

Trip of 1907

doc. 0093

New Jersey, Maryland, Virginia  
and Tennessee.

Jan 4. 450.

13-90  
+ 15

- Massachusetts 11, 12
- North Mountain 14
- Henry Rem 15
- Sanford 16
- Great Cacapon 17
- D. B. B 20
- Nipper 24

Note book of  
 Charles Schuchert  
 Yale University

New Haven, Connecticut.

1907.

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~~2932~~

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If found by any one please return to above  
 address.

Trip in Mesoria of Atlantic border  
 and Potomac - Lurray section, and  
 Tenn. with Reids.



Phil Schuchert 1202 Linden Ave. Canton Ohio.  
Miss S. Hutig 1055 Academy Ave.  
A. E. Schuchert. 119 E. Leventh. Cin. O.  
T. A. Armstrong.  
R. S. Luce.  
M. Corcoran, 59 Hall. New Haven, Conn.  
E. A. Callahan 472 Orange. " "  
T. E. Savage, Urbana, Ill.  
R. J. Basler. C. O. Mich.  
R. H. Gordon. Frank Hartley.  
L. F. Hille 192 Seneca St. Auburn N. Y.

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Mail to

July 1-2 Vincennes N. Y.  
Up to July 8 Harper Ferry, W. Va.  
Up to July 15 Martinsburg " "  
Up to July 22 Care of R. H. Gordon.  
Up to July 27 Nashville Tenn.  
Up to Aug 3 Camden Tenn.  
Up to Aug 7 Perryville, Tenn.  
After this care C. H. Gould, Norman Ill.

# Stellus Correlation Table. Am. Rep. P. G., 1904.

Remarks.	Cort 1868.		Class 1892-1894	Keller 1905
	Upper shale	and mud green mud		
Green sand. Fossils.	25'	yellow sand	Mones- green	D Monoguan Long Beach
Blauwite with sand	45'	Yellow sand	Green	C Vincentown Lewell
Layery green sand. Shale layer 7' thick. Argonza beds above at Vincentown. Lower beds have bones	40'	Yellow li-s li sand <i>Shale layer heavy mud clay at mud</i>	Red Cross	B Red Sand Sand Mansfield Mansfield Not found Vincentown
Red due to oxidized glauconite. Some fossils but few.	100'	Indurated green sand <i>Red sand.</i> Sand, micaceous clay	Mon- mouth	B Red Sand Sand Mansfield Mansfield Not found Vincentown
new sand; and sand. In one fossil or so. 3 or 4 species	30-50'	Thin Blue shale with sand mud	Mon- mouth	B Red Sand Sand Mansfield Mansfield Not found Vincentown
Fossils rather rare. Not yet worked out.	275 to 400	Laminated sand	Mon- mouth	A Vincentown Columbus Vincentown
Partian or Platyceras. Sand Upper Chetaceans. 34' - 400'		Grey sand	Mon- mouth	A Vincentown Columbus Vincentown

July first. Barrell, Reeds and Noble at Philadelphia. First train on Long Beach road to Vincentown. Stay there 3 days. Have the following papers.

- "Upper Cret. Form. of N. J., Delaware, & Ind" by Clark.
- "A Preliminary Rep. of the Cret. & Tert. Form. of N. J." by Clark.

The fossiliferous "Lower Chalk" may I think be seen about Milton and to East of Mount Holly. Ask about this at Vincentown. If so take this drive from Vincentown.

See my Cecil Court Map  
for roads and buildings.

[Look for unconformity. Said to occur  
between these formations]

Photomac



Perryville, Md.

walk along P. R. R. track from Perryville to Charleston, then across country to B. & O. R. R. and north to Perryville. Distance 14 miles.

See paper "The Stratigraphy of the Potomac Group in Md." Clark - Gibbins.

Section 1/8 miles E. of Principio Creek (P. R. R.).

Recent. Loam with gravel 5'

Paritan. Fine white sand. 11'

Loam } Brown loamy sand, bearing  
gravel and coarse towards base. 12'

Potomac

Patapsco. White clay, somewhat im-  
stained and variegated; at times  
graduating over into micaceous  
sands; changing to sand, coarse,  
conglomerate toward base — 10-20'  
Dense variegated clay. 10.

Jurassic Patuxent under ground.

Foggs Hill on B. & O. R. R.

Patapsco. Buff sands, water raising at base 25'

dense variegated clay 10-30.

In places with considerable pebbles.

The upper zone of a. - with pebbles.

See Spencer The Geol. of Massachusetts  
Ont.

Baltimore Md.

Art's School House, Hill. See 1.482.

Federal Hill. See 1.484.

Sharps Ferry. One day to see the immediate region. Then to New Market across Massanutten Mt to Luray. Enquire here about driving across to Seven Fountains, Fork Cross Roads to Woodstock. Then to

Martinsburg.

North Mountain. Stop at Martinsburg.

Great Cacapon. Hotel here.

A long section from the Medina to Hamilton.

Enquire about cuttings between here and Cumberland.

See Clark's Physical Features of Md.  
and  
Schuchert's Lower Devonian and Ontonagon  
formations of Maryland.

July 7 note under that we were  
2 to 74 days before we were

Hancock. Also see the creek on the river  
part from the other side.

Cumberland. Look after Hartleys new  
localities. Then one day

Cash Valley across to Devils Back Bone.

Pinto section. One day

Twenty-first Bridge, Kessler quarries and  
Judge Alkire. One day.

Then to Nashville, Tenn.



Manchester Tenn.

Reservoir Hill and hill just N. W.

Here Orthorhynchula linneyi and large Lepiditris.

A little to the N. and probably less than 20' lower are other quarries with an abundance of Tetradium and Chenuraria.

Naegle's Hill for an extensive section in Orthorhynchula and Tetradium horizon.

The section is as follows:

Upper Bryozoa beds or Rafinesquina zone.

Then 20' lower

Orthorhynchula horizon.

Then 12' lower

Tetradium zone.

Then 15' lower

Soft shaly limestone.

Then 8' to road at South end.

These are the Capitol Hill limestone beds of Safford, the transition beds between the Trenton and the Cincinnati beds.

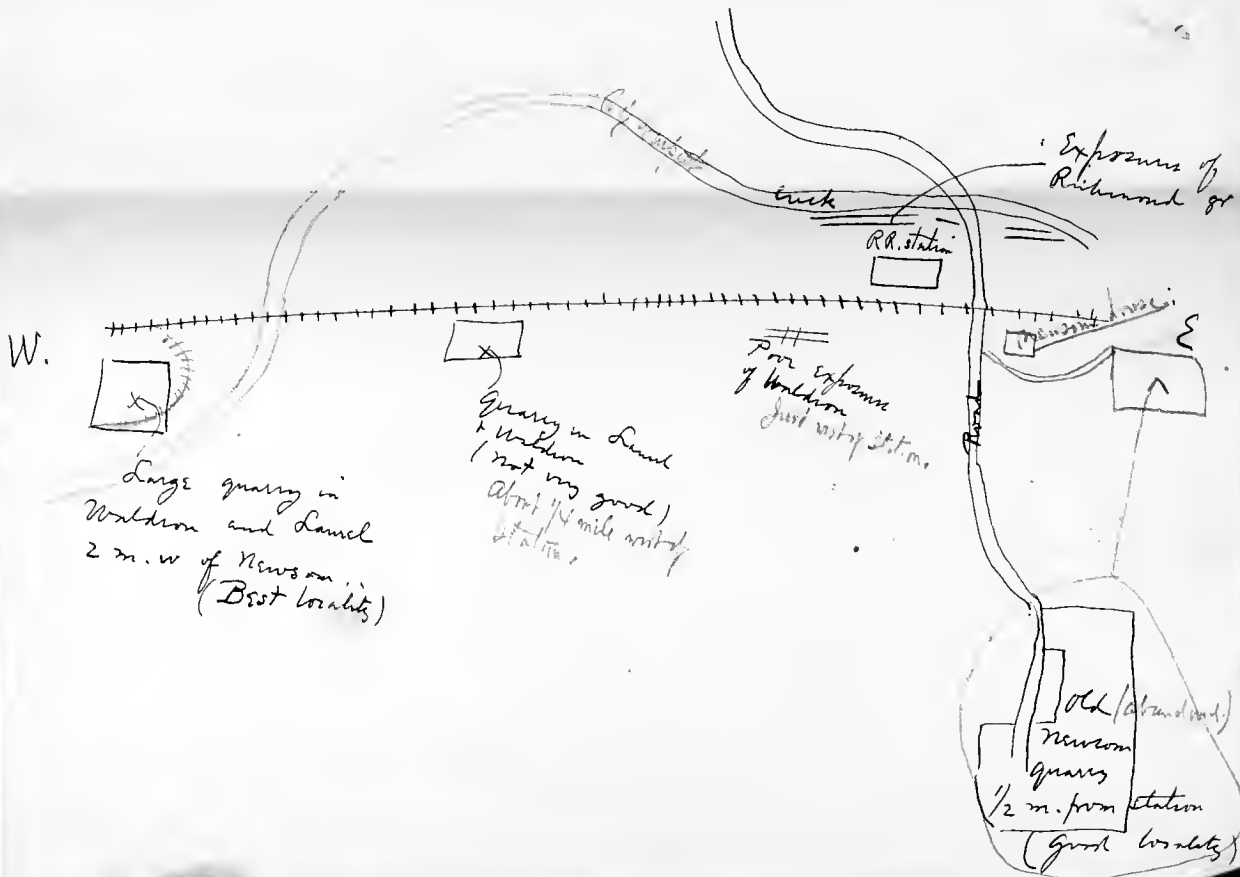
At Newson may also be seen the  
Onondaga above the Silurian or Lejo zone  
= to Louisville. Look here for Nucleoceras  
vermeili, Strophodontia demissa, D. septana  
Rhipidomella fenelope, Nucleospira concinna.

See Foerste "Silurian and Devonian  
limestones of Western Tennessee."



## Newsom localities.

There is but one house at Newsom and I was never able to get permission to stay over night there. You may have the same experience so I would advise staying at Nashville and going back and forth to Newsom by railroad. The fare is small and the train is arranged conveniently for putting in long days collecting. The sketch below gives the best localities.



## Camden, Tenn.

Best Oriskany locality just on the outskirts of the village. To the west of the village I secured nothing, or little.

