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STRATIGRAPHY AND PETROGRAPHY OF ILLINOIAN AND KANSAN DRIFT IN CENTRAL ILLINOIS

W. Hilton Johnson

ABSTRACT

The Pleistocene deposits of pre-Wisconsinan age in central Illinois include Kansan till, Yarmouth Soil, several Illinoian units (Petersburg Silt, Jacksonville till, Roby Silt, and Buffalo Hart till), and Sangamon Soil.

Petrographic evidence indicates that the Kansan till in central Illinois probably was deposited by two lobes of ice, one the Lake Erie lobe from the east and the other probably from the northeast and possibly from the Lake Michigan basin. The Lake Erie lobe till contains more calcite, more illite and less montmorillonite, and more Na-Ca feldspar and less quartz. It is coarser textured than the possible Lake Michigan lobe till.

The Petersburg Silt, primarily loessial in origin, is fossiliferous, montmorillonitic, and dolomitic. The Jacksonville till is divided petrographically into two parts. The lower part may include Mendon till and is called Jacksonville-Mendon. It is finergrained and contains more montmorillonit and less carbonate than the typical Jacksonville till. Roby Silt includes lacustrine deposits that lie stratigraphically between Jacksonville and Buffalo Hart drifts and is composed of a thin basal sand, a fossiliferous lower silt, a distinctive clay, and an upper silt. The entire unit is calcareous, and the clay-mineral fraction is largely illite and chlorite. Buffalo Hart till is similar petrographically to Jacksonville till except where incorporation of Roby Silt has modified its charcter.

The contrast between the relatively fresh glacial morphology of the Buffalo Hart drift plain and the subdued topography of the Jacksonville surface and the presence of lacustrine deposites between the two drifts suggest the possibility of a significant break in the Illinoian glacial sequence at this stratigraphic position. However, no weathered zone was found between the drifts, and the degree of development of the Sangamon Soil does not vary greatly on the two drifts. 2

INTRODUCTION

Central Illinois is the type region for the Illinoian Stage of glacial deposits, and the area between Springfield and Decatur is particularly important for the study of their stratigraphic succession. In this area occur the youngest Illinoian sediments and the type Sangamon Soil.

These deposits form the near surface materials in much of Illinois, and, therefore, a knowledge of their occurrence and petrography is important for many purposes. The principal objective of the present investigation was to determine, parily by use of mineral compositions, the stratigraphic relations within the Illinoian sequence. A particular search was made for evidence of weathering to evaluate the intervals between the Illinoian substages. The character of the Kansan drift also was studied.

The area under consideration includes a portion of the Illinoian drift plain east of the Illinois River where the Buffalo Hart is the surface drift and also a narrow strip south and west where the Jacksonville is the surface drift (fig. 1). The northeast boundary of the area is the outer margin of the Shelbyville Moraine of the Wisconsinan glaciation. The area includes portions of Christian, DeWitt, Logan, Macon, Mason, Menard, Sangamon, and Shelby Counties.

Early investigations of the glacial deposits in the area were by Bannister (1870 a; 1870 b), Worthen (1873), and Broadhead (1875). Leverett (1899) made a comprehensive study of the area and named the Buffalo Hart Moraine. Shaw and Savaqe (1913) made a small map of the prominent hills of the moraine.

In 1929 and 1930, George E. Ekblaw mapped the Illinoian moraines in this area and western Illinois, but the map was not published until later (in Wanless, 1957). Ball (1937) named the Jacksonville Moraine, and Leighton and Willman (1950) named the Payson Moraine.

On the basis of the moraines and the pro-Illinoian Loveland Loess, Leighton and Willman (1950) subdivided the Illinoian into the Loveland, Payson, Jacksonville, and Buffalo Hart Substages.

Wanless (1957) studied an area largely west of the Illinois River and found gravel, sand, and silt deposits between the Jacksonville and Buffalo Hart tills and between the Mendon and Buffalo Hart tills in the area where the Buffalo Hart ice advanced beyond the area covered by the Jacksonville ice. These intervening materlais were not leached of carbonates but showed some oxidation of the iron.

Leighton and Brophy (1961) suggested that the Illinoian ice, upon reaching its outermost extent, stagnated and essentially melted in place. The Buffalo Hart Moraine is considered by them to be composed primarily of crevasse deposits and moulin kames rather than frontal end moraine and to record a brief time of renewed snowfall and glacial activity.

Willman, Glass, and Frye (1963) introduced the name Petersburg for the calcareous pro-Illinoian loess and water-laid silts that underlie Illinoian till, restricted the name Loveland Loess to the weathered loess outside the area of Illinoian glaciation where it represents all loess deposited during Illinoian time, and extended the Payson Substage downward to include the Petersburg Silt. This reduced the Illinoian to three substages-Payson, Jacksonville, and Buffalo Hart.

Recent work (Frye, Willman, and Glass, 1964) has shown that the Illinoian drift does not extend as far as the ridge originally mapped as the Payson end moralne in western Illinois. The Mendon end moralne, about 15 miles farther east, marks the limit of Illinoian drift in that area (fig. 1). The name Liman Substage was introduced to replace Payson, and the classification into Liman, Jacksonville, and Buffalo Hart Substages is used in this report.

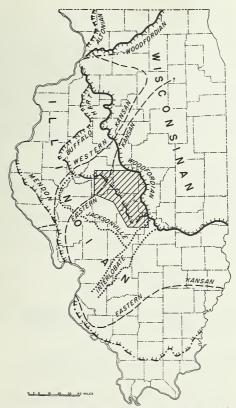


Figure 1 - Location of area studied and map showing areas of Illinois covered by the several glacial advances (after Willman, Glass, and Frye, 1963).

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Within the area actually glaciated during Illinoian time, no intra-Illinoian soil development has been identified. However, in the Great Plains region, Frye and Leonard (1954) reported four buried soils, one of which is well developed and has a moderately thick B-zone in Illinoian deposits. Frye (1962) suggested that the best developed soil may represent the interval between the Jacksonville and Buffalo Hart in Illinois. The possibility of an important break between the Jacksonville and Buffalo Hart drifts prompted this study in the type Buffalo Hart area.

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STRATIG RAPHY

Deposits of Kansan, Yarmouthian, Illinolan, Sangamonian, and Wisconsinan age are present in the area (fig. 2). The oldest Illinolan unit is the Petersburg Silt, which is primarily loessial in nature. Jacksonville and Buffalo Hart tills are present, and Mendon till may be present. The Roby Silt, between the Jacksonville and Buffalo Hart tills, includes lacustrine deposits.

The Illinoian and older deposits are covered generally by the Roxana and Peoria Loesses of Wisconsinan age. These deposits have been studied in detail (Frye and Willman, 1960; Leonard and Frye, 1960; and Frye, Glass, and Willman, 1962) and will not be considered in this report. All geologic sections referred to by name are described at the end of the report, and their locations are shown in figure 3.

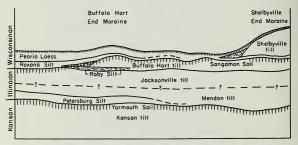


Figure 2 - Diagrammatic cross section across the study area. Length of section approximately 50 miles.

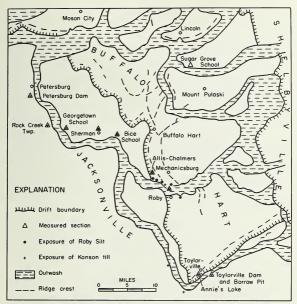


Figure 3 - Glacial map of the area, showing locations of measured geologic sections and of exposures of Roby Silt and Kansan till. (Drift boundaries after Ekblaw in Film tet al., 1959.)

Kansan Stage

Previous studies in Illinois have shown that two lobes of ice advanced into the state during Kansan time, one from the west and northwest, and the other from the east and northeast (Leverett, 1899; MacClintock, 1929, 1933; Horberg, 1956; Wanless, 1957; Flint and others, 1959; Willman, Glass, and Frye, 1963). The drift deposited by the ice from these two lobes has come to be called western Kansan and eastern Kansan, respectively. Both lobes advanced to a position near the present Illinois River Valley but are not known to have overlapped (fig. 1) and their relative ages are also unknown (Willman, Glass, and Frye, 1963). The study area is located along the western margin of the eastern lobe.

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<u>Kansan till</u>.-Till, interpreted as Kansan in age, was encountered in two separate areas, one near Taylorville and the other near Petersburg, and in two isolated outcrops (fig. 3). In most of the exposures, a weathered zone developed in the upper portion of the till is overlain by calcareous Illinoian drift (Taylorville Dam, Annie's Lake, Petersburg Dam, Rock Creek Township, and Georgetown School geologic sections). The base of the till generally is not exposed, but in a few places, it lies upon either Pennsylvanian bedrock or colluvial material overlying Pennsylvanian bedrock.

The Kansan till, where unoxidized, is usually dark gray or gray, although in the Taylorville Dam section portions of the till have a distinct pinkish cast. The oxidized till is usually yellow-brown or dark yellow-brown. The till is extremely tough and compact, and joints in the till are not uncommon.

Lacustrine silts and sands associated with the till are calcareous, carbonaceous, and in some places fossiliferous. They probably were deposited in ponds associated with or near the ice front.

With the exception of an accretion-gley exposed at the Taylorville Dam and two in-situ soil profiles exposed at the Petersburg Dam, the weathered zone in the upper portion of the till has been truncated. The weathered zone represents the first soil below the Sangamon Soil and, lacking evidence of a significant soilforming interval during Illinoian time, is the primary reason for assigning the till to the Kansan.

In several sections, the weathered zones are overlain by silts that are correlated with the widely recognized Petersburg Silt—the oldest Illinolan stratigraphic unit. Petrographic evidence, to be discussed later, also indicates a Kansan age for the till.

Yarmouthian Stage

<u>Yarmouth Soil</u>.-The soils in this area identified as Yarmouth are not as strongly developed as in some areas and, if a significant soil-forming interval is found within the Illinoian, it will be necessary to reconsider the possibility that they are intra-Illinoian soils. In the few exposures in this area, the Yarmouth Soil is not as strongly developed as the Sangamon Soil.

Two in-situ profiles that showed little evidence of truncation are exposed in excavations at the Petersburg Dam. These profiles are leached of carbonates to a depth of 5 to 6 feet. Soil structure in the B-zone is not greatly developed. Although thick, shiny clay-skins coat the fracture surfaces, the amount of clay accumulation is not great. Numerous manganese concretions are present in the lower portion of the B-zone.

The only other nontruncated profile was observed at the Taylorville Dam site, where three feet of accretion-gley overlies three feet of leached till. The accretiongley is gray with a few dark gray and yellow-brown streaks. It is very clayey and faintly laminated, and it contains a few pebbles, some sand, and locally some secondary calcium carbonate. The upper portion of the leached till contains streaks of similar gleyed clay.

