysis of the properties that inhibit ice formation and growth must be performed. Oceanographic measurements within the area in question must be secured in order to obtain information about salinity and thermal energy stored within the water mass. Unnecessary computation can be avoided by making the measurements at a time after the heat flux reverses, when thermal energy is being continually removed from the water. Otherwise the amount of heat which was added to the water column prior to the reversal of the heat flux would have to be computed before the method of Defant and Zubov could be utilized.

Establishing the ice potential of a given water column requires only the temperature and salinity at N levels within the column. The column is divided into (N-1) layers and the mean temperature (\bar{T}_n), salinity (\bar{S}_n), and density ($\bar{\sigma}_n$), are computed for each layer. The mixing model, as defined by the method of Defant and Zubov is constructed by finding the mean temperature and salinity of the surface and second layers, i.e.,

$$\frac{\bar{S}_1 + \bar{S}_2}{2} = \bar{S}_{1,2}, \quad \frac{\bar{T}_1 + \bar{T}_2}{2} = \bar{T}_{1,2}.$$

The column is then said to be partially mixed down to the second level. Complete mixing is achieved by cooling the partially mixed column until its density, $\sigma_{1,2}$, equals that of the original second layer, σ_2 . This mixing may be accomplished merely by changing the temperature of the partially mixed column, provided the water is not cooled below its freezing point. If the temperature change necessary for completely mixing the column requires a temperature below the freezing point, ice will form and the salinity of the remaining water in the column will be increased. The mixing process is carried out layer by layer until the potential heat loss is greater than the possible heat loss for a given ice season.

For convenience the expressions "partially mixed" and "completely mixed" have been used in connection with the ice potential computations. Two layers of water are termed "partially mixed" when the mixed column is defined completely by the means of the temperatures and salinities of the two original layers (i.e., no heat loss is involved). A "partially mixed" layer becomes "completely mixed" when the density of the partially mixed layer equals that of the original lower layer, with the density of the original lower layer remaining fixed throughout the process. In this latter case a heat loss is necessary, provided that the density distribution is stable. Temperature and salinity changes which result from this mixing model can be easily read from a T-S nomograph, or can be computed by use of standard hydrographic tables. (H.O. Pub. No. 615)

Complete mixing can be accomplished only by releasing q_T gram calories of sensible heat from the water column one square cm in cross section. Quantitatively, $q_T(h) = c \rho_w h \Delta T_w$ where c is the specific heat of sea water ρ_w is the density of sea water, h is the depth to which complete mixing reaches (or volume of $h \text{ cm}^3$), and ΔT_w is the amount of heat lost in changing the partially mixed column to a completely mixed column. When it is necessary to change the salinity of the column in order to achieve complete mixing, latent heat is involved in the process as well as sensible heat. Since a known percentage of the salt is frozen out of the ice, it is possible to write an expression for the latent heat as a function of the salinity change. Consequently, the ice accretion corresponding to a given salinity change can be expressed mathematically. As written by Defant, the ice thickness is given by

$$\xi_i = \frac{h \Delta S}{C \rho_w b S} \quad (1)$$

where S is the salinity of the partially mixed column and b is the percentage of the salts frozen out of the ice. The latent heat can now be evaluated from the well-known formula $q_\xi(Z) = K\xi(Z)$, where K is the latent heat of fusion, which decreases with increasing salinity of the ice. The total potential latent heat is given by

$$\int_0^h q_\xi dz = Q_l(h) \quad (2)$$

As the mixing model is developed, each layer will yield a certain amount of potential heat loss, q_T , and eventually, when the temperatures are low and the densities are high, each layer will yield a certain mass of potential ice (ξ) associated with convective mixing down through the corresponding layer. Thus it is possible to construct the very useful "potential curve" (Figure 2) from corresponding

$$\int_0^h \xi dz = l_i(h) \quad \text{and} \quad Q_T(h) = \tilde{Q}_T(h) - Q_0, \quad \text{where} \quad \tilde{Q}_T = \int_0^h q_T dz. \quad (3)$$

The potential curve is plotted, Q_T starting from zero within the layer where ice first appears. The remainder of the \tilde{Q}_T, Q_0 , which is lost before ice is formed, determines the date of ice formation. The reason for taking Q_T as zero at this point becomes evident when one considers the lower limits of the integrals in equation (13). Obviously some question arises as to the exact quantity of heat that has been removed at the time ice is first formed. There is, of course, some error introduced, but it is not large as long as the water temperatures are near the freezing point and the water column is divided into layers of sufficiently small depth, for example five to ten meters. Recent work on the same type of model as Defant's has yielded exact values of the Q_T which must be removed before ice is formed, thus determining the proper point to start accumulating the Q_T for the potential curve (Brown, 1954).

In order to illustrate the computation of the ice potential and the drawing of the potential curve, a typical oceanographic station is used as an illustration. Station 37, at 66°N , 58°W , was occupied on 3 October 1952. A plot of the oceanographic variables at this station is given in Figure 1, and the complete numerical data are shown in Table 1.

TABLE 1
Oceanographic Data for Station #37
(66°N , 58°W), 3 October 1952

Depth m.	Temperature °C.	Salinity ‰	Density Anomaly (σ_t) gm/liter
0	0.56	32.37	25.98
5	0.10	32.36	26.00
10	0.61	32.39	25.99
20	0.73	32.47	26.05
30	-0.41	32.78	26.36
50	-1.65	33.08	26.64
75	-1.61	33.69	27.13
99	-2.13	33.69	27.14
148	-0.52	34.04	27.38
197	1.52	34.21	27.40
296	0.08	34.34	27.59
394	-1.71	34.52	27.81



FIG. I OCEANOGRAPHIC PLOT FOR STATION NO. 37, 66°N 58°W, 3 OCTOBER 1952.