The chert is from 12-15 feet thick, overlain by shale 6-7 feet thick. In this is Black Shale. At the top is a 2 inch ironstone band. Above this loose sand = Hardin? Oriskany acc. to Foerste. He gives the Camden here as 60'.

## Helderberg near Camden. South.

One locality is near Chaseville on Wolf Creek up stream from the wagon road going from Chaseville to Holladay's. About six exposures in a quarter mile from the road. Thickness about 20-25' overlain by Camden chert, not exposed only saw broken chert.

Between 4-5 miles W. - E of Holladay's toward Camden along side the road are two fine localities of Helderbergian, 35' shown. Lowest beds have Michelinia and Favosites. Bygonia higher up. Phacops from upper portion.

Foerste gives this locality as 5 miles south of Big Sandy Station on the Lower Camden road.

## Big Sandy. North of Camden.

By wagon from Camden to Big Sandy and then 6 miles south to Tom Barnes is a fine ~~fine~~ <sup>about 10' thick,</sup> ~~Stel~~ <sup>Stel</sup> ~~depression~~ <sup>exposure</sup> with Camden above. Above a recent blue mud deposit. <sup>See if one can drive direct to Tom Barnes and then to Big Sandy.</sup>

The Camden is well exposed in a small quarry to the west of Big Sandy <sup>Station.</sup> About 2-3' thick. *Spirifer* and *Eatonia* common. It is often seen along the road the most western road for six miles south of Big Sandy. Forster says 5 miles S. of Big Sandy Station on the New Camden road may be seen contact between the ~~depression~~ <sup>depression</sup> and Camden.

## Pace, Henry County, Tenn.

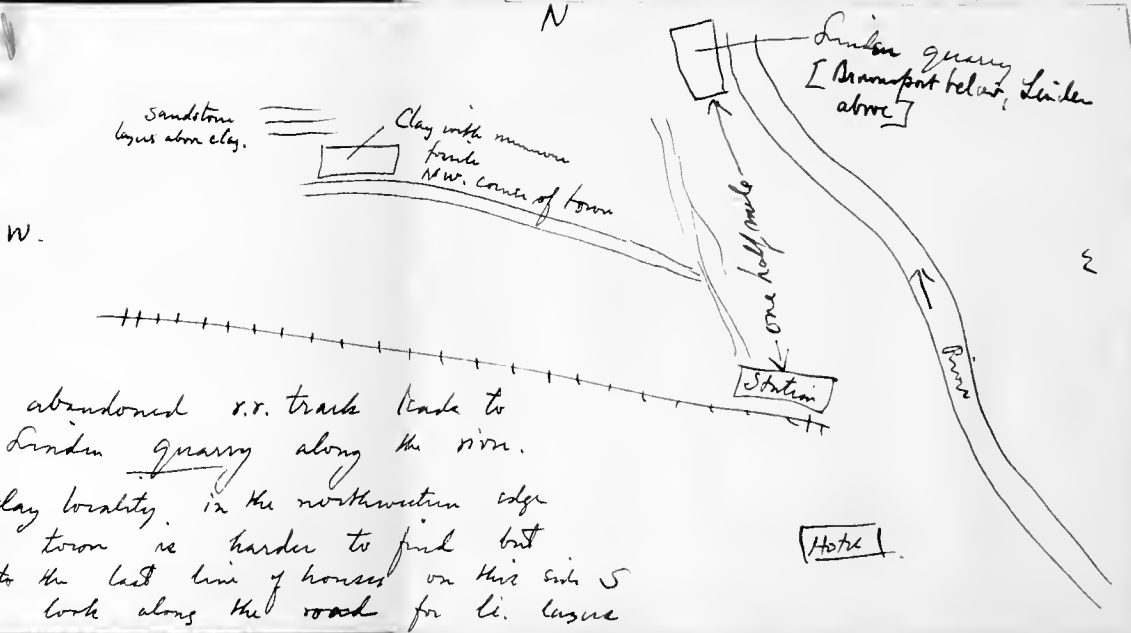
Fourteen miles north of Big Sandy Station (5 miles up from mouth of Big Sandy River) to Jeffords Williams Mill. The exposure is on Benton County side, about 30' thick. Lowest layers *Atrypa reticularis*, higher *Anastrophia* followed by *Orthis* beds, with *Orthostrophia* (these are in the center of the *Orthis* beds). *Phacops* occur from and above the *Anastrophia* beds. Also *Jophrentis*. Pace called this locality Swagnes Old Mill.

One locality is under bridge at public road near Lashleys store. See for the boy Art Lashley, Pace, who sold me fossils. at 50¢ per box cigar.

Another locality at old mill site about one mile above (? south).

Pate said get team at Big Sandy Station. (At the Williams site Forster says Camden maybe seen 50' thick and more.

Perryville. Helderbergian (Linden).  
On railroad Camden - Lexington - Perryville. [See next page?]



An abandoned r.r. track leads to the Linden quarry along the river.

The clay locality in the northwestern edge of the town is harder to find but go to the last line of houses on this side S and look along the road for li. lenses

Also see railroad cut between Perryville and Parsons for Helderbergian. It is 3 miles east of Parsons or  $3\frac{1}{2}$  miles west of Perryville, Tenn.

Fine Camden chert fossils at around Decaturville which is southwest of Perryville.

### Decaturville. or rather Dixon Spring

Poems collection was made in a radius of 5 miles around Dixon Spring, located on the old Colonel Wallis Dixon farm, 3 miles south of Perryville. This is the "glade" hills <sup>or "mounds"</sup> country. The horizon is the Brownport (above the Walden) = to the Louisville. Brownport Landing is about 5 miles S-E. of Dixon Spring, and 3 miles N. of Vice Landing.

The Brownport consists of whitest clays and soft limestones usually found in the open "mound glades". Can easily be seen in the topography. Dixon beds below are red, ~~and the~~ <sup>the</sup> ~~spanges.~~ <sup>spanges.</sup> One of these mounds a  $\frac{1}{4}$  mile N. of home of Wood Butler, on road from Vice to Perryville  $2\frac{1}{2}$  miles N. of Vice. Another  $\frac{1}{2}$  mile farther west.

93' thick here. Others eastward on road from  
Lexington Spring to Tenn. River, many more S  
and S.-m. of the Spring within a radius of a  
mile.

Those between Lexington Spring and Perryville  
are mainly in the Glenbeith and Lexington zones.

The section here is as follows:-  
Helderbergian <sup>important</sup> } = Chazy } 60' at Cander  
Cander chert

Linden li. { <sup>Upper a Pyburn</sup> Lower a Ron li with Carnegie  
Brownsport 100' upper part apt like li.

with shales below. Cathartes in above.

Lexington clay 30-45' Red clays with Fistulipora

Leop li 30-45' = Louisville li.

Waldron.

Pyburn Landing at White Sulphur  
Springs. There is a fine thick section  
of Helderbergian 100' thick. See  
Forster's p. 683. If time is at hand go  
here.

## Perryville.

In the quarry  $\frac{1}{2}$  mile N.E. of the station is found the Brownsport below followed by the Lenden = New Scotland. It is as follows:-

Lenden base has Orthothetes woolworthanus, Rhipidomella Meta, Striatopora visa.

Lower third of Lenden  $10\frac{1}{2}$  feet. Solid crinoides rock with few fossils.

Middle third  $11\frac{1}{2}$  feet. Richly fossiliferous. Light rock, fairly crinoides with clay beds. At base Strophodontata becki, Strophonella punctulifera, O. woolworthanus, Dalmanites micrurus. Very rare Uncinulus schucherti.

Above this basal zone there is more clay with many small brachiopods.

Upper third, 15 feet. Fossil clay. Exposed in the northwestern corner of the town and is overlaid by thin layers of li. full of bryozoans, in a bed  $\frac{1}{2}$  foot thick. Only bed above which may be the Hardin sandstone.

Camden fossils about Perryville are loose and from the iron ore gravels.



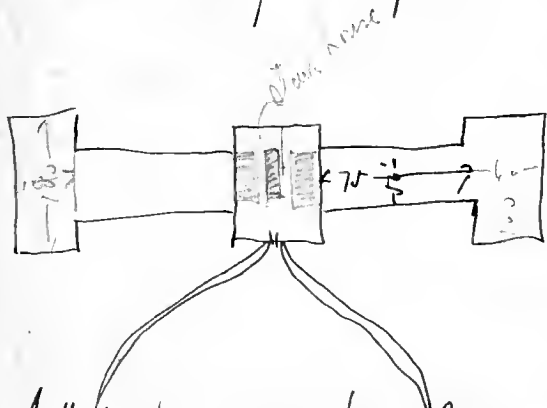


Tuesday June 25-1907.  
 Bronx Botanical Garden.

Museum building.

A grey brick building with pillars of sandstone. Approach from each side around a central fountain to central doorway. In this like the museum at Prof. Eppstein. Three stories, 2 above ground level and one below. Basement for Botany, First floor Economic Botany, second floor Systematic Botany.

Windows high above floor double with a transverse window a transverse. Piers about one-half the width of the windows. Floor plan as follows:-



Supporting pillars opposite piers, above case built round them, hollow <sup>open</sup> inside. Floor broken marble and etc. Inside color whitening. Cases oak.



on their belly. In the water which is very shallow  
they also use their legs in propulsion but it is  
here chiefly the tail. This is lateral from side  
to side just as a snake does. From the base of  
the tail they have a row of these large and  
powerful swimming organs. On the land they  
prefer to crawl if possible, rather than to  
lift on their legs.

Princeton, June 26-30-1907

Arrived here about 11 A.M. on the  
26th and each day visited some of  
the 24th field of France, now  
at Lawrenceville in Ga. Had  
34 men, 18 finches and 16 birds.  
The work is uninteresting but the  
men and the institutions are the  
main part.

Out at Lawrenceville on Sept 15  
others, the young master of the place who  
called my attention to the horse bones

in the exam near Saltun stall.  
another son is taking prelim. exam  
here for sale.

Philadelphia June 30 - 1907.

and here  
at the

the bank  
the mail change.



