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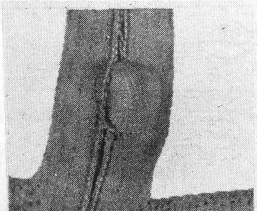
Hydatina vesicaria (Lightfoot, 1786)
Drawing by Scooter Beers

Notes on some species of the genus *Thyca*: (Eulimidae)

Donald R. Shasky

834 W. Highland Avenue, Redlands, CA 92373

Photo at right: *Thyca callista* Berry, 1958
Pta. Diablo, Baja California, Mexico
August 31, 1960



In recent months papers by Bertsch, Bratcher and DuShane have discussed *Thyca* (*Bessomia*) *callista* Berry, 1959. The purpose of this paper is to present some historical information about this species as well as additional geographic and taxonomic data. I also call attention to Warren's work on *Thyca* which has been overlooked by previous authors.

I had just finished medical school (June, 1958) and had three weeks before my intern year began. In those days, if I had free time, I was in Guaymas, Mexico, and this period was no exception.

During that trip I was collecting shells, starfish and sea urchins. A single starfish that I found, *Phataria unifasciatus* (Gray), had a small shell attached that I had never seen before.

A few days after finding this shell a stranger approached as I was working on my shells. I was between dives and was happy for someone who wanted to talk about shells. This was my first encounter with the late Dr. S. Stillman Berry.

He was returning to Redlands after collecting in Mazatlan. He inquired about what I had found, and among other things, I showed him my parasitized starfish. He was elated as he immediately recognized this as a new species of *Thyca*.

I told him that I knew someone who had previously collected great quantities of starfish, for commercial purposes, in Guaymas. I thought that perhaps he might have some specimens that had *Thyca* on them.

After returning to my home in Glendale, California, I called Leonard Bessom and, yes, he thought he did have starfish with shells attached. There was one problem — he had spray-painted all his starfish with blue paint! He carefully removed the paint and subsequently sent three specimens of the shell to Dr. Berry which later became the holotype and the two paratypes. Because my specimen came from across the bay from where Bessom collected, Dr. Berry did not include mine as a paratype when he described *Thyca* (*Bessomia*) *calista* the following year.

In 1961 I published (Veliger), that Bessom and I estimated that we found one *Thyca callista* parasitized *Phataria unifasciatus* for each 1,000 - 1,500 that we examined in the Guaymas area. That paper also included the first photograph of *T. callista*.

In 1976 while diving in one of the many bays around Isla Espiritu Santo and Isla Partida in the Gulf of California, I found that the ratio was one parasitized *P. unifasciatus* for each three I examined. These were taken snorkeling in from one to three meters.

In Isla Maria Madre in the Tres Marias Islands group I also found a fairly high density of *T. callista* on another starfish *Pharia pyramidata* (Gray).

In discussing *T. callista* in the February 14, 1985, Festivus, Dr. Bertsch failed to cite my recent paper on the shells of Manabi Province, Ecuador, where I reported having taken *T. callista* at Isla La Plata. In 1980, Warren transferred the genus *Thyca* from the family Capulidae to the family Eulimidae.

Incidentally, in my paper on the mollusks of Manabi Province, Ecuador, *Thyca callista* was inadvertently listed under both the Eulimidae and the Capulidae. I regret this oversight when I proof-read the manuscript.

In his paper, Warren assigned the following Indo-Pacific species of *Thyca* to Berry's subgenus *Bessomia*: *T. crystallina* Gould, 1846; *T. ectoconcha* Sarsin, 1887; and *T. hawaiiensis* Warren, 1980. I would like to add two additional bits of information concerning *Thyca* species.

Warren gives the range for *Thyca* (*Thyca*) *stellasteris* Koeler and Vancy, 1912, from the Gulf of Suez to West Australia. While collecting aboard a shrimp trawler between Dalrymple and Rennell Islands in the Torres Strait, N. Queensland, I took one specimen of *T. stellasteris* from an unidentified starfish. Thus the range is extended to northeastern Australia.

Lastly, I have two specimens of an undescribed *Thyca* that I collected in sightings at Isla La Plata, Ecuador. It is in manuscript and is the third known species of *Thyca* ss.

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LIVING CLOCKSPPRINGS, part 4.