Illinoian Stage

Liman Substage

Petersburg Silt.-The Petersburg Silt was defined by Willman, Glass, and Frye (1963, p. 6) to include calcareous outwash and loess deposited during the advancing stage of the Illinoian glacier. These deposits previously were considered part of the pro-Illinoian Loveland Loess and were correlated with Loveland Lovelss of the Great Plains [Leipthon and Willman, 1950; Wanless, 1957; and Leighton and Brophy, 1961]. They occur below the oldest Illinoian till in Illinois, whereas the type Loveland is considered to be largely late Illinoian in age (Frye and Leonard, 1952). As the deposits are lithologically different from the weathered Loveland Loess, the local rock-stratigraphic name, Petersburg Silt, is preferred. The Petersburg was studied in six sections near Petersburg, and silts tentatively correlated with the Petersburg were examined in three sections near Taylorville.

Near the type locality, the Petersburg Silt is variable in character. It is generally a loess-like silt that is massive, compact, and locally contains a molluscan fauna. It typically is reddish brown, but locally it is gray or tan. The loess is usually weakly calcareous, but it varies locally from calcareous to noncalcareous. In one exposure the silt is carbonaccous, dark gray, and contains wood fragments, and it appears to have accumulated in a shallow pond. The Petersburg Silt is commonly 5 to 15 feet thick, and the maximum thickness observed was 25 feet.

<u>Mendon (?) till</u>,-Mendon till may be present in the area. In several sections outside or below the Buffalo Hart drift, beds of gravel, sand, and silt or pebble pavements may represent the contact between Jacksonville and Mendon tills. Strongly contrasting clay mineral and carbonate compositions, to be discussed later, suggest the presence of two Illinoian tills older than the Buffalo Hart, but no specific change in the physical character of the till was noted.

Jacksonville Substage

<u>Jacksonville till</u>,-<u>Jacksonville till</u> (fig. 1), as used in this report, refers to the surface till south and west of the Buffalo Hart end moraine and to the till that crops out locally below Buffalo Hart till and Roby Sili in the remaining portion of the area. The unit is based on the Jacksonville end moraine (Ball, 1937; Ekblaw, in Wanless, 1957), which is located south and west of the area.

The Jacksonville ground moraine is extremely flat and exhibits very little of the sag and swell topography typical of ground moraine. Near the major drainage ways, tributaries have eroded headward and greatly dissected the drift plain.

The unoxidized Jacksonville till is gray, but in the oxidized zones it varies from dark yellow-brown or yellow-brown to grayish brown. Along joints in the gray, unoxidized portion, the till is commonly oxidized for 1 to 2 inches bordering the Joint surface, and secondary iron oxide or calcium carbonate accumulations occasionally coat these surfaces. The till is silty to sandy; thin sandy zones, lenses, and streaks are common. Locally the upper portion of the till is platy or slightly laminated.

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Pebble pavements in which the tops of the pebbles have been faceted and striated along a planned surface also are common in the facksonville till. These represent successive withdrawals and advances of the ice, the extent of which is not evident. Their common occurrence, occasionally several within a few feet, and the lack of any notable change in the character of the till above and below the pavements suggest that, in most cases, they represent minor oscillations of the ice front.

Buffalo Hart Substage

<u>Roby Silt</u>.-The Roby Silt is named herein for exposures of lacustrine silt and clay near the town of Roby, in the northwestern corner of Christian County. The type section is described in the Roby geologic section.

The unit consists of silt, clay, and sand that occur below Buffalo Hart drift and above Jacksonville drift. The deposits are not associated clearly with either the retreating Jacksonville or the advancing Buffalo Hart glaciers, but because of some exidation of the Jacksonville drift, as described later, the Roby is tentatively assigned to the Buffalo Hart Subtage. The unit is exposed in several localities in the Mechanicsburg-Roby area inside the Buffalo Hart end moraine and in two localities near Sherman outside the moraine but overlain by Buffalo Hart outwash (fig. 3). It has not been reconjized outside the area of Buffalo Hart drift.

In the type area, the Roby Silt lies upon Jacksonville till or outwash. The upper contact of the Roby Silt with the overlying Buffalo Hart drift is erosional.

A bed of sand occurs at the base of the Roby Silt in most (e.g., Roby) but not all exposures (e.g., Allis-Chalmers). The sand is thin (6-12 inches), calcareous, and strongly oxidized to yellow-brown, and it commonly contains a thin bed of brick red, sandy silt.

Above the sand is a calcareous and fossiliferous silt that is brownish gray to reddish gray-brown and has localized brown to yellow-brown streaks and specks that appear to be mostly decomposed organic matter such as stems or leaves. The silt is usually about 2 feet thick and occasionally contains thin sandy streaks. In one locality, however, it is 13 feet thick and contains 2 feet of sand, It is usually faintly laminated and locally contains secondary calcium carbonate concretions and small manganese pellets. In the Allis-Chalmers geologic section, the silt is carbonaceous and contains numerous wood fragments.

Overlying this silt is a distinctive, highly calcareous, commonly gray or brownish gray clay, $\frac{1}{2}$ to $\frac{3}{2}$ feet thick. It varies from massive clay with little or no indication of lamination to strongly laminated clay and silt with a varve-like appearance. Secondary calcium carbonate in the form of coatings and concretions is characteristic of the unit.

Another silt overlies the clay in several exposures. The silt is similar to the lower silt with the exception that no fossils were noted in it. It is 1/3 to $4\frac{1}{2}$ feet thick, but it was thicker before its upper surface was eroded severely by the Buffalo Hart glacier and its meltwaters.

The Roby Silt is lacustrine in origin, but the extent and continuity of the lake or lakes in which it was deposited are not known. In the Roby-Mechanicsburg area the silt is thought to be generally continuous, because it is present in almost all exposures of this part of the sequence and at the same general elevation.

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Exposures of the Roby Silt south and west of the Buffalo Hart end moraine were found only near Sherman, about 12 miles to the northwest. The elevation of the deposits appears slightly lower, and the deposits may not be continuous with those near Roby.

Frye (1962) suggested that the interval of time between the Jacksonville and Buffalo Hart deposits might correspond to a time of soil formation that is recognized within Illinoian deposits in the Great Plains region. An intensive search was made for a soil or evidence of a weathering break at this stratigraphic position, but other than oxidation, no weathering was noted. Jacksonville till, where lying subjacent to the Roby Silt, is usually oxidized for 2 to 3 feet, although locally oxidation may extend down from the contact for 6 to 8 feet. However, this till is sandy, and because the basal sand of the Roby Silt commonly overlies the till, some or much of this oxidation may be the result of percolating water.

Leached, clay-enriched outwash occurs locally below calcareous Roby Silt in the Bice School geologic section. The leaching and clay enrichment is interpreted to be the result of water percolating down joints in the clay zone of the Roby during Sangamonian and later time. This interpretation is based on the following observations: (1) the presence of joints in the clay; (2) the corresponding thicker development of the leached, clay-enriched zone where the clay is completely leached and its thinner development or absence where the clay is weakly to highly calcareous; (3) the local confinement of the leached, clay-enriched zone above a thin clayey sand in the outwash that appears to have been a permeability barrier to percolating water; and (4) the local, abrupt, lateral contacts between leached, clay-enriched and calcareous outwash. Similar leached, clay-enriched zones below calcareous till have been desortbed in southwestern Ohio and southeastern Indiana (Gooding, Thorp, and Gamble, 1959).

<u>Buffalo Hart till</u>.-The Buffalo Hart till is the surface drift over most of the study area and extends under Shelbyville till of Wisconsinan age to the north and east. Its southwestern boundary is marked by the outer edge of the Buffalo Hart end moraine (fig. 3).

The end moraine is well developed in some areas but poorly defined in others. It is most prominent and distinct for about 20 miles in the type area just north and south of the town of Buffalo Hart. Elsewhere, the moraine is less clear.

In sharp contrast to the generally flat Jacksonville surface, the Buffalo Hart ground moraine shows sags and swells with a relief of 5 to 15 feet. The topography generally resembles that of the Wisconsinan drift plain to the northeast more than that of the Jacksonville drift plain to the south.

In the Roby-Mechanicsburg area, where the Buffalo Hart till lies upon Roby Silt, the till is silty, contains masses of silt, and locally is more like pebbly silt than silty till. The pronounced silty character of the Buffalo Hart till is the result of erosion of the Roby Silt by the Buffalo Hart glacier. Elsewhere in the area, the till contains more sand.

The oxidized till in the Roby-Mechanicsburg area is generally grayish brown with a mottled appearance, but in places it is yellow-brown, light yellow-brown, gray, and pale olive. The unoxidized till is gray. In other areas, the unoxidized Buffalo Hart till is gray to dark gray, and the oxidized till varies from brown to yellow-brown.

In the Roby-Mechanicsburg area the till is generally 5 to 10 feet thick, except within the end moraine where it is approximately 20 feet thick. Elsewhere, the thickness of the Buffalo Hart till is not known, but it is at least 25 feet in the Sugar Grove School geologic section.

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Because of the youthful appearance of the Buffalo Hart topography, the possibility that the Buffalo Hart drift might represent an early Wisconsinan (Altonian) advance was considered. However, the presence of the Sangamon Soil on the drift was confirmed, and this eliminates the possibility of Wisconsinan age. A sample of wood from the middle clay of the Roby Silt had a radiocarbon date of >34,000 years B.P. (Isotopes, Inc., No. 1-1263).

Sangamonian Stage

<u>Sangamon Soll</u>-In this area, the Sangamon Soll is easily recognized underlying the Wisconsinan loess and developed on Buffalo Hart till, Buffalo Hart outwash, Roby Silt, Jacksonville till, Jacksonville outwash, and colluvium.

Where the Sangamon Soll has developed in a good drainage position on permeable material, it is typically reddish brown to brown. In poorly drained positions or in impermeable materials, it is commonly brownish gray to gray. Accretion-gley is present in depressions.

The soil is well developed and has strong soil structure. The B-zone is usually about 3 feet thick. The depth of leaching ranges from about 5 to 7 feet. No difference in the degree of weathering could be observed between Sangamon profiles in Jacksonville drift and those developed in Buffalo Hart drift. This suggests that the age difference between the two drifts was not sufficient to produce strongly contrasting soil profiles.

In Illinois, the Sangamon Soil has been the subject of recent investigations (Brophy, 1959; Frye, Willman, and Glass, 1960); therefore, laboratory study of the soil was not undertaken in this investigation.

PETROGRAPHY

Laboratory study was undertaken to determine the composition of the drift units, to see whether composition could be utilized to differentiate them, and to learn more about their provenances. Willman, Glass, and Frye (1963) have presented similar information for tills in Illinois on a regional basis, and comparisons are made with their data.

The grain-size distribution of the samples was determined by combined hydrometer and sieve analyses using a procedure slightly modified after that tentatively accepted by the American Society for Testing Materials (Designation: D422-547). Particles >4mm in diameter were excluded from the analyses, and the smaller fractions were classified as follows: 2-4 mm = granules; .062-2 mm = sand; .0039-.062 mm = silt; and <.0039 mm = clay.