From these data the ice potential has been calculated by the method of Defant and the results shown in Table 2A. Finally, the ice potential curve for Station 37 is shown in Figures 2. It will be noted that the ice potential curve is by no means a straight line or a smooth curve. In fact, for this station, the heat loss necessary to form a given amount of ice varies considerably. For this reason, the actual forecasting for this station is difficult, for a small additional heat loss at about 2 kg. cal. causes the formation of a large amount of ice. In this respect the example shown is not typical of many stations where the ice potential curve is quite regular and smooth; however, it is included to illustrate the kind of difficulties encountered in practice. In addition, the curve is atypical because Q_0 , the heat loss necessary to bring about formation of ice was found exactly at 25 meters when the temperature of the completely mixed layer (1-4, Table 2B) first dropped to -1.8°C . In general, Q_0 cannot be so easily determined.

C. METHOD OF COMPUTING DATE OF ICE FORMATION

There are two acceptable methods of computing the date of ice formation. Since the amount of heat that must be lost before ice is formed is known from the ice potential computations, it is only necessary to compute the time required to lose this heat. One method is to make actual observations of the heat loss per day per square cm. and thus to find the mean heat loss which is representative of a given area. The other method is to compute the total heat loss by consideration of the climatological data, the sun's altitude, and oceanographic data. For this purpose the only known formulas are those of Jacobs (1942), which yield rates of heat loss of the right order of magnitude in the Arctic, although they were specifically derived for middle latitudes.

Jacobs' formulas for computing heat losses are:

$$Q_{ab} = .025 \bar{\alpha} \left[.29 + .71(1 - \frac{\bar{c}}{100}) \right] (t - r) t, t \text{ in minutes} \quad (4)$$

$$Q_b = .94 \left[\text{eff. } Q_b (1 - .0083\bar{c}) \right] t, \quad t \text{ in minutes} \quad (5)$$

$$Q_e = 145.4 \bar{w} (e_w - e_a) \left[1 + .01 \left\{ \frac{t_w - T_a}{e_w - e_a} \right\} \right] t, t \text{ in days}, \quad (6)$$