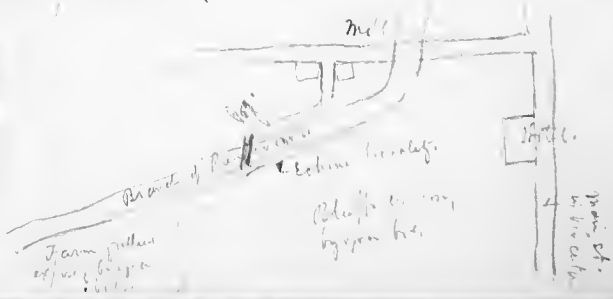
Vincennes 200 2 5

Started out early down the stream to look for other exposures. Saw no other possible places than those of yesterday. Walked to Leontine and then to Mill Creek. We look the way as if all for the same.

The country is covered with Pleistocene mantle that it is rarely seen. In the vicinity of the stream is a little hill, and a mill. A red soil cultural locality however.

Later in the day revisited the place. Found some of the same material of the same nature. Found ten more or less of the same than locality near

about 150 feet of stream over which the stream is crossed. The locality is a good one thus:-





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Vincetown July 3- 1907

Packed material and shipped me small box  
via express.

In reading Parks notes on the New Missisippian  
one is impressed with the abundance of green sand &  
glaucinite. In fact all the iron formations may be said  
to be in the main of this sand. As one goes south  
the beds become more and more massive.

Barrell thinks the <sup>Crinoid</sup> Silurian shore line may have  
extended to the Blue Ridge in which case it must have  
also existed in the Great Valley but not to the west of the  
Alleghenians. In Pennsylvania it extended to Harford.  
His argument in this matter is due to the  
way the <sup>Crinoid</sup> Silurian cuts across the Crystalline structure.  
It could not have done in and then not be so the  
Silurian rocks in which this river had cut its valley  
and as there were striped on one it began to cut  
through the ancient rock. Below to the sea.

Left at 11:30 AM a 35 mile run  
a 24.45 run at 11:30 AM. The  
seeing the hotel and hotel  
be was where we saw at 11:30 AM. The  
Hotel.





In general one gets the impression that the sea came  
in for a short time, deposited the shale, coarse,  
shaly, angular sand with rounded pebbles up to  
pebbles, later followed by more or less associated  
clays <sup>originally</sup> deposited by river.

In some places the white sands of the Pater-  
at is full of rolled pebbles. These are when  
removed from their matrix distinguishable from the  
Pleistocene pebbles in being almost entirely white,  
while the latter are stained with iron.

At Charleston we took the narrow road to  
the N. and came on the <sup>the section up to</sup> top of a little hill  
of Fogs Hill station. <sup>the section up to</sup> In a part of the Pater-  
one sees the associated clays in one place deep  
and in another and some of the sand to some  
depths are bedded to white or bluish white color  
with several hundred pebbles <sup>of the</sup>  
same size as the sand.

After some time the sea came in again to  
deposit the shales and sandstone in some  
places and the nature of the strata in air  
is generally. In one place or the same horizon  
there will be a sandy waste of much size white  
between the layers of sandstone and shale. In

Other places it will be a mottled clay and in still  
another sandy. In still other places there are considerable  
quartz pebbles not much rolled but with all the  
corners rounded. In other words it is a glacial  
deposit of fine coming off a country of metamorphic  
rocks.

In general the sequence seems to be as follows:  
Towards the close or at the close of the Cretaceous the  
eastern shore of North America was probably a low  
lying land with elevations not exceeding 100 feet and  
without any deep estuarine canyons as the present  
Delaware, Louisiana etc. The land was  
depressed some and a very shallow sea made its  
appearance depositing a basal zone of well washed  
coarse sand with well rounded pebbles. When  
this shallow sea was filled up on the sides in  
places more trace sand was deposited in streams,  
thus forming a rather deltas or a debris of fluvial  
deposits of red or variegated clays. This condition  
became general during the Tertiary and in con-  
tinental deposits occur. Now a second shallow  
sea invades in places more persistent than in other  
but nowhere are there any true marine deposits  
laid down. This is the Paritans a Plastic Stage  
of the Cretaceous. A third and decided sub-

confluence sets in bringing in an abundance of  
marine life (the first time). This is the Patuxent  
or Lower Green sands and this confluence con-  
tinues into the Eocene with the sea subsides.

Today I saw nothing in uncomformities  
described by Clarke between the Patuxent and the  
Patapsco and between the latter and the Patuxent.  
In the Patuxent deposits are based on the  
origin it seems to be a local uncomformity  
may occur nothing, but the sea was  
but only after it has been shown in evidence  
variations in the level, that it is much  
easier to find a place on the sea level  
consider its deposition and the position of water.

As mentioned in the notes, it is  
seen so much about the level it is difficult  
to find a goodly amount in some of the cut  
this described, it is a goodly amount  
the Patuxent level, but it is not like those  
of the level that it was the Patuxent level.  
The Patuxent level has been found it also as the  
Patapsco in the cut, and in the Patuxent

Arrived at the hotel at 7:30 and Barrell left on Philadelphia on the 8:05. He will pack his material and ship it by express to Monroe.

Harre de Grace July 5 - 1907

Friday

Packed box for Barrell and left on the 8:48 for Washington. P. R. R.

On the way saw almost no Potomac deposits until near Baltimore, south of Baltimore for ten miles are seen many fine exposures and most of these are near the city. Baltimore is good for six miles or so would be another fine place to visit.

Arrived at Washington at 11 A.M. Then visited the U. S. G. I. Lane Rizer, ~~and~~ Spencer. Saw of the Washington Monument and to Mount Vernon. In the evening called on Mrs. Woodley.

Washington Saturday July 6 - 1907.

Visited the U.S. G. M. Quar Merriell and  
Patterson and Stejneger. In the afternoon  
visited Great Falls.

The upper reaches of Great Falls  
is very plain above the Falls. The new gage  
is steel and an above. The iron sand  
seen here is coarse, and unsorted, showing  
no sign of any particular size. This is quite  
like the Patuxent sand seen on Thursday.

In the evening called at the home of  
and met Dr. Green.

2938

July 7 - 1907. Sunday.

Left Washington at 9 A.M. and arrived  
at Great Falls a little before eleven. We  
climbed the upper reaches of the falls and saw  
some signs of the gage. The gage is  
directly above the falls. The gage is  
at  
in and the Patuxent divides into two and  
out from it at a considerable angle.



The Staveren sandstone makes the two mountain ranges while the rocks in between are of Algonkian rocks. These are a much metamorphosed diabase making the Staveren schist and into this has been formed a granite often very coarse grained. This of these ancient rocks took the material. The Staveren is composed, a coarse grained gneiss with fine conglomeratic sandstone. This series and the Staveren shale are much altered.

As the basal layers of the Staveren limestone are said to have been Cambrian fossils, the same always the Cretaceous sandstone, the question arises how are these deposits related to the rest. Are they younger deposits upon the older? Or do they look like they were deposited on them. Or the higher Staveren limestone comes from the west and some distance from the town has resulted, the sequence and position of the Staveren limestone? These are questions of importance in paleogeography.

July 8 - Sunday Jackson Ferry.

Intended to take the 6.00 A.M. train for Washington Junction but by a curious mistake saw the train go with Leeds but Dottle and I were left. Later Leeds showed me his material. Contact between the Cambrian and Triassic he did not see but a short distance south the weathered contact became on the Potomac conglomeratic marble.

This conglomerate is made up of irregular sized, and well rounded limestone pebbles bound together by a fine cement. The pebbles vary in size from a few inches up to one foot or more as a human hand. This well cemented zone is followed by one of a sandy nature in which limestone pebbles are also found but here they are so small as to be finely and weathered out. Farther east occurs generally a red but also variegated zone of a sandy nature in which are <sup>of quartz</sup> pebbles but well rounded and well cemented. Farther west is found (pebbles from ~~the same~~ shale and quartz like remainder of ~~the~~ Washington and less of ~~the~~ pebbles.

The entire material is derived from at

hand, as the largest rocks in the conglomerate are nearest the Cambrian and become smaller toward the east.

The dip in the sandstone and red shale from the car windows look as if 20 degrees westward.

Whe and I walked along the R & O R.R. for about 2 1/2 miles to see the contact between the Cambrian (Anticline sandstone) and the Sherwood limestone. While we did not see the contact yet the fault line is well marked for on one side is the rugged relief of the Anticline sandstone with the high rounded back of it of the Sherwood shale, while on the other side is seen the very red soil of the lower ground with its undulating surface. The latter is the ground of the Valley and of the Sherwood limestone. Examined all the surfaces along the line of road but saw almost nothing of an organic nature. Nothing determinable.

Not far from the Anticline sandstone was seen something suggesting crinoid stems and the quarry near by the sections which may have been small *Crinoids*. Still farther



which now seem to me to be nothing more than the Mississippian, Cambrian and Ordovician strata under the same (therefore different physically) and in the center of elevation, more greatly metamorphosed. The local strata here are these Cambrian and Ordovician strata. In the region of the eastern Piedmont for here are seen surface flows of lava - the meta-phyolites.

Left at 5:38 P.M. for Luray. First over the R.R. to Shenandoah Park Junction and then via the Road and Luray, arrived in Luray at 8:00 P.M. Luray is a small town.

As we cross the bridge we are in the limestone and the road is on the limestone. When we see a small stream (the Shenandoah) and then to north end forms of shaly and in the evening light is dark blue. As we go on it is seen that the Mt. ... several ridges. In a small case of ... of rocks - an outlier - in a limestone valley due to a ... in the valley syncline and the final preservation of removal is due to the Massanutten ...

W. W. Co. Survey Co.

Subsidiary of the main survey

The surface is a... 50 feet... the... limestone... (a nodular... beds) seen yesterday about 1/2 mile... only one sees a... The cave... little and... as a... cave, now on the... it is... the... is a yellow... there...

The human... but... small... put...



seen on the eastern slope at levels nearly 200  
feet above the stream.

Low up the mountain and not very far  
from the (first) evidence of the Stromatolite I  
collected in the Shale and Trilobites  
beds, Trinucleus concentricus and Pleurodites.  
This, indicates the level of the and Stromatolite  
beds are less than 200 feet from the  
Massanutten and probably about 100  
feet.

Towards the crest of Massanutten Mt. we  
see the acidic and bedded white quartzite.  
For a clearly washed, white, crasse granite caste  
ite and many beds but are not so rounded.  
The evidence of a head is clear and further  
up the mountain has not been found for a great  
distance.

In the east part of the mountain we see  
a small part of the Hamilton strata and quartzite.  
In the beds of the Hamilton strata we find  
also Trilobites and Pleurodites and Trinucleus  
coronatus, Pleurodites and Trinucleus  
form, Cervicifera and Trinucleus  
Second quarters of the Hamilton strata is  
the body running the strike.