Chris Illert, 2 Tern Place, Semaphore Park, South Australia 5019

Most conchologists are pleased with themselves if they become knowledgeable about living molluscan species. Yet the vast majority of all shell forms are extinct; existing only in the fossil record. All of the early "tetrabranch" nautiloids, and the entire family of ammonoids - numbering tens of thousands of species - are gone. But, before these shelled cephalopods became extinct, the ammonoids of the Late Cretaceous era (in particular) coiled in such unconstrained wild shapes (with looping bodywhorls) that palaeontologists had given up and considered these "heteromorphs" to be a lost cause.

This is where elastic clocksprings play an important role not only in classification of "heteromorphous" geometries, but also they serve the purpose of uniting disparate geometrical shapes into a coherent and sensible whole body of evidence. Ammonite spirals can be visualised in terms of a single elastic spiral spring being wound about a spindle whilst the outermost end of the coil is held fixed at an anchor point - as in Figure 5. The individual static shell spirals can now be seen in terms of a continuous dynamic winding process. Individual shell shapes relate to the dynamic process of spring winding in much the same way that single snapshots relate to a continuous movie picture.

Previously there was no known upper limit to what shell shapes were geometrically possible in nature. How could conchologists study variations on a theme if the theme was itself unknown? Simple clockspring curves, first suggested by T.A. Cook (1908), supply the necessary geometrical theme enabling us to study not only those fossil shell shapes that exist but also those which for some reason or another never occurred.

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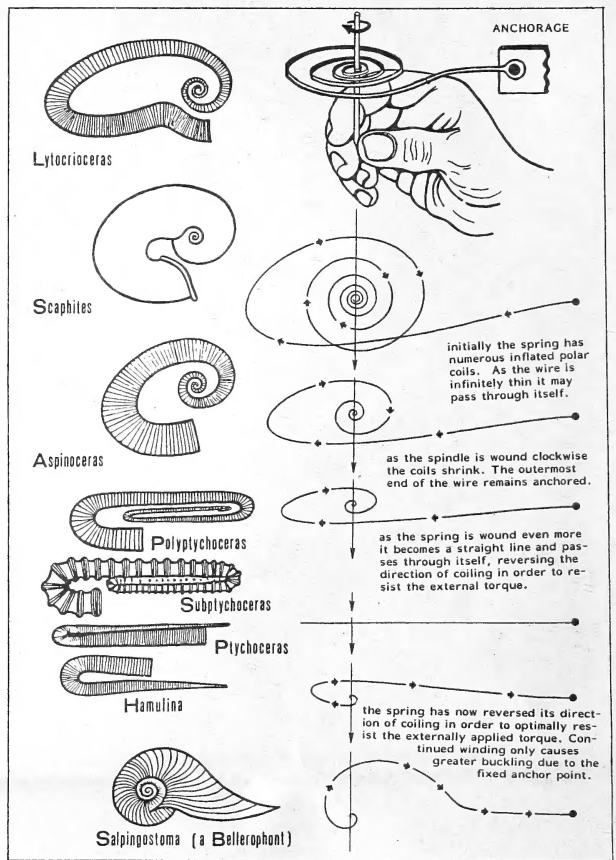
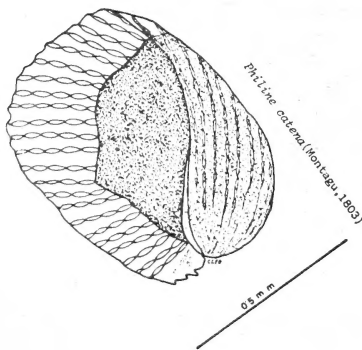


Figure 5. If we take an elastic spring with an anchored outer end, and wind it continuously (clockwise) then it goes through a variety of shapes - all of which were discovered and widely used in ages past by Ammonites.