where

Q_{ab} = total incoming radiation in gm. cal/cm.²

Q_b = total back radiation in gm.cal/cm.²

Q_e = total heat loss due to evaporation in gm.cal/cm.²

$\bar{\alpha}$ = mean altitude of the sun in degrees,

\bar{c} = mean cloud amount in percent,

\bar{w} = mean wind velocity in knots,

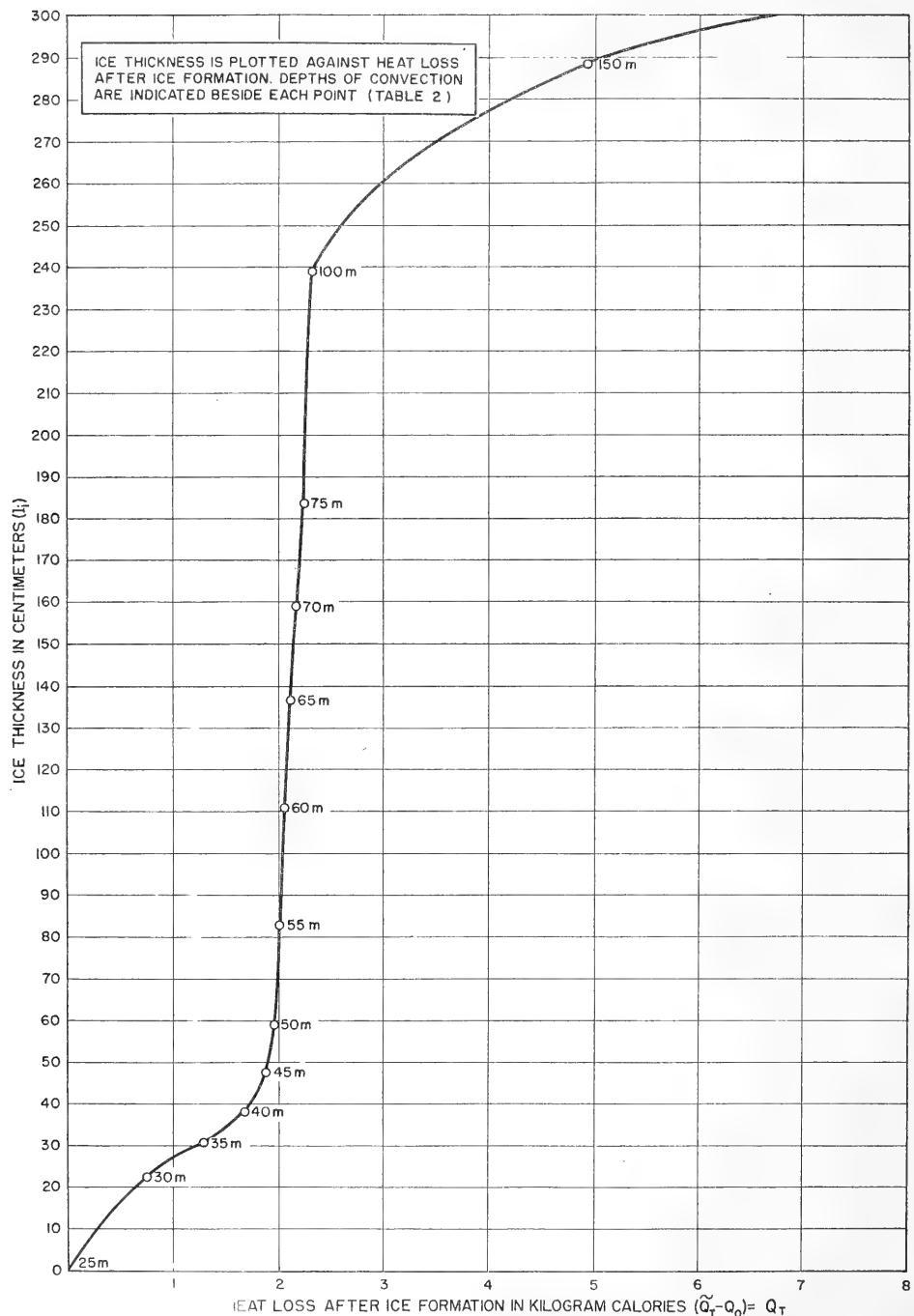


FIG. 2 ICE-POTENTIAL CURVE FOR STATION NUMBER 37.

e_w = vapor pressure of the water in inches Hg.,
 e_a = vapor pressure of the air in inches Hg.,
 T_w = temperature of the water in degrees F.,
 T_a = temperature of the air in degrees F.
 eff. Q_b (effective back radiation) = difference between
 temperature radiation of the sea surface and long-wave radiation from the
 atmosphere,

t = time in units designated above, and
 r = fractional part of the incoming radiation reflected from
 the sea surface.

The ice potential computations determine the amount of sensible heat which must be removed from the water column prior to the initial formation of ice. For example, let this value be Q_p . Then setting $Q_p = Q_b - Q_e - Q_{ab}$ and solving for t gives the number of days between the time that the oceanographic measurements were made and the time ice will form.

D. METHOD OF COMPUTING ICE GROWTH

From the time oceanographic observations are made until the time ice forms, it is assumed that the heat exchange results from incoming radiation, back radiation, and evaporation. After ice forms, the processes by which heat is removed from the water column are altered considerably. Any heat that is then removed from the water must be conducted through the solid layer of ice above it. Thus, the rate of further heat loss is determined by the temperature gradient through ice and snow. Hence, assuming steady-state conditions for finite increments of time,

$$\frac{d(Q_l + Q_T)}{dt} = k_i \frac{\Delta T_i}{l_i} = k_s \frac{\Delta T_s}{l_s} \quad (7)$$

where, Q_T is the amount of sensible heat loss (in kg. cal.),
 Q_l is the amount of latent heat loss (in kg. cal.),
 k_i is the heat conductivity coefficient of sea ice,
 k_s is the heat conductivity coefficient of snow,

ΔT_i is the temperature difference between the upper surface of the ice and freezing point of sea water (in °C.),

ΔT_s is the temperature difference between the surface of the snow and the surface of the ice (in °C.),

l_i is the thickness of the ice in cm,

l_s is the thickness of the snow in cm, and

\bar{l}_s is the mean snow cover over a finite period of time.

Since $\Delta T = \Delta T_i + \Delta T_s$, where ΔT is the temperature difference between the freezing point of sea water (taken at $-1.8^{\circ}\text{C}.$) and the snow surface temperature, and from (7)

$$\frac{\Delta T_s}{\Delta T_i} = \frac{k_i l_s}{k_s l_i} \quad (8)$$

then

$$\Delta T = \Delta T_i \left\{ 1 + \frac{k_i l_s}{k_s l_i} \right\} \quad \text{or} \quad \Delta T_i = \frac{\Delta T}{1 + \frac{k_i l_s}{k_s l_i}} \quad (9)$$

Then (7) becomes:

$$\frac{d(Q_1 + Q_T)}{dt} = \frac{k_i \Delta T}{l_i \left\{ 1 + \frac{k_i l_s}{k_s l_i} \right\}} \quad (10)$$

Substituting $Q_1 = \rho_i K l_i$ into (10) (ρ_i denotes ice density)

$$\rho_i K \frac{dl_i}{dt} + \frac{dQ_T}{dt} = \frac{k_i \Delta T}{l_i \left\{ 1 + \frac{k_i l_s}{k_s l_i} \right\}} \quad (11)$$

$$\rho_i K \left\{ 1 + \frac{k_i l_s}{k_s l_i} \right\} l_i dl_i + \left\{ 1 + \frac{k_i l_s}{k_s l_i} \right\} l_i dQ_T = k_i \Delta T dt \quad (12)$$

$$\rho_i K \int_0^{l_i(h)} l_i(Z) dl_i(Z) + \frac{\rho_i K k_i}{k_s} \int_0^{l_i(h)} l_s dl_i(Z) + \int_0^{Q_T(h)} l_i(Z) dQ_T(Z) +$$

$$\frac{k_i}{k_s} \int_0^{Q_T(h)} l_s dQ_T(Z) = k_i \int_{t_0}^t \Delta T dt \quad (13)$$

$$\frac{\rho_i K}{2} l_i^2(h) + \frac{\rho_i K k_i}{k_s} \bar{l}_s l_i(h) + \int_0^{Q_T(h)} l_i(Z) dQ_T(Z) + \frac{k_i}{k_s} \bar{l}_s Q_T(h) = k_i \int_{t_0}^t (T_F - T) dt \quad (14)$$