Light minerals (sp. gr. <2.86) of the .125-.250 mm fraction (fine sand) were analyzed according to the procedure outlined by Totten (1960). After separating the heavy minerals and removing the carbonates, approximately 800 grains were mounted on a glass slide. The grains were etched in hydrofluoric acid fumes for 12 to 14 minutes and then stained in a concentrated solution of sodium cobaltinitrite. The mineral grains on the slide were counted then and identified as follows: quartzgrains that were colorless and transparent; K-feldspar-grains with a yellow coating; and Na-Ca feldspar-grains with a white, slightly lustrous coating. Minerals and rock fragments, other than above, were grouped in a fourth class called others. The clay mineral composition of the drift was determined from oriented aggregates of the <2 microns clay by X-ray diffraction procedures. The terms illite, kaolinite, and chlorite are used as generally accepted. Montmorillonite, as used here, refers to all clay materials that expand to about 17\AA after treatment with ethylene glycol. Therefore, it includes not only montmorillonite but also mixed-layered and degraded clay minerals that are expandable. Percentage values are presented for montmorillonite, illite, and combined kaolinite and chlorite.

The amount of calcite and dolomite in the minus 200-mesh grade (<74 microns) was determined by use of a Chittick gasometric apparatus following the procedures of Dreimanis (1962).

The locations of samples used in the study are shown in figure 4 and are listed in table 1. Results of the analyses are given in tables 2 and 3 and figures 5 through 12. Averages of the petrographic data are given by stratigraphic unit in table 4. The tables are at the end of the report.

Kansan Till

Willman, Glass, and Frye (1963) have shown that eastern Kansan till can be differentiated mineralogically from western Kansan till in Illinois. The most important differences are that the eastern Kansan contains little montmorillonite, while the western Kansan contains large amounts (always over 30 percent in the claymineral fraction). The eastern Kansan contains 4 times as much garnet as epidote, while the western Kansan contains 2 to 4 times as much prior tas epidote, while the western Kansan contains 2 to 4 times as much prior at set of the ported to contain more calcite than dolomite, except in the eastern margin of the western lobe, where dolomite exceeds calcite. They suggest, on the basis of carbonate and heavy mineral compositions, that the eastern Kansan was deposited by the Lake Erie lobe.

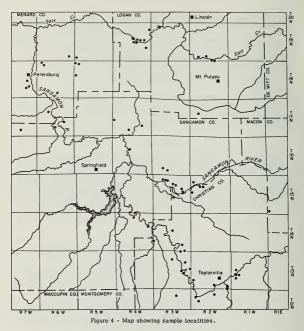
Samples of Kansan till were studied from outcrops near Taylorville and near Petersburg. Geographically both areas are in a position to be eastern Kansan. Samples from two isolated outcrops midway between the two areas are similar mineralogically to those near Taylorville and are considered with them.

Most samples from Kansan till in the Taylorville area contain about 42 percent sand and granules, 35 percent silt, and 23 percent clay, but locally the till is finer grained (fig. 5). The increase in silt and clay in the finer grained samples reflects the incorporation of local lacustrine deposits.

The unoxidized Kansan till in this area averages 65 percent illite, 32 percent kaolinite and chlorite, and 3 percent montmorillonite (fig. 6). With oxidation, the chlorite is altered, and there is a corresponding increase in expandable clay materials. The clay mineral content is typical of that derived from either the north or east in Illinois (Willman, Glass, and Frye, 1963).

The <74 microns fraction contains about 10 percent calcite and 17 percent dolomite (fig. 7). Although semi-quantitative X-ray analyses of ground whole samples indicate that eastern Kansan till contains more calcite than dolomite (Willman, Glass, and Frye, 1963), the reverse appears to be the case, with the exception of one sample, in the size fraction studied in this area. The till does contain more calcite than any of the other tills in the area. Tills with a high calcite content reflect either an eastern (Lake Erie lobe) or western source (Willman, Glass, and Frye, 1963).

This till also contains significantly more Na-Ca feldspar and less quartz than any of the other tills in the area. The average composition of the fine sand



is 69 percent quartz, 18 percent Na-Ca feldspar, 8 percent K-feldspar, and 5 percent other light minerals (fig. 8). Totten (1960) reports that Lake Erie lobe tills of Wisconsinan age in northeastern Ohio contain from 2 to 4 times as much Na-Ca feldspar as K-feldspar. On the other hand, Harrison (1959 and 1960) reports that orthoclase is the main feldspar contributor in the sand sizes of Lake Erie lobe Wisconsinan tills in central Indiana. The light mineral analyses have not provided much information about the source of the tills.

The high calcite and low montmorillonite contents indicate an eastern Lake Eric lobe source for the Kansan till of the Taylorville area. The relatively high calcite content also suggests a Kansan age for the till, because the Illinoian tills were deposited by ice entering the state from either the Lake Michigan basin or the Saginaw Bay area and they are more dolomitic (Willman, Glass, and Frye, 1963). Analyses of Kansan till from the Petersburg area shows significant and consistent textural and mineralogical differences between two groups of samples.

One group of eleven samples has textural and mineralogical properties similar to the Kansan till near Taylorville (figs. 5, 6, 7, 8). These samples, with one exception, contain relatively low amounts of montmorillonite, although slightly more than the samples from the Taylorville area. They contain essentially the same amount of calcite and dolomite as the Taylorville samples, and they also have a relatively high Na-Ca feldspar and a low quartz content. The till on the average contains 37 percent sand and granules, 35 percent silt, and 28 percent clay. Three samples from till associated with interbedded silt units are finer grained (fig. 5). This group of samples appears to be from eastern Kansan till deposited by the Lake Erie lobe.

The group of thirteen samples (figs. 5, 6, 7, 8) is from finer grained till that averages 23 percent sand and granules, 40 percent silt, and 37 percent clay. The average clay mineral composition is 35 percent montmorillonite, 44 percent tillite, and 20 percent kaolinite and chlorite. The carbonate content is low, averaging 4 percent calcite and 11 percent dolomite, and the fine sand contains approximate 98 00 percent quartz, 8 percent K-feldspar, 10 percent Na-Ca feldspar, and 3 percent other light minerals. Thus, this till differs from Lake Erie lobe Kansan till in that it is finer grained and contains more montmorillonite, less airbonate (both calcite and dolomite), less Na-Ca feldspar, and more quartz.

These differences strongly suggest that the till was derived from a source area other than that of the eastern Kansan till from the Lake Erie lobe. The two most likely alternatives are that this till is an eastward continuation of western Kansan till or that it is Kansan till of Lake Michigan lobe derivation.

The high montmorillonite content suggests a western origin, and the clay mineral composition is essentially the same as the easternmost portion of the western Kansan (Willman, Glass, and Frye, 1963). However, if there was a Lake Michigan lobe in Kansan time, it could have advanced over pro-Kansan, or older, loesses from the Ancient Missistepi Valley and incorporated the loess that has a high content of montmorillonite. Although generally eroded, pro-Kansan loess is known (Leighton and Willman, 1950; Wanless, 1957). The presence of a loess cover is suggested by the slight increase in montmorillonite in the eastern Kansan till in this area and the increase in the montmorillonite content of the eastern Kansan as it is traced to the southwest (Willman, Glass, and Frye, 1963).

Western Kansan till in northeastern Missouri contains approximately 10 percent calcite and 4 percent dolomite (Heim, 1963). As this drift is traced eatsward across western Illinois, the dolomite content increases and the calcite content decreases until, at the eastern margin of the drift, dolomite exceeds calcite. This change is across the direction of flow of the ice and results from the eastern part of the lobe passing over more dolomitic formations (Willman, Glass, and Frye, 1963). The carbonate content does not preclude the possibility that this till was derived from the Lake Michigan basin. Although the tills in this area that were derived from the Lake Michigan basin contain more than twice as much dolomite as this till (fig. 7), a relatively low dolomite content in a Lake Michigan lobe till could result from dilution of the carbonates by inclusion of large amounts of low-carbonate sufficial materials and Pennsylvanian bedrock. Interpretation of the carbonates, therefore, is not conclusive with regard to the source of the till.

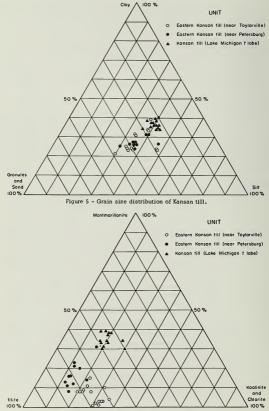


Figure 6 - Clay mineral composition of Kansan till.

Heavy minerals have proven useful in distinguishing eastern Kansan, western Kansan, and Lake Michigan lobe Illinoian tills in Illinois. Tills from a western source are characterized by considerably more epidote than garnet and relatively large amounts of kyanite, staurolite, and andalusite; tills from the Lake Michigan lobe are characterized by approximately equal amounts of garnet and epidote, although epidote is more abundant in the younger tills in the lobe; and the eastern-derived Lake Erie lobe tills are characterized by much larger amounts of garnet than epidote (Willman, Glass, and Frye, 1963).

Heavy mineral analyses were made of seven samples, three from eastern Kansan till and four from this Kansan till of questionable source (table 3).

The three samples from eastern Kansan till have heavy mineral compositions similar to eastern Kansan elsewhere in Illinois. The other four samples contain slightly more garnet than epidote. The epidote contents are greater than in eastern Kansan till but less than in western Kansan, and the garnet contents are slightly less than in eastern Kansan and slightly more than in western Kansan. The most distinguishing aspect of the heavy mineral composition of these samples is the extremely high content of opaques other than black opaques. These are primarily leucoxene and hematite (Manos, personal communication), and the significance of their great abundance is not known. The approximate equal amounts of garnet and epidote suggest that the till was deposited by an ice lobe from the Lake Michinan basin.

In two exposures in the Petersburg area, the till of questionable source occurs stratyraphically above eastern Kansan. In one exposure, it is stratyraphically below eastern Kansan, and in another, the two are interbedded. These field relations suggest deposition by two lobes that overlapped or coalesced.

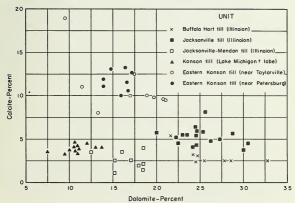


Figure 7 - Carbonate composition of Illinoian and Kansan till.

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George E. Ekblaw (personal communication) has pointed out the following field evidence, which suggests the western Kansan glacier did not cross the Illinois River Valley and thus could not have deposited the till in question. First, western Kansan till is known to thin in the direction of the valley. It is not known to crop out anywhere along the valley bluff but is always found several miles back of the bluff. Secondly, if the glacier did cross the valley and possibly coalesce with the eastern Kansan glacier, the drainage system of the Ancient Mississippi Valley would have been blocked and widespread lacustrine materials should have been deposited. Such deposits have not been found.

Consequently, field relations and the limited heavy mineral data favor deposition of the till by an ice lobe moving from the Lake Michigan basin during Kansan time. The fine texture, high montmorillonite content and low carbonate content of the till can result from incorporation of either loess or weathered material over which the ice advanced.

Illinoian Drift

The Mendon (Payson), Jacksonville, and Buffalo Hart tills contain distinct clay mineral assemblages (Willman, Glass, and Frye, 1963). Mendon till usually contains the most montmorillonite (17-64 percent), Jacksonville till intermediate amounts (11-28 percent), and Buffalo Hart till extremely low amounts. Also the Jacksonville contains less kaolinite and chlorite than the other units.