Philine catena (Montagu, 1803)
 Drawing by Casto L. Fernández-Ovies

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Bivalve Molluscan paleoecology of northern exposures of the marine Neogene Imperial Formation in Riverside County, California

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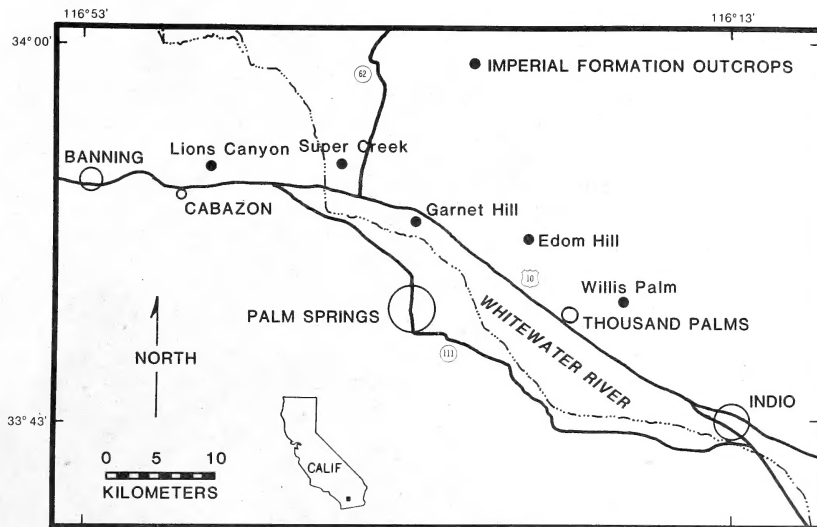


Figure 1. Index map showing localities of the Imperial Formation in the study area.

INTRODUCTION

The marine Imperial Formation was deposited at the head of the proto-Gulf of California during the latest Miocene through late (?) Pliocene time. The formation crops out discontinuously northward from just south of the California-Mexico border to San Geronio Pass, Riverside County, California. Exposures of the Imperial Formation in Riverside County occur along the San Andreas-Banning fault zone. Important exposures occur at Lions Canyon, Super Creek, Garnet Hill, Edom Hill and Willis Palm (fig. 1). The faunas from Super Creek, Garnet Hill and Willis Palm each represent different environments and form the basis for this note. The Imperial Formation in Riverside County attains a maximum thickness of 105 m and is divided into two members separated by a time transgressive boundary from north to south.

The lower, Latrania Sand Member (Keen and Benton, 1944), whose type section is in Imperial County, is generally composed of coarse- to medium-grained sandstone with scattered conglomeratic beds; it attains a maximum thickness of 30 m in Riverside County. The molluscan assemblage from the Latrania Sand Member in Riverside County is characterized by species that lived in a euhaline, moderate- to high-energy, intertidal (?) to inner shelf, rocky to sandy habitat. Taxa that commonly occur in Riverside County include *Spondylus victoriae* (Sowerby), *Pycnodonte heermanni* (Conrad) [= "*Ostrea*" *heermanni* Conrad], *Dendostrea angelica* (Rocheburne), and *Argopecten* spp.

The unnamed, upper member of the Imperial Formation is about 75 m thick in Riverside County. In this area it represents a low- to moderate-energy, outer shelf facies and is composed of medium- to fine-grained sandstone and siltstone. Characteristic bivalve mollusks in this member in Riverside County include *Cyclopecten* sp. cf. *C. pernomus* (Hertlein), *Dendrostrea? vespertina* (Conrad), and *Anomia peruviana* d'Orbigny.

PALEOECOLOGY Super Creek area

Both members of the Imperial Formation are exposed along Super Creek. The Latrania Sand Member occurs in small discontinuous synclines (fig. 2), and appears to have been deposited in small pocket beaches in rocky headlands. The outcrops are characterized by abundant *Spondylus victoriae* and *Dendostrea angelica* that are found attached to large clasts in conglomeratic stringers. Bernard (1983) reports *D. angelica* in water depths of 1 to 5 m and *S. victoriae* from 10 to 40 m. Thus the Latrania Sand Member along Super Creek was probably deposited in shallow subtidal depths that deepened upsection to between 10 and 40 m at the top of the member. These beds are overlain by a 70-cm thick "worm tube" marl at the top of the Latrania Sand Member. The environment of deposition of the marl was not determined because of the poor preservation of its fossils.