The right-hand expression in equation 14 is usually denoted by the title "degree-days of frost." A degree-day of frost is defined as a day with a mean temperature 1-degree Centigrade below the freezing point of sea water. While the freezing temperature of sea water varies with salinity, it is here assumed to be -1.8°C . A degree-day of frost is then 1 day with a mean temperature of -2.8°C . Degree-days of frost usually are accumulated for forecasting over periods of 15 or 30 days. By giving the following values to some of the terms in equation 14:

$$\begin{aligned} l_i &= 0.9 \\ K &= .080 \quad \frac{\text{Kg Cal}}{\text{gm}} \\ k_i &= .389 \quad \frac{\text{Kg Cal}}{\text{cm }^{\circ}\text{C day}} \\ k_s &= .062 \quad \frac{\text{Kg Cal}}{\text{cm }^{\circ}\text{C day}}, \end{aligned}$$

the formula can be used directly to forecast the ice in terms of degree-days of frost

$$\int_{t_0}^t (T_F - T) dt.$$

Thus, for a given value of \bar{l}_s all terms in the equation are known except l_i and Q_T . A given l_i and likewise the corresponding Q_T , can be selected from the potential curve. By substituting l_i and Q_T into equation (14) and by evaluating the area under the potential curve between the proper limits for the expression

$$\int_0^{Q_T(h)} l_i^T(Z) dQ_T(Z)$$

the number of degree-days of frost required to form l_i centimeters of ice can be computed.

To eliminate the necessity of computation, the end results of the derivation have been presented in the form of tables (Tables 3-12). In addition the integral

$$\int_0^{Q_T(h)} l_i^T(Z) dQ_T(Z)$$

had to be evaluated by approximation methods. It was found from the majority of potential curves that the integral could be approximated as

$$\frac{l_i}{2} Q_T(h),$$

The error introduced was found to be small unless the water temperature was unusually high, and in most cases ice will not form under such conditions. Since the potential curve $l_1(Q_T)$ must always be plotted, it is readily apparent when the above approximation is invalid; in this situation the formula can be used. In order to facilitate substitution of the actual integral for the approximation, an auxiliary table (Table 13) is included which gives the degree-days of frost associated with various values of l_1 and Q_T . If the potential curve is very irregular, Table 13 is entered, and by comparison with the degree-day figure computed from an average l_1 , one can determine the amount of error introduced into the total. Most frequently this will be found to be small, since the potential term is only one of four separate factors contributing to the total; however, the actual value can be added or subtracted and a more accurate figure secured.

The form of equation 14 which was used in the tables is

$$\int_{t_0}^t (T_F - T) dt = \frac{1}{k_1} \left[\frac{\rho_1 K}{2} l_1^2 + \left\{ \frac{K k_1 l_s}{k_s} + \frac{Q_T}{2} \right\} l_1 + \frac{k_1 l_s}{k_s} Q_T \right]. \quad (15)$$

The degree-days of frost for fixed $l_s = 0, 2.5, 5, 7.5, 10, 15, 20, 30, 40$, and 50 cm. are given in Tables 3 to 12. The concept of degree-days has an advantage over one of calendar days because of its versatility as far as forecasting is concerned. An ice forecast in days is difficult because it requires a forecast of air temperatures as well as ice growth, while the degree-day forecast permits ice growth forecasts which fit all conceivable temperature variations.

E. USE OF THE ICE GROWTH TABLES IN FORECASTING

Forecasting the ice growth in a given area requires the following information: (1) oceanographic data taken after the water mass begins to cool and prior to ice formation, (2) the date of ice formation, and (3) an expression relating ice growth to degree-days of frost for the given area. With this information the forecaster can then proceed to make a forecast for the entire season. The procedure is as follows: (1) the ice potential of each oceanographic station is calculated and the ice potential curve (example, Figure 2) is drawn; (2) from available meteorological and oceanographic data, the date of ice formation is determined; and (3) the forecast is made in terms of degree-days of frost.

Since the calculation of the ice potential, the date of first ice formation, and the ice potential curve have been previously explained, the remaining discussion will describe the preparation of an actual forecast. As before, Station #37 is used as an example. The ice potential curve for this station is given in Figure 2. In order to draw the ice forecast curves for this station it is necessary to read the ice thickness figures corresponding to integral values of Q_T from the ice potential curve. The degree-days of frost corresponding to each combi-

nation of l_1 and Q_T values are read from each of the tables 3 to 12. A plot of ice thickness vs degree-days of frost is made from each set of values obtained from the graphs, and a smooth curve is drawn for each value of snow depth (l_s). An example of a completed forecast is given in Figure 3. In Figure 3, $Q_T = 1$ corresponds to an ice thickness of 27 cm. and $Q_T = 2$ to an ice thickness of 72 cm. on the ice potential curve of Figure 2. To convert the completed forecast into a time forecast the meteorologist must provide temperature and snow depth forecasts in order to secure an ice thickness forecast of 120 to 150 days for this station.

Now let us suppose that the date of first ice formation at Station #37 is 12 November and that the meteorologist provides a forecast, including the following data for the succeeding months:

Period	Mean Temperature (°C)	Degree-Days of Frost	Snow Depth (cm.)
12 - 15 Nov.	-5.8	16	0
16 - 30 Nov.	-4.8	45	0
1 - 15 Dec.	-8.8	105	2.5
16 - 31 Dec.	-7.8	96	2.5
1 - 15 Jan.	-9.8	120	7.5
16 - 31 Jan.	-18.8	272	7.5

The ice thickness for any given date may be found from the above data and the forecast curve.

During the period when the snow depth is zero, the growth of ice will be along the upper curve on Figure 3; on 30 November, the final day with no snow, the total of degree-days of frost is $16 + 45$ or 61 and the ice thickness will be 17 cm. (point A on Figure 3).

During the month of December the ice grows under a snow depth of 2.5 cm. from a starting point of 17 cm. This point is marked A'. During December a total of 105 plus 96 or 201 degree-days of frost is counted, so that it is necessary to follow the curve $l_s = 2.5$ for 201 degree-days of frost beyond point A', or to point B. At point B, reached 31 December, the ice thickness will be 35 cm.