Petersburg Silt

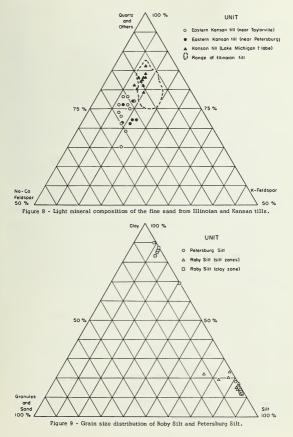
In the type area, loess from the Petersburg Silt contains little or no sand, about 86 percent silt, and about 13 percent clay (fig. 9), and the median diameter is about 21 microns. The mean size is essentially the same as for samples of Peoria Loess in this area (Smith, 1942), which suggests that the source area for the loess of the Petersburg, like the Peoria Loess, was at or near the position of the present Illinois Valley.

The Petersburg Silt has a high content of montmorillonite, which reflects the contribution of clays from a western source to the sediments in the Ancient Mississippi Valley (Willman, Glass, and Frye, 1963). The average clay mineral composition in the type area is 49 percent montmorillonite, 36 percent lilite, and 15 percent kaolinite and chlorite (ifg. 10), but the considerable variation in clay mineral composition suggests that the unit may have a more complex source and depositional history than can now be deciphered.

The calcite content of the Petersburg Silt is usually low (average 2 percent), and the dolomite content varies from 0 to 24 percent. These variations also suggest a complex history for the unit and probably reflect both variations in the rate of deposition and possibly time breaks during deposition of the loess.

Jacksonville till

Samples of Jacksonville till fall into two distinct petrographic groups that are referred to as Jacksonville and Jacksonville-Mendon in this report. The two groups are distinguished by their clay mineral compositions but also differ in texture and carbonate composition. Samples with more than 20 percent montmorillonite in the clay mineral fraction are termed Jacksonville-Mendon, those with less than 20 percent are Jacksonville.



In geologic sections containing both Jacksonville and Jacksonville-Mendon samples, Jacksonville samples always occur above Jacksonville-Mendon, but a definite stratigraphic break has not been recognized between them.

The Jacksonville is the coarsest till in the area and on the average contains 44 percent sand and granules, 32 percent silt, and 24 percent clay. Although there is some overlap (fig. 11), samples of Jacksonville-Mendon are finer grained and average 26 percent sand and granules, 46 percent silt, and 28 percent clay.

The Jacksonville-Mendon till contains significantly more montmorillonite than the Jacksonville (fig. 12). Montmorillonite ranges from 21 to 51 percent in the Jacksonville-Mendon till, but only to 16 percent in the Jacksonville. Both contain approximately the same amount of kaolinite and chlorite and do not have the low kaolinite and chlorite content as noted for Jacksonville till by Willman, Glass, and Frye (1963).

Jacksonville till samples contain approximately 10 percent more dolomite and 2 percent more calcite in the <74 micron fraction than the Jacksonville-Mendon (fig. 7). Both contain considerably more dolomite than calcite. With the exception of the Lake Michigan (?) lobe kansan till, the Jacksonville-Mendon till contains less carbonate than the other tills.

The lightweight minerals of the fine sand from these two groups are similar, averaging 78 percent quartz, 9 percent K-feldspar, 10 percent Na-Ca feldspar, and 3 percent others.

The higher montmorillonite content, lower carbonate content, and finer texture of the Jacksonville-Mendon reflects incorporation of the widespread, thick Petersburg Silt. Mendon till in western Illinois contains large amounts of montmorillonite that is also attributed to the incorporation of Petersburg Silt (Willman, Glass, and Frye, 1963).

The lack of a definite stratigraphic break between the Jacksonville-Mendon and Jacksonville tills suggests that the Mendon glacier did not retreat this far to the northeast following its advance, and thus Mendon till would not be separated from Jacksonville till by a physical break. The gradual decrease upward in the montmorillonite content of till overlying Petersburg Silt in several geologic sections (e.g., samples 446-453) is attributed to a lessening influence of the Petersburg Silt as the unit was progressively buried by till deposition. However, there is an abrupt change in the montmorillonite content in other exposures, and in these areas, the lower part of the unit probably is equivalent to Mendon till in western Illinois.

Roby Silt

Sixteen samples of Roby Silt were analyzed-5 from the lower silt, 8 from the middle clay, and 3 from the upper silt.

The middle clay is extremely fine grained and contains little or no sand, about 12 percent silt, and 88 percent clay (fig. 9). The median diameter usually is less than 1 micron. The silt zones, although variable, usually contain about 2-3 percent sand, 80 percent silt, and 17 percent clay, and the median diameter ranges from 6 to 18 microns.

The clay mineral composition of the middle clay is uniform, but the silts have a variable composition. Illite is the predominant clay mineral in the middle clay, but small amounts of chlorite, montmorillonite, and kaolinite are present (fig. 10). In the silts, the montmorillonite ranges from 0 to 38 percent, illite from 40 to 80 percent, and kaolinite and chlorite from 9 to 32 percent (fig. 10). These variations occur in both silt zones and probably result from uneven mixtures of wind blown material with outwash.

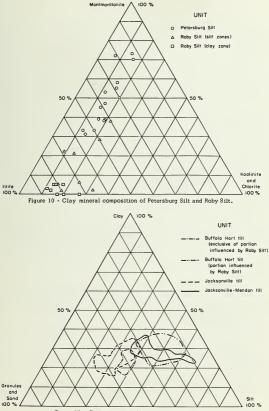


Figure 11 - Grain size distribution of Illinoian tills.

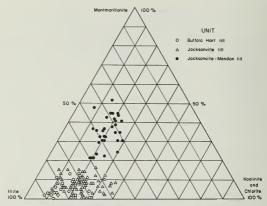


Figure 12 - Clay mineral composition of Illinoian tills.

The middle clay zone and the silt zones contain approximately the same amount of calcite, but the silt zones contain about three times more dolomite. The clay zone averages 6 percent calcite and 8 percent dolomite, while the silt zones average 5 percent calcite and 24 percent dolomite. The difference reflects variations in the amount of dolomite in different size fractions.

Buffalo Hart till

The Buffalo Hart till generally has about the same grain-size distribution as Jacksonville till. It contains approximately 38 percent sand and granules, 37 percent silt, and 25 percent clay. In areas where the Buffalo Hart glacier overrode Roby Silt, the sand content decreases about 20 percent and the silt and clay contents increase correspondingly (fig. 11).

The only notable mineralogical variation in the Buffalo Hart till is a slight increase in the montmorillonite content in areas where Roby Silt has been overridden.

The average clay mineral composition of the till is 4 percent montmorillonite, 78 percent illite, and 18 percent kaolinite and chlorite (fig. 12). Dolomite is the predominate carbonate (fig. 7), and the lightweight mineral composition is the same as other Illinoian tills in the area (fig. 8).

Mineralogical differences between Buffalo Hart till and Jacksonville till are slight, but Buffalo Hart till generally contains slightly less montmorillonite, more illite, and less calcite than Jacksonville till (figs. 7 and 12, table 4).

CONCLUSIONS

The principal interpretations and conclusions are as follows:

- Petrographic differences between two groups of samples suggest that the Kansan till was deposited by two different lobes of ice. One lobe entered the area from a generally easterly direction, probably from the Lake Erie basin. The direction of advance of the other lobe is less certain, but the evidence favors its origin in the Lake Michigan basin.
- 2. The Yarmouth Soil in this area is not as strongly developed as the Sangamon Soil.
- The Petersburg Silt is primarily loessial in origin but in part is water laid. The loess is fossiliferous, montmorillonitic, and usually dolomitic. It was probably blown off the valley trains of the first Illinolan glacier.
- The Jacksonville and Buffalo Hart tills definitely are present, and the Mendon till probably is present.
- The glacial morphology of the Buffalo Hart drift plain is considerably fresher than that of the Jacksonville drift plain. It may indicate a significant age difference, but the differences can be related to the character of the original depositional surface.
- 6. The Roby Silt is composed of lacustrine silts and clays and lies stratigraphically between Jacksonville and Buffalo Hart tills. It is fossiliferous, calcareous, illitic, chloritic, and locally carbonaceous. The Roby Silt indicates a significant retreat of the ice front before the Buffalo Hart advance.
- The Sangamon Soil appears to be as well developed on Buffalo Hart till as on Jacksonville till. This suggests that the time between the deposition of the two tills is short when compared to that represented by the Sangamonian Stage.
- 8. Physical and compositional differences between the tills reflect variations in the mineral composition and physical character of the bedrock in the source areas and along the route of ice advance. The most important differences are:
 - (a) Eastern Kansan till from the Lake Erie lobe differs from the Kansan till that probably was deposited by the Lake Michigan lobe. It contains more carbonates, especially calcite, more illite, less montmorillonite, more Na-Ca feldspar, and less quartz, and it is coarser textured.
 - (b) Eastern Kansan till from the Lake Erle lobe differs from Illinoian tills of the Lake Michigan lobe in that it contains more calcite, less dolomite, more Na-Ca feldspar, and less quartz.
 - (c) The undifferentiated Jacksonville-Mendon till is finer textured and contains more montmortllonite and less carbonate (both calcite and dolomite) than the definitely Jacksonville till.
 - (d) The Jacksonville-Mendon till and the Kansan till of the Lake Michigan lobe are similar petrographically. Both apparently incorporated much loess and, as a result, are finer grained and contain more montmorillonite and less carbonate than the other tills in the area.
 - (e) The Jacksonville and Buffalo Hart tills are similar, but Jacksonville till contains slightly more montmorillonite and calcite than Buffalo Hart. Texturally the two tills are similar, except the Buffalo Hart till is much finer grained where the glacier overrode Roby Silt.