The Latrania Sand Member is overlain by medium- to fine-grained sandstone and siltstone of the unnamed upper member. About 10 m above the top of the Latrania Sand is a bed of *Atrina* n. sp. preserved in life position. *Atrina* is found in water depths from 1 to 125 m in the eastern Pacific, but only *A. texta* Hertlein & Strong, is reported to occur below 30 m. Benthic foraminifers from this bed suggest inner neritic water depths (0-50 m) (K. A. McDougall, written commun., 1981). Together these data suggest water depths of less than 50 m for the lower part of the unnamed upper member of the Imperial Formation. Near the top of the upper member, *Cyclopecten* sp. cf. *C. pernomus* and many shell fragments have been concentrated into small shell lag deposits, possibly by currents. Water depth during deposition of this unit is unclear, but the occurrence of *C. sp. cf. C. pernomus* limits the depth to 355 m (Bernard, 1983). Benthic foraminifers throughout the upper member indicate that water depths increase to middle and outer neritic depths (50-150 m) near the top of the formation (K. A. McDougall, written commun., 1981).

In the Super Creek area, the Imperial Formation is overlain by medium- to coarse-grained sands and conglomeratic beds of the Painted Hill Formation, previously interpreted as alluvial fan deposits by Allen (1957). The interfingering of the Painted Hill Formation with neritic silts of the Imperial Formation indicates that the base of the Painted Hill Formation probably is not of alluvial fan origin.

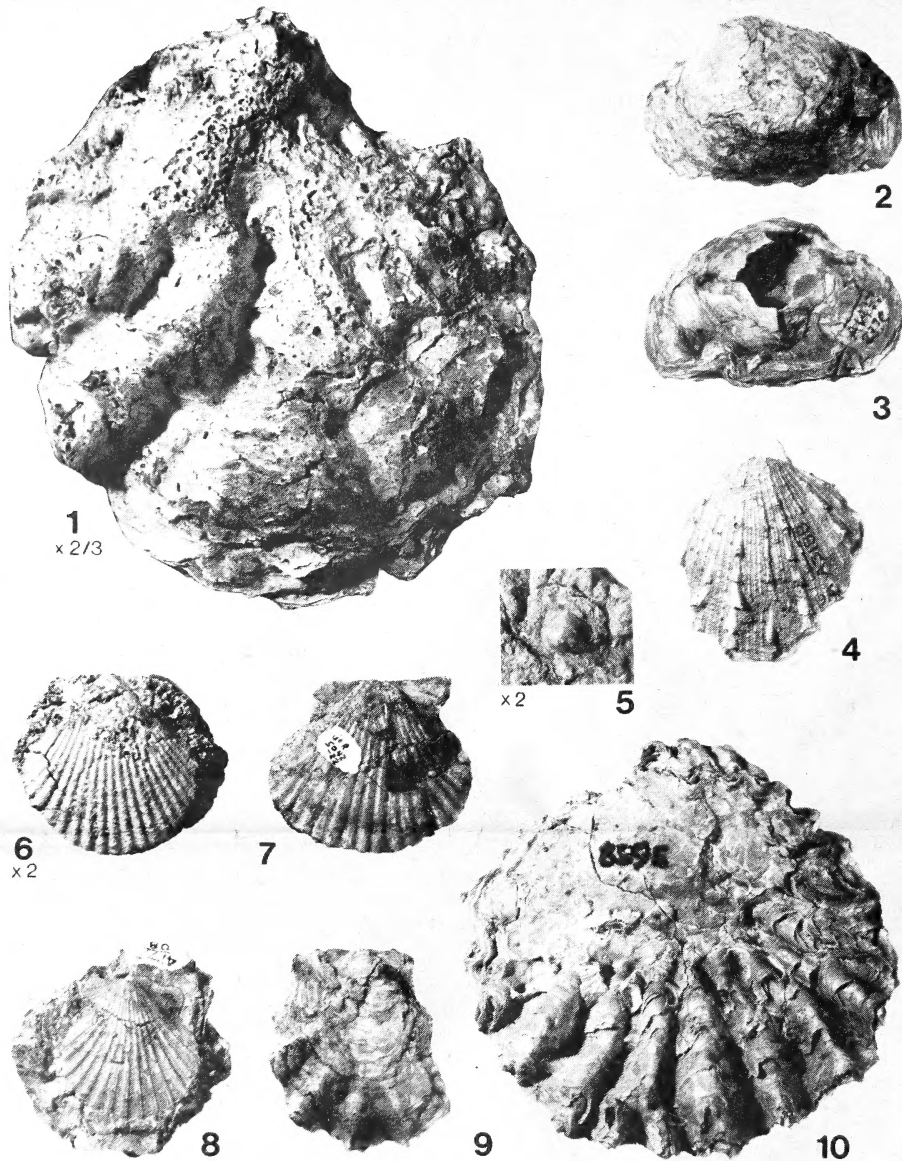


PLATE 1

(All figures x 1 except as indicated)