At 2 P.M. started out on an exploration trip.  
on the Seven Fountains road  
Loom saw some *Stromatolites* but did not collect  
these in the present.

At Seven Fountains saw excellent exposures on  
two hillsides of the limestone limestone. This is  
locally designated by Loom. The soil over these  
limestone is red as usual for limestone  
in this country. In the soil over one portion of the  
fine limestone we saw siliceous fossils - a number of  
bacteria, a *Stromatolite* and 2 other species  
In the soil over the coarse limestone we saw  
fossil *Stromatolites* and *Stromatolites*.

The fossils of these limestones are 93 ft or all  
"long in the *Stromatolites* -

the same as the *Stromatolites* that are found  
where about 1000 ft and below. The fossils are  
all evidence of a *Stromatolite* as in the  
plates.

Above these limestones  
we found some *Stromatolites* in the *Stromatolites*  
limestone and *Stromatolites*. There may be some  
*Stromatolites*. I saw no evidence of the  
true *Stromatolites*. (See later notes)  
(This zone is the *Stromatolites* zone of *Stromatolites*).

July 11  
Fork Cross Road, Virginia

One species in east of the fork ...  
in presence of Hamillia, among Troneta cerasiata,  
C. vitifolia but not of C. ...  
common, S. med'alis (one with ...)  
its occurrence, S. granulosa (rare), Troneta  
... common). Micro...  
common). antacul...

This locality ...  
takes one into ... of the ...  
...  
Hamilton.

... walked over to seven mountains ...  
in a ...  
All of the seven mountains ...  
if it ... into no ...  
... the fossils.

Above it ... in a ...  
...  
...  
... from 1000 to ...  
...  
... some



of natural ... some ...  
 (It had ... like ... of the ... any ...  
 layers. It had ... a ... mud ...  
 ... of ... or ...  
 ... Can this ...  
 ... Later - this certainly  
 was a piece of ... li. indicating that they were eroded  
 out here. The fossils attached to it were like those at Cherry Run.

... overall ...  
 ... the ...  
 ...  
 ...  
 ...

(Faulted down Upper Massanutten - see above)

This is the Upper Massanutten faulted down. The  
 sandy shale beneath must be the Lower Massanutten,  
 the fossils of which indicate ...

...  
 ...  
 ...

Livingi (some very large ones, *Platystrophia*  
but have too many plications on side and sinus, bar),  
Orthis testudinaria, <sup>white</sup> Orthis maerata and Rafinesquina alternata.

These rocks are ... with the  
lower <sup>down faulted</sup> Mananthee sandstone. The latter comes in  
sharp, a white sandstone ...  
base with fossils of Orthis spanish.

Just above this <sup>down faulted</sup> Mananthee sandstone  
I have collected Trinacrus concentricus, Polymeria  
callicoides, Plectambonites ... Orthis  
testudinaria (looks like Orthis with the ...  
on the same line). The horizon ... must be  
in the <sup>upper or lower</sup> ... with the ...  
lower <sup>down faulted</sup> ... is certainly Lorrainei  
... Taronurus  
... are ... just east of the  
street ...

...  
... (smaller corallites) of 5  
8-10 inches in diameter.



He said the patterns out and we see nothing to de-  
 termine until we come to the very bottom of the Mass. str.  
 The Mt. road then crosses the creek and we see on the  
 dip the other way, i.e. west. In the very narrow valley we  
 soon see the <sup>se</sup> travertine beds, red as green, of the Mass. str.  
 Then comes the last portion of the Mass. str. when we  
 again come upon the Rockwood. The rise here is low  
 and gentle to the Mass. str. when the Massanuttet  
 outcrops and we have a <sup>grand</sup> view of the Mass. str. and beautiful  
 valley of Mass. str.

[2948]

The Rockwood just above the Massanuttet  
 gives one the impression of a stratified deposit in thin  
 masses of blue, peculiar coloring, and  
 and the unusual presence of some reddish  
surfaces. We saw large cracks in all places with these  
cracks in a more or less than 10 feet thick.  
 The lower part is made up of sandstone and red shales.

In the Massanuttet when the top is seen  
it is made up of masses which are either in  
horizontal or vertical positions and no other  
places.

In the Mass. str. one  
sees the fine bedding with the Massanuttet is open  
to the south and west. Between the strata for  
the most part is up to the base of the Mass. str.





clay continues down to the water.

At 1950 feet <sup>altitude of</sup> pick up east of the inlet on

beds.

altitude.

At 2500 feet <sup>altitude of</sup> see

of a Diplopoda and see Calymene beds  
and Trilobites see specimens. The locality is west of  
the mountain. The road on the road of one  
has a view over the hills to the over the Skanaw-  
dash. There is no well shown at the bridge.  
(Later, the same horizon found as (Mantoung of the top of  
the (Mantoung shales))

There is no fossil in more continuous section  
of the Mantoung, since than this one on the  
road from Yordalstock in Mananattun Mt. of  
one exposure beside the road <sup>(see later)</sup>  
to the Mananattun sandstone in the  
see the Mananattun shales and  
in the Mananattun shales and  
the highest part of the shales is  
sandy shale that changes into yellowish red, white  
or mottled white and blue limestone.  
See and write down the names of the more early  
strata of the Mananattun shales.  
one piece was found at black, to  
the road side; the road about 1/2 mile from  
the bridge. See the limestone in the road







June 13 - 1908  
Woodstock to O. O. ...

Left Woodstock at 16 O.M. ...  
here to the Valley ...  
Ferry, and then over to ...

Did more than ...  
contact between ... (=Chenestuy)  
fun. The ...  
St. John ... the ebescis

or more additional ...  
be ... up ...  
did not ...  
slab on the top ...  
all ... with ...  
the stack ...  
from ...

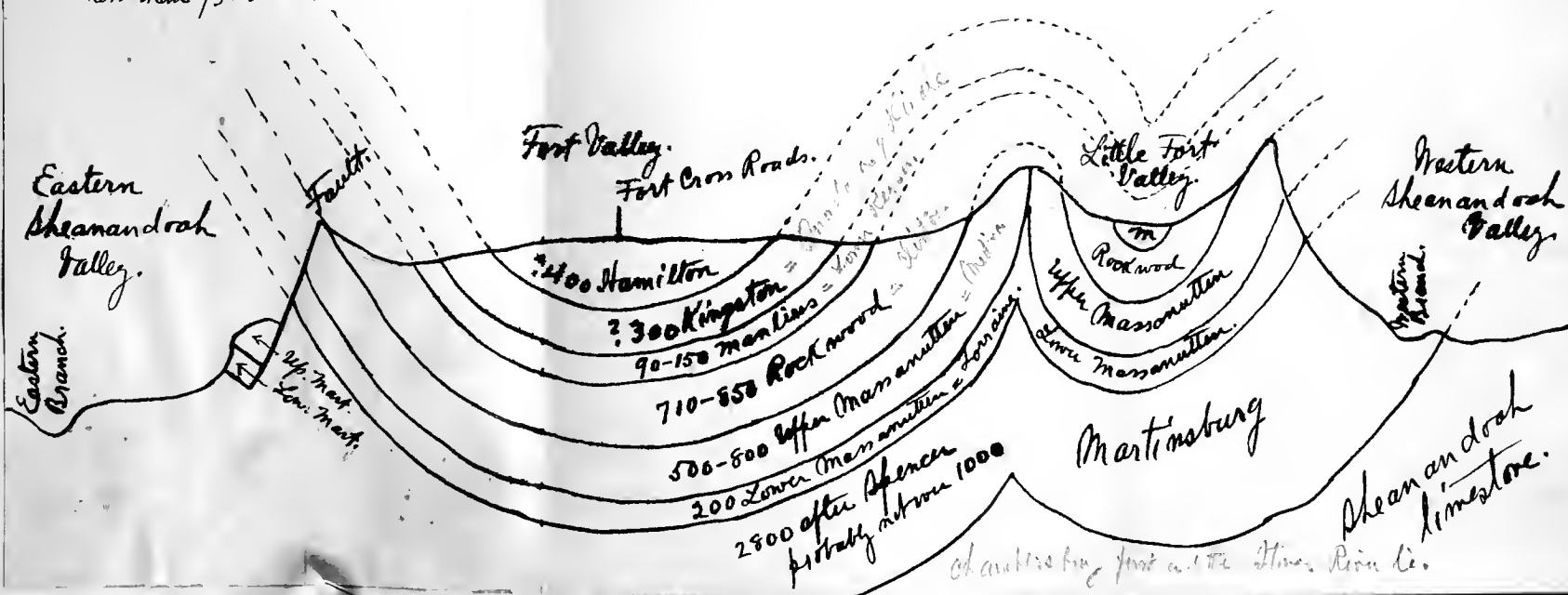
The ... here ...



July 13 - 1909

I designed section across Massanutten Mt. from east of Seven Fountains to Goodstock.

This side also has a small valley as on the west but less than 1/3 as wide.



of one  
Del c  
regul  
a; at  
main  
the  
m  
creek





Combined with the <sup>Lower</sup> Mananuttet results it runs up into the Lorraine.

About 3 miles west of Martinsburg limestone appears again. Its lithological aspect indicates it to be the same series that underlies the Martinsburg near the city. If this is true the Martinsburg lies in a syncline and its thickness is probably that given it by Leitch 700-1000 feet.

Martinsburg to New York

July 17-1907

Went west from Martinsburg to New York

appears dipping little to the west.

Leitch's measurements 700-1000 feet.

at Martinsburg.

at New York.

at New York.

at New York.

at New York.

at New York.

at New York.

at New York.

at New York.

at New York.

at New York.

at New York.

at New York.



deposits. The country rises considerably toward the mountain and the B. & O. goes through the lowest part of it.

[F673]  
[F672] In the cut through North Mountain we started our work in the Becraft limestone where I collected considerable loose fossils. The lot with the greatest abundance of Rhipidomella assimilis is from about ten feet of the top of the limestone, while the smaller lot is from a zone about five <sup>feet</sup> lower. The Becraft limestone is continued upward by a shale series in which few fossils can now be had due to <sup>these fossils become more calcareous during the life.</sup> the mountain making. Phacops and a small cep. corals are the only ones readily seen, but on breaking some of the limy pieces from the <sup>top</sup> near the top of the series collected Styliolina, ostracods, an Orthis and Aenolia. <sup>They are the same phase of the Onondaga.</sup> Have taken these fossils. The limestone has at least one zone of conglomerate.