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	Location	Stratigraphic		Location	Stratigraphic
Sample	<pre>½ Sec., T-R, County,</pre>	Unit	Sample	<pre>½ Sec., T-R, County,</pre>	Unit
No.	(geol. sec. no.)	(feet below top)	No.	(geol. sec. no.)	(feet below top)
1	NW NW SW 5, 18N-5W,	JacksMendon	178	do.	Jacksonville (13)
	Menard		181	NW NE SW 10, 15N-3W,	Kans, (East)
7	SE SW NE 36, 13N-2W,	Jacksonville (.7)		Sangamon	
	Christian		183	NE SW NW 17, 17N-2W,	Buffalo Hart (6)
8	do.	Jacksonville (4.3)		Logan	
9	do.	Jacksonville (5.5)	193	NW SW SE 5, 15N-3W,	Jacksonville (5)
10	do. do.	Jacksonville (9)	195	Sangamon do.	Jacksonville (11.5)
14	do. do.	Kans. (East) (4) Kans. (East) (8.5)	201	GO. SW NW SE 5, 15N-3W,	Jacksonville (11.5) Jacksonville (6)
19	do.	Kans. (East) (0.5) Kans. (East) (14)	201	Sangamon	Jacksonville (6)
46	SE SW NE 36, 13N-2W,	Jacksonville (2.3)	211	SE NW NW 10, 15N-3W,	Jacksonville (4)
-10	Christian (11)	000000000000000000000000000000000000000		Sangamon	000000000000000000000000000000000000000
48	do.	Jacksonville (6.5)	212	do.	Buffalo Hart (.5)
50	do.	JacksMendon (9.5)	213	do.	Buffalo Hart (2.5)
51	do.	Petersburg (?) (0.5)	214	do.	Buffalo Hart
58	do.	Kans. (East) (2)	220	SW SE SW 12, 14N-4W.	Jacksonville (4)
60	do	Kans. (East) (8)		Sangamon	
61	do.	Kans. (East) (10.5)	221	do.	Kans. (East) (4)
63 64	do. do.	Kans. (silt) (12.5) Kans. (silt) (13.5)	222	SW SW NW 30, 14N-3W, Christian	Jacksonville (3)
65	do.	Kans. (East) (13.5) Kans. (East) (14.5)	226	do.	Jacksonville (8)
66	do.	Kans. (East) (16)	226	NE NE NW 2, 13N-3W,	Jacksonville (8)
76	NW SE SE 11, 15N-3W,			Christian	
	Sangamon		230	NW SW SW 18, 15N-2W,	Buffalo Hart (.5)
86	NW SE SE 11, 15N-3W,	Jacksonville (5)		Christian	
	Sangamon		231	do.	Buffalo Hart (2)
106	NE SW SE 6, 15N-2W,	Jacksonville (3.5)	232	do.	Roby (.7)
	Sangamon		233	do.	Roby (3)
112	NW NE SE 8, 15N-2W,	Jacksonville (.5)	234	do.	Roby (5.5)
	Christian		235	do.	Roby (7)
113	do. NE SW SE 4, 15N-3W,	Jacksonville (3.5) Buffalo Hart (2.5)	241	SW SW NE 34, 13N-2W, Christian (2)	Jacksonville (1)
11/	Sangamon	Builaio Hart (2.5)	243	do,	Jacksonville (4)
120	do.	Jacksonville (1)	245	do.	Kans. (East) (1)
124	NW SE SE 11, 13N-3W,	JacksMendon	247	do.	Kans. (East) (4)
	Christian		2.68	SW SE NW 22, 15N-3W,	Jacksonville (2)
125	NW NE SE 11, 13N-3W,	Jacksonville (6)		Christian	
	Christian		256	NW NE SE 8, 15N, 2W,	Jacksonville (14)
127	do.	Jacksonville (11)		Christian	
128	do.	Jacksonville (15)	259	SW NW SW 3, 12N-2W,	JacksMendon (4.5)
129		Jacksonville (4)		Christian	
136	Christian		260	NE NE SE 7, 12N-2W,	JacksMendon
130	NE NW SW 23, 13N-3W, Christian	Jacksonville (1)	263	Christian do.	Kans, (East) (5)
138	do.	Jacksonville (2.5)	264	do.	Kans. (East) (5)
141	SE NW NW 23, 13N-3W.	Jacksonville (3)	266	SE SE SE 3, 12N-3W,	Jacksonville (3)
	Christian	Successfulling (b)	200	Christian	Successive (5)
144	do.	Jacksonville (11.5)	267	do.	Jacksonville (7)
147	SW NE SE 23, 13N-3W,	Kans. (East)	272	NE NE SW 30, 13N-2W,	JacksMendon (4)
	Christian			Christian	
148	do.	Kans. (silt)	273	SW NE SE 23, 13N-3W,	Jacksonville (2)
150	SE NW SE 13, 13N-3W,	Jacksonville (5)		Christian	
	Christian		274	do.	Jacksonville (8.5)
151 156	do.	Jacksonville (8)	283		Buffalo Hart (1)
130	NE SW NW 10, 13N-3W, Christian	Jacksonville	284	Sangamon do.	P. (C. 1. 11
162	NE SW SE 26, 14N-3W,	Jacksonville (2.5)	284	do.	Buffalo Hart (6) Buffalo Hart (11)
	Christian		285	do. (7)	Roby (.5)
165	do.	Jacksonville (8)	287	do.	Roby (2)
167	SW SE NW 15, 14N-3W,	Jacksonville (4)	288	do.	Jacksonville (1)
	Christian		291	do.	Jacksonville (12)
168	do.	Jacksonville (7.5)	292	do.	Jacksonville (4.5)
171		Jacksonville	293	do.	Jacksonville (8.5)
	Christian		297	SE SW SE 9, 19N-4W,	JacksMendon (7)
173	NE NE SW 10, 15N-3W,	Roby (1)		Logan	
173 176	NE NE SW 10, 15N-3W, Sangamon NE NE SW 10, 15N-3W,		298	Logan SE SE SW 14, 19N-4W, Logan	JacksMendon (5)

TABLE 1 - LOCATION AND DESCRIPTION OF SAMPLES

	Location	Stratigraphic		Location	Stratigraphic
Sample	<pre>k Sec., T=R, County,</pre>	Unit (feet below top)	Sample No.	E Sec., T-R, County, (geol. sec. no.)	Unit (feet below top)
No.	(geol. sec. no.)	(feet below top)		(geo1. sec. no.)	(leet below top)
299	NW SE NE 12, 19N-4W, Logan	Buffalo Hart (3)	389 397	do. SE SE NW 32, 19N-2W,	Buffalo Hart (18) Buffalo Hart (3)
302	SE SW SE 16, 19N-4W, Logan	JacksMendon	398	Logan do. (10)	Buffalo Hart (8)
307	SW NE SE 29, 16N-3W,	Buffalo Hart (0.5)	399	do.	Buffalo Hart (13)
	Sangamon (1)		400	do.	Buffalo Hart (18)
308	do.	Buffalo Hart (2)	402	SW NW NE 33, 19N-2W,	Buffalo Hart (1)
309	do.	Buffalo Hart (4)		Logan	
310	do.	Buffalo Hart (6)	410	SE NW SW 5, 18N-6W,	Jacksonville (5.5)
311	do .	Buffalo Hart (8)	413	Menard SE NW NW 16, 17N-6W,	Toole North (10)
312 314	do. do.	Buffalo Hart (8) Buffalo Hart (11)	415	Menard	Jacksmendon (12)
315	do.	Buffalo Hart (11.5)	414	do.	JacksMendon (22)
320	do.	Roby (1)	415	do.	JacksMendon (25.5
324	do.	Roby (3)	418	do. (8)	Petersburg (4)
325	NE SW SE 29, 16N-3W,		421	do.	KansL.Mich? (4)
	Sangamon		422	do.	KansL.Mich? (11)
326	do.	Buffalo Hart (2.5)	423	SW SW NE 34, 18N-6W,	JacksMendon (10)
327	do.	Roby (.2)	424	Menard	
329	do.	Roby (2)	424	do. do.	JacksMendon (16) JacksMendon (28)
333	NW SW SW 5, 19N-3W, Logan	Buffalo Hart (4)	420	SE SE SW 23, 17N-6W,	Jacksonville (4.5)
337		Buffalo Hart (4)	467	Sangamon	Jacksonville (4.5)
337	Logan	builato nare (4)	434	do.	Kans. (East) (2)
338	do.	Jacks,-Mendon (8)	435	do.	KansL.Mich? (8)
340	NE NE NW 26, 19N-4W,	JacksMendon (8)	436	do. (4)	KansL.Mich? (14)
	Logan		437	do.	KansL.Mich? (20)
341	NW NW NE 16, 16N-4W,	Jacksonville (5)	438	do.	Kans. (East) (26)
	Sangamon		439	do.	KansL.Mich? (32)
343	NE SE NW 19, 17N-4W,	Roby (2)	440 441	do.	KansL.Mich? (38)
344	Sangamon (9)		441	SW SE SE 23, 18N-7W, Sangamon	Kans. (Last)
344	do. SW NW NW 15, 19N-4W, Logan	Roby (3.5) JacksMendon (3)	446	SE NE SE 28, 17N-6W, Sangamon	Jacksonville (0.5)
346	NW SW SE 15, 19N-4W,	Jacks Mendon (4)	447	do.	Jacksonville (2.5)
540	Logan	Guenor-Included (4)	448	do.	Jacksonville (3.5)
347	do.	JacksMendon (14)	449	do.	Jacksonville (10.5)
348	SE SE NW 15, 19N-4W,	Buffalo Hart (3)	450	do.	Jacksonville (16.5)
	Logan		451 452	do. do.	Jacksonville (22.5)
349	do.	Buffalo Hart (8)	452	do.	JacksMendon (28.5) JacksMendon (34)
350	SE SE NW 15, 19N-4W, Logan	Buffalo Hart	473	NE NE NW 8, 18N-6W,	Petersburg (1)
355	NE SE NW 19, 17N-4W,	Taaksamii 11a (2)	47.5	Menard	receisourg (1)
555	Sangamon	Jacksonville (2)	474	do.	Petersburg (3)
356	SE SE NW 26, 17N-4W,	Jacksonville	475	do.	Petersburg (6)
	Sangamon		476	do.	Petersburg (11)
357	NE SW SW 28, 16N-3W,	Jacksonville (2)	480	do.	Petersburg (16)
	Sangamon		498		Kans. (East) (3)
362	do.	Buffalo Hart (5.5)	499	Menard do.	No
369	NW NE SW 31, 19N-5W, Menard	Jacksonville (5.5)	499	do. do.	KansL.Mich? (7) KansL.Mich? (15)
370	SW SE NE 36, 19N-6W, Menard	JacksMendon (4.5)	505		Jacksonville (2.8)
374	NW SE NE 36, 19N-6W,	Jacks Mendon (8)	507	do.	Petersburg (8)
	Menard	outer menden (c)	512	do.	Kans. (East) (8)
376	do.	Jacks Mendon (19)	514	do.	Kans. (East) (10.5)
378	SW SW NE 6, 18N-5W,	JacksMendon (12)	515	do.	KansL.Mich? (15.5)
380	Logan SE NW SW 18, 18N-5W,	Jacksonville (4)	516	SE NW NE 23, 18N-7W, Menard	
	Logan		517 518	do.	Jacksonville (9)
384	SE SW SE 34, 19N-2W,	Buffalo Hart (5)	518 519	do.	JacksMendon (17)
385	Logan do.	Buffels Hank (0)	319	SE NW NE 23, 18N-7W, Menard	Jacksnendon (10)
385	do.	Buffalo Hart (9) Buffalo Hart (15)	520	do.	Petersburg (1)
387	do.	Buffalo Hart (15) Buffalo Hart (16)	522	do.	Petersburg (6.5)
388	SE SE SE 34, 19N-2W,		523	NE NW SW 31, 15N-1E,	Buffalo Hart (10)
	Logan		520	Christian	
			528	SW SE NW 20, 13N-1W,	Jacksonville (2)