- Figure 1. *Pycnodonte (Pycnodonte?) heermanni* (Conrad 1855). Left valve. LACMIP 7162; length 131 mm, height 148 mm. CIT loc. 192
- Figures 2 and 3. *Anomia peruviana* d'Orbigny, 1846. Articulated valves. UCMP 37531; length 49.9 mm, height 35.1 mm. UCMP loc. A-1417.
- Figure 4. *Spondylus victoriae* Sowerby, 1859. Right valve. UCMP 37532; length 35.4 mm, height 39.6 mm. UCMP locality A-3188.
- Figure 5. *Cyclopecten* sp. cf. *C. pernomus* (Hertlein, 1935). Right valve. UCMP 37533; length 4.4 mm, height 4.8 mm. UCMP loc A-1407.

- Figure 6. *Argopecten* sp. cf. *A. mendenhalli* (Arnold, 1906). Left valve?. UCR 5035/11; length 17.1 mm, height 18.1 mm. UCR loc 5035
- Figure 7. *Argopecten sverdrupi* (Durham, 1950). Right valve. UCR 5042/23; length 39.1 mm, height 34.6 mm. UCR loc 5042.
- Figure 8. *Leptopecten (Leptopecten) palmeri* (Dall, 1897). Left valve. UCR 4777/11; length 26.3 mm, height 27.6 mm. UCR loc 4777.
- Figure 9. *Dendostrea? vespertina* (Conrad, 1854). Right valve. LACMIP 7163; length 31.1 mm, height 38.1 mm. CIT loc 187.
- Figure 10. *Dendostrea angelica* (Rocheburne, 1895). Left valve. LACMIP 7164; length 79.5 mm, height 80.5 mm. UCR loc 8595.

PASTELS



HB 602
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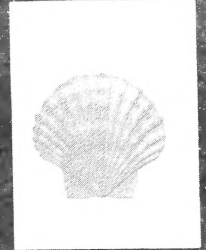
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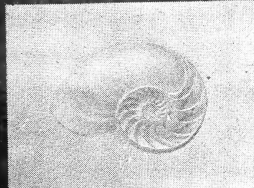
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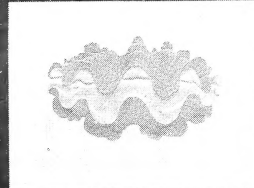
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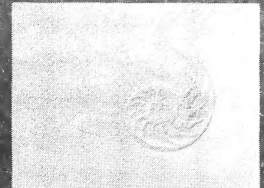
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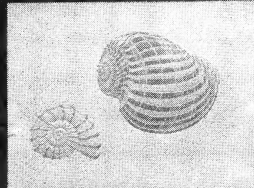
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P 202 (Peach)



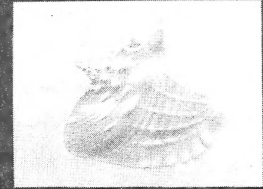
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P 218 (Peach)



P 150 (Peach)

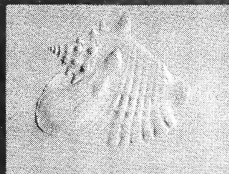
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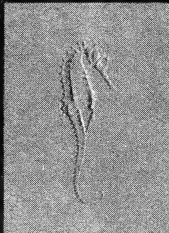


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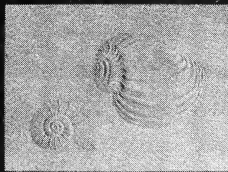


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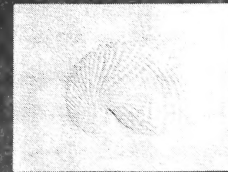
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BL 636



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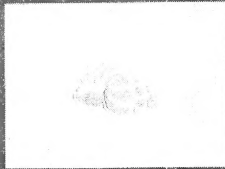


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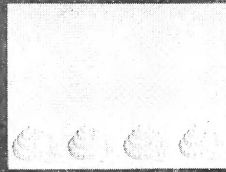