The Marcellus Haell shale rests directly on these <sup>and probably of distinct periods.</sup> Helderbergian strata. The two seem to be conformable but there is here absent all of the Onondaga and Onondaga. <sup>(the calcareous zone)</sup>

Above, comes in the Hamilton, near the base of which Reed collected a slab with Spizifer zigzag, Orthothetes and Chonetes (see

(in the concretions beds)

slab). Higher up, Hoke collected Tropidoleptus  
carinatus, a large Panella, and saw Spinifer  
mucronatus.

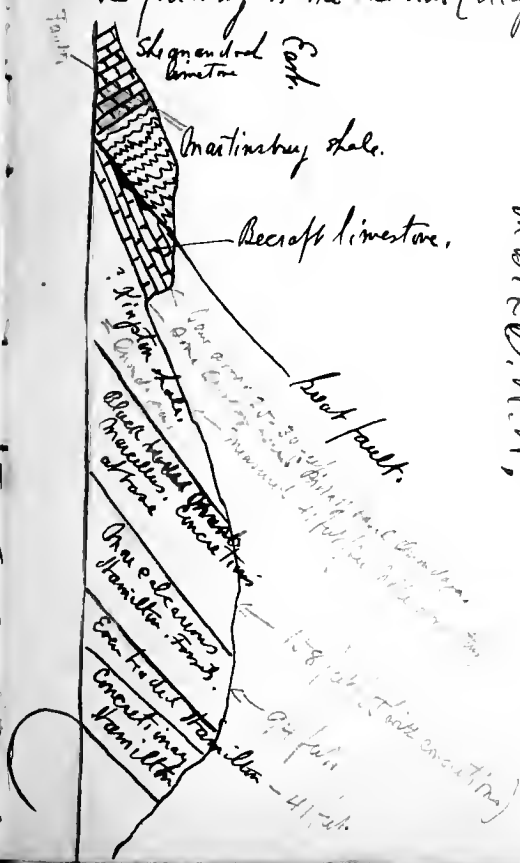
Hoke soon determined the overturned  
nature of the syncline but at first drew it as  
an anticline. Reeds saw nothing more than  
an descending series ~~but~~ <sup>and</sup> did not make  
use of the overturned syncline.

Resting on the Beecraft limestone is  
another greatly crumpled shale series which  
at first I thought may be Salina. After  
some work collected on the dump to the  
north of the tracks near a farm house  
several specimens of Climacograptus. The  
horizon is therefore Martinsburg and the great  
overthrust fault here lies between the Mar-  
tinsburg and the Beecraft. At this point  
there is therefore cut out some of the Helder-  
bergian, all of the Salina, Rockwood and  
Massanutten.

Driving south from North Mountain sta-  
tion to Hedgeville I saw in a stone wall  
of the boarding house on the Hill, blocks  
of what appeared to be the Massanutten  
and certainly blocks of the red continental

lets deposit of the basal Rockwood. The ridge of North Mountain is considerable high here and rises to the south and gives the appearance of having on its crest the Massanutten quartzite. In other words the great over thrust fault of the valley reposes on any Silurian horizon up to the Becroft.

The following is the section (diagrammatic):-



Continued local syncline  
at North Mountain, one mile  
west of North Mountain station  
on D. & O. R. R.,



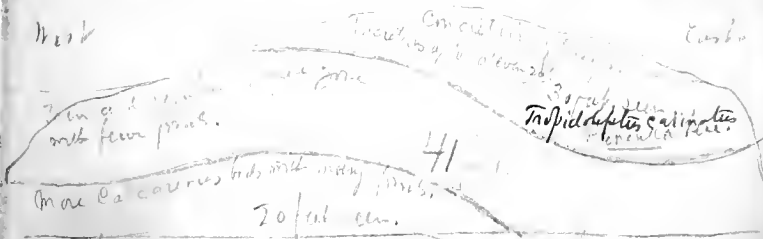
the Becraft are are continuous and there is no  
 essential difference in lithic appearance <sup>except for the slaty</sup> between them.  
 The thickness by stepping is 84 yards with a dip to the  $\nearrow$   
 east of  $30^\circ$ . This makes the thickness 100' at.

675 The Hamilton is not collected on the base as  
 here determined. Leptæoclema, Leptæoclema,  
Streptelasma rectum etc.

The section is continued down with the thin  
 bedded limestone, but the <sup>in 100 yds dip</sup> section is not well exposed.

These strata are all on the same side  
 and are i.e. Hamilton Hamilton Hamilton  
 in the rocks and above the Hamilton Hamilton  
 the Marcellus Marcellus Marcellus  
 considerable Marcellus Marcellus Marcellus  
 the Marcellus Marcellus Marcellus

In a cut on the side of the Hamilton Hamilton  
 the Hamilton Hamilton Hamilton



The Hamilton has an abundance Cometes portales  
C. portales, C. portales, Tentaculites,  
Strophomena

... cut ...  
This is the zone that the two limbs of the syncline rest against each other

... a measured ... 94'

... the ... a measured ... basal ... (Foot) ... column

... All the ...

... the ...

... the ...

... a Beograd.

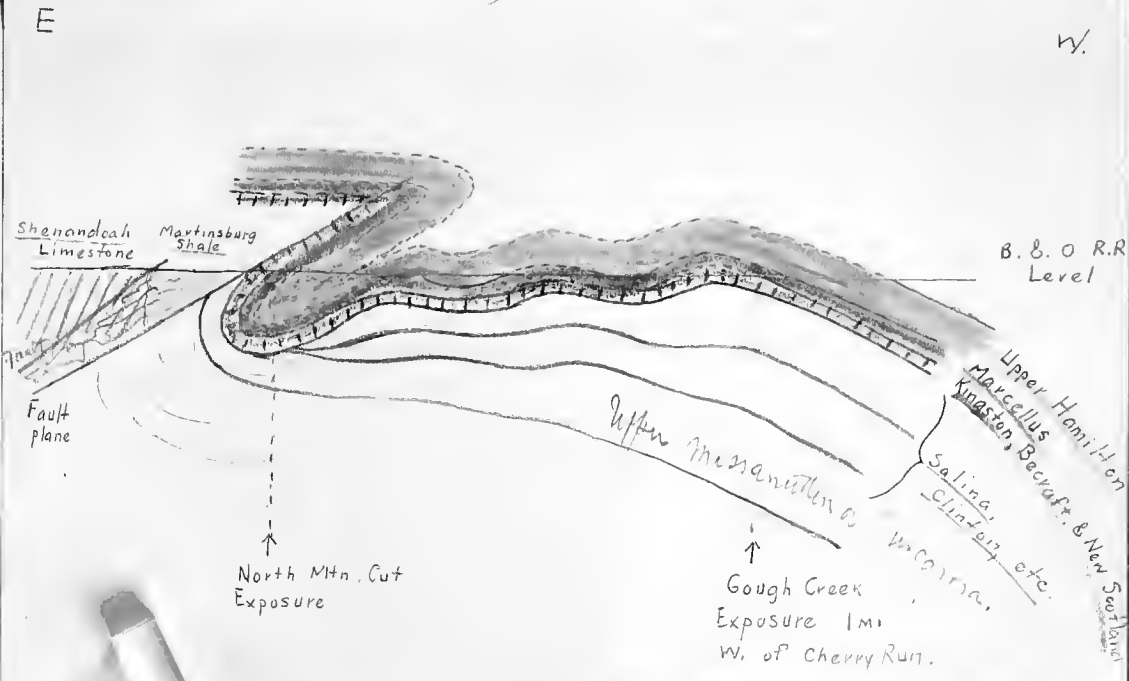
... (Comparison) ...



into a Jan  
 but I can  
 thru lines  
 doubtedly  
 limestone  
 Mountain  
 (1921  
 Co. of p.

Dec. 93  
 July 15

Idealized Section  
 on B & O. R.R. level between  
 North Mountain and Gough Creek.  
 (After L. F. Noble).



San Vito de los Rios

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100

like a sandstone, the thickness could not be determined  
but it can not be less than 4 feet thick. Therefore  
three times as thick as at Lough Remy. It is un-  
doubtedly this layer that represents the massive  
limestone of Seven Fountains in Grassmount  
Mountain.

(1921. That in the complement at the base of the  
Carrigrohane lies on the lower ledge.





By rail-road it is about 10-15 miles from  
New Rochelle to Flushing Meadows and while here there is  
no Christiana <sup>at the latter place</sup> - the population is  
great and of mixed blood - a few present at  
Harrods for the sea - one from the west and  
a few of the same at Harrods. Of course  
in different places there is plenty of  
of minerals.

In the west quarry in Flushing Meadows  
is a layer of <sup>of</sup> which is up in the  
mountain side, at least 100 feet in height

75 feet wide and 100 feet deep one is un-  
dermined by the action of the sand in state of its  
being porous and of a nature of  
This is the coarsest water of that is used as  
the water of Flushing Meadows is about  
one third of the water is thrown away, the  
remainder is washed and then  
is used for the same. It is also pumped  
and used for the same and is  
used for the same of water  
in day  
about 4

is also used for the same at this locality  
is used for the same of water  
about 4



By out-road it is about 10-11 miles from  
Dunrobin to Slary Glen and while here there is  
Dunrobin at the latter place.



about 10-11 miles from  
Dunrobin to Slary Glen and while here there is  
Dunrobin at the latter place.



Originally it must have been considered with  
lime because of the many corals. None of  
the living shells or life, all are casts. Thin  
sandstone in the great anticlines is a water zone  
and it is the sandstone that has taken out  
all the lime. The view that we get now is  
it was probably formed in the fossiliferous  
strata.

When we go again to the great  
limestone quarries, interesting. They are a length  
of 1/2 mile. The material is hard, a good  
metal and flux in steel making. Some is used  
for lime, and some for...

In all the quarries the material is the light  
blue (weathers white) limestone of even grain in  
beds. Usually the thin bedded uneven  
grained limestone. It has fossils, but we could  
be made out. I was quite surprised to find  
in some a limestone at the base of the  
decidedly was a fossil. These are the  
thin beds, the ones in which  
and see in Sunday... in the  
section.