TABLE 1 - Continued

	Location	Stratigraphic		Location	Stratigraphic
Sample	Ł Sec., T-R, County,	Unit	Sample	E Sec., T-R, County	Unit
No.	(geol. sec. no.)	(feet below top)	No.	(geol. sec. no.)	(feet below top)
529	SW SE NW 20, 13N-1W, Christian	Jacksonville (8)	581	NW NE NE 26, 18N-6W, Menard (6)	Kans. (East) (1)
532	NE SE NW 20, 13N-1W,	Jacksonville (1)	582	do.	Kans. (East) (2.5)
	Christian		586	do.	Kans, (East) (9)
533	do.	Jacksonville (6)	591	NW NE NE 26, 18N-6W,	
535	SW SE SE 6, 12N-2W,	Jacks Mendon (4)		Menard	
	Christian		592	do.	JacksMendon (5.5)
538	do.	Petersburg (4)	593	do.	Petersburg (2)
546	NW NW SE 36, 13N-2W,	Jacksonville (12)	594	do.	Petersburg (5)
	Christian		595	do.	Petersburg (8)
548	do.	Petersburg (?) (1)	596	do.	Petersburg (11)
549	do. (12)	Petersburg (?) (2)	597	do. (5)	Petersburg (13)
550	do.	Petersburg (?) (2.3)	598	do.	Yarmouth (.5)
551	do.	Petersburg (?) (3.3)	599	do.	Yarmouth (1.8)
554	SE NW NW 17, 13N-1W,	Jacksonville (3)	600	do.	Yarmouth (3.2)
	Christian		601	do.	KansL.Mich? (1.7)
556	SE NE NE 30, 13N-1W,	Jacksonville (2.5)	602	do.	KansL.Mich? (3.3)
	Christian		603	do.	KansL.Mich? (7.2)
565	NW NW SW 3, 14N-4W, Sangamon	Jacksonville (11)	610	NW NE NE 26, 18N-6W, Menard	Yarmouth (0.3)
566	NE NW NW 3, 14N-4W,	Jacksonville (10)	611	do.	Yarmouth (1)
	Sangamon		612	do.	Yarmouth (1.7)
567	S1 SW SE 28, 17N-4W,	Roby (0.5)	613	do.	Yarmouth (2.8)
	Sangamon (3)		614	do. (6)	Yarmouth (4)
569	do.	Roby (1.8)	615	do.	Yarmouth (4.8)
570	do.	Roby (3.0)	616	do.	Kans. (silt) (0.4)
572	NW NE NE 26, 18N-7W,	KansL.Mich?	617	do.	Kans. (East) (0.7)
	Menard		618	do.	Kans. (East) (2)
574	SE NE NE 34, 17N-7W, Sangamon	Jacksonville	619	do.	Kans. (East) (4)

TABLE 1 - Continued

											-	
		Size cent)		1 (y Miner percent	t)		ght Minerals carbonate-f:	ree basis		(per	onates (cent)
	Granules			Montmor-		Kaolinite			Na-Ca			Dolo-
No.	and sand	Silt	Clay	illonite	Illite	+ chlorite	Quartz	K-feldspar	feldspar	Others	cite	mite
17	25	49	26	30 15	49 77	21 8						
8	51	30	19	15	68	17	77	12	9	2		
9	43	35	22	14	67	19	80	10	8	2		
10	-10			9	66	25						
14	42	34	24	8	72	20						
15	29	48	23	3	63	34	70	8	20	2	9.5	21.1
19	28	39	33	4	60	36	71	9	16	4		12.5
46	47	29	24	9	73	18	79	9	11	1	8.0	25.6
48				8	69	23						28.8
50	40	33	27	21	60	19	77	10	10	3		
51				36	44	20						
58				8	72	20						
60	40	37	23	3	65	32	73	6	18	3	10.0	
61	29	48	23	2	67	31	69	7	19	5	9.7	19.7
63				15	60	35						
64				2	66	32						
65	24	37	39		69	31	61	12	23	4	19.0	
66	28	39	33	1	66	33	64	11	20	5	11.1	11.5
76 86	50 47	28 29	22 24	6	70 78	18 16	77	9	13	1		
106	47	30	24	4	78	24	78	12	9	1		
112	47	31	20	11	72	17	78 78	13 10	10	2		
112	49	32	19	*1	12	17	/0	10	10	2	6.0	25.3
117	16	47	37								0.0	23.3
120	38	30	32	4	80	16	78	11	10	1	1. 2	24.6
124	32	42	26	30	50	20	70		10	-	4.5	24.0
125	49	30	21	15	73	12	78	10	11	1		
127	47	32	21	8	68	24	10	10		-		
128	39	35	26	3	59	38	78	10	9	3	5.5	23.1
129	54	28	18	5	74	21	78	11	10	ĩ	5.5	23.1
136	50	29	21	8	73	19	77	11	10	2		
138	32	34	34	2	63	35	76	11	10	3	6.0	24.9
141	52	22	26	11	78	11	78	10	10	2		
144	42	32	26	6	64	30	75	11	11	3	5.6	24.6
147	40	37	23	3	64	33	73	8	16	3		
148				8	58	34						
150	50	30	20									
151	48	29	23	7	79	14						
156	53	29	18	11	70	19						
162	52	29	19	10	71	19	80	11	8	1		
165	34	39	27	2	75	25	78	11	8	3		
167												
168 171	11	26	-	6	62	32						
171 173	41	36	23	6 2	71 85	23 13						
173	47	30	23	4	00	13						
178	47	23	23	6	62	32	80	9	9	2		
181	35	35	30	8	66	28	66	10	22	2	13.0	12 6
183	45	37	18		00	20	00	10	22	2	10.0	42.0
193	46	31	23									
195	46	30	24		72	28	81	9	9	1		
201	45	31	24		74	26	79	10	é	2	3.8	30.0
211	48	30	22	4	64	32	80	10	10	9	5.0	20.0
212	30	49	21	10	73	17	78	11	9	2		
213	22	49	29			_,				-		
214	24	49	27	3	72	25	79	9	10	2	2.5	28.4
220	46	35	19	16	60	24	75	11	11	3		
221	35	34	31	9	69	22					12.6	17.4
222	55	26	19									
224	40	35	25	3	63	34					6.2	24.5

TABLE 2 - GRAIN SIZE AND MINERAL ANALYSES

	Grain				ay Miner			ght Mineral				onates
	(perc	cent)			percent	t)	(carbonate-fi	ree basis)	(per	cent)
	Granules			Montmor-		Kaolinite			Na-Ca			Dolo-
No.	and sand	Silt	Clay	illonite	Illite	+ chlorite	Quartz	K-feldspar	feldspar	Others	cite	mite
226	39	37	24	3	68	29						
230	23	46	31	9	74	17	76	14	9	1		
231		-10		13	73	14			,	-		
232	0	87	13	21	65	14					3.0	26.5
233	3	77	20	5	80	15					5.5	23.0
234	2	13	85		83	17						
235	7	77	16	38	40	22					6.5	16.0
241	49	28	23	13	77	10						
243	50	32	18	6	70	24	81	8	10	1		
246	42	33	25	8	70	22	73	9	14	4		
247 248	49 43	30 34	21 23	11	66 69	23 24	71	6	16	7	9.5	21.1
248	43	34	17	/	69	24						26.2
259	25	50	25	44	39	17	77	11	11	1	4.0	20.2
260	27	45	28	34	40	26	78	11	7	4		
263	42	35	23	20	63	17	72	7	15	6		
264	39	37	24	10	59	31	67	ŝ	19	6	10.0	19.0
266	45	31	24									
267	28	31	41	10	67	23						
272	34	36	30	23	56	21	78	9	10	3		
273				5	72	23						
274	42	35	23	8	57	35					4.3	24.6
283	20	57	23									
284	24	47	29	11	80	9	75	14	9	2		
285 286	24	44 13	32 83	12 3	69 86	19	77	11	11	1		21.6
280	4	83	15	28	50	11 22						8.5 23.0
287	45	32	23	4	85	11	78	12	8	2	9.5	23.0
291	41	35	24	4	65		70	12	0	2		
292	45	34	21	5	85	10	80	12	7	1		
293	42	34	24	9	70	21	80	10	7	3		
297	35	37	28	32	50	18	81	9	8	2		
298	26	39	35	42	42	16	82	9	8	1		
299	45	35	20	8	79	13	76	11	11	3		
302	31	39	30	37	42	21	80	8	10	2		
307	22	42	36									
308	23	45	32									
309 310	25 21	45 49	30 30	4	78	18	79	10	9	2		
311	17	51	32	5	81	10	/9	10	9	2		
312	17	52	31	2	01	14						
314	17	55	28								24	24.6
315	14	58	28		77	23	78	10	11	1		24.7
320	ò	31	69		73	27		10		-	5.5	9.0
324	2	84	14	2	66	32						30.0
325	12	60	28									
326	14	57	29		70	30	81	9	9	1	2.6	25.4
327				2	80	18						26.0
329					80	20					5.5	10.5
333	36	37	27	2	84	16	80	9	9	2		
337 338	36 31	38 36	26 33	25	87	13						
	27			35	50	15	81	8	9	1		
340 341	44	43 28	30 28	44 6	36 74	20 18	82	8	8	2	3.6	16.1
341	44	13	28 87	5	81	18						
345	13	65	22	6	70	24					5.5	25.0
345	36	36	28	47	45	8					5.5	25.0
346	27	41	32	36	46	18						
347	28	41	31	44	38	18	80	7	11	1		
348	37	37	26	5	76	19						
349	37	40	23	2	76	22						
350	25	57	18	4	69	27	75	13	10	2	2.4	27.7

TABLE 2 - Continued

							-					
	Grain	Size cent)			y Mines			ght Mineral: carbonate-f:				cent)
Sample	Granules	cont)	-	Montmor-	Percell	Kaolinite		car 001/8 L0-1	Na-Ca	í		Dolo-
No.	and sand	Silt	Clay		Illite	+ chlorite	Quartz	K-feldspar	feldspar	Others	cite	
	48	28	24	8	81	11						
355 356	48 48	28	24	8	73	24					5.0	27.3
357	46	29	25	3	87	10	76	14	8	2	5.0	27.5
362	12	54	34	6	83	11	74	14	11	1		
369	45	32	23	8	74	18	80	9	9	2	4.4	22.7
370	19 20	53	28	35	47	18	78	10	11	1		
374 376	20	54 55	26 26	32 38	48 40	20	82 80	8	7 10	3 1		
378	24	49	27	29	52	19	75	16	10	2	25	15.1
380	42	32	26	3	77	20	76	11	10	3	2.5	1.7.2
384	38	34	28	3	84	13	74	15	9	2		
385	37	35	28									
386	39	35	26		73	27	79	9	11	1		
387	40	37	25								3.4	24.0
388 389	37 44	35 39	28 17		70	30	78	12	8	2	2 5	32.7
397	40	34	26		70	30	70	12	•	2	2.5	32.1
398	39	35	26	3	88	9	77	11	11	1		
399	40	34	26									
400	39	28	23	9	79	12	77	10	12	1		
402	37 42	35 32	28	3	85	12	79	10	10	1		
410 413	42	32 51	26 24	8 21	69 59	23 20	74	12	11	3	4.6	30.5
414	10	67	23	40	36	20	75	14	8	3	1 5	18.5
415	10	07		41	39	20	/5	14	0	2		18.2
418				39	45	16					0.5	3.5
421	26	36	38	31	49	20						10.6
422	24	42	34	34	39	27	79	8	12	1	3.8	11.0
423	29	39	32	40	39	21	82	7	9	2		
424 426	26 25	39 39	35 36	32 30	48 41	26 29	80	7		2		
429	31	38	31	14	66	29	78	11	11	2	3.5	12.4
434	38	36	26	14	65	21	71	7	17	5	13.0	14.7
435	22	39	39	34	46	20	79	8	9	4		10.0
436	23	42	35	33	49	18	80	8	10	2	3.2	9.4
437	20	40	40	33	47	21	79	9	10	2	3.4	7.5
438 439	37 22	36 40	27 38	6 37	76 44	16 19	70	11	18	1	11.0	
439	30	36	38 34	37	44 43	19	80	8		,		10.4
441	40	34	26	20	65	15	80	8	8	4	4.0	10.8
446	40	34	26	7	79	14	81	7	7	5	10.5	10.9
447	44	32	24	10	71	19	79	10	10	1		
448	44	32	24	11	68	21	83	8	8	1		
449	39	35	26	12	62	26	74	11	12	3		20.0
450	37	35	28	8	55	37	79	10	8	3		22.2
451 452	33 26	37 46	30 28	12 24	58 46	30 30	78	8	12			23.5
452	26	46 68	28	24 43	46 36	30 21	/8	8	12	2		17.3 15.2
473	·			70	14	16					0.0	0.0
474				50	30	20					0.0	0.0
475	1	84	15	72	21	7					0.4	2.0
476	1	87	12	73	13	14						19.0
480 498	1	87	12	45	41	14						16.5
498	38 26	35 39	27 35	16 34	71 49	13	79	-	10		11.9	
500	26	39	35	34	49	17 20	79	7	12 10	2 3		13.0
505	32	39	29	11	70	19	77	12	9	2	4.6	12.0
507				58	27	15			-	-	0.5	6.5
512	27	47	26	12	72	16	70	11	17	2		
514	37	37	26	11	75	14	69	9	20	2	13.2	
515	24	39	37	30	50	20	82	8	9	1	3.9	11.2