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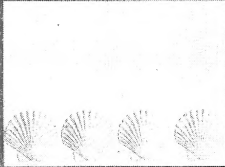


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PP 626 (Pink)

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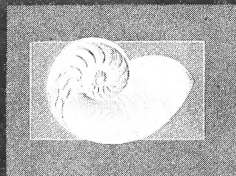
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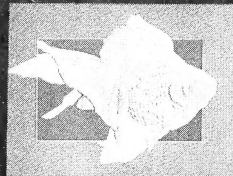
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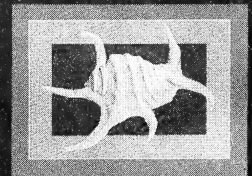
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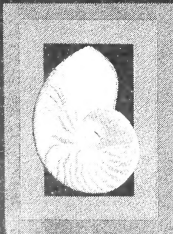


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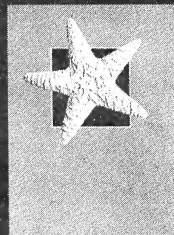


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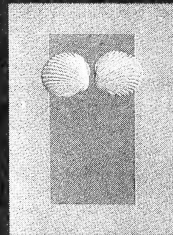
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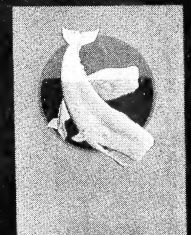
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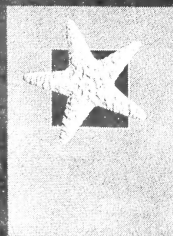
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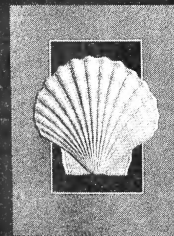
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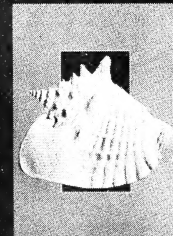
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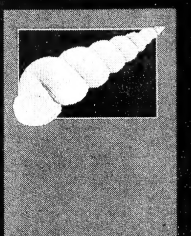
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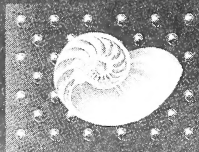
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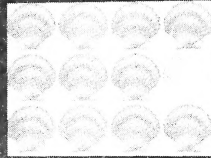
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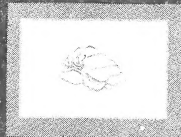
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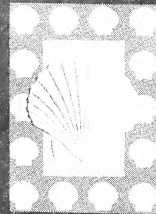
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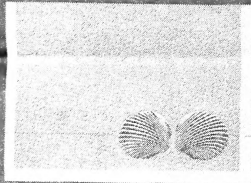
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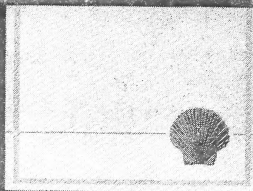
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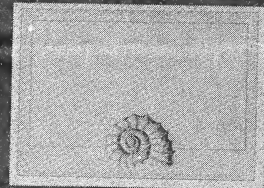
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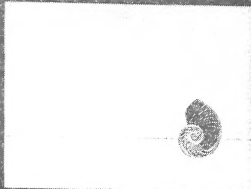
PR 331



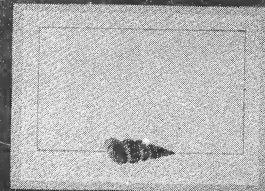
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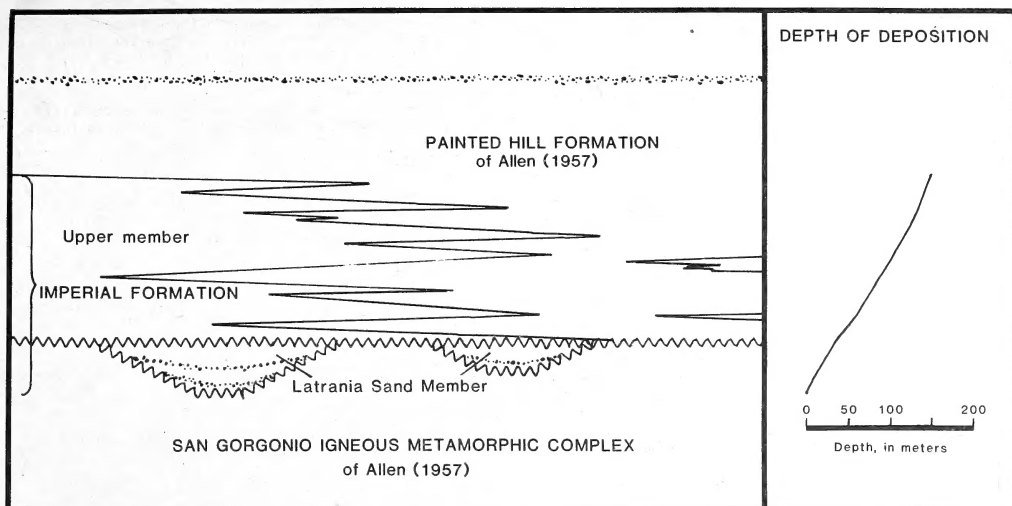