These light-blue limestones are worked in the

soil divide ... some ... then  
latter part the red part to a depth of ...  
which ... the surface has ...  
and ... This common to see the  
prints in the limestone well ...  
and ... In some cases, even  
... These conditions are ... for the  
upper ...

(about 1/2 ... on some ...  
holes of a ... (very much like ...  
then upon a shale ... and ...  
... of these ...  
... these ...  
Took a photo of ...  
out in place.  
... next ...





One is impressed with the beds above the cement  
in being <sup>to</sup> decidedly more cradled. They are  
of a thickness of about 200 ft. and consist  
there are at times shales in part with spots of sand  
all seen cradled. The bedding is not distinct  
in a continental nature and at the same cradling  
is the great mass of it referred to the same.  
The beds are and sand has a sea floor of sand  
around it in a somewhat form. In the same  
we may see the jet at the island or point here  
an abundance of the beds. The beds with  
the peculiar reddish green markings are the sand  
series below the cement beds and a sand beds  
with sand rocks are laid down. Later the sea  
beds are the great red massive series of beds  
and the beds are laid down.

red beds for  
17 pages

Faint handwritten notes at the top of the page, possibly including a date or location.

# Cumberland, Md.

below

hot  
offer the  
eyes  
a  
distance of 30

at D.B.B.  
distance

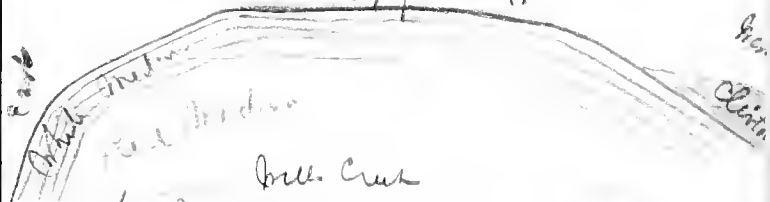
the  
distance  
here

Wells Mt. of mass

Cum gratia and Sep 1-1959

The red Medina sandstone shows much cross bedding but is not extensive. Cut out one one foot deep. Sand equal in grain throughout. In some layers irregularly sided vertical holes that were once round then irregular. There are of organic origin but not of the present.

Tog of Wells Mt.



The Red Medina is a series of alternating sandstone and red shaly shales.

Red Medina passes into White Medina. The latter some at Clifton

The White Medina is not very bedded like fine grained quartzite with some cross bedding. It becomes thinner bedded with large beds and the more shaly and fine into the Clifton shales.

Red shales 6"

etc etc etc 14"

Red shales







7  
The ... of ...  
The ... of ...  
The ... of ...

As one goes up the hill ...  
one ...  
one ...  
one ...

July 20th Saturday

... ..

... ..  
Spent the day with ...  
... ..

The bed of my ... section ...  
... ..  
... ..

The New ...  
thick here as on the ...  
... ..  
... ..  
... ..

shales of the Lower Oriskany appear. In these beds I saw Leptocoelia, Phyllites, L. imbricata, Oitholites (see specimen) and the genus Spizya described by me. The corals were in situ.

The erosion of the bed in question beneath the Lower Oriskany is decidedly different between Lee and Capt Valley. At D.B.B. only 10 feet of the New York and is left. At Conjansville a part of the same mountain is present. At Capt Valley all of the bed in question is gone as the Lower Oriskany has been eroded to the mantle.

Out of the sea cracked layers 3C Rocks they are Styria, Styria, Styria, Styria and Styria (see spec). These were in a rolled mass as if dropped into the limestone, not rolled and bed. It is difficult to understand the various cracks and the various positions in the rock beds, a more mature. The sea was in a state of boiling and the various rocks were deposited. Certain of the rocks (see specimens) seen in the sea were for one seen - these were in the limestone (see spec) which were into normal marine deposits with the fossils.



July 23. Tuesday. New Ireland.  
Pinto = P. Tomaleson, New Ireland.

Took the school train. <sup>the same page 141</sup>  
Just at the station we see an iron ore vein  
which has a thickness of about 15 feet, about one mile, and  
iron ore. Beneath is a shale zone & also a zone where  
another form of iron appears but on a red ore. This zone  
has a thickness of about 100 feet. (Note on)

The iron ore bed has an abundance of Orthoceras  
Orthoceras and some <sup>small</sup> Pelecypoda. In  
places an abundance of Forams. These ore beds are  
due to subsequent deposition on the iron  
ore. Orthoceras and Forams are inter-  
mitted with shale and iron ore. Orthoceras  
beds of a thickness of about 10 feet  
with iron ore and shale veins of various  
sizes. Many of the Orthoceras are small.

Passing over a gravel bottom we reach the Clinton  
mill and continue up the BC To mine Clinton.  
The Clinton is separated from the Hippen Member by  
a sandstone, about 2 feet thick, and by  
shales and sandstones about 3 feet, and another sandstone  
about 18 inches thick.

The Clinton passes other than the sandstone  
Hippen Member and beneath these sandstones about  
3 feet the beds are red and black = base of bed



looking like that of the New England. It does not seem possible that there are corals here for the yellow (weathered) stone of the Lower Oolite any more is seen on the road of the hills. A little farther up the hill is a red metal ore, indicating the presence of the Lower oolite. The entire thickness of the Oolite any here can not be much more than for the Romney is seen a little farther on in the road.

Valley ... .. is 1080

Maajan ... .. 345

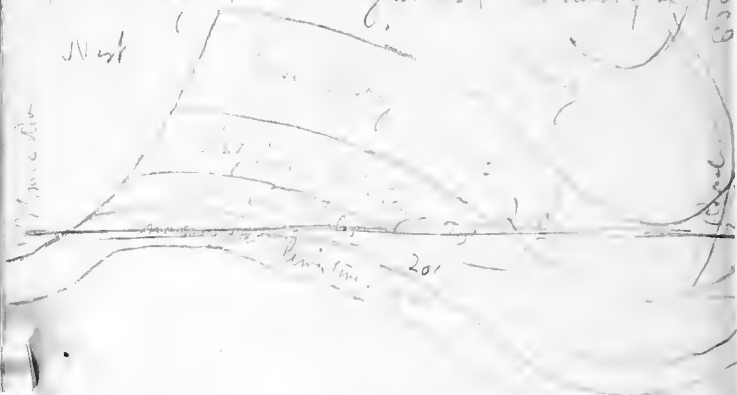
Make these ... ..

The ... .. is first ... ..  
 turns probably ... ..  
 ... ..  
 ... ..

... .. and examined ... ..  
 section ... ..  
 Saw nothing to add to my ... ..  
 a few fossils ... ..  
 ... limestone.

We ... ..  
 Lower Oolite ... ..  
 ... ..

100, 110, 120  
 the cost of the  
 Any or total state cost of 24¢ per  
 100



Secured for little part the State  
 and P.P. in the area, in the  
 position of the rocks in the  
 the middle regions of the  
 lbs.

measured the for in the  
 and found the shale in the  
 The right side of the  
 all 17' b. <sup>thick</sup>  
 when the  
 limestone  
 seems  
 in the  
 cont. in  
 but these have

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Loma section at Leesee, Va. (doc. 93)  
September 8-1908. July 23

Base. Thin bedded dark blue waterlime. Beds nearly  
2 inches thick when weathered. About 18'

A similar zone with shale beds at top and bottom, each  
about 5' thick. Center dark blue li. in beds 2"-8"  
About 20'

Dark blue waterlime in beds 6"-24" About 18'  
Shale zone " 4-5'

Dark blue waterlime in beds 2"-24" Becomes  
less pure below, passes almost into a magnesian  
shale. Horizon of Judge Alders R. lamellosa  
across the hills. About 120-135'

End of Loma quarry. At the base of this quarry  
a small fauna reminding of the Carl Valley  
fauna. Rhynchonella neglecta type, R. lamellosa,  
Bygonia, Rhynchospira, Ostracoda.

Have this fauna from the quarry.  
It is about 100 feet to the base of the large  
quarry higher in section and of which  
a section has been published.

---

See Hanging Rock section at Gapacomor  
Va. at end of this book.

1. at the top of the page. Let these have

no macropleura but have other New Westland fossils.  
These chert beds then soon give way to heavy bedded  
dense crystalline and crinoidal limestones of the Eri-  
mons. Balance of section measured by Reeds. but  
I did not take them down because the contacts  
were not so readily determinable as some years ago  
when I did the work for the same group.

At Twenty Mile Bridge the lower limestone beds  
of the macropleura zone as in the Reeser gossamer  
do not have the fossil. In both places the shale  
immediately above belongs to the New Westland.  
The Lower Gristley that comes in clearly and  
distinctly. There is no transition zone.

July 25 Monday Cumberland

Packed our collection in tin cases and with  
some packed boxes and today have to  
go to the station.

Left London at 11:15 AM. Arrived  
at the station at 4:30 PM.  
continued to the city.

2939 ↓

[E5622]

July 26 Tuesday To Cincinnati

Left London at 11:15 AM for Cincinnati on  
the 5:12 PM.

At Clarkburg of Pa we are on the 11:15 AM  
and we are in the city of Cincinnati when this  
arrives.

continued to the city.

[Jul 27]

July 27 Wednesday  
continued to the city.

[C0520] + C9218





July 30 - Tuesday, Madison.

The section at North is now down to the  
 to ... (in ...)

Heavy bedded Niagara dolomite with casts of  
Strophomena radiata, Alipha reticularis, and Strophomena  
 ... much thicker ...

Thin bedded ... dolomite, almost ...  
 ... (that glacial ...)  
 ... Alipha reticularis  
 ... 0 or 12 feet (beds 2-5 inches).

Labrador  
 ... normal ...  
 ... (has 18 ... 3 feet) ... here  
 ... (has 18 ... 3 feet) ... here  
 ... (has 18 ... 3 feet) ... here

... 25 ... The ... at the top.

Following the same line ...  
 (2-6 inches). After 8 feet ...  
 ... about a reef ...  
Calapogon  
Strophomena ... and some Strophomena ...

... the same ...  
 ... with ...  
 Then the thin bedded ...

The result in Ohio is that the  
Rhyolite is conspicuous. These beds are seen  
in two cities representing a thickness of  
not less than 150 feet.

As far as this section goes the  
all appear to be conformed and an unbroken  
sequence. The conformity becomes more  
evident, which is not only due to the  
stratigraphic position, but also the bedding  
into a nice conformation, to shales with  
of beds. The conformity is to these  
shales and conformity to be a sequence.