Similarly, during January the ice grows under a snow depth of 7.5 cm. for 120 plus 272 or 392 degree-days beyond the point B' corresponding to an ice thickness of 35 cm. On 31 January the point C is reached, and the ice thickness will be 53 cm. This constitutes an operational forecast, as the user of the forecast can determine the ice thickness on any day between 12 November and 31 January by counting the number of degree-days of frost and noting the snow depth applicable to the situation.

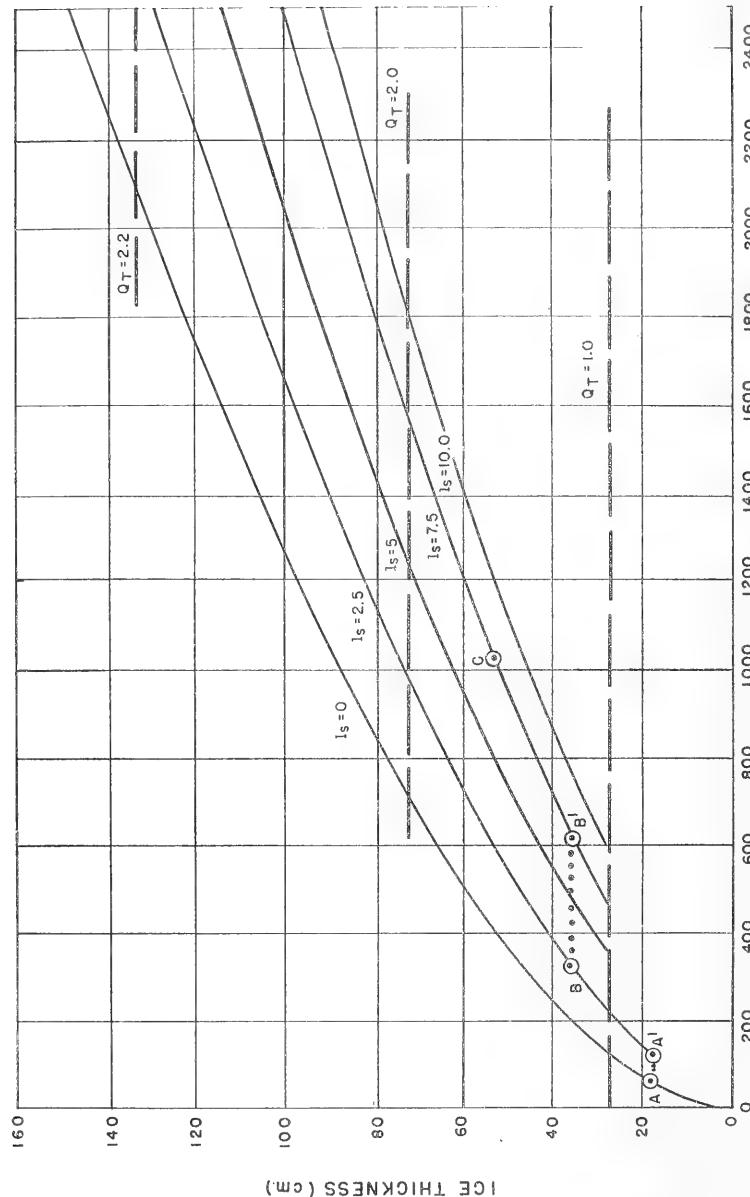


FIG. 3. ICE GROWTH FORECAST FOR STATION # 37

F. BAFFIN BAY-DAVIS STRAIT LONG-RANGE ICE FORECAST FOR SEASON 1952-53.

During early autumn, 1952, oceanographic observations were taken at the locations shown on Figure 4. The method explained previously was employed to forecast the ice formation and growth of these locations. The dates of ice formation were computed and are shown beside each station. Since the thermohaline structure at some of the stations was such that no ice would be formed, it was possible to delineate the theoretical extreme ice boundary shown as a heavy line in Figure 4. This boundary agrees very well with the boundary which was actually observed in late March and early April 1953, (shown as a dashed line in Figure 4) except for the occurrence of ice south of Stations 39 and 40. It is possible that ice which was observed at these two stations was not formed there, but had drifted from nearby areas of ice formation. The dispersion of ice near the observed boundary in this area indicates that the ice actually was formed outside the area.

CONCLUSIONS

The methods of Zubov and Defant for ice potential calculation have proved in practice to give reasonable answers for open seas and for in-shore areas where local variations in the physical properties of the water are not large. In harbors and areas where runoff is important, changes in salinity and density are so rapid that it is difficult to attach a meaning to an average value of these parameters. Hence any ice potential calculation based on one or a few observations in a harbor may give a potential ice growth which is unreasonably high or low for a single season.

The forecast of ice growth based on the ice potential can be used only during the period when ice thickness is increasing. No theory has been included which accounts for decreasing ice thicknesses during the breakup period. Indeed, no mathematical theory has yet been developed for the breakup of ice.

Many refinements of the method presented in this paper are possible. Some of them are discussed in a paper by Brown (1954). The method described here was used to forecast during the 1952-53 ice season; the forecast proved to be satisfactory. The most important advancement discussed in this report is the use of the ice potential as a basis for forecasting ice growth.