TABLE 2 - Continued

	Grain (per				y Mine			ght Mineral carbonate-f				nates cent)
Sample	Granules	cenc)	_	Montmor-	percen	Kaolinite		lai bonace-1	Na-Ca			Dolo-
No.	and sand	Silt	Clay		Illite	+ chlorite	Quartz	K-feldspar		Others		
516	44	30	26	10	78	12	73	9	14	4		
517				11	66	23						
518	14	60	26	34	42	24	76	10	12	2	2.0	18.0
519	27	46	27	22	53	25	75	9	10	6		
520				38	43	19						19.0
522				33	55	12					7.0	26.0
523	31	37	32		79	21	80	10	8	2		
528				6	75	19						
529	49	31	20	2	77	21	78	9	11	1		
532				8	71	21			13	5		
533	39	36	25 26	7	75 51	18 13	71	11	13	2		
535	26	48 84			38	13	77	11	12	2	1.0	13.0
538	1		15 26	47			80	10	8	2	1.0	13.0
546 548	40 14	34 36	26	10	59 15	31 16	80	10	0	2		
548	14	20	30	33	38	29						
550				58	21	29						
551				52	28	20						
554	46	30	24	9	62	29						
556	40	50	24	6	74	20						
565	35	40	25	7	58	35	80	12	6	2		
566	29	48	23	8	65	27	77	11	8	4		
567	1	76	23	22	69				-			
569	ī	12	87		78	17					5.5	4.0
570	ī	- 9	90	5	82	13					3.5	7.0
572	23	41	36	30	39	30	80	7	10	3	4.0	13.8
574				10	70	20						
581				21	68	11						
582	30	39	31	13	75	12						
586	39	37	24	16	72	12					12.6	17.0
591				21	59	20						
592				51	35	14						
593	1	86	13	33	50	17						2.5
594	1	87	12	38	43	19						24.5
595	1	86	13	53	35	12						13.5
596	1	85	14	31	35	14					0.5	7.5
597	1	81	18	55	29	16					0.5	4.0
598	18	47	35	48	34	18	70	11	18	1		
599	29	36	35	42	43	15	72	12	15	1		
600	23	41	36	49	35	16	73	11	14	2		
601	20	41	39	49	31	20	82	7	8	3		
602	23	41	36	38	42	20	76	10	11	3	1.0	11.0
603	23	43	34	39	44	17	83	7	7	3	4.0	11.2
610	18	45	37	42	31	27	75 72	11	12	4		
611 612	19 17	38 46	43 37	46	37 38	1/	72	11	13	2		
612	17	46 47	37	44 47	38	18	73	12	13	4		
613	29	4/ 36	38	4/	38	15	68	10	15	6		
614	29	36 50	33	33	55 40	13	72	10	13	5		
616	16	50 61	23	4/	40 51	13	70	10	15	3		
617	44	30	36	21	68	14	68	11	18	3		
618	44	30	25	21	66	11	69	12	16	3	11.5	16.4
619	29	40	31	31	52	17	74	9	14	3		16.0
		-			52							
-			_									

TABLE 2 - Continued

	Opa	que				Т	rans	pare	nt H	eavy	Min	eral	s (%)			
Sample No.	Black	Others	Tourmaline	Zircon	Garnet	Epidote	Rutile	Titanite	Kyanite	Staurolite	Andalusite	Actinolite	Hornblende	Enstatite	Hypersthene	DiopAug.	Others
60	11	13		1	14	5		1	1		1	1	64	1	6	5	
247	8	20	1	1	19	3		1	1	1			61	3	3	2	4
422	11	56	4		14	10		2			1		55		8	3	1
499	8	51	1		13	12		1		3		2	60	1	1	2	4
500	5	51	4	1	18	8			1	1	1		62		2		2
514	9	8	2	1	25	6					1	1	61			2	1
515	6	51	1	1	20	12			1				57	1	1	2	4

TABLE 3 - HEAVY MINERAL ANALYSES* (carbonate-free basis)

* Analyses by Constantine Manos.

		Grain	Grain Size		0	Clay Minerals	liners	als		Ligh	t Min	Light Minerals		Car	Carbonates	es
Stratigraphic Unit	safqms2 io radmuW	bnaS bna selunard	aile	Clay	saiqmaS io radmuN	Montmortlionite	TIIIce	Kaolinite + chlorite	ssigmeS io rsdmuN	Suartz	Tsqsb191-X	Na-Ca feldspar	Огретя	Number of Samples	Galcite	Dolomite
Illinoian																
Buffalo Hart	38	28	44	28	24	4	78	18	19	77	11	10	73	80	e	26
Roby Silt																
silt zones	7	4	78	18	80	16	64	20	1	ı	1	1	I	7	ŝ	24
clay zone	9	٦	15	84	80	2	82	16	1	1	1	1	1	S	9	8
Jacksonville	66	44	32	24	68	80	70	22	64	78	6	10	e	17	ŝ	25
Jacksonville-Mendon	28	26	46	28	31	34	46	20	21	79	6	10	2	6	e	17
Petersburg Silt (loess only)	6	-	85	14	15	67	36	15	1	1	1	I	ı	14	2	11
Kansan																
Eastern (Taylorville area)	14	36	37	27	15	7	99	27	11	69	8	18	ŝ	10	11	16
Eastern (Petersburg area)	11	37	35	28	12	17	69	14	7	70	10	17	e	80	12	16
Lake Michigan? lobe	14	23	40	37	14	35	44	21	12	80	8	10	e	12	4	11

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GEOLOGIC SECTIONS

The locations of the geologic sections are shown on figure 3. Color notations and color names refer to those of the Munsell color charts and are for moist samples. Soil zone nomenclature follows that of Frye, Willman, and Glass (1960). Numbers in parentheses at the end of descriptions are sample numbers. The sections are arranged alphabetically by name.

> Thickness (feet)

NO. 1 - ALLIS-CHALMERS SECTION

Measured in two lifts of a road cut on the Allis-Chalmers Testing Ground, in SW4 NE4 SE4 sec. 29, T. 16 N., R. 3 W.; Sangamon County, Illinois (1961).

PLEISTOCENE SERIES

WISCONSINAN STAGE WOODFORDIAN AND ALTONIAN SUBSTAGES Peoria Loess and Roxana Silt undifferentiated

Sangamon Soil on Buffalo Hart drift 7. Soil, A-zone, yellow-brown (10 YR 4/4) to

 Soll, B-zone, yellow-brown (10 YR 5/4) with dark brown (10 YR 4/3) clay-skins; clay loam texture; medium subangular blocky structure; Fe and Mn pellets in lower 2 feet.... 3.3 ILLINGIAN STAGE

BUFFALO HART SUBSTAGE

- 5. Till, siliy upper 3 feet leached; upper portion gray-broom (10 YR 5/2) with yellowbrown (10 YR 5/4) and gray (10 YR 5/1) motcles; grades to a light yellow-brow (2.5 YR 6/4) or pale olive (3 YR 6/2; locally contains streaks of silt, fine sand, or clay in the lower portion; uppermost 9 feet exposed in the lower lift (307 to 315)..... 14.3 Roby Silt
- Silt, yellow-brown (10 YR 5/4) to gray-brown (10 YR 5/2), calcareous; few bands of brownish yellow (10 YR 6/6-8) iron stain... 0.5

- Outwash, sand and gravelly sand, calcareous; upper 2 feet yellow-brown (10 YR 5/4) grading

Total 36.0

NO 2 - ANNIE'S LAKE SECTION

Measured 20 yards SE of Annie's Lake and near the spillway in SW% SW% NE% sec. 34, T. 13 N., R. 2 W.; Christian County, Illinois (1961).

PLEISTOCENE SERIES

WISCONSINAN AND SANGAMONIAN STAGES

Yarmouth Soil on Kansan till

- Till, leached, oxidized, yellow-brown; contains secondary CaCO₃, gray pebbly clay, and gleyed inclusion; upper contact erosional..... 2.0 KANSAN STAGE
- Till, calcareous, oxidized, grayish brown; secondary CaCO₃ down joints; lower portion only slightly oxidized (246 and 247)... 5.0

Total 18.0

NO. 3 - BICE SCHOOL SECTION

Measured along the north bank of an abandoned gravel pit in S¹/₂ SW¹/₂ SE¹/₂ sec. 28, T. 17 N., R. 4 W.; Sangamon County, Illinois (1962).

PLEISTOCENE SERIES WISCONSINAN STAGE WOODFORDIAN SUBSTAGE Peoria Loess

Thickness

(feet)

ILLINOIAN STAGE

BUFFALO HART SUBSTAGE

- Silt, thin, discontinuous, noncalcareous; variable shades of gray and brown; pockets of gray-brown clay; upper contact erosional (567)......0-1.0

Total 27.5

NO. 4 - GEORGETOWN SCHOOL SECTION

In SEX SEX SWX sec. 23, T. 17 N., R. 6 W.; Sangamon County, Illinois (1961).