Higher conformity is to be seen above the  
conformity beds of the conformity.

^ conformity beds of the conformity  
is conformity beds of the conformity.

Left about 15 feet conformity beds of the conformity  
Lancaster.





Adelphi, Joplin, Ansett, via, \_\_\_\_\_ and  
at \_\_\_\_\_

Stopping at \_\_\_\_\_ and \_\_\_\_\_  
\_\_\_\_\_ day. A letter was sent to \_\_\_\_\_  
at 3 P.M. for \_\_\_\_\_

The railroad track goes over the \_\_\_\_\_ plain  
and then through the low but rough hills of the \_\_\_\_\_  
\_\_\_\_\_ a long ascent into the \_\_\_\_\_  
\_\_\_\_\_ town \_\_\_\_\_ the \_\_\_\_\_  
and remains in it up to at least \_\_\_\_\_  
a distance of \_\_\_\_\_ miles. Through \_\_\_\_\_  
\_\_\_\_\_ and but for a \_\_\_\_\_  
The \_\_\_\_\_ is almost the \_\_\_\_\_  
\_\_\_\_\_ (the \_\_\_\_\_ miles to the \_\_\_\_\_). The  
country is \_\_\_\_\_ and \_\_\_\_\_  
first in the hills, a few \_\_\_\_\_  
\_\_\_\_\_ is  
deep as the \_\_\_\_\_  
that \_\_\_\_\_

Arrived at \_\_\_\_\_ at 5:30 \_\_\_\_\_  
\_\_\_\_\_





Peppan Section showing (20 miles W. of Fuller quarry  
Hanson)

Hamilton above Anderson  
Photo by Frouster.

See S. I. A., XII, ~~page~~  
p. 408, fig 1.2.

{ Mace's Hole

1-2 feet sandy shale with Barroisella  
= Hardy's section

1-2' heavy bedded *St. americana* ls.

*Anchyrocerinus bulbosus* etc

8' heavy bedded *Trilobites* *Primitivus* *Crinoid*

ls. with *Murchisonia furcata*.

Orulof

Leicester Niagara.

In conversation with Lander April 1908

(Huron? Saturday, Hanson Tenn.

Took the 7 A.M. train for Hanson. I looked up  
the old quarry near the station but soon  
found it was the *Fuller quarry* only present.  
The Fuller quarry but rain soon set in and stopped  
all work. Returned at 1.32 P.M. and looked  
through boxes. The first lot was...



Secured a part quantity of Walden crinoids  
and few shells.

The dip. to the west and ...  
the quantities in beds dip at a  $20^{\circ}$  S. E.

The Walden fauna is as well preserved as in India-  
na and so far as the brachiopods are concerned there  
appears to be no local species. The only about form is  
Rhynchoteta americana, among the crinoids there are  
several local (and) forms (Circulopora) and a  
? Caryoceras quite different from another. None of  
the crinoids are complete all have been broken before  
excavation. Many of large valves lie with the broad  
typical side downward and the basal side more  
or less vertical. Many brachiopods are separated  
valves. The higher are less common here than in  
Indiana and this is especially true of the Tremat-  
ites.

August 2 Saturday, Dawson Tenn.

Took the 7 A.M. train for Dawson. I looked up  
the old quarry near the station but secured  
for less than in the <sup>Samuel Smith</sup> ~~Walden~~ <sup>only present</sup> ~~quarry~~.  
The Fuller quarry but saw some Spirifer and Stroph-  
all sorts. Returned at 1.32 P.M. and loaded  
three boxes. The hotel ...

project and mail bill of today to Camden.

2949

Account Sunday. Camden.

Left New York on the 2:15 P.M. train for Camden. Arrived about 5 P.M.

Went to the gas station, 2 p.m., my father gave me a ride to New York. We through the back state and other into the city.

Account Monday. Camden

(4) Collected in the Camden Lower part of the village near the village to the south.

Went on the hill slope back of the houses in the northern part of the village. Several were collected there on my way to the top.

So far as I can see the hill is a part of the base of the Pleistocene. It certainly is not the Pleistocene. The shale in the lower part is a layer of dark shale with some small pieces of wood. There are no fossils and it is not a Pleistocene. The shale is a part of the Pleistocene. All of these objects were collected on the hill. The shale is a part of the Pleistocene.



The Camden has many dips but did not measure them. They all dip toward Lower Devonian appear above on the west side of the Tennessee.

August Tuesday 6. Holidays.

Started out with a heavy fog in the morning and after driving for more than one hour found that we were on the wrong road. In driving westward to Hollidays we came to the bed of a stream about 1/2 miles to the east of Sis Mondays in a small stream bottom. Here we saw a few feet of the top of the bed of the immediate south the Camden also. The former is of the same character as that near Hollidays, more a thin bedded limestone than shale. These beds are full of bryozoa - the Trenton type. The Camden comes in sharply - there being no transition between the two the Camden and the bed of the stream.

(1) About four miles north of Hollidays beside the road we spent the day in the upper beds of the bed of the stream. Here it is a thin bedded limestone series with shale partings bearing a mass of bryozoa. At least 25 feet is shown to the road. Limestone these beds consist of a series of layers, etc.

Palamedes, Quadrifidus and Lyellia galeata  
are seen. In other words it was New York basal  
New Zealand.

(2) From base limestone, to the spring the road  
down to the stream may be seen the shaly series  
of the New Zealand reptile with Stictolites  
Palamedes Stictolites, Palamedes Stictolites. These  
Stictolites Palamedes, Palamedes, Stictolites  
The middle of the thickness may be seen part  
a Palamedes Stictolites Palamedes. This zone is mostly  
of iron 15 feet thick. From the  
stream may be seen still Stictolites Palamedes  
the top layers of Palamedes Stictolites in  
abundance and stems, which in New Zealand  
is called Palamedes Stictolites. This is  
undoubtedly the same as the Stictolites  
in my hand. It is Palamedes Stictolites - at the  
of the Palamedes Stictolites is where Palamedes  
may be seen in the Palamedes Stictolites  
layers of Palamedes Stictolites in the  
the Palamedes Stictolites are the Palamedes  
of the Palamedes Stictolites and Palamedes Stictolites are  
seen by them

August, Tuesday, Perryville.

Packed car for tomorrow and left at 9:58 AM, and left for Perryville via New River and Lexington. Arrived at 6:30 P.M. after a week of 40 hours of driving.

2947

August 26, Wednesday, Perryville.

Went up at 5:30 and climbed a part of the mountain. A specimen about the middle of the ridge.

Here may be seen the top 20 feet of the Brown-  
part upon which with others. It begins with a limestone member, a  
a circular limestone quarry an  
abundance. May be seen stones and  
Camarevilia. These same beds are seen in the  
New Zealand in a far in it I saw maclurei  
fulvipes, tristis, tristis, tristis,  
and tristis.

These limestones are followed by blue shales that weather  
out as a trap. I saw many fossils of a few trilobal  
types. One of the fossils of tristis was seen. There  
appear the first tristis and tristis.

Above is a debris of modern corals and shells  
and all some of the (2) beds of the clay.



August 10 Saturday. Return to  
Louisville

Shipped one box from Perryville.

Left at 7 A.M. Arrived at Lexington 9 A.M.  
Departed at once for Paris where I arrived at 11.30.  
Left Paris at 5.10 P.M. and arrived at Louisville  
at 2.10 A.M., having had a sleep on a sleeper  
did not have light up until 6 A.M.

August 11 Sunday.

Left Louisville at 9 A.M. Arrived at Cincinnati  
at 1.50. Had dinner at the Grand Union and  
left on the Big Four at 3 P.M.

Went 25 miles east of Louisville on the Big Four  
the last. The Niagara plain is seen. In the  
foreground is undulating and the hills of Richmond  
appear seen on the horizon, Ind. Then continued  
the road dips constantly into lower Cincinnati  
beds.

Arrived at Cleveland at 6.40.  
Stopping at American Hotel.



of the Hamilton fossils. The lower 10 feet are of  
heavier beds and apparently more magnesian  
and less fossiliferous. It is the zone that is  
preferred for lime. The upper 10 feet are ~~thin~~  
beds - even shaly. However no separation can  
be made at this point for the ~~lower~~ <sup>lower</sup> ~~and~~ <sup>and</sup> ~~upper~~ <sup>upper</sup> ~~beds~~  
are identical.

In places the beds are fissiled and deeply  
channeled (not so pronounced as at Kellys Island).

The ~~side~~ quarries seem to be identical  
with the Kellys Island Lime Co.

These beds carry more fossiliferous than the same  
beds on the small side of the lake at <sup>not shown</sup> ~~there~~ they are more  
bedded heavy, there <sup>are</sup> coral reefs. While the  
same corals occur in both places they are  
more abundant in the north. In the north  
the beds are more cherty and the fossils really all  
silicious. This the ~~stratigraphic~~ <sup>stratigraphic</sup> ~~is~~ <sup>is</sup> ~~due~~  
rather a share or is this alteration due to  
leaching in the north at <sup>here</sup> ~~some~~ <sup>place</sup> ~~where~~ <sup>because</sup>  
of the Black Hall cover?

It seemed to me that ~~the~~ <sup>the</sup> ~~same~~ <sup>same</sup> ~~cuts~~ <sup>cuts</sup> are  
most abundant in the coal layers. They are  
present in good abundance and especially in the  
lower 10 feet. They are far removed from a <sup>land</sup>

-14- Wednesday

August 13 Tuesday, Canton, C.

Packed the Sardine boxes and shipped by Express. Left at 11:00 A.M. for Canton at 11:00 A.M. - see diary.

Met him and went with me to a place and stayed. Left for a nice town with him. Remained at her home until the next day at 11:30. Returned on the same P. & O. Cleveland. Left the latter place at 2 P.M. and arrived at 10 A.M. Falls at 10 A.M. Stayed at the Grand Hotel.

August 15 Thursday, Cincinnati.

Started on along the Central line - hand  
to the Cincinnati.

The following describes a series of  
them and many of the details in which many are  
the fossil corals and some a few other fossils.  
Alternating along the section are also corals  
and shells. The fossils themselves are small  
and some of the shells are small. Towards the base  
of the section are some of the fossils -  
the corals and shells are small - but one will



with layered. Beneath the cross bedded layer about  
 about 3 feet thick shows more action of some force for  
 the material is considerably thrown up in waves thus



Beneath that layer the strata  
 is also somewhat shaly and  
 is then cut as a bedded side

The zone is about 3 feet thick and then appears the  
 regular Rochester shale, as far as general appearance  
 goes almost base of point.