Figure 2. Stratigraphic relationships of the unnamed upper member and the Latrania Sand Member of the Imperial Formation, and their inferred depths of deposition along Super Creek, Riverside County, California.

Garnet Hill Area

About 17 m of the Latrania Sand Member are poorly exposed on Garnet Hill. These beds are composed of coarse- to fine-grained sandstone with abundant gneissic and marble clasts as much as 2 m in diameter. Some of these clasts have been bored by lithophagid bivalves (identified by G. L. Kennedy, oral commun., 1984). Also present on some of the clasts are poorly preserved specimens of *Spondylus* sp. and *Dendostrea* sp. cf. *D. angelica*. *Pycnodonte heermanni* is abundant as float on the south slope of Garnet Hill. These species, although inconclusive as to their environmental parameters, suggest shallow water depths with at least some areas of exposed rocks.

Willis Palm Area

The Imperial Formation at Willis Palm consists of about 53 m of generally fine-grained sandstone and siltstone. Three shell beds are exposed in this section. The first contains unidentifiable mollusk fragments, crab chelipeds, and the gastropod *Melampus* sp., which suggests a salt marsh environment. About 2 m above this bed is a 2-m-thick oyster bioherm composed of flat to slightly plicate *Dendostrea? vespertina*. This bioherm suggests a moderately shallow bay or lagoonal environment. The third shell bed, about 15 m above the base of the formation, is 20 cm thick and composed of abundant well preserved shells. This bed is characterized by the bivalve mollusks *Leptopecten palmeri* (Dall), *Cyclopecten* sp. cf. *C. pernomus*, *Cyrtopleura costata* (Linnaeus), and abundant corbulids. These species suggest shallow water depths between 1 and 90 m. Benthic foraminifers from this section indicate shallow water (K. A. McDougall, written commun., 1984). Fossils from throughout the Willis Palm section indicate that water depths increased from very shallow inner neritic (< 10 m) at the first shell bed to a maximum depth of inner neritic (10 - 50 m) just above the top shell bed, and then shallowed to less than 10 m at the top of the formation.

CONCLUSIONS

Other exposures of the Imperial Formation are present in Riverside County but were either inaccessible during the present study (Lions Canyon) or contained such a meager fauna that paleoecological analysis was impossible (Edom Hill).

The sections at Super Creek and Garnet Hill appear to be older than those exposed at Willis Palm. These relative-age estimates are based on biostratigraphic data and on an unpublished K/Ar date on a basalt flow in the Painted Hill Formation along Super Creek (J. C. Matti, personal commun., 1984). The Super Creek section seems to represent a deepening basin: the intertidal or shallow subtidal rocky shore facies of the Latrania Sand Member at the base of the formation grades into middle and outer neritic, fine-grained, continental-shelf facies of the upper member at the top of the formation. Sediments exposed at Garnet Hill appear to represent the same environment as represented by the Latrania Sand Member at Super Creek. The Willis Palm section represents a transgressive-regressive sequence starting with a salt marsh environment, which grades upsection into a shallow lagoon or bay, characterized by an oyster bioherm. This is followed further upsection by 2 slightly deeper, possibly open marine environment, which again changed to shallow or possibly intertidal depths at the top of the formation.

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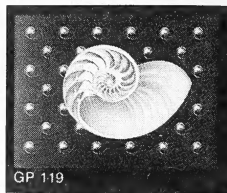
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