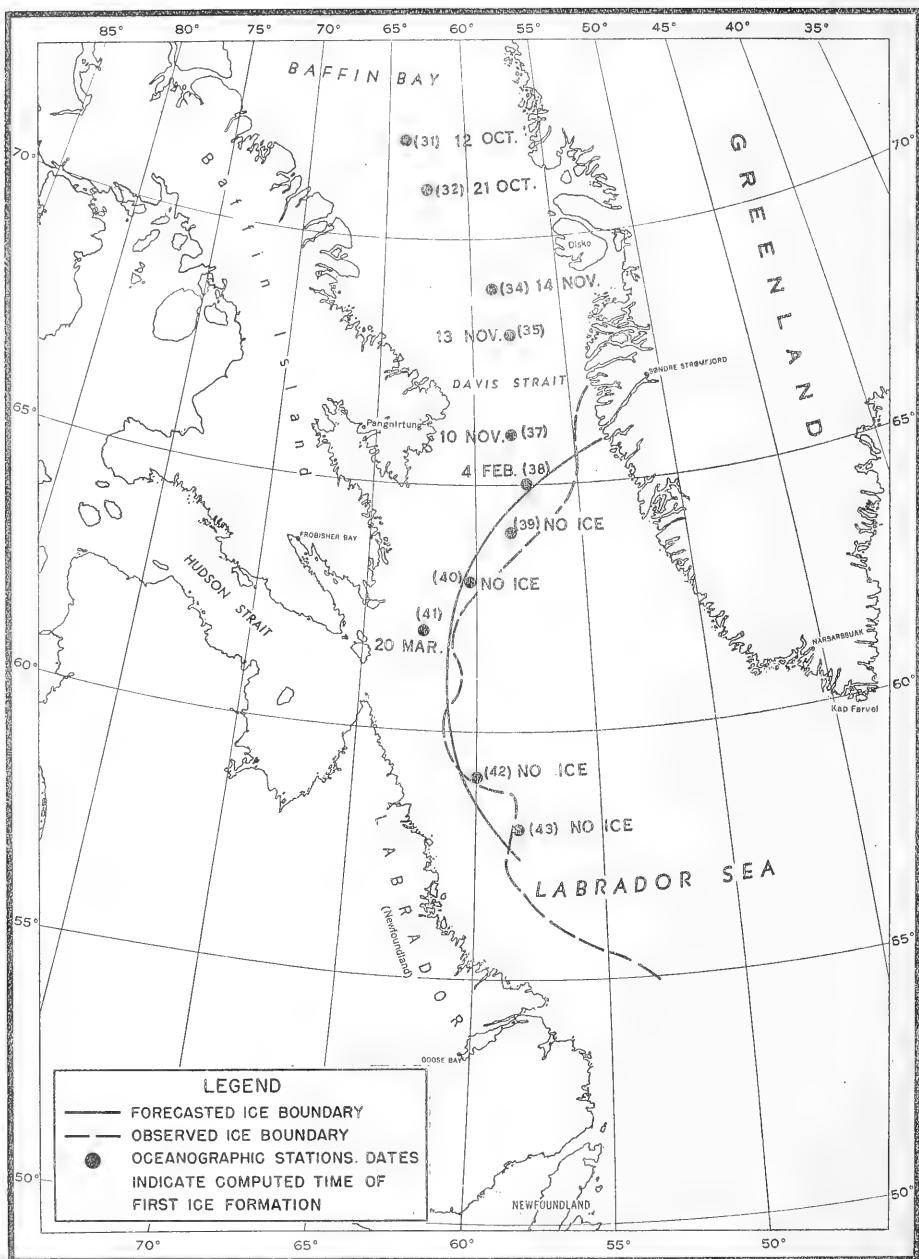


FIG. 4. FORECASTED AND OBSERVED ICE BOUNDARIES IN BAFFIN BAY AND DAVIS STRAIT FOR THE SEASON 1952-1953

TABLE 7

DEGREES-DAYS OF FROST ASSOCIATED WITH GIVEN ICE THICKNESSES AND HEAT LOSSES FOR SNOW COVER = 10 cm.

Heat Loss in kg. cal. ($\frac{1}{cm}$)										
	6	7	8	9	10	11	12	13	14	15
0	162	223	314	415	506	597	687	777	867	957
5	60	925	1396	1754	2131	2399	2657	2917	3174	3432
10	125	195	276	354	431	509	587	665	743	821
15	195	375	556	737	917	1098	1279	1460	1641	1822
20	265	455	635	820	1007	1197	1387	1573	1765	1957
25	332	632	932	1231	1531	1831	2131	2431	2731	3031
30	400	725	1119	1315	1511	1707	1904	2101	2301	2501
35	470	925	1398	1751	2121	2484	2847	3207	3567	3927
40	540	1025	1495	1855	2215	2575	2935	3295	3655	4015
45	610	1145	1597	1906	2255	2613	2971	3329	3687	4045
50	680	1263	1719	2014	2315	2675	3034	3394	3754	4114
55	750	1151	1313	1454	1707	2074	2435	2794	3154	3514
60	1230	1263	1507	1715	1984	2222	2480	2699	2937	3175
65	1145	1391	1636	1830	2125	2370	2650	2904	3199	3494
70	1266	1518	1769	2020	2272	2523	2774	3025	3276	3526
75	1325	1625	1897	2145	2422	2650	2923	3195	3471	3749
80	1521	1785	2050	2311	2573	2842	3116	3370	3634	3934
85	1656	1926	2197	2457	2711	2908	3279	3550	3820	4091
90	1795	2072	2349	2626	2903	3180	3457	3731	4011	4288
95	2222	2505	289	3072	3255	3536	3826	4115	4405	4695
100	2378	2666	2956	3246	3536	3826	4115	4405	4695	4985
110	3003	3305	3637	3938	4230	4523	4821	5121	5421	5721
120	2726	3012	3357	3673	4010	4359	4795	5250	5656	6051
130	3074	3102	3170	3559	4337	5011	5717	6501	7309	8105
140	3440	3424	3453	3811	4505	5175	5857	6536	7193	7876
150	3824	3824	3824	3824	4595	5264	5934	6564	7193	7876
160	4227	4594	528	5695	6303	6957	7611	8217	8827	9433
170	4645	5029	5605	5988	6629	7295	7950	8555	9151	9757
180	5069	5462	5974	6367	6960	7552	8145	8747	9349	9951
190	5547	5953	6358	6764	7110	7575	8130	8623	9015	9408
200	6022	6411	6459	6778	7096	7495	7890	8286	8677	9062
210	6520	6951	7382	7914	8245	8714	9214	9614	10008	10396
220	7034	7478	7922	8366	8937	9459	9958	10401	10896	11391
230	7567	8024	8480	8937	9391	9851	10308	10767	11227	11686
240	8118	8728	9057	9327	9653	10135	10621	11111	11596	12074
250	8687	9110	9505	9827	10266	10757	11255	11756	12254	12753
260	9255	9771	10266	10757	11255	11756	12254	12753	13252	13751
270	9882	10390	10877	11365	11853	12341	12838	13326	13814	14303
280	10507	11014	11501	11998	12485	12972	13460	13947	14434	14921