The first abundant fossil zone in the Rochester is  
 to appear about 15 feet from top. This zone goes down to 257 feet  
 thick and has some thin layers of limestone filled with the  
 characteristic fossils.

Then comes a mass of heavy bedded  
 fossiliferous sandstone, some has about 15 inches thick  
 of this bedded sandstone. *When quadrifid, trigonum*

Then a zone of fine grained sandstone about 4 feet thick  
 of open, very thin bedded sandstone slightly crossbedded  
 about 6 feet. Followed by variegated red and grey cross

bedded sandstone. This zone passes into alternating shales and  
 thin zones of sandstone. Then the regular *leptod* *leptod* series  
 for *Artisphyschus laurani* appears a mass in these thin

beds of sandstone. Sometimes the sandstones have thin zones  
 of *irregular* bedded beds up to 1/8 to 1/4 inch. All other thin flat  
 shale inclusions up to 2 inches across. In other places



for the sandstone near the center. This zone is about 25 feet thick.

Then heavy bedded sandstone probably ten feet thick (Could not see beneath the cliff. This 25' thick

Below are red shales alternating with sandstones or hard beds. Some seem to go down to the river and may be over 100 feet thick in some exposures.

From the evidence today it seems to me that the Medina is far more closely related to the Niagara as here shown than to the Richmond of any area.

There is not a single Richmond fossil present in this Medina ~~rock~~ while all seem clearly related to the Niagara.

The Upper or fossiliferous Medina is clearly a marine deposit for about two-thirds of it is regularly bedded, one always sees traces of the Lejopoda and at about 27 feet from the top in a zone not over 4 inches thick is where I collected the pelecypoda.

This Upper Medina clearly goes over into non-typical marine deposits in which there is much cross bedding, some scouring out of deposits and subsequent filling of the cavities, rolling of mud into irregular pellets and brecciating of shale into the sandstones, and occasionally a little evidence of irregular

sun-cracking. It is in these deposits that one finds the Arthropods and other so-called fossiliferous markings. Finally the red color leaves these deposits and the last 6 feet or so are a white sandstone.

Above these sandstones follows at once a different colored shale - a green shale and then the Clinton dolomites. There is hardly any thing about these Clinton fossils to distinguish them from the Rochester shale - it is only the presence for few things as the Preistonia cylindrica and the Athyrid quadrifidata.

Can it be that the top of the Medina represents a land <sup>interior</sup> and that this Clinton is the final or topmost member invading from the east. It is at Rochester that one finds the Pentamerus Morysi in the Clinton but it is not present here in the Niagara River pass. Then too it is at the close of the Clinton that one finds the reefs with the oyster species.

At the top of the Clinton here there are no reefs and all the fossils seen are those of the Rochester.

The next highly fossiliferous zone is near the center of the Rochester shale and for about 20 feet one sees the regular trigona and associated fauna. Then practically barren shales passing rapidly over into the fossiliferous dolomites.

It was greatly surprising to find beneath Anthropus bearing beds and the regular Medina fauna the same Bygonia beds seen by me some years ago at Hamilton and now at Washington. These fossils which will certainly pronounce Clinton as it has one of the characteristic Clinton Bygonia <sup>? Bygonia</sup> of this. Further it is from this zone that came the stapes sent me by Grant. Amongst these Bygonia I saw a valve of Orthis and the slots taken will indicate the balance of the fauna. If these fossils correlate with the Ohio Clinton then clearly the true Medina fauna is above and is of Silurian age.

August 16 Fairport. Left at  
departed at 8.30 in the boat arriving there at 9.00  
Called on George T. Mac and he took me to  
the Westmore quarry. Here about 20 feet of Medina  
is shown. The basal layers for feet are thin white  
sandstones with little conchoidal. These follow  
about 30 feet of thin bedded sandstones with some  
shales. These sandstones are decidedly conchoidal.  
This towards the top. The bedded sandstone  
the Medina fossils. Then  
came in mostly shales with some thin sandstones  
some of which are well ripple marked, all others









Hollow beneath the base of the quarry these make the top of the Medina. ~~The sandstone has a certain amount of more than 10 feet above the section of the first quarry takes of as far down (down)~~

To the south east of the mountain there is another quarry. How the water in it is to be seen in one wall. This is a red sandstone. The base is where the dolomite begins on the Medina side. The height of this face is about 32 feet high.

The Medina red sandstone is not so coarse as those at the top of the first visit to the quarry. The concretions larger in them, but less well developed, and about eight feet beneath the quarry is a layer of red sandstone. This fine texture is also seen in the quarry.

A little higher up the road one sees that the dolomite series make a precipitous cliff. And the Medina beneath is a strong red sandstone.

The Medina red sandstone is not so thick as the first quarry, and about 8 feet beneath the dolomite.

Very low hills of gray limestone are seen on the west side is of a red sandstone, just like the Medina shale seen in the Niagara gorge. From its level up to the top of the hills the elevation may be 200 feet.



## Boxes shipped

- |   |      |             |       |          |          |
|---|------|-------------|-------|----------|----------|
| 1 | from | Vincennes   | by    | Express. | Shet.    |
| 5 | "    | Cumtland    | and   | Ind      | Freight. |
| 1 | "    | Louisville  | Ky.   | Express  |          |
| 3 | "    | Knoxville   | Tenn. | Freight. |          |
| 2 | "    | Camden      | Tenn  | "        |          |
| 1 | "    | Springville | Tenn  | "        |          |
| 1 | "    | Cleveland   | Co.   | Express  |          |
|   |      | Landusky    |       | Freight. |          |
| 1 | "    | Highland    | Ind.  | Express  |          |

are no exposures and the road is some upon  
the upper plane that extends to Niagara Falls and  
Buffalo.

The Onondaga is therefore a land interval  
over which the Clinton and Salem Silurian seas trans-  
gressed.

The Lorraine is found in those borders of  
Hamilton, showing up by the lake. At Toronto the  
Lorraine is exposed. See Parks for a section.

At 10:30. Leaving Niagara Falls.  
Leaving the falls today. Will leave on the 7:30 P.M.  
for New York. Shipped a box by express.  
In the afternoon visited the grounds of

Buffalo, Lehigh Co. Took three pictures. The birds are nests rather nicely on the ground, but in regular direction of the light and eyes as deep as feet. Corniferous woods may be seen directly in the distance.

The walk returns to the hotel and takes a number of pictures for the morning. Several birds seen.

Left Buffalo at 7:30 A.M. Buffalo 8:35 and should have been at New York at 8 A.M. rest of the afternoon until 10. Then by the "Great Northern Railway, America". The Porter said this was the reason they were late. Several local trains stop by us.



Hanging Rock section Sep. 9-1908 At Trapacomo N. Va.

Upper Crinoid.

loc. 93

①

Heavy bedded whitish sandstone with  
masses of impure lime near top. In other rocks  
the layers become gray toward the  
top. Crinoidal zones with rounded  
trifurcate valves, small green chert  
occasionally up to 1/2 inch, rounded.

Lower Crinoid

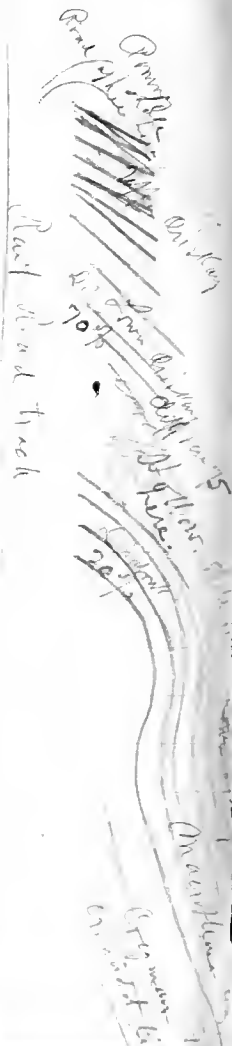
Distinctly bedded here but upper  
by the top into chert interbedded  
with the limestone. The lower part  
is a thin bedded limestone with fine zones  
of chert.

Lower Crinoid in one cliff near  
middle end of section about 1/2 mile high  
- 6' and 10' zones  
Plectambonites abundant, also  
Leptaecidia flabellata, most common  
near here

Acledonites, Macropleura

Small black chert near

at the base of the lower part.





A little farther south the red size higher  
as the Moerlein appears down to the Parson  
house junction as Bygonia cross of the D.S.B.  
section.

Shaly zone at south end of cut about  
32 yards across, dip at one side being vertical  
at the other about 50 degrees.

Lower Gristley. At south end of sect on the cliff  
about ~~285~~ 285 feet high above South  
Branch. Of this about 260 can be seen.

Here the separation of the chert from the clay  
from the upper Gristley is marked. Fully 70 feet  
of L.O. can be seen. This then would make  
the upper series a 185 feet thick and all of the  
Gristley about 260 feet thick. This can be  
seen & interpreted as the clay bed being  
above the Moerlein red.

At the south end of section just south of  
the Parson house (hotel) may be seen 20 feet  
more of the well bedded shaly upper Gristley  
so that all the Gristley here measures 280 feet.  
It may even be 285 feet.



As one goes to the top of the mountain back of the Parson house the coarse sandstone with bands of conglomerate are very noticeable. The pebbles cover the ground and some are as large as half an inch. (4)

The Cumberland Gulf continues the entire Diluvian time passing on the whole into the Manlius. A low topography prevails to the east for the deposits are essentially limestone, though somewhat impure. This is continued into Croghan's bed towards the base a higher topography begins to appear seen in the highest bed of the higher Croghan's. This becomes more decided in the Proterozoic bed and then the sea becomes very muddy and finally on land.

The land period in the western Cumberland Gulf prevails while in the east was deposited the Becraft and with the latter the water again becomes general throughout the Cumberland Gulf. Elevation of the land continuing to the east on the Piedmont slope and the Onondaga limestones more and more of s. lime and becomes sandy and fine a conglomeratic sandstone is deposited. Elevation continues and all becomes land.

At Cumberland on the 4<sup>th</sup> June I  
saw the grass and water appear as  
here. The conditions are good for  
the Cumberland Gulf.