- PLEISTOCENE SERIES WISCONSINAN STAGE WOODFORDIAN SUBSTAGE
 - Peoria Loess
- - YARMOUTHIAN STAGE Yarmouth accretion-gley
- 3. Soil, G-zone, gray and dark brownish gray

Thickness (feet)

with local yellow-brown mottling, blocky, clayey; more brownish toward base.... 2.0

- Till, calcareous, oxidized, yellow-brown (10 YR 5/4) near top grading to gray-brown (10 YR 5/2) toward base; several silty streaks and pinkish zones near top (434-440)...... 38.0

Total 51.0

NO. 5 - PETERSBURG DAM SECTION NO. 1

Composite section measured on the nose of first bluff west of the dam and on the south side of the lake in SWŁ NWŁ NEŁ NEŁ sec. 26, T. 18 N., R. 6 W.; Menard County, Illinois (1962).

PLEISTOCENE SERIES

- WISCONSINAN STAGE WOODFORDIAN AND ALTONIAN SUBSTAGES Peoria Loess and Roxana Silt undifferentiated
- Loess, upper 10 feet in slope and vegetated; lower 4 feet brown to yellow-brown, colluvial 14.0
 - SANGAMONIAN STAGE Sangamon Soil on colluvium and Jacksonville
- Soil, B-zone, dark reddish brown to dark brown in the lower 12 inches, clayey, quite pebly and sandy; Mn concentration; clay skins down fractures; structure weak.. 3.0 ILLINGIAN STACE

JACKSONVILLE SUBSTAGE

- Till, calcareous, upper portion gray-brown with Mn and Fe staining down joints; grades into gray till with oxidation confined to area along joints and to thin sandy streaks; lower 2 feet locally very silty and brown; fossils in lower portion (591 and 592)

LIMAN SUBSTAGE

Petersburg Silt

- Silts, loessial, variable in color and carbonate content, locally fossiliferous, pink, buff, and shades of gray and brown, compact to friable, oxidized down joints with Fe and some Mn stain; locally contains small Mn pellets; thickens to vest (593-597).. 2-24.0 VARMUTHLAN STACE
- Yarmouth Soil on Kansan till 4. Soil, A-zone, dark yellow-brown with local
- yellow-brown mottles, platy (598).... 1.0
 3. Soil, B-zone, compact, yellow-brown; dark Mn staining and thin clay-skins down fractures; no typical soil structure (599-600)... 3.0 KNNSAN STAGE

Total 44.5

NO. 6 - PETERSBURG DAM SECTION NO. 2

Measured 40 yards east of section no. 1 and on the west side of the south dam abutment in SE& NW% NE% NE% sec. 26, T. 18 N., R. 6 W.; Menard County, Illinois (1962).

The upper part of the section consists of approximately 18 feet of Peoria Locess and Roxana Silt, and the Sangamon Soil developed into either colluvium or till with approximately 6 to 7 feet of leaching. All units thin to the north where the Yarmouth Soil is best exposed.

PLEISTOCENE SERIES

ILLINOIAN STAGE

JACKSONVILLE SUBSTAGE

Petersburg Silt

 Silt, calcareous, gray to buff, discontinuous; irregular thickness; contains small Mn pellets; probably not in place............0-3.0 YARMOUTHIAN STAGE

Yarmouth Soil on Kansan till

 Soil, B-zone, locally cut out; where most complete the upper 2 feet is clayey, tough, Thickness (feet)

yellow-brown with local dark Hm staining; contains thick (up to 1/16 inch), shiny, continous clay-skins which thin toward base as Mm concentration increases; no apparent structure; lower 3 feet not as tough, light yellow-brown with may dark Hm pellets (1/8 inch diameter) and Hm stains on fracture surfaces; slight gray mottling; and thin, discontinuous clay-skins (610 to 615).... 5.0 KNNAN STACE

- Till, leached, brown to yellow brown; locally lower 12 to 18 inches is calcareous; Mn concentration down joints and locally Mn pellets in the upper portion (581 and 582)..... 2.0
- 2. Silt, calcarcous except in the leached zone of the Yamouth Solly Here leached is light yellow-brown and contains Mn pellets; where calcarcous has a zone of secondary CaCo₂ at top, the upper 6 inches is brownish gray with bright yellow-brown iron streaks, then 6 inches is uniform gray silt, and 24 inches the transform state of the second state of the second

Total 17.0

NO. 7 - ROBY SECTION

Measured along the south bank of an abandoned meander channel in NW½ SE½ NE½ sec. 14, T. 15 N., R. 3 W.; Sangamon County, Illinois (1961).

PLEISTOCENE SERIES

- WISCONSINAN STAGE WOODFORDIAN AND ALTONIAN SUBSTAGES Peoria Loess and Roxana Silt undifferentiated

Sangamon Soil on Buffalo Hart till 10. Soil, A-zone, leached, pebbly silt, yellow-

- light mottles, clay loam; moderately strong subangular blocky structure; Fe and Mn concentrations and pellets in the lower portion 3.0

ILLINOIAN STAGE BUFFALO HART SUBSTAGE

 Till, leached, gray and yellow-brown, mottled, silty; considerable Mn concentration.. 2.5

35

 Till, calcareous, oxidized, grayish brown, locally gray or yellow-brown, silty silt inclusions common in lower portion (283-285).
 Roby Silt (upper contact distorted and unit locally

repeated)

- Clay, calcareous, weakly laminated; upper 3 inches brownish gray, remainder gray; contains much secondary CaCO₃ (286).......5-1.0
 Silt, calcareous, brownish gray to reddish
- Silt, calcareous, brownish gray to reddish gray-brown, fossiliferous; contains small Mm pellets and locally sand inclusions (287)... 1-2.0

- Till, slightly oxidized, brownish gray; upper surface is faceted, striated pebble pavement; local, thin, calcareous sand zones; lower 8 feet examined by augering (293).....
- Silt, calcareous, upper 2 feet gray-brown, lower 2 feet deep brown; base not reached... 4.0

Total 51.0

NO. 8 - ROCK CREEK TOWNSHIP SECTION

Measured along the south bank of a tributary to the Sangamon River in center W½ NW½ sec. 16, T. 17 N., R. 6 W.; Menard County, Illinois (1961).

PLEISTOCENE SERIES ILLINOIAN STAGE JACKSONVILLE SUBSTAGE

- Till, calcareous, oxidized except for the lower 6 feet, grades from grayish yellowbrown to gray; more oxidized down joints; lower 4 feet slightly oxidized, silty, and contains organic fragments (413-415).. 24.0 LIMAN SUBSTAGE Petersburg Silt

Kansan till

 Till; upper 12 inches leached; oxidized yellow-brown in the upper 6 feet, lower portion gray with oxidation down joints; not well exposed; base covered (421 and 422)... 10.0

Total 46.0

Thickness (feet)

NO. 9 - SHERMAN SECTION

Measured in road cuts along Interstate 55 in NE& SE& NM& sec. 19, T. 17 N., R. 4 W.; Sangamon County, Illinois (1962).

PLEISTOCENE SERIES

WISCONSINAN STAGE WOODFORDIAN AND ALTONIAN SUBSTAGES Peoria Loess and Roxana Silt undifferentiared

- Loess; upper portion yellow-brown to dark yellow-brown grading to dark reddish brown in the lower portion; locally sandy; lower 5 feet is yellow-brown sand...... 16.0 ILLINGIN STAGE
 - BUFFALO HART SUBSTAGE

- Clay and thinly laminated silt and clay, noncalcareous except locally in the lower portion, gray to brownish gray, seconday CaCO3 concretions near base (343)...... 1.5-2.5

- Colluvium, sandy, calcareous, not well exposed; rests on Pennsylvanian sandstone..... 5.0

Total 33.5

NO. 10 - SUGAR GROVE SCHOOL SECTION

Measured in stream cuts along Salt Creek in SEX SEX NMX sec. 32, T. 19 N., R. 2 W.; Logan County, Illinois (1961). Upper part contains Peoria Loess which is slumped and poorly exposed.

PLEISTOCENE SERIES

SANGAMONIAN STAGE

Sangamon Soil on Buffalo Hart till 3. Soil, A-zone, dark gray-brown (10 YR 4/2 to

3/3), strong, medium, subangular blocky

- structure, grades to B-zone..... 1.0 2. Soil, B-zone; upper 12 inches is dark graybrown (10 YR 4/2) to dark yellow-brown (10

Total 25.0

NO. 11 - TAYLORVILLE DAM SECTION

Measured at the west end of the center line of the dam in SE₂ SM² NE¹₂ sec. 36, T. 13 N., R. 2 W. Christian County, Illinois (1961). This section is no longer exposed.

- PLEISTOCENE SERIES
 - WISCONSINAN STAGE WOODFORDIAN SUBSTAGE Peoria Loess
- Loess, yellow-brown (10 YR 5/4), noncalcareous; contains surface soil profile... 3.3

- Soil, CL-zone, clay, pebbly and locally sandy gray (5 Y 6/2), with lower portions mottled gray, yellow-brown, and dark brown; locally contains Mn and Fe concentrations.... 0.8

 Soil, CL-zone, sandy silt, massive, noncalcareous, slight concentrations of Fe and Mn, lower 6 inches is sand...... 2.0

ILLINOIAN STAGE

JACKSONVILLE SUBSTAGE

- Till, calcareous except for localized leaching in the upper 2 inches, oxidized, yellowbrown (10 YR 5/4); lower 2 feet is only slightly oxidized and pale brown (10 YR 5/3) to grav-brown (10 YR 5/2) (46 and 47). 6.0
- Till, calcareous, oxidized, dark brown (10 YR 4/3); upper surface is faceted, striated pebble pavement; thin, sandy, and silty zones in the lower 2 feet (48 and 50)..... 4.0 LHAN SUBSTAGE Petersburg (7) Silt
- Silt, calcareous, dark yellow-brown (10 YR 4/4) with gray streaks; contains a few pebbles (51)..... 1.0

YARMOUTHIAN STAGE Yarmouth accretion-gley

- KANSAN STAGE
- Till, calcareous, oxidized; upper part is dark yellow-brown (10 KR 4/4); at 5 feet grades down to unoxidized, very dark gray (10 FR 3/1) with a few oxidized and and gravel pockets and 2-4 inches oxidized along joints; lower 2 feet carbonaceous with abundant wood fragments (58, 60, and 61)... 12.0

Total 45.1

NO. 12 - TAYLORVILLE DAM-BORROW PIT SECTION

Borrow pit in NW1 NW1 SE1 sec. 36, T. 13 N., R. 2 W., Christian County, Illinois (1961).

	ISTOCENE SERIES
I	LLINOIAN STAGE
	JACKSONVILLE SUBSTAGE
6.	Till, calcareous, gray; oxidized down joints
	and along sandy seams; top covered (546)
5.	Till, silty, calcareous, carbonaceous, dark
	grayish brown; wood fragments common 2-5.0
	TTUNK OUTCOM OF
	LIMAN SUBSTAGE
	Petersburg (?) Silt
4.	Clay, dark gray to black, leached, hackly;
	slickensided fracture (548) 1-1.5
3.	Silt, calcareous, dark grayish brown, car-
	bonaceous; contains many wood fragments (549)

2. Silt, leached, dark blackish gray, carbon-

- aceous; laminated to fissle; contains wood fragments (550).....0.3
- Silt, calcareous, peaty, and carbonaceous; wood abundant; locally noncalcareous; dark grayish brown in the upper portion grading into dark bluish gray silt which is fossiliferous; base not exposed (55)..... 1-1.5

Total 10.0

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