(°C.)

TABLE 10
DEGREE-DAYS OF FROST ASSOCIATED WITH GIVEN DEPTHSSESSES AND HEAT LOSSES FOR SNOW COVERS = 30 cm.

		Heat Loss kg. cal. / °C.																
		1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	20	
0	0	134	362	1152	1936	2619	2903	3387	3971	4355	4839	5323	5906	6290	6774	7353	9677	
5	5	176	667	1117	1664	2138	2622	3118	3509	3999	4539	5080	5570	6060	6550	7041	7531	9982
10	10	258	851	1848	2344	2911	3338	3735	4314	4823	5325	5822	6313	6815	7312	7808	16292	
20	20	734	1213	1753	2262	2772	3282	3701	4191	4710	5320	5830	6339	6849	7353	7868	5377	
30	30	1126	1651	2173	2696	3213	3741	4242	4736	5233	5330	5853	6875	7398	7920	8442	9965	
40	40	1512	2077	2612	3147	3653	4218	4753	5232	5721	6259	6894	7130	7665	8500	9036	9571	
50	50	1923	2522	3070	3618	4166	4771	5252	5720	6253	6906	7406	8003	8551	9292	9647	10325	
60	60	2424	2981	3515	4126	4747	5228	5739	6250	6742	7472	8033	8594	9155	9716	10277		
70	70	3166	4010	4611	5216	5813	6312	6812	7311	7811	8311	8811	9311	9811	10312	10778		
80	80	3372	3966	4553	5110	5716	6313	6912	7511	8111	8711	9311	9811	10311	10830	10420		
90	90	3885	4485	5081	5681	6283	6883	7483	8083	8683	9283	9883	10383	10983	11480	10430		
100	100	4409	5032	5634	6247	6859	7472	8081	8681	9281	9881	10381	10981	11581	12180	11420		
110	110	4922	5577	6203	6828	7453	8076	8706	9326	9926	10526	11126	11726	12326	12926	13526		
120	120	5513	6151	6779	7428	8066	8701	9342	9942	10542	11142	11742	12342	12942	13542	14142		
130	130	6109	6714	7395	8016	8697	9318	9929	10550	11150	11750	12350	12950	13550	14150	14750		
140	140	6691	7354	8019	8683	9317	9921	10513	11105	11705	12305	12905	13505	14105	14705	15305		
150	150	7368	7935	8661	9333	9915	10512	11102	11702	12302	12902	13502	14102	14702	15302	15902		
160	160	7913	8533	9232	9900	10502	11102	11702	12302	12902	13502	14102	14702	15302	15902	16502		
170	170	8270	9085	9790	10502	11202	11902	12602	13302	14002	14702	15402	16102	16802	17502	18202		
180	180	8940	9750	10550	11350	12150	12950	13750	14550	15350	16150	16950	17750	18550	19350	20150		
190	190	9240	10050	10850	11650	12450	13250	14050	14850	15650	16450	17250	18050	18850	19650	20450		
200	200	10610	11610	12610	13610	14610	15610	16610	17610	18610	19610	20610	21610	22610	23610	24610		

(T) = Effective time

TABLE II
DEGREE-DAYS OF FROST ASSOCIATED WITH GIVEN ICE THICKNESSES AND HEAT LOSSES FOR SNOW COVER = 10 cm.

	0	1	2	3	4	5	6	7	8	cal. (70°)	9	10	11	12	13	14	15
0	615	1290	1936	2571	3226	3971	4716	5516	6316	6152	6906	7652	7097	7742	8387	9032	9677
5	395	1558	2189	2811	3422	4134	4846	5547	6250	6099	6796	7547	7402	8054	8705	9357	10008
10	1174	1132	1790	2448	3106	3764	4422	5080	5738	6396	7054	7712	8370	9028	9686	10358	
15	966	1637	2308	2979	3549	4230	4991	5662	6333	7004	7675	8346	9016	9687	10357		
20	1177	2151	2811	3528	4212	4895	5579	6263	6947	7630	8314	8998	9681	10355			
30	2066	2703	3399	4095	4792	5489	6186	6882	7579	8275	8972	9669	10365				
50	2554	3264	3973	4682	5392	6101	6811	7520	8229	8939	9648	10358					
60	3120	3813	4565	5287	6009	6732	7454	8176	8899	9621	10313						
70	3705	4410	5175	5910	6616	7321	8036	8851	9586	10321							
90	4308	5056	5804	6552	7300	8048	8796	9541	10292								
90	4920	5691	6452	7213	7974	8734	9495	10256									
100	5571	6344	7118	7892	8665	9439	10213										
110	6230	7016	7803	8589	9376	10152											
120	6907	7706	8506	9305	10104												
130	7603	8445	9227														
140	8317	9142	9967														
150	9050	9888															
160	9302	10652															
170	10571																
ICE THICKNESS IN CM. (L)																	

TABLE 12
DEGREE-DAYS OF FROST ASSOCIATED WITH GATE LINE THICKNESSES AND HEAT LOSSES FOR SHOT COVER = 50 cm.
(L.F.)

		Heat Loss kg. cal. (Q_1)												
		1	2	3	4	5	6	7	8	9	10	11	12	13
0	0	806	1612	2019	3226	1032	1839	5645	6152	7258	8061	8871	9677	10484
5	5	293	1105	1918	2731	3314	1357	5110	5233	6796	7608	8121	9231	10047
10	10	590	1409	2228	3018	3867	1686	5568	6325	7144	7961	8783	9632	10352
20	20	1198	2030	2937	3695	4527	5359	6191	7023	7856	8638	9520	10375	
30	30	1825	2670	3515	4360	5205	6050	6895	7740	8585	9430	10275		
50	50	3135	4005	4876	5914	5902	6760	7618	8476	9334	10191			
60	60	3817	4701	5717	6613	7483	8359	9230	10100					
70	70	4518	5514	6511	7511	8511	9511	10511	11511					
80	80	5238	6147	7056	7965	8815	9784	10693	11693					
90	90	5975	6898	7820	8742	9664	10472	11350	12222					
100	100	6732	7667	8602	9537	10472	11350	12222	13122					
110	110	7507	8455	9323	10222	11122	12022	12922	13822					
120	120	8300	9261	10122	11022	11922	12822	13722	14622					
130	130	9112	10086	10956	11822	12722	13622	14522	15422					
140	140	9913	10886	11756	12622	13522	14422	15322	16222					
150	150	10792	11766	12636	13502	14402	15302	16202	17102					

TABLE 13

DEGREE-DAYS OF FROST FOR DIFFERENT HEAT LOSS AND ICE THICKNESS ($\frac{Q_T}{2k_1}$)

$\frac{Q_T}{2k_1}$	$\frac{I^2}{c_1}$	I^2	c_1	Heat Loss (Hg. Cal.)											
5	6	13	19	26	32	39	45	51	58	64	71	77	84	90	96
10	13	26	39	51	64	77	90	103	116	129	141	154	167	180	193
20	26	51	77	103	129	154	180	206	231	257	283	308	334	360	386
30	39	77	116	154	193	231	270	308	347	386	424	462	501	540	578
40	51	103	154	207	257	308	360	411	463	514	566	617	668	720	771
50	64	129	193	257	321	386	450	514	578	643	707	771	835	900	964
60	77	154	231	308	386	463	540	617	694	771	848	925	1003	1080	1157
70	90	180	270	360	450	540	630	720	810	900	990	1080	1170	1260	1350
80	103	206	308	411	514	617	720	823	923	1029	1131	1234	1337	1440	1542
90	116	231	347	463	578	694	810	925	1041	1157	1272	1388	1504	1620	1735
100	129	257	386	514	643	771	900	1029	1157	1285	1414	1542	1671	1799	1928
110	141	283	424	566	707	848	990	1131	1272	1414	1555	1697	1838	1979	2121
120	154	308	463	617	771	925	1080	1234	1388	1542	1697	1851	2005	2159	2314
130	167	334	501	668	835	1003	1170	1337	1504	1671	1838	2005	2172	2339	2506
140	180	360	540	720	900	1080	1260	1440	1620	1799	1979	2159	2339	2519	2699
150	193	386	578	771	964	1157	1350	1542	1735	1928	2121	2314	2506	2699	2892
160	206	411	617	823	1029	1234	1440	1645	1851	2057	2262	2468	2674	2879	3085
170	219	437	656	874	1093	1311	1530	1748	1967	2185	2404	2622	2844	3059	3278
180	231	463	694	925	1157	1388	1620	1851	2082	2314	2545	2776	3008	3239	3470
190	244	488	733	977	1221	1465	1710	1954	2198	2442	2686	2931	3175	3419	3663
200	257	514	771	1028	1285	1542	1799	2057	2314	2571	2828	3085	3342	3599	3896
210	270	540	810	1080	1350	1620	1889	2159	2429	2699	2969	3229	3509	3779	4049
220	283	566	848	1131	1414	1697	1979	2262	2545	2828	3111	3393	3676	3959	4212
230	296	591	887	1183	1478	1774	2069	2365	2661	2956	3252	3548	3843	4139	4434
240	308	617	925	1234	1542	1851	2159	2468	2776	3085	3393	3702	4010	4319	4627
250	321	643	964	1285	1607	1928	2249	2570	2892	3213	3535	3856	4177	4499	4830
260	334	668	1003	1447	1671	2085	2339	2674	3008	3342	3676	4010	4344	4679	5013
270	347	694	1041	1388	1735	2082	2429	2776	3123	3470	3817	4165	4512	4859	5206
280	360	720	1080	1440	1799	2159	2519	2879	3239	3599	3959	4319	4679	5039	5398
290	373	746	1118	1491	1864	2237	2609	2982	3335	3728	4100	4473	4856	5219	5591
300	386	771	1157	1542	1928	2314	2699	3085	3470	3856	4242	4627	5013	5398	5784

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A technique for predicting ice formation and growth is presented, using the ice potential calculations and heat budget equations of various authors, new formulas are derived for computing ice growth in terms of known or predicted oceanographic and meteorological data. The technique is applied to forecasting ice conditions in Baffin Bay-Davis Strait area for 1952-53.

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- iii. H. O. TR-4





